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# International geomagnetic reference field: the fourteenth generation

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**Abstract** In November 2024, the International Association of Geomagnetism and Aeronomy (IAGA) Division V Working Group (V-MOD) adopted the fourteenth generation of the International Geomagnetic Reference Field (IGRF). This IGRF updates the previous generation with a definitive main field model for epoch 2020.0, a main field model for epoch 2025.0, and a predictive linear secular variation for 2025.0 to 2030.0. We describe the equations defining the IGRF, the spherical harmonic coefficients for this fourteenth generation model, maps of magnetic declination, inclination, and total field intensity for the model at 2025.0 and maps of the predicted rate of change of the field for the 2025.0 to 2030.0 time period.

**Keywords** IGRF, Magnetic field modeling, Geomagnetism

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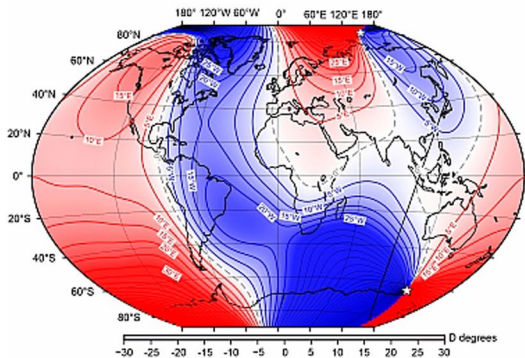
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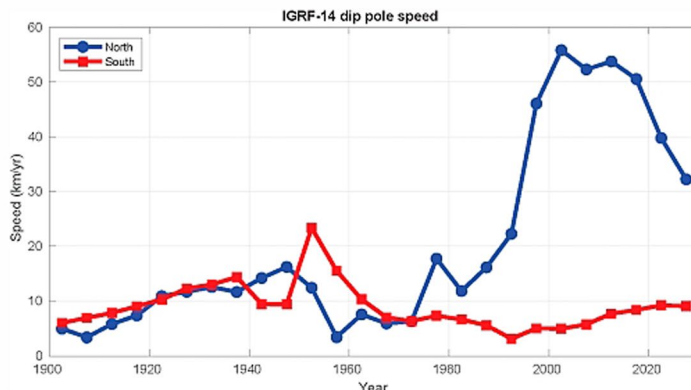
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**Graphic Abstract**

**IGRF-14 Declination for 2025.0**



**IGRF-14 Dip pole velocity 1900-2030**



**1 Introduction**

The International Geomagnetic Reference Field (IGRF) is a set of spherical harmonic coefficients which describe the large-scale, time-varying portion of Earth’s internal magnetic field between epochs 1900.0 A.D. and 2030.0. The IGRF is produced and maintained by an international task force of scientists under the auspices of the International Association of Geomagnetism and Aeronomy (IAGA) Working Group V-MOD (ROR: <https://ror.org/013ym9476>). This fourteenth generation IGRF has been derived from observations recorded by satellites, ground observatories, and magnetic surveys.

The IGRF acts as a geophysical standard reference model and is routinely used by the scientific community, for example, to study Earth’s large-scale time-varying field and as a baseline to investigate local magnetic anomalies in the lithosphere. It is widely used in satellite attitude determination and control systems and other applications requiring orientation information. In the past two decades, the IGRF and its derivatives have been incorporated into space weather modeling and forecasting services.

Earth’s core field changes continuously and unpredictably on timescales ranging from months to millions of years due to the chaotic nature of its generation mechanism (e.g., Aubert 2023; Hulot et al. 2010; Rappaldini et al. 2021). As secular variation can be up to ±150 nT/year in some locations, ignoring the change could incur an error of 1500 nT after a decade. In order to account for temporal changes on timescales of a few years, the IGRF is regularly revised, typically every five years to limit this error to below the sensitivity level of most users.

The latest updates to the IGRF were created from data captured at both ground observatories and from dedicated magnetic satellite missions. The temporal variation is extremely well measured by the fixed ground network while satellite data are excellent for spatial coverage globally. With the extensive availability of satellite data from the past 25 years, remote locations such as the oceans and polar regions are well characterized; though the model would certainly be improved by the addition more observatories over the oceans in the southern Pacific for example.

The IGRF models are continuous in space and have a resolution of  $2\pi R_E / \sqrt{n(n+1)}$ , where  $R_E$  is the Earth’s radius and  $n$  is the maximum degree and order. As a spherical harmonic degree and order 13 model, the IGRF-14 has an approximate spatial resolution of  $40,000 \text{ km} / \sqrt{13 \cdot 14} = 2965 \text{ km}$ , i.e. around 3000 km at Earth’s surface.

The modeling of the core field uses modern versions of Gauss’ method for determining the internal (core and crust) part of the field from the external (ionosphere and magnetosphere) contributions. Satellite and ground observatory data captured at different altitudes are combined in this manner, reducing them to a common level at the Earth’s reference radius ( $R_E = 6371.2 \text{ km}$ ). The technical papers from several previous generations such as Alken et al. (2021) and references therein explain the detail. The IGRF will provide valid estimates of the field from the core mantle boundary at a nominal radius of 3485 km to around 10,000 km altitude where the values of the field become small. Beggan (2022) examined the fit of the model to repeat station measurements from around the world to determine the approximate

uncertainty limits. The nominal resolution of an individual value from the model is given as 0.1 nT, though the uncertainties depend on the component of the magnetic field being evaluated.

Table 1 summarizes the current and past generations of IGRF. Each generation is composed of a set of model coefficients representing the internal time-varying geomagnetic field, which are provided in 5-year intervals. The years for which coefficients are provided are called *model epochs*. The coefficients of a certain epoch represent a snapshot of the geomagnetic field at that time, and can be labeled either as a Definitive Geomagnetic Reference Model (DGRF) or as an IGRF. DGRF models are deemed to be definitive because they have been built from the highest quality ground observatory and satellite data available from that time period and therefore are unlikely to be improved in future IGRF revisions. Models labeled as IGRF are non-definitive, and will likely be revised in the future as more data are collected and made available. DGRF models require sufficient data to compute a representative field and have only been built for the period starting in 1945. Prior to this insufficient data were available to make a high quality model so these remain labeled as IGRF. Details of the history of IGRF can be found in Barton (1997) and Macmillan and Finlay (2011). Past generations of IGRF models are archived at <https://www.ncei.noaa.gov/products/international-geomagnetic-reference-field>. Since later IGRFs can revise model parameters for past epochs, it is important to record which generation of IGRF was used to process a particular dataset, so that the original values can be recovered and reprocessed with the latest generation of IGRF, if needed.

In this letter, we describe the fourteenth generation of IGRF, known hereafter as IGRF-14. IGRF-14 provides a DGRF model for epoch 2020.0, an IGRF model for epoch 2025.0, and a predictive IGRF secular variation model for the 5-year time interval 2025.0–2030.0. For epochs 1900.0–2015.0, the IGRF-14 model coefficients are unchanged from IGRF-13. IGRF-14 was finalized in November 2024 by a task force of IAGA Working Group V-MOD. In the following sections, we give an overview of the IGRF model, provide the final set of IGRF-14 coefficients, and briefly discuss large-scale features of the geomagnetic field at Earth’s surface as observed in the updated model.

## 2 Mathematical formulation of the IGRF model

The IGRF describes the main geomagnetic field  $\mathbf{B}(r, \theta, \phi, t)$  which is produced by internal sources primarily inside Earth’s outer core. The IGRF is valid on and above Earth’s surface, where the main geomagnetic field can be described as the gradient of a scalar potential,  $\mathbf{B} = -\nabla V$ , and the potential function  $V(r, \theta, \phi, t)$  is represented as a finite series expansion in terms of spherical harmonic coefficients,  $g_n^m, h_n^m$ , also known as the Gauss coefficients:

$$V(r, \theta, \phi, t) = a \sum_{n=1}^N \sum_{m=0}^n \left(\frac{a}{r}\right)^{n+1} [g_n^m(t) \cos m\phi + h_n^m(t) \sin m\phi] P_n^m(\cos \theta). \tag{1}$$

Here,  $r, \theta, \phi$  refer to coordinates in a geocentric spherical coordinate system, with  $r$  being radial distance from the center of the Earth, and  $\theta, \phi$  representing geocentric co-latitude and longitude, respectively. A reference radius

**Table 1** Summary of IGRF generations, validity periods, release years, and references

| Full name            | Short name | Validity period | Definitive period | Release year | Reference                                       |
|----------------------|------------|-----------------|-------------------|--------------|---|
| IGRF 14th generation | IGRF-14    | 1900.0–2030.0   | 1945.0–2020.0     | 2024         | This paper                                      |
| IGRF 13th generation | IGRF-13    | 1900.0–2025.0   | 1945.0–2015.0     | 2019         | Alken et al. (2021)                             |
| IGRF 12th generation | IGRF-12    | 1900.0–2020.0   | 1945.0–2010.0     | 2014         | Thébault et al. (2015)                          |
| IGRF 11th generation | IGRF-11    | 1900.0–2015.0   | 1945.0–2005.0     | 2009         | Finlay et al. (2010a)                           |
| IGRF 10th generation | IGRF-10    | 1900.0–2010.0   | 1945.0–2000.0     | 2004         | Maus et al. (2005); Macmillan and Maus (2005)   |
| IGRF 9th generation  | IGRF-9     | 1900.0–2005.0   | 1945.0–2000.0     | 2003         | Macmillan et al. (2003)                         |
| IGRF 8th generation  | IGRF-8     | 1900.0–2005.0   | 1945.0–1990.0     | 1999         | Mandea and Macmillan (2000)                     |
| IGRF 7th generation  | IGRF-7     | 1900.0–2000.0   | 1945.0–1990.0     | 1995         | Barton (1997)                                   |
| IGRF 6th generation  | IGRF-6     | 1945.0–1995.0   | 1945.0–1985.0     | 1991         | Langel (1992)                                   |
| IGRF 5th generation  | IGRF-5     | 1945.0–1990.0   | 1945.0–1980.0     | 1987         | Barraclough et al. (1987); Langel et al. (1988) |
| IGRF 4th generation  | IGRF-4     | 1945.0–1990.0   | 1965.0–1980.0     | 1985         | Barraclough (1987)                              |
| IGRF 3rd generation  | IGRF-3     | 1965.0–1985.0   | 1965.0–1975.0     | 1981         | Peddie (1982)                                   |
| IGRF 2nd generation  | IGRF-2     | 1955.0–1980.0   | –                 | 1975         | IAGA Division I Study Group (1975)              |
| IGRF 1st generation  | IGRF-1     | 1955.0–1975.0   | –                 | 1968         | Cain and Cain (1971); Zmuda (1971a, 1971b)      |

$a = 6371.2$  km is chosen to approximate the mean Earth radius. Temporal changes are denoted by  $t$ . The  $P_n^m(\cos \theta)$  are Schmidt semi-normalized associated Legendre functions of degree  $n$  and order  $m$  (e.g., Winch et al. 2005). The parameter  $N$  specifies the maximum spherical harmonic degree of expansion, and was chosen to be 10 up to and including epoch 1995, after which it increases to 13 to account for the smaller scale internal signals which can be captured by high-resolution satellite missions such as Ørsted, CHAMP and Swarm. The Gauss coefficients  $g_n^m(t)$ ,  $h_n^m(t)$  change in time and are provided in units of nanoTesla (nT) in IGRF-14 at 5-year epoch intervals. The time dependence of these parameters is modeled as piecewise linear, and is given by:

$$g_n^m(t) = g_n^m(T_t) + (t - T_t)\dot{g}_n^m(T_t), \quad (2)$$

$$h_n^m(t) = h_n^m(T_t) + (t - T_t)\dot{h}_n^m(T_t), \quad (3)$$

where  $g_n^m(T_t)$ ,  $h_n^m(T_t)$  are the Gauss coefficients at epoch  $T_t$  (in years), which immediately precedes time  $t$  (also in years). The model epochs in IGRF-14 are provided in exact multiples of 5 years starting in 1900 and ending in 2025 (see Table 2), so that  $T_t \leq t < T_t + 5$ . For  $T_t < 2025$ , the parameters  $\dot{g}_n^m(T_t)$ ,  $\dot{h}_n^m(T_t)$  represent the linear approximation to the change in the Gauss coefficients over the 5 year interval spanning  $[T_t, T_t + 5]$ . They may be computed in units of nanoTesla per year (nT/year) as:

$$\dot{g}_n^m(T_t) = \frac{1}{5}(g_n^m(T_t + 5) - g_n^m(T_t)), \quad (4)$$

$$\dot{h}_n^m(T_t) = \frac{1}{5}(h_n^m(T_t + 5) - h_n^m(T_t)). \quad (5)$$

The main field coefficients are not yet known for  $T_t = 2030$ , and so for the final 5 years of model validity (2025–2030 for IGRF-14), the coefficients  $\dot{g}_n^m(2025)$ ,  $\dot{h}_n^m(2025)$  are explicitly provided (see last column of Table 2) in units of nT/year. Details on the individual candidate secular variation forecasts and the procedure used to combine them into a final set of  $\dot{g}_n^m(2025)$ ,  $\dot{h}_n^m(2025)$  may be found in Beggan et al. (2026) and references therein.

### 3 The 14th generation IGRF

In July 2023, during an IAGA V-MOD Working Group meeting held in Berlin, Germany, a task force of volunteer geomagnetic modelers was assembled to oversee the call for IGRF-14 candidate models and their evaluation. In March 2024, the task force issued an international call for three candidates:

- a DGRF main field model for the epoch 2020.0,
- an IGRF main field model for the epoch 2025.0,
- an IGRF linear secular variation model for the time period 2025.0–2030.0.

Nineteen teams representing 27 international institutes responded to the call. The number of teams and institutions who participated in IGRF-14 exceeded that of any previous generation. The task force received 14 DGRF main field candidates for epoch 2020.0, 15 IGRF main field candidates for 2025.0, and 18 IGRF secular variation candidates for 2025.0–2030.0. Following recent IGRF conventions, the main field candidates for IGRF-14 describe the spatial variation of the field to a maximum spherical harmonic degree and order of 13, while the secular variation candidates extend to a maximum degree and order of 8. Each of the 19 teams was managed by a team leader from the *lead institution*, and many teams also included personnel from supporting institutions.

The 19 lead institutions for IGRF-14 are: (1) British Geological Survey (United Kingdom), (2) Technical University of Denmark (Denmark), (3) National Institute of Natural Hazards (China), (4) Universidad Complutense de Madrid (Spain), (5) University of Colorado Boulder (United States), (6) GFZ Helmholtz Centre for Geosciences (Germany), (7) Institut de physique du globe de Paris (France), (8) Institut des Sciences de la Terre (France), (9) Geophysical Center of the Russian Academy of Sciences (Russia), (10) Kyoto University (Japan), (11) University of Leeds (United Kingdom), (12) University of Edinburgh (United Kingdom), (13) NASA Goddard Space Flight Center (United States), (14) Université de Strasbourg (France), (15) Technical University of Berlin (Germany), (16) Macau Institute of Space Technology and Application (China), (17) Wuhan University (China), (18) Observatoire de Physique du Globe de Clermont Ferrand (France) and (19) University of Science and Technology Houari Boumediene (Algeria). Further details of their submissions are described in Beggan et al. (2026).

Some of the lead institutes listed above also acted as supporting institutions to other teams. The supporting institutes which are not listed above include: The Institute of Statistical Mathematics (Japan), Institute of Science Tokyo (Japan), Nantes Université (France), University of Tokyo (Japan), CEA-Leti (France), University of Grenoble (France), SHOM (France) and University of Liverpool (UK).

Data recorded by the Swarm satellite mission (Friis-Christensen et al. 2006) and the ground observatory network (see Table 4) played a crucial role in the development of many of the IGRF-14 candidate models. Swarm ASM-V experimental data from the CNES-funded

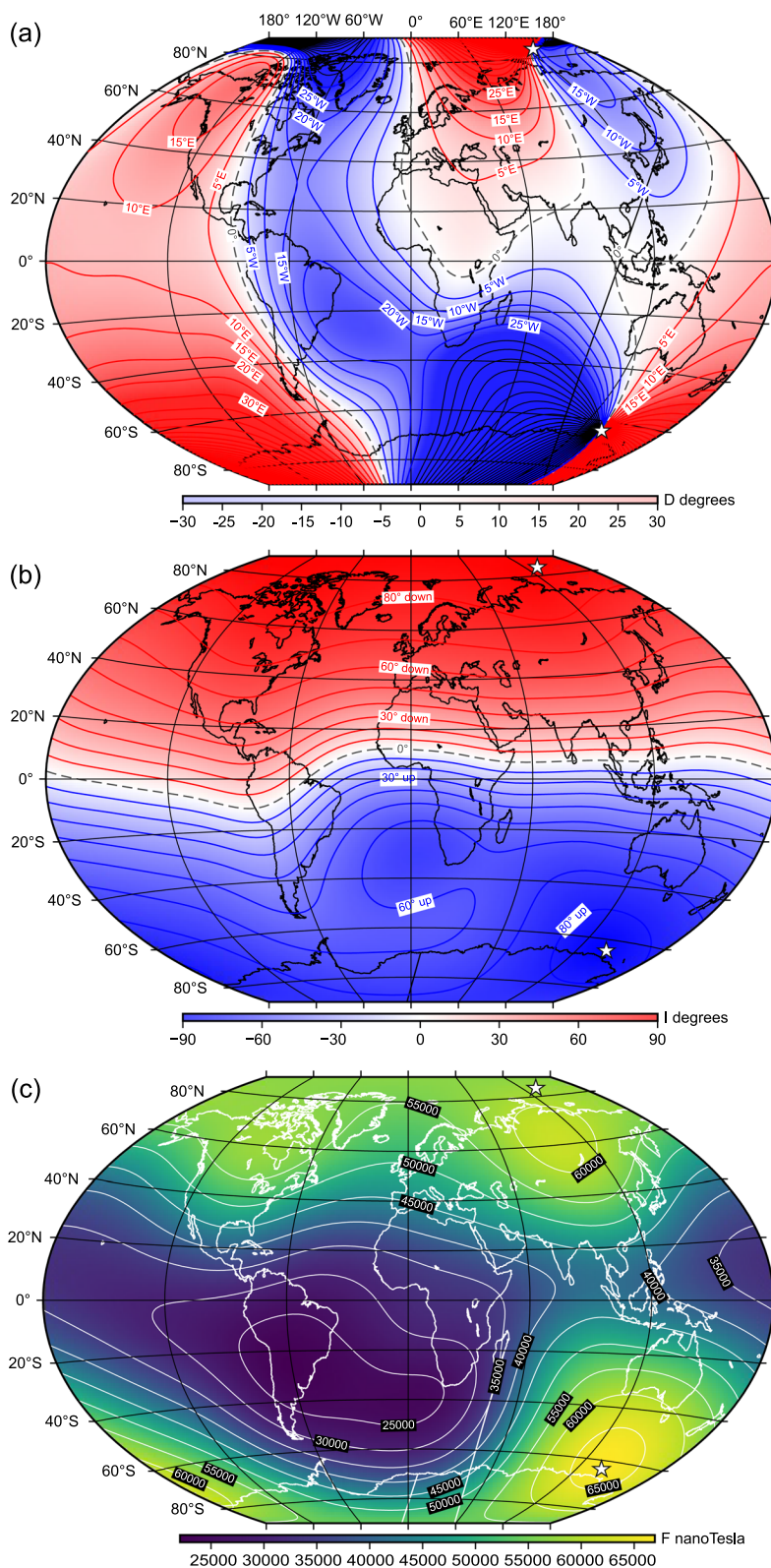




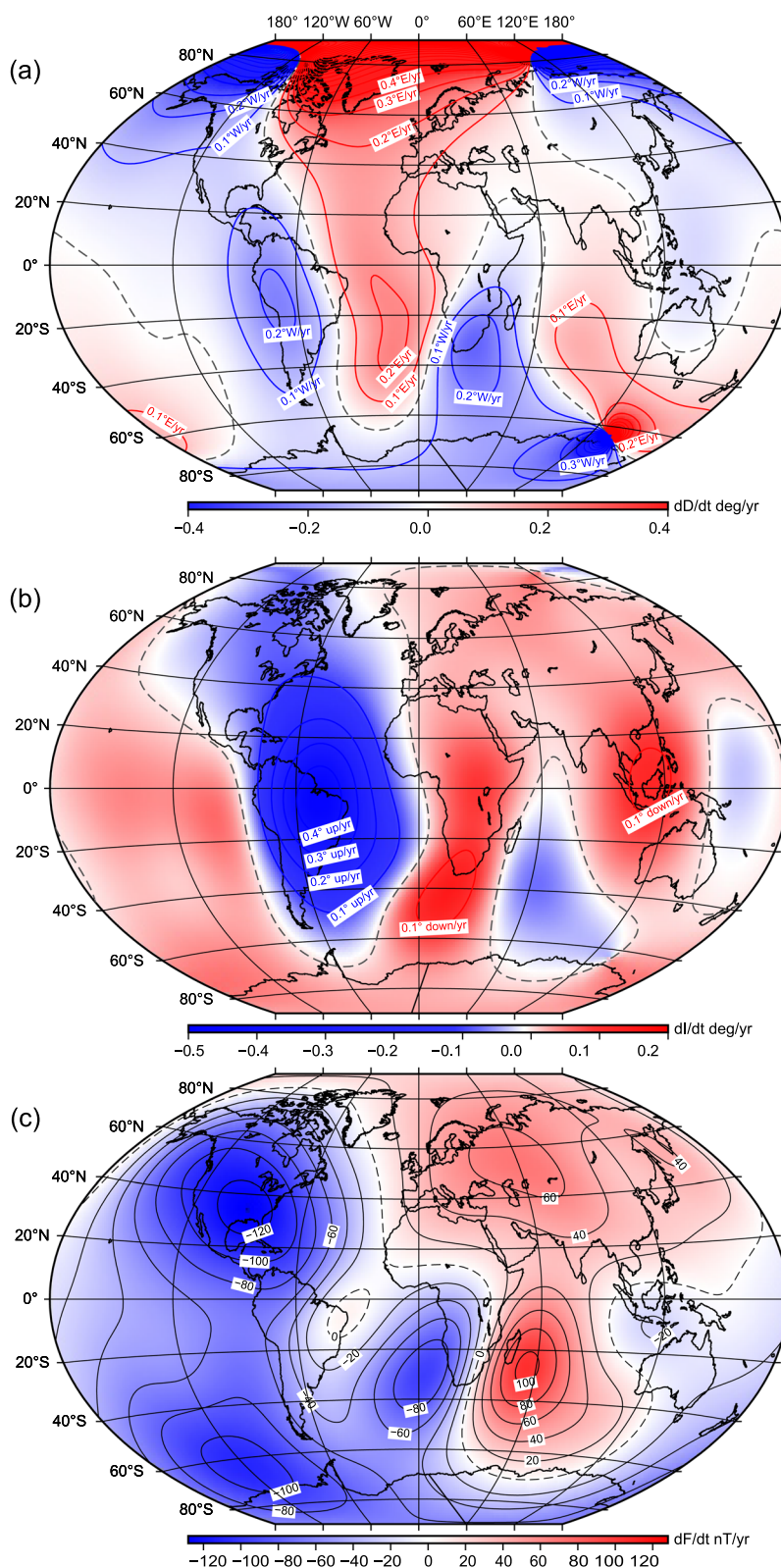








**Fig. 1** Maps of (a) declination and (b) inclination angle and (c) total field intensity at the WGS-84 ellipsoid surface for epoch 2025.0. For (a) and (b) the zero contour is shown as a dashed black line, positive contours in red, and negative contours in blue. White stars indicate the approximate locations of the north and south magnetic dip poles. Projection is Winkel Tripel



**Fig. 2** Maps of predicted annual secular variation in (a) declination, (b) inclination and (c) total field at the WGS-84 ellipsoid surface averaged over 2025–2030. The zero contour is shown as the dashed black line, positive contours in red, and negative contours in blue. Projection is Winkel Tripel

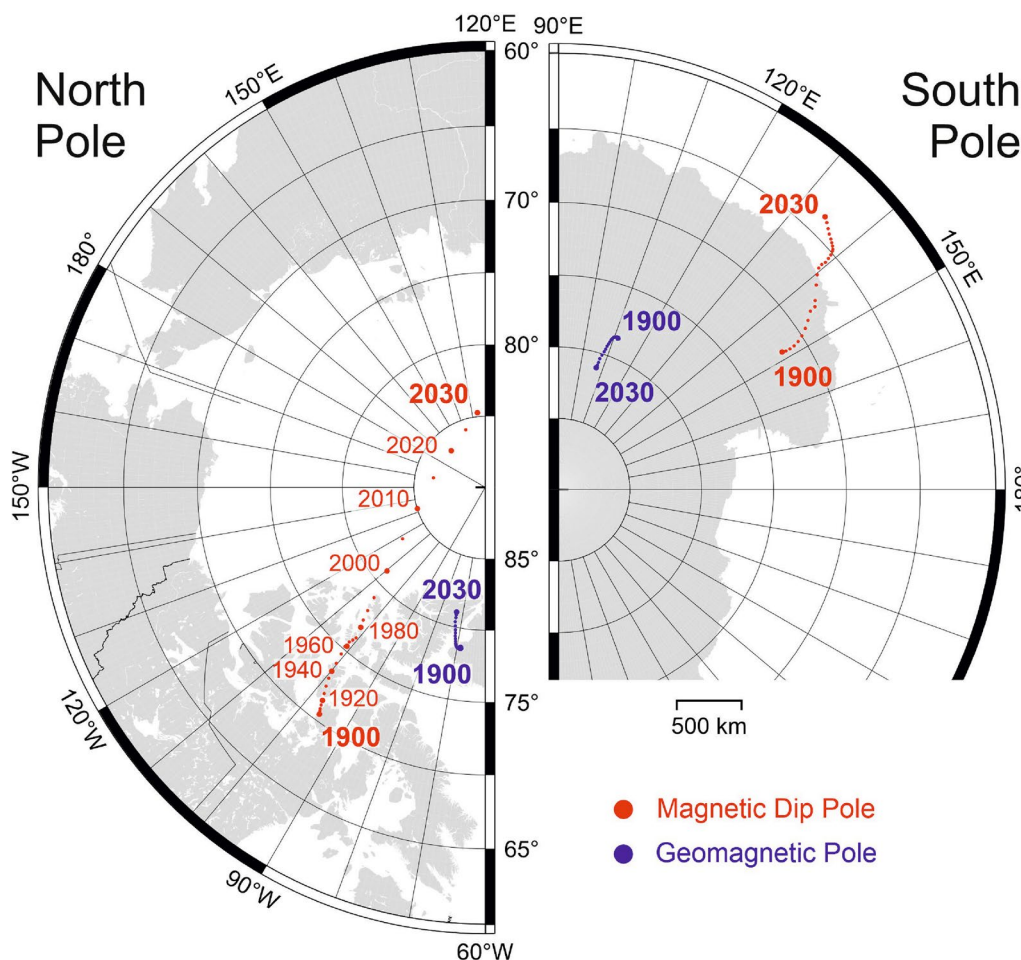
Absolute Scalar Magnetometer instrument (Hulot et al. 2015; Léger et al. 2015) as well as data from the Ørsted (Olsen 2007), CHAMP (Maus 2007), SAC-C (Colomb et al. 2004), Cryosat-2 (Olsen et al. 2020), and CSES (Shen et al. 2018) missions were also used by some of the teams. For the first time, data from the Macau Science Satellite 1 (MSS-1) (Zhang 2023; Jiang et al. 2024) were included in several IGRF candidates.

The IGRF-14 Evaluation Taskforce voted to calculate the final main field model for epoch 2020.0 as the median of the Gauss coefficients of all the candidate models. The task force voted to use a robust Huber weighting in space to determine the final coefficients for the 2025 main field epoch and for the secular variation model for 2025.0–2030.0. Further details of the candidate models, the evaluation process, and the final model determination are provided in Beggan et al. (2026).

#### 4 IGRF-14 model coefficients and maps

Table 2 lists the IGRF-14 spherical harmonic Gauss coefficients, which can be used with Eq. (1) to determine the geomagnetic potential (and vector geomagnetic field) anywhere on or above Earth’s surface. This table serves as a published record of IGRF-14, which should allow users to ensure they use the correct model coefficients for a particular epoch compared with previous generations. The main field coefficients are given in units of nT, and the predictive secular variation coefficients (last column) are given in units of nT/year. These coefficients are available in digital form from <https://www.ncei.noaa.gov/products/international-geomagnetic-reference-field> along with software to compute magnetic field components at different times and spatial locations, in both geocentric and geodetic coordinate systems.

A new innovation in IGRF-14 has been to establish an open access record of the candidates and final coefficient



**Fig. 3** Motion of the magnetic dip pole (red) and geomagnetic pole (blue) since 1900 from IGRF-14 in the northern hemisphere (left) and southern hemisphere (right). The scale provides an indication of distance on the WGS-84 ellipsoid that is correct along lines of constant longitude and also along the middle lines of latitude shown. Note the left and right panels use different longitude ranges. The maps use stereographic projection

files in a GitHub repository (Beggan et al. 2024) and to deposit the final coefficients in a Zenodo repository (International Association of Geomagnetism and Aeronomy (2024)). These make finding and reproducing the IGRF-14 coefficients compliant with FAIR standards (e.g., Ayala et al. 2022).

Using Equation 1 allows the North ( $X$ ), East ( $Y$ ) and Vertical ( $Z$ ) components of the magnetic field vector to be computed. These three orthogonal components can be converted to the angles more familiar to compass users of Declination ( $D$ ) and Inclination ( $I$ ), along with the Total Field intensity ( $F$ ). Taken together, these three quantities fully describe the vector magnetic field at Earth's surface:

$$D = \tan^{-1}(Y/X), \tag{6}$$

$$I = \tan^{-1} (Z/\sqrt{X^2 + Y^2}), \tag{7}$$

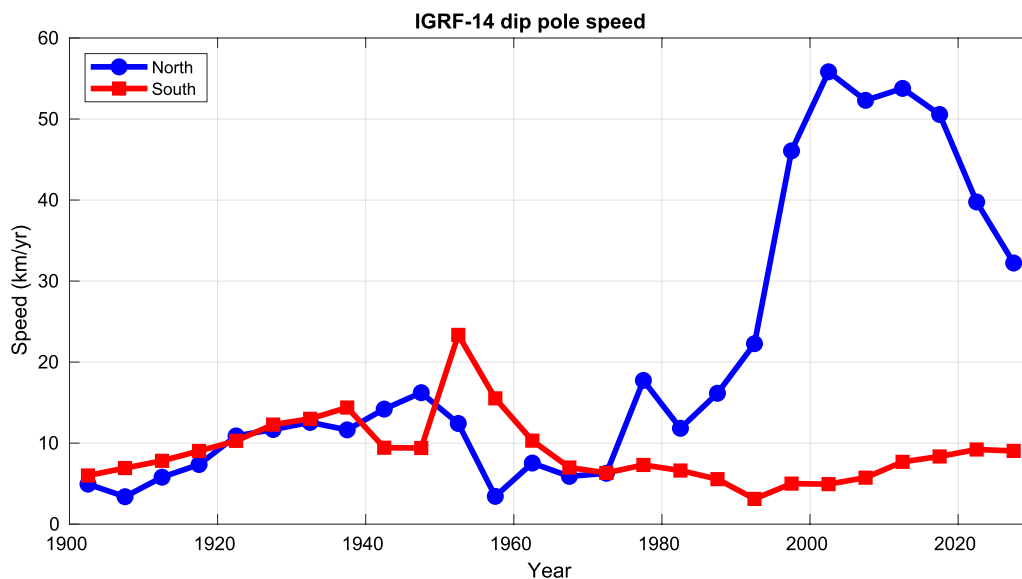
$$F = \sqrt{X^2 + Y^2 + Z^2}. \tag{8}$$

In the convention of geomagnetism, a deviation of the compass needle west of True North is considered to be negative Declination while east of True North is positive. A needle with its north magnetic pole pointing into the ground is considered to have positive Inclination while negative Inclination is when the north pole of the needle points upward away from the ground. Compasses generally have positive Inclination in the northern hemisphere and negative in the south, though this is not true everywhere. The total field is a scalar value which is always

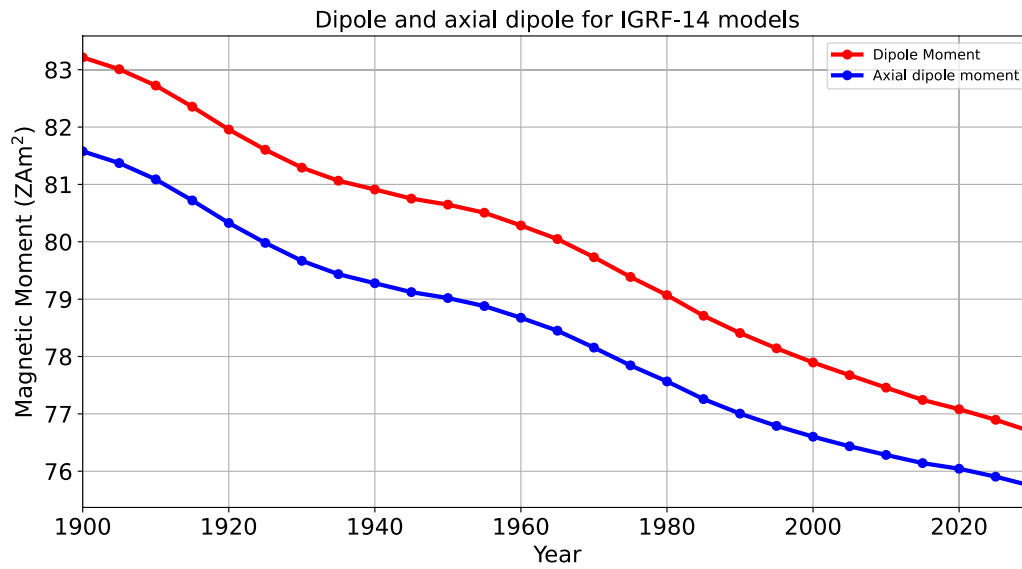
positive and varies between 22,000 nT and 70,000 nT on the Earth's surface.

The IGRF-14 provides magnetic values in geocentric spherical coordinates which must be converted to geodetic coordinates using a reference ellipsoid, usually WGS-84. This involves computing the equivalent geodetic latitude and rotating the  $X$  and  $Z$  magnetic components to compensate for the change in the normal vector.

Figure 1 shows global maps of the geodetic IGRF-14 declination ( $D$ ), inclination ( $I$ ), and total field intensity ( $F$ ) on Earth's surface in 2025.0 in Winkel Tripel projection. For the declination component (top panel), the dashed lines are the agonic lines (zero declination angle) on which a magnetic compass needle points to true geographic north. For the inclination map (middle panel), the dashed contour line of zero inclination angle and is known as the magnetic dip equator. It approximately aligns with the geographic equator except for a large, well-known southward deviation over South America. The total field ( $F$ ) map (bottom panel) shows that the highest field intensities occur in Siberia in the northern hemisphere and in the Southern Ocean between Australia and Antarctica in the southern hemisphere. We also see a region of significantly weaker field (compared to an idealized dipole), centered over South America, which is known as the South Atlantic Anomaly. The South Atlantic Anomaly has deepened by around 150 nT over the last five years and moved westward at around 20 km/year. The minimum total field value at 2025.0 on the Earth's surface is 22,071 nT and lies at 26.22°S, 60.03°W in a rural area of northeastern Argentina. We



**Fig. 4** Average speed of the magnetic dip poles over each five year epoch, plotted at the midpoint between epochs (i.e. the speed over 2020–2025 is shown at 2022.5). The value for 2025–2030 is a forecast



**Fig. 5** Dipole and axial dipole moment time series derived from IGRF-14 in units of Zetta ( $10^{21}$ ) Am<sup>2</sup>. The value for 2030 is a forecast

**Table 3** Magnetic pole position since 1900 determined from IGRF-14 in units of degrees

| Epoch  | North dip pole |           | South dip pole |           | North geomagnetic pole |           | South geomagnetic pole |           |
|--------|----------------|-----------|----------------|-----------|------------------------|-----------|------------------------|-----------|
|        | Latitude       | Longitude | Latitude       | Longitude | Latitude               | Longitude | Latitude               | Longitude |
| 1900.0 | 70.46          | -96.19    | -71.72         | 148.32    | 78.68                  | -68.79    | -78.68                 | 111.21    |
| 1905.0 | 70.66          | -96.48    | -71.46         | 148.54    | 78.68                  | -68.75    | -78.68                 | 111.25    |
| 1910.0 | 70.79          | -96.72    | -71.15         | 148.64    | 78.66                  | -68.72    | -78.66                 | 111.28    |
| 1915.0 | 71.03          | -97.03    | -70.80         | 148.54    | 78.64                  | -68.57    | -78.64                 | 111.43    |
| 1920.0 | 71.34          | -97.38    | -70.41         | 148.20    | 78.63                  | -68.38    | -78.63                 | 111.62    |
| 1925.0 | 71.79          | -97.99    | -69.99         | 147.62    | 78.62                  | -68.27    | -78.62                 | 111.73    |
| 1930.0 | 72.27          | -98.68    | -69.52         | 146.79    | 78.60                  | -68.26    | -78.60                 | 111.74    |
| 1935.0 | 72.80          | -99.33    | -69.06         | 145.76    | 78.57                  | -68.36    | -78.57                 | 111.64    |
| 1940.0 | 73.30          | -99.87    | -68.57         | 144.59    | 78.55                  | -68.51    | -78.55                 | 111.49    |
| 1945.0 | 73.93          | -100.24   | -68.15         | 144.44    | 78.55                  | -68.53    | -78.55                 | 111.47    |
| 1950.0 | 74.64          | -100.86   | -67.89         | 143.55    | 78.55                  | -68.85    | -78.55                 | 111.15    |
| 1955.0 | 75.18          | -101.42   | -67.19         | 141.50    | 78.54                  | -69.16    | -78.54                 | 110.84    |
| 1960.0 | 75.30          | -101.03   | -66.70         | 140.23    | 78.58                  | -69.47    | -78.58                 | 110.53    |
| 1965.0 | 75.63          | -101.34   | -66.33         | 139.53    | 78.60                  | -69.85    | -78.60                 | 110.15    |
| 1970.0 | 75.88          | -100.97   | -66.02         | 139.40    | 78.66                  | -70.18    | -78.66                 | 109.82    |
| 1975.0 | 76.15          | -100.64   | -65.74         | 139.52    | 78.76                  | -70.47    | -78.76                 | 109.53    |
| 1980.0 | 76.91          | -101.68   | -65.42         | 139.35    | 78.88                  | -70.76    | -78.88                 | 109.24    |
| 1985.0 | 77.40          | -102.61   | -65.13         | 139.18    | 79.04                  | -70.90    | -79.04                 | 109.10    |
| 1990.0 | 78.10          | -103.69   | -64.91         | 138.90    | 79.21                  | -71.13    | -79.21                 | 108.87    |
| 1995.0 | 79.04          | -105.29   | -64.79         | 138.73    | 79.39                  | -71.42    | -79.39                 | 108.58    |
| 2000.0 | 80.97          | -109.64   | -64.66         | 138.30    | 79.61                  | -71.57    | -79.61                 | 108.43    |
| 2005.0 | 83.19          | -118.22   | -64.55         | 137.85    | 79.82                  | -71.81    | -79.82                 | 108.19    |
| 2010.0 | 85.02          | -132.84   | -64.43         | 137.32    | 80.09                  | -72.21    | -80.09                 | 107.79    |
| 2015.0 | 86.31          | -160.34   | -64.28         | 136.60    | 80.38                  | -72.61    | -80.38                 | 107.39    |
| 2020.0 | 86.49          | 162.80    | -64.08         | 135.87    | 80.65                  | -72.68    | -80.65                 | 107.32    |
| 2025.0 | 85.73          | 138.65    | -63.86         | 135.07    | 80.85                  | -72.76    | -80.85                 | 107.24    |
| 2030.0 | 84.73          | 125.94    | -63.66         | 134.27    | 81.05                  | -72.96    | -81.05                 | 107.04    |

Latitudes are provided in the WGS-84 geodetic system

**Table 4** Magnetic observatories contributing data used in the construction of IGRF-14.

| Supporting agencies   | Country      | Observatory IAGA code  |
|---|--------------|--|
| Centre de Recherche en Astronomie, Astrophysique et Geophysique | Algeria      | TAM  |
| Universidad Nacional de la Plata                                | Argentina    | TRW  |
| Servicio Meteorologico Nacional                                 | Argentina    | PIL, ORC   |
| Geoscience Australia  | Australia    | ASP, CKI, CNB, CSY, CTA, DVS<br>GNA, GNG, KDU, LRM, MAW, MCQ                 |
| Geosphere Austria   | Austria      | WIC  |
| Institut Royal Météorologique                                   | Belgium      | DOU, MAB   |
| CNPq-Observatorio Nacional                                      | Brazil       | VSS, TTB   |
| Bulgarian Academy of Sciences                                   | Bulgaria     | PAG  |
| Geological Survey of Canada                                     | Canada       | ALE, BLC, BRD, CBB, FCC, IQA<br>MEA, OTT, RES, STJ, VIC, YKC                 |
| Dirección Meteorológica de Chile                                | Chile        | IPM  |
| Academy of Sciences   | China        | BMT, SSH   |
| China Earthquake Administration                                 | China        | CDP, CNH, GLM, GZH, KSH, LZH<br>MZL, QGZ, QIX, QZH, THJ, WHN                 |
| Instituto Geographico Agustin Codazzi                           | Columbia     | FUQ  |
| University of Zagreb  | Croatia      | LON  |
| Academy of Sciences of the Czech Republic                       | Czechia      | BDV  |
| Technical University of Denmark, DTU Space                      | Denmark      | BFE, NAQ, GDH, THL   |
| Addis Ababa University  | Ethiopia     | AAE  |
| Finnish Meteorological Institute                                | Finland      | NUR  |
| Geophysical Observatory   | Finland      | SOD  |
| Institut de Physique du Globe de Paris                          | France       | AAE, BOX, CLF, DLT, KOU, IPM<br>LZH, MBO, PHU, PPT, TAM                      |
| Ecole et Observatoire des Sciences de la Terre                  | France       | AMS, CZT, DMC, DRV, PAF, TAN   |
| Universität München   | Germany      | FUR  |
| Alfred-Wegener-Institute for Polar Marine Research              | Germany      | VNA  |
| GFZ Helmholtz Centre for Geosciences                            | Germany      | ABG, BFO, GAN, KMH, LRV, NGK, PAG, SHE,<br>SUA, STT, TDC, TTB, VNA, VSS, WNG |
| Universität Stuttgart and KIT                                   | Germany      | BFO  |
| Institute of Geology and Mineral Exploration                    | Greece       | PEG  |
| Academy of Sciences   | Hungary      | NCK  |
| Mining and Geological Survey of Hungary                         | Hungary      | THY  |
| University of Iceland   | Iceland      | LRV  |
| Indian Institute of Geomagnetism                                | India        | ABG, JAI, NGP, PND, SIL<br>SHL, TIR, UJJ, VSK                                |
| National Geophysical Research Institute                         | India        | CPL, HYB   |
| Meteorological and Geophysical Agency                           | Indonesia    | KPG, PLR, TND, TUN   |
| Met Éireann   | Ireland      | VAL  |
| Survey of Israel  | Israel       | AMT, BGY, ELT  |
| Instituto Nazionale di Geofisica e Vulcanologia                 | Italy        | AQU, CTS, DMC  |
| Japan Meteorological Agency                                     | Japan        | CBI, KAK, KNY, MMB   |
| Geospatial Information Authority of Japan                       | Japan        | ESA, KNZ, MIZ  |
| Institute of the Ionosphere                                     | Kazakhstan   | AAA  |
| Korean Meteorological Administration                            | Rep of Korea | CYG  |
| Institut et Observatoire Géophysique d'Antananarivo             | Madagascar   | TAN  |
| Maldives Meteorological Service                                 | Maldives     | GAN  |
| Institute of Geological and Nuclear Sciences                    | New Zealand  | API, EYR, SBA  |
| University of Tromsø  | Norway       | BJN, DOB, TRO  |

**Table 4** (continued)

| Supporting agencies   | Country        | Observatory IAGA code  |
|---|----------------|--|
| Instituto Geofísico del Peru  | Peru           | HUA  |
| Academy of Sciences   | Poland         | BEL, HLP, HRN  |
| Universidade de Coimbra   | Portugal       | COI  |
| Instituto Português do Mar e da Atmosfera                                   | Portugal       | STT  |
| Geological Survey of Romania  | Romania        | SUA  |
| AARI  | Russia         | VOS  |
| GC RAS  | Russia         | ARS, KLI, MHV, SPG   |
| IPE RAS   | Russia         | BOX  |
| IG UB RAS   | Russia         | ARS  |
| IDG RAS   | Russia         | MHV  |
| IKIR FEB RAS  | Russia         | KHB, MGD, PET  |
| IPGG SB RAS   | Russia         | NVS  |
| ISTP SB RAS   | Russia         | IRT  |
| IZMIRAN SpB Branch  | Russia         | SPG  |
| SHICRA SB RAS   | Russia         | YAK  |
| Dept. of Agriculture, Forestry, Fisheries & Meteorology                     | Samoa          | API  |
| Geomagnetic College Grocka  | Serbia         | GCK  |
| Slovenska Akademia Vied   | Slovakia       | HRB  |
| National Research Foundation  | South Africa   | HBK, HER, KMH, TSU   |
| Observatori de l'Ebre   | Spain          | EBR, LIV   |
| Real Instituto y Observatorio de la Armada                                  | Spain          | SFS  |
| Instituto Geográfico Nacional   | Spain          | GUI, SPT   |
| Sveriges Geologiska Undersökning  | Sweden         | ABK, LYC, UPS  |
| Swedish Institute of Space Physics  | Sweden         | KIR  |
| ETH Zurich  | Switzerland    | GAN  |
| Boğaziçi University   | Türkiye        | IZN  |
| Academy of Sciences   | Ukraine        | AIA, LVV   |
| British Geological Survey   | United Kingdom | ASC, ESK, HAD, JCO, KEP, LER, PST, SBL                                 |
| US Geological Survey  | United States  | BRW, BOU, BSL, CMO, DED, FRD, FRN<br>GUA, HON, NEW, SIT, SJG, SHU, TUC |
| Institute of Geophysics of the Vietnamese Academy of Science and Technology | Vietnam        | DLT, PHU   |

also point out that a second low point in the South Atlantic Ocean, offshore of South Africa, continues to grow as noted by Finlay et al. (2020) and Rother et al. (2021).

Figure 2 shows the predicted average change of the  $D$ ,  $I$ , and  $F$  components on Earth's surface during the interval between 2025 and 2030. At low and middle latitudes, the map of  $dD/dt$  (top panel) predicts the largest declination changes in the South Atlantic Anomaly region and also in the polar regions, with northern polar declination angle changing more than in the southern polar region. The  $dI/dt$  map (middle panel) predicts the largest changes over Brazil, where the magnetic dip equator has moved relatively rapidly over the past few decades. The largest changes in field strength ( $dF/dt$ ) are modeled in North America (of  $-128$  nT/yr) while the Indian Ocean also has a variation of  $104$  nT/yr.

Figure 3 presents the positions of the *geomagnetic* poles and *dip* poles as given by IGRF-14 for 1900–2025, and their predicted positions in 2030. The geomagnetic poles are calculated from the three dipole ( $n = 1$ ) Gauss coefficients and correspond to where the magnetic dipole axis intersects a sphere of mean Earth radius (6371.2 km). These poles are antipodal and are also known as centered dipole poles (Laundal and Richmond 2017, Eq. 14). The geomagnetic poles can be used to specify the relative orientation of Earth's magnetic field with respect to the Sun, and they are often used in magnetospheric studies for this purpose.

The magnetic *dip* poles are defined as the locations where the main magnetic field is normal (inclination angle is  $90^\circ$ ) to Earth's surface, as represented by the WGS-84 reference ellipsoid. Here, we use the full set

of IGRF-14 coefficients to spherical harmonic degree and order  $N = 13$ . The location of the dip pole is found by searching for the position where the horizontal field component ( $H = \sqrt{X^2 + Y^2}$ ) is zero. We use a gradient descent search algorithm to find this point on a grid of  $0.01^\circ$  resolution to give a modeled location to within 1 km accuracy. In reality, the influence of secular variation and the external magnetic field will cause the true location to move continuously.

Magnetic dip poles provide a key reference for local orientation when navigating on or close to Earth's surface at high-latitudes. However, note that in general a compass needle does not point to the dip pole but aligns along the local magnetic field direction. For a perfect dipole field, the geomagnetic and dip poles would nearly coincide, but not exactly since the geomagnetic poles are defined with respect to a sphere of mean Earth radius, while the dip poles are defined with respect to the WGS-84 ellipsoid. As can be seen in Figure 3, there are significant differences between the two due to the non-dipolar structure of Earth's magnetic field. The geomagnetic and dip pole locations are provided in Table 3. Interestingly, IGRF-14 predicts the magnetic north and south dip poles will briefly have the same longitude (around  $136^\circ\text{W}$ ) in early 2026.

Figure 4 shows the speed of the two magnetic dip poles. The north magnetic dip pole experienced a strong acceleration from about 1960–2000, but has seen a modest deceleration over the past 25 years, peaking at 56 km/year in 2002.5 and slowing to 40 km/year in 2012.5. IGRF-14 forecasts continued deceleration with a speed of 31 km/year in 2027.5. This equates to a dip pole motion of about 100 m/day. However, past IGRF forecasts have often over- or underestimated the dip pole velocity (Finlay et al. 2010b). As with any model, uncertainties are present in IGRF and are further discussed by Lowes (2000).

Using the ratio of the dipole to non-dipole terms from the Lowes–Mauersberger power spectrum (Beggan et al. 2026) at Earth's surface in 2025, the contribution from the dipole terms  $g_1^0, g_1^1, h_1^1$  account for 92.31% of the power in the main geomagnetic field, a slight fall from 92.52% in 2020. Figure 5 shows the decline in the dipole moment of the geomagnetic field since 1900 as predicted by IGRF-14 (red). The dipole moment is defined as:

$$M(t) = \frac{4\pi}{\mu_0} a^3 \sqrt{g_1^0(t)^2 + g_1^1(t)^2 + h_1^1(t)^2}, \quad (9)$$

where  $\mu_0$  is the permeability of free space ( $4\pi \times 10^{-7}$  H/m). Archaeomagnetic and palaeomagnetic studies often estimate the dipole strength along the rotation axis, ignoring the off-axis terms  $g_1^1, h_1^1$ . This axial dipole moment is defined as  $M_A(t) = 4\pi a^3 |g_1^0(t)|/\mu_0$  and is consequently lower, as shown in blue in the figure. The dipole field is forecast to continue to fall in strength over the next five years and weaken relative to the non-dipolar contributions.

## 5 Conclusion

We present the 14th generation of the International Geomagnetic Reference Field. The model is valid from 1900.0 to 2030.0, describing the internal magnetic field of the Earth using spherical harmonic Gauss coefficients to degree and order 13 with snapshots every five years. The final snapshot between 2025.0 and 2030.0 is a forecast of secular variation to degree and order 8.

The IGRF is a truly international effort involving thousands of observers, scientists and engineers from around the world. Without their contributions, this would not be possible. This generation introduced the use of open source software tools including data repositories and automated evaluation code. The IGRF-14 coefficients, model code and maps are freely available for use. They can be found both within and at the online sources noted in this paper.

## 6 IGRF-14 online data products

Coefficients of IGRF-14 in ASCII format, other data formats, programs, software and further general information about IGRF:

<https://www.ncei.noaa.gov/products/international-geomagnetic-reference-field>

Coefficients of IGRF-14 (Zenodo repository): <https://doi.org/10.5281/zenodo.14218973>

IGRF-14 candidate models and evaluation code (GitHub repository): <https://github.com/IAGA-VMOD/IGRF14eval>

Python software to compute magnetic field components from coefficients (Laundal et al. 2024): <https://doi.org/10.5281/zenodo.14231854>

Online calculation of magnetic field components for IGRF-14:

<https://www.ncei.noaa.gov/products/international-geomagnetic-reference-field>

[http://geomag.bgs.ac.uk/data\\_service/models\\_compass/igrf\\_calc.html](http://geomag.bgs.ac.uk/data_service/models_compass/igrf_calc.html)

<http://wdc.kugi.kyoto-u.ac.jp/igrf/point/index.html>

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**Author contributions**

C.D. Beggan is chair of the IAGA DIVV-MOD (2023-2027) and initiated, coordinated and organized the call and delivery of the 14th generation of the IGRF. C. Kloss is co-chair (2023-2027). CDB and CK wrote the manuscript based on the analyses of the contributing co-authors. All other authors contributed modelling results and/or detailed technical analyses for IGRF-14. All co-authors have read and approved the manuscript.

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**Data availability**

Swarm and Cryosat-2 data are available from <https://earth.esa.int/web/guest/swarm/data-access>. Swarm ASM-V experimental data are available from <https://dataverse.ipgp.fr/dataverse/swarm-data>. CHAMP data can be obtained from <http://isdg.gfz.de>. Ørsted and SAC-C data are available from [https://www.space.dtu.dk/english/research/scientific\\_data\\_and\\_models/magnetic-satellites](https://www.space.dtu.dk/english/research/scientific_data_and_models/magnetic-satellites). CSES data are available from <http://www.leos.ac.cn>. Macau Science

Satellite data are available from <https://mss.must.edu.mo/>. INTERMAGNET data are available from <https://www.intermagnet.org>.

## Declarations

### Competing interests

The authors declare that they have no competing interests.

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