

1 Modelled connectivity of Antarctic krill 2 spawning and nursery grounds along the 3 Western Antarctic Peninsula

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23 Abstract

24 Understanding the connectivity between spawning and nursery grounds of Antarctic krill
25 (*Euphausia superba*) is essential for elucidating population dynamics and informing
26 ecosystem-based management of krill. This study employed a regional ocean model and
27 particle tracking to investigate larval krill transport and connectivity along the western
28 Antarctic Peninsula (wAP), a key region for Antarctic krill, over three austral summers
29 (2016–2019). Simulations incorporated the descent-ascent cycle of early larval development
30 based on initial embryo sizes, along with varying larval behavior, to examine how physical
31 and biological processes affect habitat connectivity. Results revealed that nursery ground use
32 and retention times were closely linked to local bathymetry. Deeper regions, such as the
33 Bransfield Strait and Marguerite Bay, retained larvae for longer durations (>50 days), while
34 shallower nursery grounds (Gerlache Strait, Grandidier Passage) acted as secondary nursery
35 habitats in simulations with smaller initial embryo sizes. Simulations with larger initial
36 embryo sizes led to notable changes in retention, connectivity, and spawning locations, as
37 more embryos successfully hatched before reaching the seafloor. Interannual wind
38 variability emerged as a key driver of larval supply pathways, influencing connectivity
39 between spawning areas and nursery grounds. These results highlight how a shallower
40 hatching depth expands the range and reliability of the connectivity between spawning and
41 nursery grounds in the wAP. By illustrating how physical forcing and biological traits jointly
42 shape connectivity and retention between spawning and nursery grounds this study

43 emphasizes the importance of integrating biological parameterization with physical modeling
44 to better understand krill population dynamics.

45 1 Introduction

46 Antarctic krill (*Euphausia superba*, hereafter krill) are vital to the Southern Ocean
47 ecosystem, linking upper trophic levels including penguins, seals, and whales, to energy from
48 primary producers (Nicol et al. 2012, Trathan & Hill 2016, Kawaguchi et al. 2024). While
49 krill are found at a circumpolar scale, the population center is located at the northern tip of
50 the western Antarctic Peninsula (wAP) in the Southwest Atlantic (Atkinson et al. 2017,
51 Perry et al. 2019). The region is both ecologically significant supporting high levels of
52 biodiversity, as well as anthropologically significant with concentrated areas of scientific
53 research, tourism, and commercial fishing (Nicol & Foster 2016, Meyer et al. 2020). The
54 largest fishery, by biomass, in the Southern Ocean, specifically targets krill (Nicol & Foster
55 2016).

56
57 Climate-driven changes, such as warming sea temperatures (Vaughan et al. 2003, Meredith &
58 King 2005, Turner et al. 2022), shifting ocean currents and reduced sea ice coverage (Turner
59 et al. 2013, Stammerjohn et al. 2015), have been linked to declines in krill recruitment and a
60 southward shift in krill distribution in the region over the past 90 years (Atkinson et al.
61 2019). The impacts of climate change on the large and connected marine ecosystem linked
62 by advection have implications for population structure and the relative connectivity

63 between areas of the Antarctic. Understanding the connections between where krill are
64 spawned and where they develop can illuminate how krill connectivity patterns may be
65 affected by climate change.

66

67 The early life cycle of krill begins with an embryo-larvae descent-ascent cycle. After
68 successful fertilization and spawning near the ocean surface, embryos sink to depths
69 determined by their density relative to ambient conditions as influenced by the embryo
70 diameter and developmental stage (Ikeda 1984, Hofmann et al. 1992). After hatching, larvae
71 develop through several stages as they ascend back towards the surface as nauplii (Marschall
72 & Hirche 1984). Spawning typically occurs during the austral summer (December–February,
73 Quetin et al. 1994), though females can spawn multiple times per season under favorable
74 food and thermal conditions (Quetin & Ross 2001, Kawaguchi 2016). Timing and frequency
75 vary among regions and years, influenced by sea ice, food availability, and female condition
76 (Tarling 2020).

77

78 After developing through calyptope stages, larvae become furciliae at which point they can
79 undertake diel vertical migrations (DVM): feeding in productive surface waters at night and
80 spending daylight hours at depth, or vice versa according to environmental conditions
81 (Makarov 1982, Meyer et al. 2017). Despite this vertical control, early-stage krill swimming
82 strength remains limited, rendering their horizontal movement largely passive and subject to
83 ocean currents (Hofmann & Lascara 2000). Sea ice can further modulate larval transport by

84 providing refuge from predation and supplemental food sources (Kohlbach et al. 2017,
85 Schaafsma et al. 2017). Incorporating sea ice into models can markedly alter predicted
86 connectivity pathways relative to ocean-only simulations (Thorpe et al. 2007, Young et al.
87 2023).

88

89 The wAP is a focal region for understanding these processes, as it hosts multiple nursery
90 grounds where larval krill aggregate in early winter such as the Bransfield Strait, Gerlache
91 Strait, Grandidier Passage, and Marguerite Bay (Figure 1) (Reiss et al. 2017, Meyer et al. 2017,
92 Perry et al. 2019). Notably, the Bransfield Strait harbors greater krill abundance in winter
93 than summer, suggesting that this region is critical for overwintering (Reiss et al. 2017,
94 Meyer et al. 2017). The Gerlache Strait, meanwhile, supports overwintering *E.*
95 *crystallorophias* and adult *E. superba* (Loeb & Santora 2015), and the Grandidier Passage
96 hosts continuous bands of feeding larvae in summer with recruits concentrated on the
97 continental shelf (Perry et al. 2019). Marguerite Bay is notable for large aggregations of larval
98 krill and observations of gravid females at depth, indicating a productive spawning habitat
99 (Lascara et al. 1999, Ashjian et al. 2008). Despite this broad recognition of nursery hotspots,
100 the mechanisms linking embryonic krill transport from spawning sites to these nurseries
101 remain poorly resolved.

102

103 Changes in water mass properties and circulation along the wAP may redistribute spawning
104 grounds, influencing recruitment success (Reiss 2016, Atkinson et al. 2019). As water-mass

105 movements shift, connectivity to suitable nursery sites could be modified, potentially
106 reducing survival if krill fail to encounter suitable habitats or if sea ice coverage diminishes
107 (Meyer et al. 2017, Young et al. 2023). Resolving where embryos originate, how they move,
108 and which factors control their ultimate success is vital for both understanding krill
109 population dynamics and informing sustainable fisheries management. To address these
110 uncertainties, numerical modeling offers a powerful platform for simulating ocean–ice–krill
111 interactions at spatial and temporal resolutions difficult to achieve through field studies
112 (Johnston et al. 2022).

113

114 Here, we employ the Regional Ocean Modeling System (ROMS) configured for the Southern
115 Ocean to investigate how ocean circulation and larval behaviors, such as DVM and ice-
116 associated advection, shape connectivity pathways for early-stage krill. Our particle tracking
117 experiments integrate the embryo–larval descent–ascent cycle, capturing the transition from
118 sinking embryos to nauplii stages and culminating in the onset of DVM in the furcilia phase.
119 By focusing on the wAP, we aim to illuminate the link between spawning locations and
120 known nursery grounds, exploring how future shifts in wind, temperature, and sea ice may
121 reroute these connections. Identifying the sources and sinks of krill larvae has direct
122 relevance to fisheries and conservation policies, as decision-makers require information on
123 where to allocate resources for monitoring and what habitats merit protection under
124 dynamic environmental conditions. Through our modeling approach, we offer new insights
125 into the spatial and temporal aspects of krill larval transport, setting the stage for more

126 targeted research and management actions that support the continued productivity and
127 resilience of the Southern Ocean ecosystem.

128 2 Materials and Methods

129 2.1 Ocean Circulation Model

130 The circum-Antarctic circulation model used in this study, described by Dinniman et al.
131 (2020), is an application of ROMS, which is a primitive equation finite volume model with a
132 terrain-following vertical coordinate system (Haidvogel et al. 2008, Shchepetkin &
133 McWilliams 2009). The model includes a dynamic sea-ice model (Budgell 2005), and
134 explicitly simulates the mechanical and thermodynamic interactions between floating ice
135 shelves and the water underneath (Holland & Jenkins 1999, Dinniman et al. 2011). The
136 dynamic sea-ice model includes two-layer ice thermodynamics (Mellor & Kantha 1989,
137 Häkkinen & Mellor 1992), a snow layer, a simple estimate of frazil ice production in the
138 water column (Steele et al. 1989), and an elastic-viscous-plastic rheology (Hunke &
139 Dukowicz 1997, Hunke 2001).

140

141 The model domain includes the entire Antarctic continental shelf, including the portion
142 (~35%) of the continental shelf ocean underneath the floating ice shelves, and extends
143 northward from the continent past the annual maximum sea ice extent and across Drake
144 Passage to South America (Dinniman et al. 2020). The model grid uses a polar stereographic
145 projection with a 5-km horizontal resolution and has 32 terrain-following vertical levels

146 with smaller spacing near the surface and bottom. The bathymetry for the ocean and ice
147 shelf cavities (bottom and top ocean surfaces) is from the RTopo-2 dataset (Schaffer et al.
148 2016), with updates in the Amundsen Sea Embayment (Millan et al. 2017).

149

150 The circum-Antarctic simulation used in this study starts after a six-year model spin-up.
151 Lateral boundary conditions for temperature and salinity, velocity, and sea surface height are
152 the same as in Dinniman et al. (2015, 2020). There is no need to specify sea ice boundary
153 information as the model domain fully encompasses any sea ice observed during the satellite
154 era (1979 to present). Following the six-year spin-up simulation, the model was run for four
155 years using atmospheric forcing for winds, 2-m air temperatures, sea level pressure, and
156 relative humidity obtained from the ERA5 reanalysis (Hersbach et al. 2020) covering the
157 years 2016-2019. Precipitation and cloud fraction were specified as described in Dinniman et
158 al. (2020) and ice shelves were assumed to be static (no thinning, thickening or iceberg
159 calving). Characteristics of the model physical representation near Antarctica have been
160 presented in Dinniman et al. (2020, 2023) and Galton-Fenzi et al. (2025).

161 2.2 Lagrangian Drifter Simulation Design

162 A set of online Lagrangian particle tracking simulations was conducted to incorporate unique
163 drifter behavior across three consecutive austral summers: 2016/17, 2017/18, and 2018/19. In
164 each simulation, one drifter was released in every 25 km by 25 km area (625 km²) across all
165 areas of the Antarctic continental shelf and slope, extending to a maximum seabed depth of
166 3,000 m, but excluding regions beneath ice shelves (Figure 1A). A total of 7,707 drifters were

167 released per deployment. Drifters were released every 14 days from November 1 through
168 March 21, resulting in 11 releases and 84,777 drifters per simulation. The release window
169 was selected to span the known timing of krill spawning across the Southern Ocean.
170 Additionally, it allows for temporal investigation of early and late season spawning events,
171 making the overall dataset more useful for other study adaptations. This approach represents
172 an idealized reality with a uniform spatial and temporal distribution of spawning. Drifters
173 were tracked until Dec. 31, 2019, with the first six months of the drifter trajectories used for
174 subsequent analyses in this study.

175
176 While this paper focuses on transport pathways between spawning and nursery grounds, the
177 full scale of the research encompasses an effort to assess the interplay between larval krill
178 behavior and interannual oceanographic variations across the entire Southern Ocean. As part
179 of that effort, additional simulations were conducted that varied drifter behavior parameters
180 such as diel vertical migration behaviors and sea ice advection rules. The alternative
181 configurations are fully described in the Supplementary Materials along with sensitivity
182 analyses of behavioral differences (see Supplementary Material). Model output from all
183 configurations are available for public download (see Data Availability).

184

185 2.3 Lagrangian Drifter Configurations

186 To isolate the effects of early developmental dynamics on transport from spawning grounds,
187 we compared two simulations that were identical except for the configuration of the initial

188 embryo size assigned to each drifter (620 μm vs. 640 μm). This section outlines the shared
189 behavioral configuration used in both simulations, including: (1) temporal parameters and
190 vertical motion behaviors, (2) embryo size parameterization, (3) horizontal advection rules,
191 and (4) diel vertical migration and sea ice interactions.

192

193 Drifters were advected every model baroclinic time step (180 s) using either the three-
194 dimensional ocean velocity or the two-dimensional sea ice velocity (see below). Their
195 motion included a random vertical walk, as described by Hunter et al. (1993) and Visser
196 (1997), which was a function of the model-computed vertical turbulence. This random walk
197 was designed to represent the transport effects of parameterized vertical turbulent motions,
198 thereby enhancing the realism of drifter trajectories under varying oceanographic
199 conditions.

200

201 Vertical motions to represent early life stage behavior and DVM of krill were added to the
202 drifters. Drifters were initially released at 5 m below the surface and given an additional
203 vertical descent rate to represent the initial descent of the embryos after spawning until
204 hatching occurred at depth (Marr 1962, Quetin & Ross 1984). Drifters that reached the
205 seafloor before hatching were assumed to die (Hofmann & Hüsrevoğlu 2003, Tarling et al.
206 2007). On successful hatching, the larvae were given an additional vertical ascent rate to
207 simulate the ascent of larvae (Hofmann et al. 1992) until they reached the surface (Figure 2).
208 Vertical descent and ascent rates and hatching time were calculated off-line using the

209 embryo-larva descent-ascent model of Hofmann et al. (1992) for November–March in three
210 regions with variable water column physical properties: the Bransfield Strait, Marguerite
211 Bay, and the Bellingshausen Sea.

212

213

214

215 Embryo-larva descent-ascent rates were calculated for two different initial embryo sizes (620
216 μm and 640 μm) to span the observed range of newly released single-celled embryos (Fraser
217 1936, Marschall 1983, Quetin & Ross 1984, 2003). The average descent and ascent rates and
218 hatching times across the three locations and five months were applied to all drifters
219 according to initial embryo size (Table 1). While the average hatching time was the same
220 between initial embryo sizes, there was substantial variance in descent-ascent rates.

221 Specifically, although 620 μm embryos sank until they hatched in the off-line simulations,
222 640 μm embryos sank initially but shortly afterwards became positively buoyant and
223 ascended to the surface where they ultimately hatched. This aligns with Hofmann and
224 Hüsrevolğlu (2003), who described the positively buoyant gastrula stage for larger embryos
225 but also noted that such buoyancy does not match existing observations of early krill life
226 stages, which are not typically found at the surface (Hofmann & Hüsrevoğlu 2003 Section
227 3.4.1). Furthermore, nauplii stages have not been observed at the sea surface, implying that
228 hatching likely occurs at depth (Hempel 1979, 1985, Schnack et al. 1985). These

229 observations, while limited in spatial and temporal coverage, suggest that the extent to which
230 near-surface hatching occurs remains uncertain.

231

232 Following the completion of the initial descent–ascent cycle, drifters representing newly
233 hatched krill larvae remained near the surface and were advected horizontally by either
234 ocean currents or sea ice drift, depending on local conditions. Drifters transitioned to sea ice
235 advection when three criteria were met: they were located within 2 meters of the surface,
236 more than 120 hours had elapsed since release, and the local sea ice concentration exceeded
237 25%. This behavioral rule was designed to reflect the tendency of larval krill to associate
238 with sea ice when it is present, which can influence both their transport pathways and
239 survival prospects. If these conditions were not met, drifters were advected according to
240 ocean currents.

241

242 After 63 days, drifters were assumed to have reached a life stage capable of completing diel
243 vertical migration (DVM) (Ikeda, 1984) and were assigned the following
244 vertical behavior as determined dynamically based on local sea ice concentration. When sea
245 ice concentration was below 25%, drifters followed a standard DVM pattern: they descended
246 at 0.016 m s^{-1} during daylight hours (when the sun was above the horizon) to a maximum
247 depth of 150m and ascended at the same rate during nighttime until reaching within 1.5m of
248 the surface (Hudson et al. 2022) (Figure 3A). When sea ice concentration exceeded 25%,
249 drifters switched to a reverse DVM pattern, adapted from observations of overwintering krill

250 behavior in the northwestern Weddell Sea (Meyer et al. 2017). Under this regime, drifters
251 descended at 0.016 m s^{-1} during nighttime to a maximum depth of 20m and ascended during
252 daylight hours until they were again within 1.5m of the surface (Figure 3B). This behavioral
253 switching allowed the model to capture seasonal and spatial variability in krill vertical
254 positioning in response to sea ice conditions.

255 2.4 Evaluation and Analysis

256 Spawning ground analysis along the Antarctic Peninsula was conducted by isolating the
257 starting locations of drifters that were retained within regions that represented known
258 nursery grounds. For this paper, four core nursery grounds (NG) were selected along the
259 wAP: the Bransfield Strait (BS), the Gerlache Strait (GL), the Grandidier Passage (GP), and
260 Marguerite Bay (MB) (Figure 1). Locations were selected as potential nursery grounds based
261 on prior research identifying them as hotspots for larvae and juveniles (see Introduction).
262 Each NG was defined as a fixed polygonal region using latitude and longitude bounds (shown
263 in Figure 1 and listed in Supplementary Table S1).

264
265 For each NG, 'larval influx' was identified as the number of larvae that entered an NG for
266 any duration over the simulated lifespan. 'Retention duration' was then calculated as the
267 number of days a particle remained within a specific NG. To reduce the influence of outlying
268 retention spans, the duration was reported as the median along with the interquartile range
269 (IQR), representing the middle 50% of larvae retained in each NG. Nursery ground use, or
270 'larval retention', was then determined by the retention duration within the NG. Drifters

271 were identified as retained in an NG by residing within the NG for at least 21 days within
272 their first 180 days of life. This criterion ensured that retained larvae spent three weeks of
273 larval development within a specified NG and did not merely transit through the region. The
274 'retention rate' was then expressed as the percent of larval influx retained for at least 21 days
275 after entering.

276

277 To characterize the composite spatial distribution of potential spawning locations for larvae
278 retained in each nursery, we combined the release locations across all simulation years, for
279 each embryo size. Release locations were binned into a uniform grid (0.25° latitude \times 0.5°
280 longitude) and 2D histograms were computed for each nursery. Each composite histogram
281 was then normalized by the total number of retained larvae for that nursery. The normalized
282 values were then used to create maps in which each grid cell shows the percentage of that
283 nursery's retained larvae originating from that location, with all cells summing to 100. This
284 approach enables direct visual comparison of spatial distributions across nurseries.

285

286 To identify the spatial patterns of connectivity, larval trajectories of retained larvae were
287 categorized for each NG. The trajectory between spawning location and the first entrance of
288 the NG was considered the import route, while export routes were identified as the
289 trajectory after which the larvae did not re-enter the NG. For each initial embryo size, year,
290 and NG, the import and export segments were compiled across all larvae. Trajectory points
291 within the analysis region were binned into a uniform 0.25° grid to generate raw 2D

292 histograms of point density. These were then normalized by the number of contributing
293 larvae to produce maps of mean trajectory point density per larva. To facilitate visual
294 comparison across NGs, density values were scaled to percent-of-maximum and then
295 compared with year-specific animations of import and export trajectories to assess dominant
296 transport pathways. Directional summaries were produced by synthesizing dynamic patterns
297 (e.g. looping and reversals in the trajectories) not readily captured in the density maps alone.
298

299 3 Results

300 The results are presented in three parts. First, we characterize patterns of larval krill
301 retention within defined nursery grounds along the wAP (Section 3.1: Nursery Ground
302 Retention). Second, we identify the spawning locations that give rise to successful
303 recruitment in each nursery ground and describe the connectivity pathways linking them
304 (Section 3.2: Spawning Grounds and Connectivity). These first two sections are based on the
305 simulation configured with 620 μm embryos. Lastly, we assess how these patterns shift when
306 embryo size is increased to 640 μm (Section 3.3), allowing us to evaluate the influence of
307 initial embryo size on spawning locations and connectivity.

308 3.1 Nursery Ground Retention

309 The number of larvae that entered each NG (influx) varied considerably, with BS-NG
310 consistently receiving approximately three times more larvae than the other NG (Table 2,
311 embryo size 620 μm). However, retention durations were relatively consistent across the

312 three years simulated but showed variation between nursery grounds (Table 2). Annually,
313 both the BS-NG and MB-NG retained 74% of larvae that passed through their boundaries.
314 The annual median retention duration was similar in the BS and MB nursery grounds (48
315 and 50 days, respectively), but both nurseries showed substantial variability in duration
316 range with IQRs exceeding 60 days (Table 2). GP-NG retained just under half of all entrants,
317 while GS-NG had notably lower retention and the shortest median duration (Table 2). Across
318 all nursery grounds, large IQRs indicated substantial variability in retention durations.
319

320 3.2 Spawning Grounds and Connectivity

321 Viable spawning locations for 620 μm embryos were restricted to areas with depths sufficient
322 for embryos to hatch above the seafloor, typically >700 m. Of the four nursery grounds, only
323 BS-NG and MB-NG had suitable local spawning habitats for the smaller sized embryos with
324 54% and 64% (respectively) of retained larvae originating locally (Table 2). Within BS-NG,
325 much of the strait itself supported local spawning, with the southern edge of the strait
326 contributing the highest relative share of spawning locations across all three years (Figure
327 4A). Spawning locations for larvae retained in BS-NG spanned a broad longitudinal range
328 from $\sim 80^\circ$ W to nearly 50° W. In contrast, MB-NG had the most spatially restricted set of
329 spawning locations, concentrated within deep-trough habitats of the Marguerite Trough and
330 with no spawning occurring east of 70° W (Figure 4D).

331

332 In contrast, no local spawning occurred in GS-NG or GP-NG. Spawning locations for GS-NG
333 overlapped substantially with those of BS-NG, with the highest relative concentrations
334 occurring around the Boyd Strait (Figure 4B). GP-NG spawning locations also included sites
335 within the Bransfield Strait and BS-NG, but were more diffuse along the continental shelf,
336 with a few areas contributing higher proportions than others (Figure 4C).

337

338 Larval transport patterns from spawning locations into the GS-NG and GP-NG indicated that
339 most larvae entered the NGs from further north except for some larvae being transported
340 across the continental shelf and into the GP-NG (5A). While there was some interannual
341 variability in the distribution of spawning locations, the transport pathways into both NGs
342 remained consistent. The same was true for transport of externally spawned larvae into the
343 MB-NG (Figure 4D). The larvae spawned off the continental shelf (Figure 4D) traveled north
344 along the continental shelf until reaching the Marguerite Trough and were then transported
345 into the MB-NG (Figure 5A).

346

347 The trajectories of larvae from non-local spawning locations into the BS-NG varied
348 interannually (Figure 6A). In certain years, like the 2016–2017 and the 2017–2018 season,
349 larvae were advected from areas along the shelf as far south as Adelaide Island along the
350 northeastward along-slope current (Figure 5A). In the 2018–2019 season, larval transport was
351 observed from the northwest Weddell Sea into the BS-NG.

352

353 Within the model, February near-surface winds over the study region also varied
354 interannually (Figure 7). During 2016–2017, strong northerlies occurred over the shelf break
355 near Boyd Strait in February (Figure 7A). In the 2018-2019 season, the wind regime shifted,
356 with weaker westerlies at the tip of the Antarctic Peninsula and a more southerly component
357 blowing northward (Figure 7C)

358

359 Despite interannual variability in spawning ground locations for the BS-NG, trajectories of
360 larvae exiting BS-NG within the 180 day-simulation remained mostly consistent across all
361 three years (Figure 6B). The primary dispersal trajectories either sent larvae north around
362 Elephant Island and toward the South Scotia Ridge, or southward through the Gerlache
363 Strait and into GS-NG and even as far as the GP-NG (Figure 5B, Figure 8A). Export of larvae
364 from the GS-NG followed the same spatial pattern, with trajectories consistently moving into
365 the GP-NG (Figure 8A) either via the Gerlache Strait or along the continental shelf and
366 around Anvers Island (Figure 5B). If larvae were exported from GP-NG, they primarily
367 moved onto the continental shelf (Figure 5B). However, in one of the three years simulated,
368 larvae were exported along the coastline of Adelaide Island (Figure 5B) and into the MB-NG
369 from the GP-NG (Figure 8A). Lastly, export of larvae from the MB-NG either exited via the
370 trough, or, more commonly, were transported further south around the tip of Alexander
371 Island (Figure 5B). Of the larvae exported out of the region, trajectories did not transport
372 larvae into any of the northern NGs (Figure 8A).

373

374 3.3 Embryo Size

375 Increasing the initial embryo size from 620 μm to 640 μm altered the patterns in larval
376 retention, potential spawning locations, and connectivity between NGs. The number of
377 successful spawning events was much higher and vastly increased the larval influx into all
378 nurseries such as a nearly four-fold increase in the BS-NG and 15-fold increase in the MB-
379 NG (Table 2). Most notably, successful spawning occurred within GS-NG and GP-NG (Table
380 2). These changes were accompanied by changes in retention patterns where GS-NG and GP-
381 NG both saw median retention durations more than double (11 vs 25 days and 20 vs 44 days,
382 respectively) thereby supporting higher retention rates (37 % vs 52 % and 46 % vs 69 %,
383 respectively) (Table 2). In contrast, MB-NG exhibited only minor changes in retention (50 vs
384 58 days) and an increase in retention rate to 82% (from 74%), while BS-NG retention rates
385 remained stable (down to 70% from 74%) and duration slightly decreased (48 days down to
386 41 days) (Table 2).

387
388 Spawning grounds expanded across the continental shelf in the larger embryo simulations, as
389 evidenced by localized spawning accounting for over a quarter of retained larvae in GS-NG
390 (26%) and GP-NG (34%) (Table 2). The spatial distribution of potential spawning locations
391 was more diffuse with individual 0.5° grid cells rarely contributing more than 5% of a
392 nursery's retained larvae across the wAP (Figure 9). The broader spread of potential
393 spawning areas coincided with a decline in locally spawned larvae in BS-NG (35%) and MB-
394 NG (34%) (Table 2), resulting in nursery populations composed of both locally derived and

395 externally supplied larvae. Despite the increase in potential spawning locations, the
396 dominant larval transport routes into each nursery remained consistent.

397

398 Connectivity between nursery grounds increased in simulations with larger embryos, with
399 both the number of inter-nursery connections and their persistence across years increasing
400 (Figure 8). This was most notable for MB-NG, where connections from GP-NG were present
401 in all three years of the 640 μm simulations (Figure 8B) compared to only one year in the 620
402 μm simulations (Figure 8A). MB-NG also received larvae from GS-NG in two of the three
403 years in the larger embryo simulations (Figure 8B). In contrast, connections from MB-NG to
404 northern nursery grounds occurred in only one of the three years (Figure 8B), and
405 connectivity between BS-NG and MB-NG did not change with embryo size (Figure 8A, B).

406 4 Discussion

407 4.1 Nursery Use and Retention

408 Bathymetry and circulation in the wAP differentially shape simulated larval residence across
409 NGs. The assumption that embryos do not survive contact with the seafloor has ruled out
410 much of the wAP as suitable spawning habitat in prior studies (Hofmann and Hüsrevoğlu,
411 2003). While this assumption is commonly modelled (Hofmann & Hüsrevoğlu 2003, Tarling
412 et al. 2007, Thorpe et al. 2019), little empirical research on krill embryos and the seafloor has
413 been conducted. Much of the habitat within the four NGs is shallower than the ~ 700 m
414 needed for the simulated 620 μm embryos to hatch above the seafloor and survive (Table 1).

415 As a result, NGs with deeper areas, such as the BS-NG, facilitated local spawning, had higher
416 larval influx and exhibited high retention rates and long retention durations. While the MB-
417 NG had less influx than the BS-MB, larvae were also locally spawned and exhibited high
418 retention rates with longer retention periods. This was likely due to a combination of the
419 bathymetric features of the Marguerite Trough and circulation dynamics that allow larvae to
420 remain in these regions after hatching. These results align with earlier findings that regions
421 exceeding 1000 m in depth provide viable habitat for larval krill development (Siegel &
422 Watkins 2016, Atkinson et al. 2019). In contrast, shallower grounds like the GS-NG and GP-
423 NG offer briefer occupancy (<21 days) potentially due to insufficient depth for localized
424 hatching combined with circulation patterns that promote faster export. Further analysis
425 isolating locally hatched drifters or examining local circulation differences could clarify the
426 drivers of these contrasting retention times.

427

428 Local spawning and prolonged residence in BS-NG and MB-NG indicate that these nurseries
429 can support early larval development, provided sufficient food availability. In contrast, GS-
430 NG and GP-NG did not host early larval stages but instead provided habitat for later-stage
431 larvae advected from other areas such as BS-NG. Nursery grounds that support later stage
432 larvae may play an important role in facilitating overwinter survival, a prerequisite for local
433 recruitment (survival of spawned krill through their first winter, Constable 2011, Meyer et
434 al. 2020). Our results show how physical transport and subsequent retention along the wAP

435 influence the distinct roles that different nursery grounds can play in supporting larval
436 development.

437

438 Climate trends, including warming, acidification, and declining sea ice (Flores et al. 2012,
439 Rogers et al. 2020), may alter these patterns, reducing habitat suitability in northern wAP
440 but possibly expanding southern habitats (Constable et al. 2014, Veytia et al. 2020, Sylvester
441 et al. 2021). Concurrent declines in adult body size (Atkinson et al. 2019) could lower
442 fecundity (Tarling et al. 2007, Siegel & Watkins 2016) and reshape nursery use. Nevertheless,
443 BS-NG remains relatively resilient thus far: historical hotspots have persisted despite regional
444 warming (Perry et al. 2019), suggesting that local bathymetry, circulation, and krill behavior
445 may collectively buffer some nursery grounds against broader environmental shifts.

446

447 Linking nursery use to larval survival offers valuable insights into krill recruitment
448 dynamics. Our simulations emphasize the role of seafloor depth and circulation in successful
449 embryonic and larval development, although unmodeled factors, such as prey availability,
450 predation, and variability in physiological behavior, may further shape outcomes.

451 Incorporating these drivers into models with higher spatial resolution will clarify how krill
452 populations respond to ongoing environmental change, especially in nearshore habitats that
453 are understudied, where early-stage overwintering may be crucial.

454 4.2 Spawning Grounds and Connectivity

455 In simulations using 620 μm embryos, successful spawning in the BS-NG and the MB-NG
456 resulted in two distinct larval populations that occasionally exchanged with other nurseries,
457 GS-NG and GP-NG, as seasonal winds and circulation modified transport paths. Interannual
458 variability allowed some larvae from BS-NG to reach GS-NG and GP-NG as secondary
459 nursery grounds, whereas MB-NG remained largely self-sustained and disconnected from
460 northern areas.

461
462 Among the four nursery grounds, BS-NG displayed the most diverse spawning origins in our
463 simulations. Across all years, a substantial portion of larvae was spawned locally, while
464 varying wind and circulation patterns occasionally delivered additional inputs from the wAP
465 shelf edge and the northwest Weddell Sea. These findings align with Perry et al. (2019) and
466 the conceptual models of Conroy et al. (2020), both of which propose that localized krill
467 productivity and wind-driven exchange shape larval distributions in slope, shelf, and coastal
468 waters.

469
470 Transport analysis revealed how wind forcing mediates interannual variability in larval
471 supply routes to the BS-NG. In summer 2016-2017, strong northerlies over the shelf break
472 near Boyd Strait resulted in an anomalous (compared to the other two years) on-shelf Ekman
473 transport that allows drifters that normally would continue to advect to the northeast in the
474 along-slope current (Moffat & Meredith 2018) to be transported into Bransfield Strait via

475 Boyd Strait. By 2018–2019, the westerlies present at the tip of the Antarctic Peninsula were
476 weaker and had a more southerly component, in agreement with previous observations by
477 Renner et al. (2012), favoring transport from the northwest Weddell Sea instead. This
478 demonstrates how local wind fields can either enhance or redirect larval inflows, ultimately
479 influencing BS-NG connectivity. Despite these shifts, final dispersal routes remained
480 consistent: export occurred north toward the South Scotia Ridge or south into GS-NG and
481 GP-NG. MB-NG stayed relatively isolated, receiving larvae mainly from the south via the
482 Marguerite Trough.

483

484 Future projections of strengthening and poleward-shifting westerlies under climate change
485 (Goyal et al. 2021) suggest that wind driven circulation dynamics may undergo significant
486 alteration to the end of the 21st century. Enhanced westerlies could reduce the frequency of
487 southerly wind-driven transport from the Weddell Sea into BS-NG, instead favoring stronger
488 along-slope currents that bypass Boyd Strait. Such changes could disrupt the balance of larval
489 supply and connectivity patterns observed in this study, with implications for krill
490 recruitment and ecosystem dynamics. Enhanced westerly-driven circulation patterns have
491 been documented near the Antarctic Peninsula (Thompson et al. 2015, Reiss et al. 2017),
492 where changing water-mass properties can reshape connectivity and recruitment (Atkinson
493 et al. 2019).

494

495 As described above, the BS-NG served as a source of later-stage furcilia dispersing
496 downstream to the GS-NG and GP-NG, often via the coastal current (Moffat et al. 2008). In
497 GS-NG, simulated larvae had already resided in the BS-NG for at least two months indicating
498 that they arrived as more advanced furcilia during late summer and fall. This passive
499 transport route aligns with observed active migration routes of adult krill, which migrate
500 inshore to the Gerlache Strait (Cleary et al. 2016) or aggregate in Bransfield Strait during
501 winter (Reiss et al. 2017). While adult krill are capable of active migration, the overlap
502 between active and passive routes highlights the importance of these transport corridors in
503 maintaining connectivity between regions. In GP-NG, simulations revealed a mixing ground
504 where larvae from BS-NG converged with those transported off-shelf from areas further
505 south toward Marguerite Bay. This convergence highlights the role of GP-NG as a junction
506 point for multiple transport pathways, increasing the diversity of larval origins in the
507 nursery.

508

509 Although our circulation model is too coarse to resolve fjords and smaller-scale bathymetric
510 features, these findings suggest that broad circulation patterns, such as the coastal current
511 connecting BS-NG to GS-NG and GP-NG, can act as consistent larval transport routes.
512 However, further modeling and observational studies are needed to evaluate the stability of
513 these transport pathways and their sensitivity to environmental changes. Notably, the
514 Gerlache Strait is a key area for winter krill fisheries (Johannessen et al. 2024). Given its
515 ecological and economic significance, higher-resolution modeling and targeted field studies

516 are essential to better understand the interplay between larval transport, nursery ground
517 connectivity, and potential anthropogenic pressures in this region.

518

519 Under smaller embryo simulations, MB-NG remained a self-sustaining nursery. Previous
520 studies show that krill populations in Marguerite Bay exhibit complex connectivity, with
521 gravid females feeding at depth (Clarke & Tyler 2008) and high larval abundances in fall
522 along the shelf break (Ashjian et al. 2008). Piñones et al. (2013) also found that Circumpolar
523 Deep-Water intrusions maintain local populations via on-shelf spawning and remote inputs
524 from the Bellingshausen Sea. Our simulations suggest that most larvae in the MB-NG are
525 spawned locally or advected from farther south on the shelf via the Marguerite Trough, and
526 do not advect northward within their first 180 days (Figure 8). This also implies that any
527 Bellingshausen Sea connection likely involves adult or juvenile migration rather than larval
528 drift. Recent research indicates partial isolation between krill populations north and south of
529 Low Island and the southern Bransfield Strait (Gallagher et al. 2023), which may explain
530 why MB-NG acts as a mostly closed system in our simulations, reinforcing the role of
531 localized circulation in modulating connectivity.

532

533 Notably, the model omits biogeochemical processes, ocean temperature, chlorophyll-a
534 concentrations, and other ecological factors known to influence krill development and
535 habitat suitability (Ross & Quetin 1989, Piñones & Fedorov 2016) and tests a particular set of
536 behavioral assumptions. Although we account for some interannual variability in circulation

537 and larval transport, we do not capture the full range of processes affecting spawning success
538 and overwintering. Future studies that incorporate wintertime surveys and additional
539 environmental variables could refine our understanding of how circulation, behavior, and
540 survival might be affected by changing climate conditions and productivity regimes.

541 4.4 Impact of Embryo Size on Spawning and Connectivity

542 The introduction of larger embryos in our simulations altered both the spatial distribution of
543 successful spawning and the connectivity among nursery grounds, emphasizing the
544 sensitivity of early-life ascent–descent cycles to poorly constrained biological parameters.
545 Compared to smaller embryos, for which viable hatching was confined primarily to deeper
546 habitats (e.g., the BS-NG and MB-NG) due to our assumption of seafloor mortality, larger
547 embryos completed development in shallower waters, reducing seafloor contact and thus
548 expanding potential spawning grounds (Hofmann et al. 1992, Capella et al. 1992).

549
550 This shift toward on-shelf hatching is consistent with earlier work suggesting that the gap
551 between hatching depth and the ocean floor is often minimal along many parts of the WAP
552 shelf (Hofmann & Hüsrevoğlu, 2003). Although our simulations did not incorporate seasonal
553 changes in embryo buoyancy or descent–ascent timing (as might occur from December to
554 March), the results raise the possibility that early-season spawning combined with larger
555 embryo sizes could facilitate on-shelf hatching under real-world conditions. Warming
556 temperatures may further accelerate embryonic development (Thorpe et al. 2019), suggesting
557 that temperature-driven variability in descent–ascent cycles could be even more pronounced

558 than shown here. Incorporating spatially and temporally varying environmental conditions
559 fields into future online ascent–descent models would help clarify these dynamics.

560

561 With larger embryos, overall connectivity increased among nursery grounds. GS-NG and
562 GP-NG, previously secondary refuges used mostly by older furcilia that had drifted from BS-
563 NG or MB-NG, became prominent nurseries supporting earlier life stages for longer periods.
564 In some years, MB-NG, which had appeared isolated under smaller-embryo scenarios,
565 exhibited greater inflows from GS-NG and GP-NG, reducing its apparent isolation. These
566 findings parallel observational evidence that larval krill can occupy broad shelf regions, even
567 while adults often aggregate nearshore in late fall (Siegel 1988, 2005, Ashjian et al. 2004).
568 That said, we considered only seafloor-interaction mortality in these simulations. Real-world
569 conditions likely impose additional constraints, such as predation, starvation, or
570 physiological stress (cf. Thorpe et al. 2019), which could dampen the increased connectivity
571 patterns observed here, especially if mortality is elevated in shallow or cold waters.

572

573 Notably, our results challenge the long-standing assumption that deep offshore spawning is
574 the primary route for successful hatching. Early studies (Hofmann et al. 1992, Capella et al.
575 1992) posited that embryos contacting the seabed were non-viable, an idea reinforced by
576 Hempel and Hempel (1986), who found deeper (>1000 m) hatching more likely based on
577 limited sampling from seven deepwater stations. However, that limited dataset did not
578 capture near-shelf conditions, and the paucity of comparable information from shallower

579 regions leaves open the possibility that on-shelf hatching is more common than historically
580 recognized. Expanding observational efforts, especially in on-shelf environments, could
581 either validate or refine these model outcomes.

582

583 By facilitating on-shelf spawning, larger embryos effectively blurred distinctions between
584 nearshore and oceanic habitats, intensifying cross-shelf transport and enabling more larvae
585 to reach inner-shelf areas. This outcome aligns with Piñones et al. (2013), whose model
586 simulations suggested significant on-shelf movement of shelf-break-spawned larvae, and
587 with Conroy et al. (2020), who found that spatial separation between younger and older
588 stages can be disrupted by variations in early-stage survival or behavior. If future conditions,
589 such as ongoing ocean warming, further favor on-shelf hatching, recruitment patterns along
590 the wAP may increasingly hinge on nearshore processes, influencing metapopulation
591 resilience and overall ecosystem dynamics.

592

593 Finally, we note that large uncertainties persist regarding embryo-size distributions,
594 buoyancy profiles, and stage-specific mortality rates under different environmental
595 conditions (Ross et al. 2004). Empirical data on variable sinking rates, hatching success, and
596 larval survival remain sparse, particularly in the shallow shelf zones of the wAP. Addressing
597 these gaps through laboratory experiments, long-term field campaigns, and multidisciplinary
598 collaborations will be vital to refine krill recruitment models (Johnston et al. 2022, Meyer et
599 al. 2020). By integrating more realistic representations of mortality related to temperature,

600 food availability, and additional sources and additional behavior strategies as necessary,
601 future modeling efforts can better capture the full complexity of krill embryonic and larval
602 connectivity, ultimately informing ecosystem-based management in the face of rapid
603 environmental change.

604 4.5 Conclusions

605 Our modeling experiments indicate that bathymetry, regional circulation, and larval traits
606 (particularly embryo size) determine connectivity between spawning areas and nursery
607 grounds along the wAP. Interannual wind variability is a key driver of larval supply
608 pathways, altering inflows from year to year. Embryo size influences on-shelf spawning
609 success and connectivity among nursery regions, producing complex patterns of larval
610 distribution, retention, and dispersal that challenge older assumptions of strictly deep-water
611 spawning.

612

613 These findings derive from idealized models that exclude temperature-dependent mortality,
614 acidification, and other biogeochemical factors, which may further moderate larval success.

615 Future work integrating finer-scale hydrodynamic modeling, laboratory data on embryo

616 buoyancy and mortality, and in situ observations, especially in poorly sampled winter or

617 nearshore habitats, will help to further refine these mechanisms. Such efforts will enhance

618 predictive capabilities for krill distribution in a changing climate and inform fisheries

619 management and conservation in this ecologically and economically important region of the

620 Southern Ocean.

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630 Data Availability

631 Model output used in this study is publicly available from the Biological and Chemical
632 Oceanography Data Management Office (BCO-DMO) and should be cited as:

633 Brooks CM, Dinniman M, Sylvester Z, Bernard KS, Thorpe S (2025) Modeled locations
634 of drifters representing transport of early life stages of Antarctic krill from 2016
635 to 2019. <https://doi.org/10.26008/1912/bco-dmo.964861.1>
636

637 Literature Cited

638 Ashjian CJ, Davis CS, Gallager SM, Wiebe PH, Lawson GL (2008) Distribution of larval krill
639 and zooplankton in association with hydrography in Marguerite Bay, Antarctic
640 Peninsula, in austral fall and winter 2001 described using the Video Plankton
641 Recorder. *Deep Sea Research Part II: Topical Studies in Oceanography* 55:455–471.
642 Ashjian CJ, Rosenwaks GA, Wiebe PH, Davis CS, Gallager SM, Copley NJ, Lawson GL,
643 Alatalo P (2004) Distribution of zooplankton on the continental shelf off Marguerite
644 Bay, Antarctic Peninsula, during Austral Fall and Winter, 2001. *Deep Sea Research*
645 *Part II: Topical Studies in Oceanography* 51:2073–2098.
646 Atkinson A, Hill SL, Pakhomov EA, Siegel V, Anadon R, Chiba S, Daly KL, Downie R,
647 Fielding S, Fretwell P, Gerrish L, Hosie GW, Jessopp MJ, Kawaguchi S, Krafft BA,
648 Loeb V, Nishikawa J, Peat HJ, Reiss CS, Ross RM, Quetin LB, Schmidt K, Steinberg

- 649 DK, Subramaniam RC, Tarling GA, Ward P (2017) KRILLBASE: a circumpolar
 650 database of Antarctic krill and salp numerical densities, 1926–2016. *Earth System*
 651 *Science Data* 9:193–210.
- 652 Atkinson A, Hill SL, Pakhomov EA, Siegel V, Reiss CS, Loeb VJ, Steinberg DK, Schmidt K,
 653 Tarling GA, Gerrish L, Salliey SF (2019) Krill (*Euphausia superba*) distribution
 654 contracts southward during rapid regional warming. *Nature Clim Change* 9:142–147.
- 655 Budgell WP (2005) Numerical simulation of ice–ocean variability in the Barents Sea region.
 656 *Ocean Dynamics* 55:370–387.
- 657 Capella JE, Quetin LB, Hofmann EE, Ross RM (1992) Models of the early life history of
 658 *Euphausia superba*—Part II. Lagrangian calculations. *Deep Sea Research Part A*
 659 *Oceanographic Research Papers* 39:1201–1220.
- 660 Clarke A, Tyler PA (2008) Adult Antarctic Krill Feeding at Abyssal Depths. *Current Biology*
 661 18:282–285.
- 662 Cleary A, Durbin E, Casas M, Zhou M (2016) Winter distribution and size structure of
 663 Antarctic krill *Euphausia superba* populations in-shore along the West Antarctic
 664 Peninsula. *Mar Ecol Prog Ser* 552:115–129.
- 665 Conroy JA, Reiss CS, Gleiber MR, Steinberg DK (2020) Linking Antarctic krill larval supply
 666 and recruitment along the Antarctic Peninsula. *Integrative and Comparative Biology*
 667 60:1386–1400.
- 668 Constable AJ (2011) Lessons from CCAMLR on the implementation of the ecosystem
 669 approach to managing fisheries: Lessons from CCAMLR on EBFM. *Fish and Fisheries*
 670 12:138–151.
- 671 Constable AJ, Melbourne-Thomas J, Corney SP, Arrigo KR, Barbraud C, Barnes DKA, Bindoff
 672 NL, Boyd PW, Brandt A, Costa DP, Davidson AT, Ducklow HW, Emmerson L,
 673 Fukuchi M, Gutt J, Hindell MA, Hofmann EE, Hosie GW, Iida T, Jacob S, Johnston
 674 NM, Kawaguchi S, Kokubun N, Koubbi P, Lea M-A, Makhado A, Massom RA,
 675 Meiners K, Meredith MP, Murphy EJ, Nicol S, Reid K, Richerson K, Riddle MJ,
 676 Rintoul SR, Smith WO, Southwell C, Stark JS, Sumner M, Swadling KM, Takahashi
 677 KT, Trathan PN, Welsford DC, Weimerskirch H, Westwood KJ, Wienecke BC, Wolf-
 678 Gladrow D, Wright SW, Xavier JC, Ziegler P (2014) Climate change and Southern
 679 Ocean ecosystems I: how changes in physical habitats directly affect marine biota.
 680 *Global Change Biology* 20:3004–3025.
- 681 Dinniman MS, Klinck JM, Bai L-S, Bromwich DH, Hines KM, Holland DM (2015) The Effect
 682 of Atmospheric Forcing Resolution on Delivery of Ocean Heat to the Antarctic
 683 Floating Ice Shelves. *Journal of Climate* 28:6067–6085.
- 684 Dinniman MS, Klinck JM, Smith WO (2011) A model study of Circumpolar Deep Water on
 685 the West Antarctic Peninsula and Ross Sea continental shelves. *Deep Sea Research*
 686 *Part II: Topical Studies in Oceanography* 58:1508–1523.
- 687 Dinniman MS, St-Laurent P, Arrigo KR, Hofmann EE, van Dijken GL (2020) Analysis of Iron
 688 Sources in Antarctic Continental Shelf Waters. *Journal of Geophysical Research:*
 689 *Oceans* 125:e2019JC015736.

- 690 Dinniman MS, St-Laurent P, Arrigo KR, Hofmann EE, van Dijken GL (2023) Sensitivity of
691 the Relationship Between Antarctic Ice Shelves and Iron Supply to Projected Changes
692 in the Atmospheric Forcing. *Journal of Geophysical Research: Oceans*
693 128:e2022JC019210.
- 694 Flores H, Atkinson A, Kawaguchi S, Krafft B, Milinevsky G, Nicol S, Reiss C, Tarling G,
695 Werner R, Bravo Rebolledo E, Cirelli V, Cuzin-Roudy J, Fielding S, van Franeker J,
696 Groeneveld J, Haraldsson M, Lombana A, Marschoff E, Meyer B, Pakhomov E, Van de
697 Putte A, Rombolá E, Schmidt K, Siegel V, Teschke M, Tonkes H, Toullec J, Trathan P,
698 Tremblay N, Werner T (2012) Impact of climate change on Antarctic krill. *Mar Ecol*
699 *Prog Ser* 458:1–19.
- 700 Fraser FC (Francis C (1936) On the development and distribution of the young stages of Krill
701 (*Euphausia superba*). *Discovery reports* 14:1–192.
- 702 Gallagher KL, Dinniman MS, Lynch HJ (2023) Quantifying Antarctic krill connectivity
703 across the West Antarctic Peninsula and its role in large-scale *Pygoscelis* penguin
704 population dynamics. *Sci Rep* 13:12072.
- 705 Galton-Fenzi BK, Porter-Smith R, Cook S, Cougnon E, Gwyther DE, Huneke WGC,
706 Rosevear MG, Asay-Davis X, Boeira Dias F, Dinniman MS, Holland D, Kusahara K,
707 Naughten KA, Nicholls KW, Pelletier C, Richter O, Seroussi HL, Timmermann R
708 (2025) Realistic ice-shelf/ocean state estimates (RISE) of Antarctic basal melting and
709 drivers. *EGUsphere*:1–27.
- 710 Goyal R, Gupta AS, Jucker M, England MH (2021) Historical and Projected Changes in the
711 Southern Hemisphere Surface Westerlies. *Geophysical Research Letters*
712 48:e2020GL090849.
- 713 Haidvogel DB, Arango H, Budgell WP, Cornuelle BD, Curchitser E, Di Lorenzo E, Fennel K,
714 Geyer WR, Hermann AJ, Lanerolle L, Levin J, McWilliams JC, Miller AJ, Moore AM,
715 Powell TM, Shchepetkin AF, Sherwood CR, Signell RP, Warner JC, Wilkin J (2008)
716 Ocean forecasting in terrain-following coordinates: Formulation and skill assessment
717 of the Regional Ocean Modeling System. *Journal of Computational Physics* 227:3595–
718 3624.
- 719 Häkkinen S, Mellor GL (1992) Modeling the seasonal variability of a coupled Arctic ice-
720 ocean system. *J Geophys Res* 97:20285.
- 721 Hempel I (1979) Vertical distribution of eggs and nauplii of krill (*Euphausia superba*) south
722 of Elephant Island. *Meeresforschung* 27:119–123.
- 723 Hempel I (1985) Vertical Distribution of Larvae of Antarctic Krill, *Euphausia superba*. In:
724 *Antarctic Nutrient Cycles and Food Webs*. Siegfried WR, Condy PR, Laws RM (eds)
725 Springer Berlin Heidelberg, Berlin, Heidelberg, p 308–310
- 726 Hempel I, Hempel G (1986) Field observations on the developmental ascent of larval
727 *Euphausia superba* (Crustacea). *Polar Biol* 6:121–126.
- 728 Hersbach H, Bell B, Berrisford P, Hirahara S, Horányi A, Muñoz-Sabater J, Nicolas J, Peubey
729 C, Radu R, Schepers D, Simmons A, Soci C, Abdalla S, Abellan X, Balsamo G,
730 Bechtold P, Biavati G, Bidlot J, Bonavita M, Chiara G, Dahlgren P, Dee D,

- 731 Diamantakis M, Dragani R, Flemming J, Forbes R, Fuentes M, Geer A, Haimberger L,
 732 Healy S, Hogan RJ, Hólmi E, Janisková M, Keeley S, Laloyaux P, Lopez P, Lupu C,
 733 Radnoti G, Rosnay P, Rozum I, Vamborg F, Villaume S, Thépat J (2020) The ERA5
 734 global reanalysis. *QJR Meteorol Soc* 146:1999–2049.
- 735 Hofmann E, Capella JE, Ross RM, Quetin LB (1992) Models of the early life history of
 736 *Euphausia superba*—Part I. Time and temperature dependence during the descent-
 737 ascent cycle. *Deep Sea Research Part A Oceanographic Research Papers* 39:1177–
 738 1200.
- 739 Hofmann E, Hüsrevoğlu Y (2003) A circumpolar modeling study of habitat control of
 740 Antarctic krill (*Euphausia superba*) reproductive success. *Deep Sea Research Part II:*
 741 *Topical Studies in Oceanography* 50:3121–3142.
- 742 Hofmann E, Lascara C (2000) Modeling the growth dynamics of Antarctic krill *Euphausia*
 743 *superba*. *Mar Ecol Prog Ser* 194:219–231.
- 744 Holland DM, Jenkins A (1999) Modeling Thermodynamic Ice–Ocean Interactions at the Base
 745 of an Ice Shelf. *Journal of Physical Oceanography* 29:1787–1800.
- 746 Hudson K, Oliver MJ, Kohut J, Cohen JH, Dinniman MS, Klinck JM, Reiss CS, Cutter GR,
 747 Statscewich H, Bernard KS, Fraser W (2022) Subsurface Eddy Facilitates Retention of
 748 Simulated Diel Vertical Migrators in a Biological Hotspot. *JGR Oceans* 127.
- 749 Hunke EC (2001) Viscous–Plastic Sea Ice Dynamics with the EVP Model: Linearization
 750 Issues. *Journal of Computational Physics* 170:18–38.
- 751 Hunke EC, Dukowicz JK (1997) An Elastic–Viscous–Plastic Model for Sea Ice Dynamics.
 752 *Journal of Physical Oceanography* 27:1849–1867.
- 753 Hunter JR, Craig PD, Phillips HE (1993) On the use of random walk models with spatially
 754 variable diffusivity. *Journal of Computational Physics* 106:366–376.
- 755 Ikeda T (1984) Development of the larvae of the Antarctic krill (*Euphausia superba* Dana)
 756 observed in the laboratory. *Journal of Experimental Marine Biology and Ecology*
 757 75:107–117.
- 758 Johannessen ED, Krafft BA, Donovan CR, Wiff R, Caneco B, Lowther A (2024) Sensitivity of
 759 the stock assessment for the Antarctic krill fishery to time-varying natural and fishing
 760 mortality. *Fisheries Management and Ecology* 31:e12721.
- 761 Johnston NM, Murphy EJ, Atkinson A, Constable AJ, Cotté C, Cox M, Daly KL, Driscoll R,
 762 Flores H, Halfter S, Henschke N, Hill SL, Höfer J, Hunt BPV, Kawaguchi S, Lindsay D,
 763 Liszka C, Loeb V, Manno C, Meyer B, Pakhomov EA, Pinkerton MH, Reiss CS,
 764 Richerson K, Jr. WOS, Steinberg DK, Swadling KM, Tarling GA, Thorpe SE, Veytia D,
 765 Ward P, Weldrick CK, Yang G (2022) Status, Change, and Futures of Zooplankton in
 766 the Southern Ocean. *Frontiers in Ecology and Evolution* 9.
- 767 Kawaguchi S (2016) Reproduction and Larval Development in Antarctic Krill (*Euphausia*
 768 *superba*). In: *Biology and Ecology of Antarctic Krill*. Advances in Polar Ecology,
 769 Siegel V (ed) Springer International Publishing, Cham, p 225–246

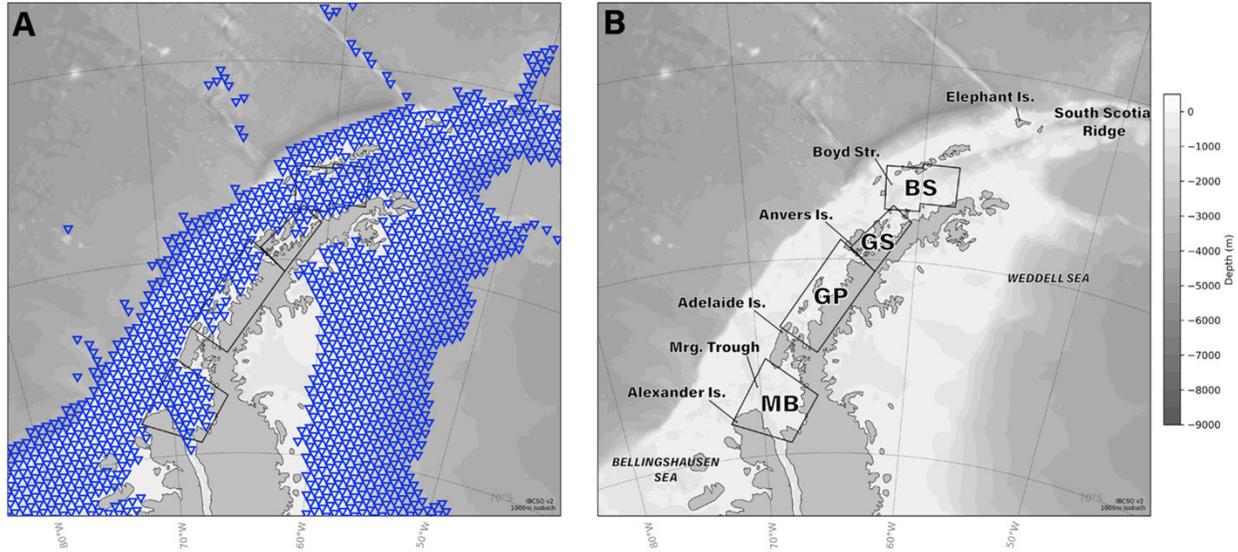
- 770 Kawaguchi S, Atkinson A, Bahlburg D, Bernard KS, Cavan EL, Cox MJ, Hill SL, Meyer B,
 771 Veytia D (2024) Climate change impacts on Antarctic krill behaviour and population
 772 dynamics. *Nat Rev Earth Environ* 5:43–58.
- 773 Kohlbach D, Lange BA, Schaafsma FL, David C, Vortkamp M, Graeve M, van Franeker JA,
 774 Krumpfen T, Flores H (2017) Ice Algae-Produced Carbon Is Critical for Overwintering
 775 of Antarctic Krill *Euphausia superba*. *Front Mar Sci* 4.
- 776 Lascara CM, Hofmann EE, Ross RM, Quetin LB (1999) Seasonal variability in the distribution
 777 of Antarctic krill, *Euphausia superba*, west of the Antarctic Peninsula. *Deep Sea*
 778 *Research Part I: Oceanographic Research Papers* 46:951–984.
- 779 Loeb VJ, Santora JA (2015) Climate variability and spatiotemporal dynamics of five Southern
 780 Ocean krill species. *Progress in Oceanography* 134:93–122.
- 781 Makarov RR (1982) Ontogenetic migrations in Antarctic euphausiids: “The Antarctic”.
- 782 Marr J (1962) *The Natural History and Geography of the Antarctic Krill (Euphausia superba*
 783 *Dana)*. *Discovery Reports* 32, pp. 33–464. Cambridge: University Press 1962. 10, – £.
- 784 *Int Revue ges Hydrobiol Hydrogr* 32:33–637.
- 785 Marschall HP (1983) Sinking speed, density and size of Euphausiid eggs. *Meeresforsch* 30:1–
 786 9.
- 787 Marschall H-P, Hirche H-J (1984) Development of eggs and nauplii of *Euphausia superba*.
 788 *Polar Biol* 2:245–250.
- 789 Mellor GL, Kantha L (1989) An ice-ocean coupled model. *Journal of Geophysical Research:*
 790 *Oceans* 94:10937–10954.
- 791 Meredith MP, King JC (2005) Rapid climate change in the ocean west of the Antarctic
 792 Peninsula during the second half of the 20th century: RAPID OCEAN CLIMATE
 793 CHANGE AT THE WAP. *Geophys Res Lett* 32:n/a-n/a.
- 794 Meyer B, Atkinson A, Bernard KS, Brierley AS, Driscoll R, Hill SL, Marschoff E, Maschette
 795 D, Perry FA, Reiss CS, Rombolá E, Tarling GA, Thorpe SE, Trathan PN, Zhu G,
 796 Kawaguchi S (2020) Successful ecosystem-based management of Antarctic krill should
 797 address uncertainties in krill recruitment, behaviour and ecological adaptation.
 798 *Communications Earth & Environment* 1:1–12.
- 799 Meyer B, Freier U, Grimm V, Groeneveld J, Hunt BPV, Kerwath S, King R, Klaas C,
 800 Pakhomov E, Meiners KM, Melbourne-Thomas J, Murphy EJ, Thorpe SE,
 801 Stammerjohn S, Wolf-Gladrow D, Auerswald L, Götz A, Halbach L, Jarman S,
 802 Kawaguchi S, Krumpfen T, Nehrke G, Ricker R, Sumner M, Teschke M, Trebilco R,
 803 Yilmaz NI (2017) The winter pack-ice zone provides a sheltered but food-poor habitat
 804 for larval Antarctic krill. *Nature Ecology & Evolution* 1:1853–1861.
- 805 Millan R, Rignot E, Bernier V, Morlighem M, Dutrieux P (2017) Bathymetry of the
 806 Amundsen Sea Embayment sector of West Antarctica from Operation IceBridge
 807 gravity and other data. *Geophys Res Lett* 44:1360–1368.
- 808 Moffat C, Beardsley RC, Owens B, van Lipzig N (2008) A first description of the Antarctic
 809 Peninsula Coastal Current. *Deep Sea Research Part II: Topical Studies in*
 810 *Oceanography* 55:277–293.

- 811 Moffat C, Meredith M (2018) Shelf–ocean exchange and hydrography west of the Antarctic
812 Peninsula: a review. *Philosophical Transactions of the Royal Society A: Mathematical,*
813 *Physical and Engineering Sciences* 376:20170164.
- 814 Nicol S, Foster J (2016) The Fishery for Antarctic Krill: Its Current Status and Management
815 Regime. In: *Biology and Ecology of Antarctic Krill*. Advances in Polar Ecology, Siegel
816 V (ed) Springer International Publishing, Cham, p 387–421
- 817 Nicol S, Foster J, Kawaguchi S (2012) The fishery for Antarctic krill - recent developments:
818 Krill fishery review. *Fish and Fisheries* 13:30–40.
- 819 Perry FA, Atkinson A, Sailley SF, Tarling GA, Hill SL, Lucas CH, Mayor DJ (2019) Habitat
820 partitioning in Antarctic krill: Spawning hotspots and nursery areas. *PLOS ONE*
821 14:e0219325.
- 822 Piñones A, Fedorov AV (2016) Projected changes of Antarctic krill habitat by the end of the
823 21st century: Changes in Antarctic Krill Habitat. *Geophysical Research Letters*
824 43:8580–8589.
- 825 Piñones A, Hofmann E, Daly K, Dinniman M, Klinck J (2013) Modeling the remote and local
826 connectivity of Antarctic krill populations along the western Antarctic Peninsula.
827 *Mar Ecol Prog Ser* 481:69–92.
- 828 Quetin LB, Ross RM (1984) Depth distribution of developing *Euphausia superba* embryos,
829 predicted from sinking rates. *Mar Biol* 79:47–53.
- 830 Quetin LB, Ross RM (2001) Environmental Variability and Its Impact on the Reproductive
831 Cycle of Antarctic Krill. *Am Zool* 41:74–89.
- 832 Quetin LB, Ross RM (2003) Episodic recruitment in Antarctic krill *Euphausia superba* in the
833 Palmer LTER study region. *Mar Ecol Prog Ser* 259:185–200.
- 834 Quetin LB, Ross RM, Clarke A (1994) Krill energetics: seasonal and environmental aspects of
835 the physiology of *Euphausia superba*.
- 836 Reiss CS (2016) Age, Growth, Mortality, and Recruitment of Antarctic Krill, *Euphausia*
837 *superba*. In: *Biology and Ecology of Antarctic Krill*. Advances in Polar Ecology, Siegel
838 V (ed) Springer International Publishing, Cham, p 101–144
- 839 Reiss CS, Cossio A, Santora JA, Dietrich KS, Murray A, Mitchell BG, Walsh J, Weiss EL,
840 Gimpel C, Jones CD, Watters GM (2017) Overwinter habitat selection by Antarctic
841 krill under varying sea-ice conditions: implications for top predators and fishery
842 management. *Marine Ecology Progress Series* 568:1–16.
- 843 Renner AHH, Thorpe SE, Heywood KJ, Murphy EJ, Watkins JL, Meredith MP (2012)
844 Advective pathways near the tip of the Antarctic Peninsula: Trends, variability and
845 ecosystem implications. *Deep Sea Research Part I: Oceanographic Research Papers*
846 63:91–101.
- 847 Rogers AD, Frinault BAV, Barnes DKA, Bindoff NL, Downie R, Ducklow HW, Friedlaender
848 AS, Hart T, Hill SL, Hofmann EE, Linse K, McMahan CR, Murphy EJ, Pakhomov EA,
849 Reygondeau G, Staniland IJ, Wolf-Gladrow DA, Wright R (2020) Antarctic Futures:
850 An Assessment of Climate-Driven Changes in Ecosystem Structure, Function, and

- 851 Service Provisioning in the Southern Ocean. Annual Review of Marine Science
852 12:null.
- 853 Ross RM, Quetin LB (1989) Energetic cost to develop to the first feeding stage of *Euphausia*
854 *superba* Dana and the effect of delays in food availability. Journal of Experimental
855 Marine Biology and Ecology 133:103–127.
- 856 Ross RM, Quetin LB, Newberger T, Oakes SA (2004) Growth and behavior of larval krill
857 (*Euphausia superba*) under the ice in late winter 2001 west of the Antarctic
858 Peninsula. Deep Sea Research Part II: Topical Studies in Oceanography 51:2169–2184.
- 859 Schaafsma FL, Kohlbach D, David C, Lange BA, Graeve M, Flores H, Franeker JA van (2017)
860 Spatio-temporal variability in the winter diet of larval and juvenile Antarctic krill,
861 *Euphausia superba*, in ice-covered waters. Marine Ecology Progress Series 580:101–
862 115.
- 863 Schaffer J, Timmermann R, Arndt JE, Kristensen SS, Mayer C, Morlighem M, Steinhage D
864 (2016) A global, high-resolution data set of ice sheet topography, cavity geometry,
865 and ocean bathymetry. Earth Syst Sci Data 8:543–557.
- 866 Schnack SB, Marschall S, Mizdalski E (1985) On the distribution of copepods and larvae of
867 *Euphausia superba* in Antarctic waters during February 1982. Meeresforschung
868 30:251–263.
- 869 Shchepetkin AF, McWilliams JC (2009) Correction and commentary for “Ocean forecasting
870 in terrain-following coordinates: Formulation and skill assessment of the regional
871 ocean modeling system” by Haidvogel et al., J. Comp. Phys. 227, pp. 3595–3624.
872 Journal of Computational Physics 228:8985–9000.
- 873 Siegel V (1988) A Concept of Seasonal Variation of Krill (*Euphausia superba*) Distribution
874 and Abundance West of the Antarctic Peninsula. In: *Antarctic Ocean and Resources*
875 *Variability*. Sahrhage D (ed) Springer Berlin Heidelberg, Berlin, Heidelberg, p 219–
876 230
- 877 Siegel V (2005) Distribution and population dynamics of *Euphausia superba*: summary of
878 recent findings. Polar Biol 29:1–22.
- 879 Siegel V, Watkins JL (2016) Distribution, Biomass and Demography of Antarctic Krill,
880 *Euphausia superba*. In: *Biology and Ecology of Antarctic Krill*. Advances in Polar
881 Ecology, Siegel V (ed) Springer International Publishing, Cham, p 21–100
- 882 Stammerjohn SE, Maksym T, Massom RA, Lowry KE, Arrigo KR, Yuan X, Raphael M,
883 Randall-Goodwin E, Sherrell RM, Yager PL (2015) Seasonal sea ice changes in the
884 Amundsen Sea, Antarctica, over the period of 1979–2014. Elementa: Science of the
885 Anthropocene 3:000055.
- 886 Steele M, Mellor GL, McPhee MG (1989) Role of the Molecular Sublayer in the Melting or
887 Freezing of Sea Ice. J Phys Oceanogr 19:139–147.
- 888 Sylvester ZT, Long MC, Brooks CM (2021) Detecting Climate Signals in Southern Ocean
889 Krill Growth Habitat. Front Mar Sci 0.

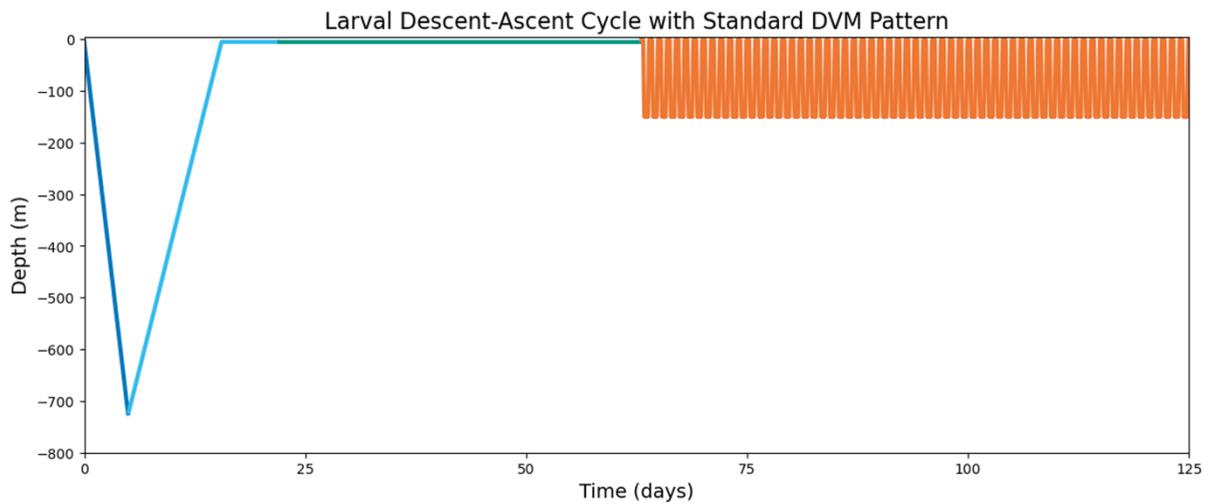
- 890 Tarling GA (2020) Routine metabolism of Antarctic krill (*Euphausia superba*) in South
891 Georgia waters: absence of metabolic compensation at its range edge. *Mar Biol*
892 167:108.
- 893 Tarling GA, Cuzin-Roudy J, Thorpe SE, Shreeve RS, Ward P, Murphy EJ (2007) Recruitment
894 of Antarctic krill *Euphausia superba* in the South Georgia region: adult fecundity and
895 the fate of larvae. *Marine Ecology Progress Series* 331:161–179.
- 896 Thompson DWJ, Barnes EA, Deser C, Foust WE, Phillips AS (2015) Quantifying the Role of
897 Internal Climate Variability in Future Climate Trends. *Journal of Climate* 28:6443–
898 6456.
- 899 Thorpe S, Tarling G, Murphy E (2019) Circumpolar patterns in Antarctic krill larval
900 recruitment: an environmentally driven model. *Marine Ecology Progress Series*
901 613:77–96.
- 902 Thorpe SE, Murphy EJ, Watkins JL (2007) Circumpolar connections between Antarctic krill
903 (*Euphausia superba* Dana) populations: Investigating the roles of ocean and sea ice
904 transport. *Deep Sea Research Part I: Oceanographic Research Papers* 54:792–810.
- 905 Trathan PN, Hill SL (2016) The Importance of Krill Predation in the Southern Ocean. In:
906 *Biology and Ecology of Antarctic Krill*. Advances in Polar Ecology, Siegel V (ed)
907 Springer International Publishing, Cham, p 321–350
- 908 Turner J, Holmes C, Caton Harrison T, Phillips T, Jena B, Reeves-Francois T, Fogt R, Thomas
909 ER, Bajish CC (2022) Record Low Antarctic Sea Ice Cover in February 2022.
910 *Geophysical Research Letters* 49:e2022GL098904.
- 911 Turner J, Maksym T, Phillips T, Marshall GJ, Meredith MP (2013) The impact of changes in
912 sea ice advance on the large winter warming on the western Antarctic Peninsula. *Int J*
913 *Climatol* 33:852–861.
- 914 Vaughan DG, Marshall GJ, Connolley WM, Parkinson C, Mulvaney R, Hodgson DA, King
915 JC, Pudsey CJ, Turner J (2003) Recent rapid regional climate warming on the
916 Antarctic Peninsula. *Climatic Change* 60:243–274.
- 917 Veytia D, Corney S, Meiners KM, Kawaguchi S, Murphy EJ, Bestley S (2020) Circumpolar
918 projections of Antarctic krill growth potential. *Nature Climate Change*:1–8.
- 919 Visser A (1997) Using random walk models to simulate the vertical distribution of particles
920 in a turbulent water column. *Mar Ecol Prog Ser* 158:275–281.
- 921 Young EF, Thorpe SE, Renner AHH, Murphy EJ (2023) Environmental and behavioural
922 drivers of Antarctic krill distribution at the South Orkney Islands: A regional
923 perspective. *Journal of Marine Systems*:103920.
- 924

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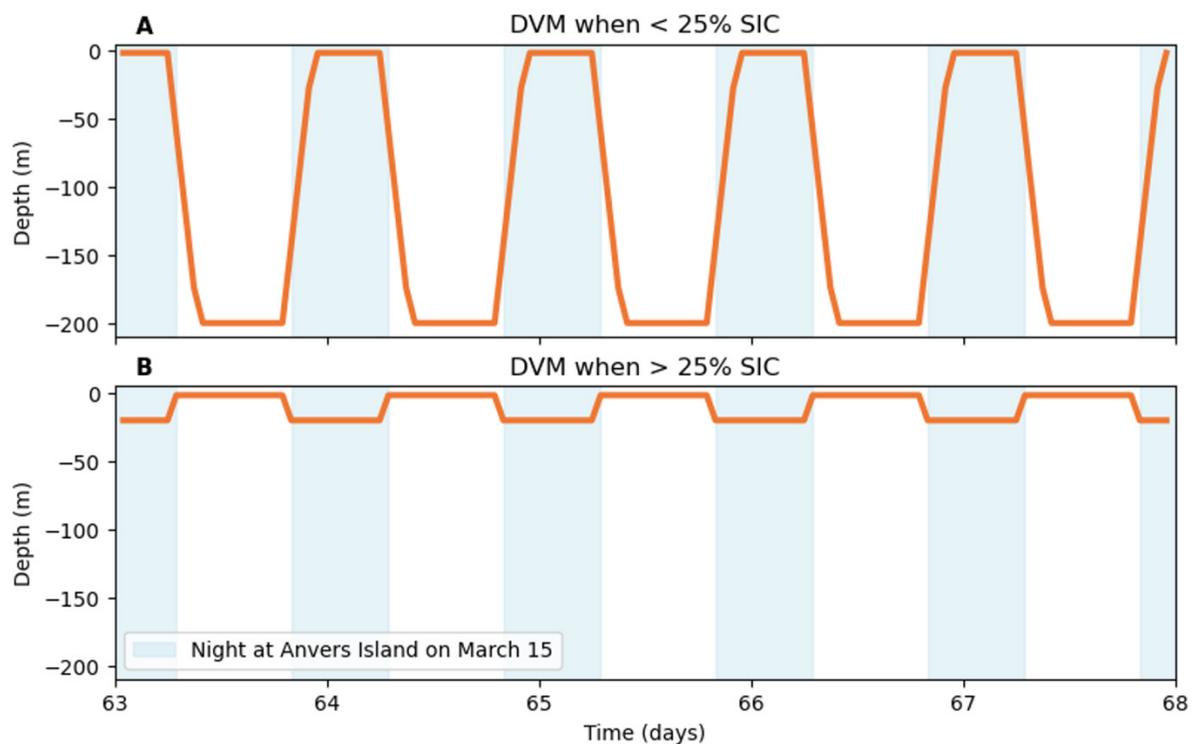
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927 **Figure 1:** Model Scope and Analysis Reference Region – Panel A identifies the starting
 928 locations for all drifters across all scenarios as blue triangles representing all possible
 929 spawning locations within the model. Panel B illustrates the four identified nursery grounds:
 930 Bransfield Strait [BS-NG], Gerlache Strait [GS-NG], Grandidier Passage [GP-NG], and
 931 Marguerite Bay [MB-NG] (see supplementary material for coordinates). Geographic features
 932 are labeled (Mrg. Trough - Marguerite Trough).



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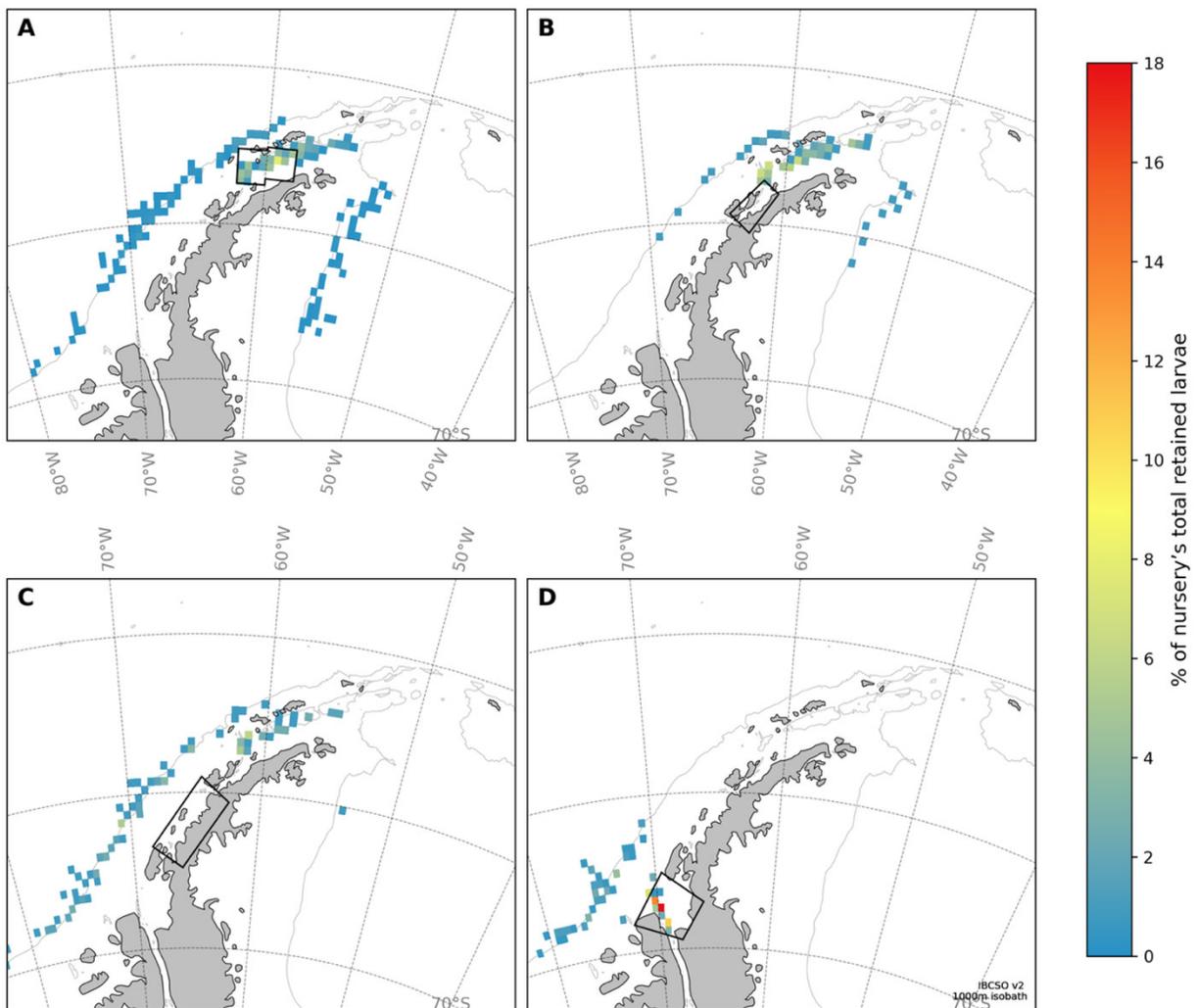
934 **Figure 2:** Example vertical behavior of a Lagrangian particle representing the early life stages
 935 of Antarctic krill, shown here for an embryo with initial diameter of 620 μm . Life stages in
 936 the circulation model are defined by elapsed time since embryo release (Hempel, 1979,
 937 Ikeda, 1984, Jia et al. 2014, Makarov, 1982). Stages are color coded to illustrate the vertical
 938 behavior associated with each phase of development. The schematic depicts the offline-
 939 parameterized descent of the embryo in blue, hatching at depth, and subsequent ascent of
 940 the larva (nauplii stages: cyan). Early-stage behavior differs between 620 μm and 640 μm
 941 embryos (see Table 1), affecting the vertical range and timing of development prior to the
 942 onset of diel vertical migration (calyptope stages: green). Diel vertical migration begins after
 943 63 days of development (furcilia stages: orange).



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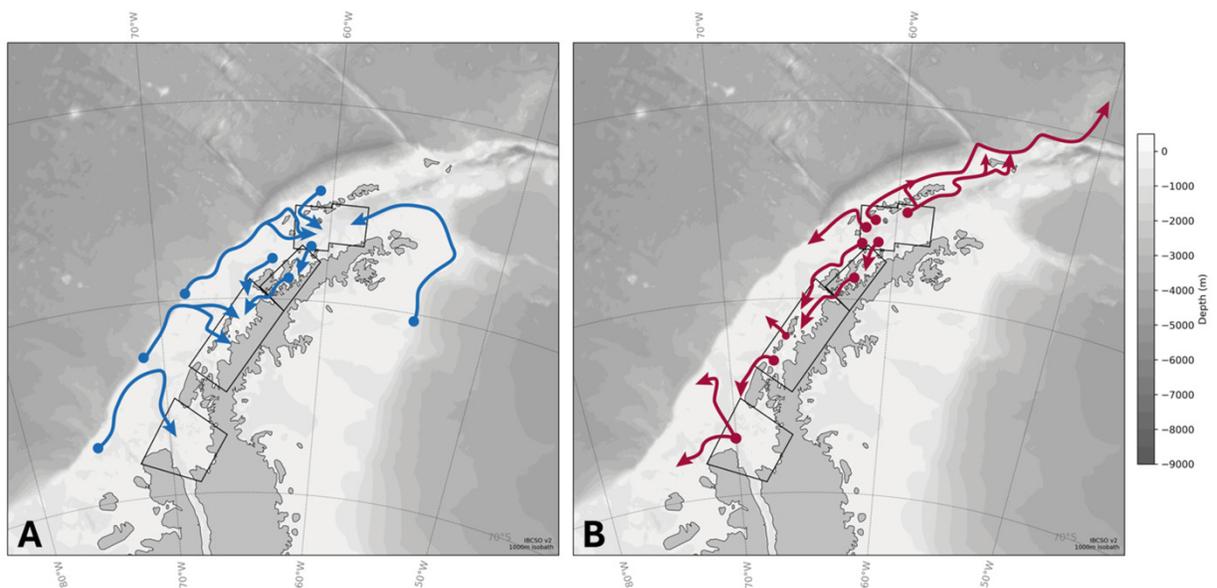
945 **Figure 3.** DVM behavior applied to particles beginning on day 63, once larvae reach the
 946 furcilia stage. Panels A and B depict a cross-sectional example of vertical position at Anvers
 947 Island, showing differing behavior based on local sea ice concentration. Panel A shows
 948 standard DVM in open water (<25% SIC), where particles descend during daylight and
 949 ascend at night. Panel B shows reversed DVM under sea ice (>25% SIC), with nighttime
 950 descent and daytime ascent.

951

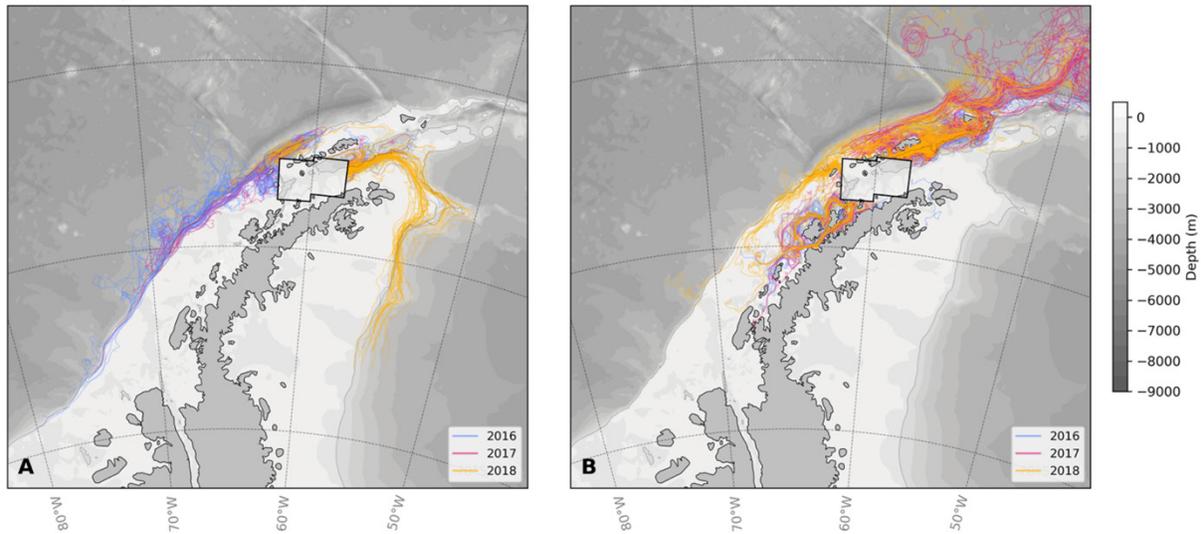


952

953 **Figure 4:** Composite maps of spawning locations of larval krill (620 μm initial embryo size)
 954 retained in each nursery ground, aggregated across 2016–2018. Each panel shows the relative
 955 spatial distribution of particle release locations that led to successful retention in the
 956 respective target nursery, expressed as a percentage of that nursery’s total retained larvae.
 957 The color scale is shared across panels to enable direct comparison of origin concentration.
 958 Total particle counts per nursery (N) are listed in Table 2.

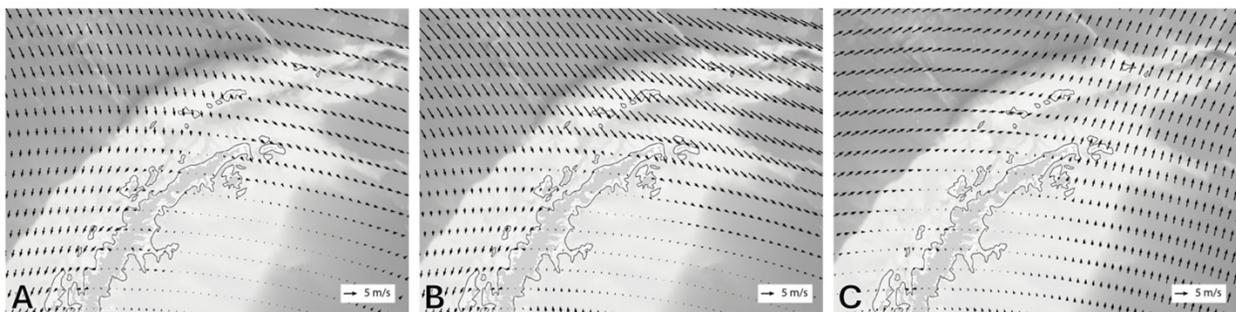


959
 960 **Figure 5:** Panels A and B provide a summarization of the trajectories of larvae (620 μm
 961 embryos) into and out of all nursery grounds (NGs). Panel A summarizes the primary
 962 trajectories that advected larvae into NGs while Panel B summarizes the primary trajectories
 963 larvae took after exiting an NG.



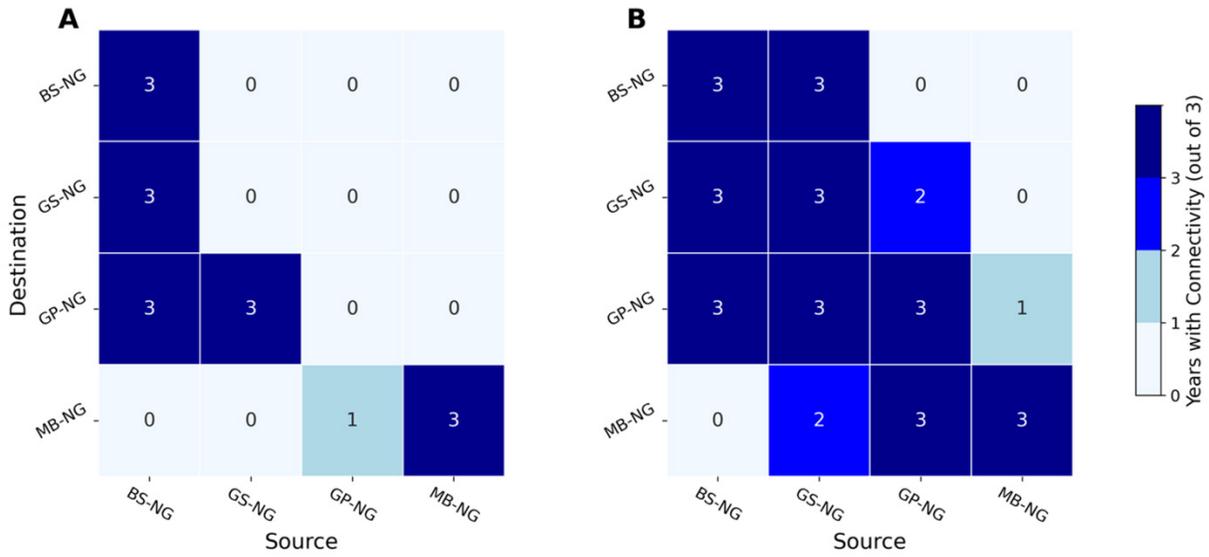
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965 **Figure 6:** Larval krill trajectories for the Bransfield Strait nursery ground (BS-NG) across
 966 three consecutive summers (2016–2019). Panel A shows particle trajectories into BS-NG for
 967 larvae not spawned locally. panel B shows trajectories of particles after exiting from BS-NG.
 968 Pathways are colored by simulation start year: 2016 (blue), 2017 (magenta), and 2018 (gold).
 969 Background shading represents seafloor bathymetry (IBCSO v2).



970

971 **Figure 7:** Variability in local near surface (10 m) winds for the month of February are shown
 972 for each year: A) February 2017, B) February 2018, C), February 2019. Based on ERA5
 973 reanalysis data used to force the ocean circulation model. Size of arrows represents wind
 974 speed (see scale arrows on each panel).



975

976 **Figure 8:** Connectivity matrices of larvae retained by nursery grounds across simulations for

977 (A) 620 µm embryos and (B) 640 µm embryos. Each heatmap shows the number of distinct

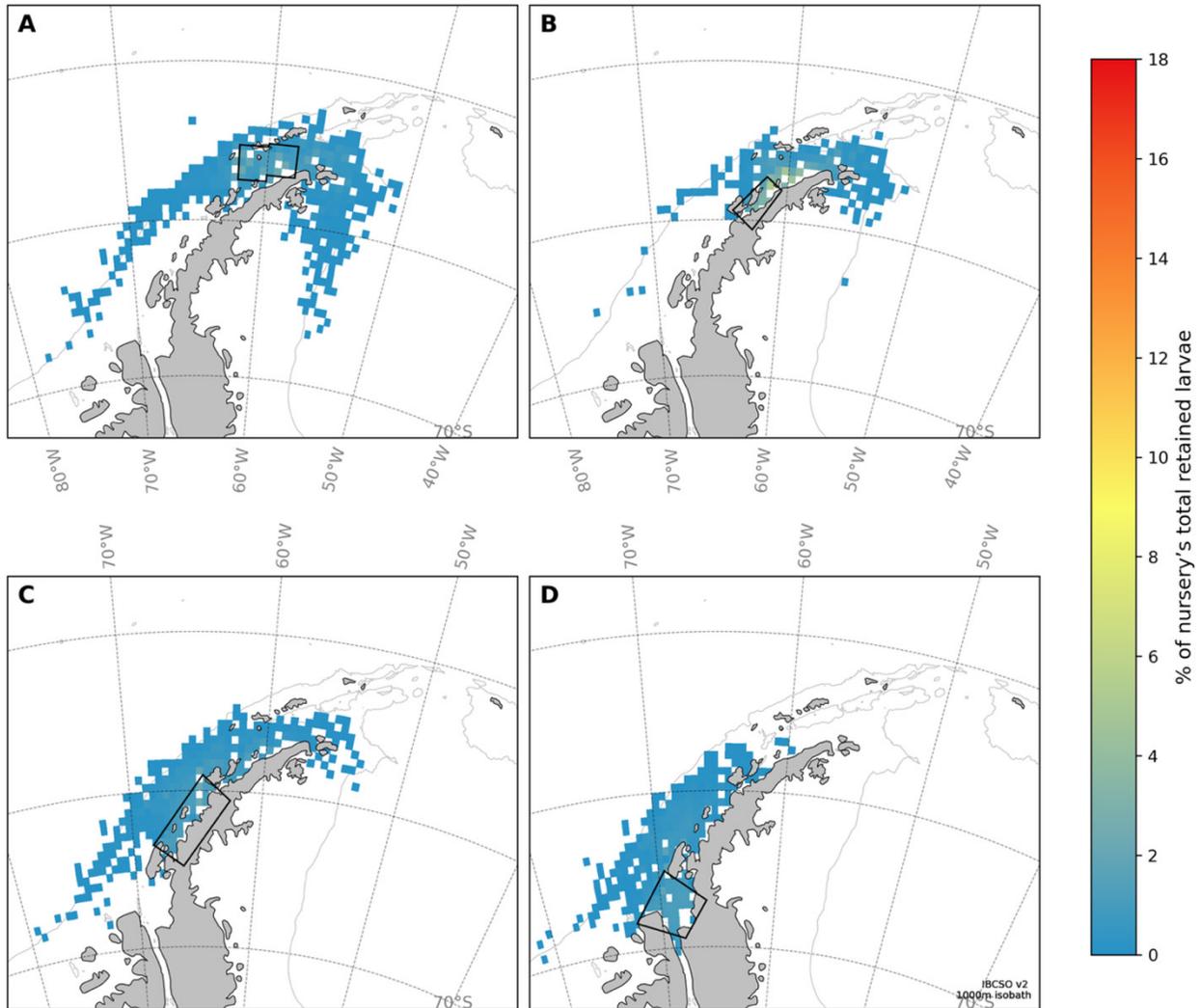
978 simulation years (out of 3) in which embryos were spawned in a nursery (source region:

979 columns), before being retained for at least 21 days in the destination NG (rows). Diagonal

980 elements represent self-retention (within the same region), while off-diagonal elements

981 indicate inter-regional connectivity.

982



983

984 **Figure 9.** Composite spawning origin maps for larval krill (640 μm initial embryo size)
 985 retained in each nursery ground, aggregated across 2016–2018. Each panel shows the relative
 986 spatial distribution of particle release locations that led to successful retention in the
 987 respective target nursery, expressed as a percentage of that nursery's total retained larvae.
 988 Color indicates the fraction of all contributing particles for that nursery that originated in
 989 each spatial bin. Values are unitless and reflect relative contributions, not physical densities

990 or temporally resolved information. The color scale is shared across panels to enable direct

991 comparison of origin concentration. Total particle counts per nursery are listed in Table 2.

992

993 **Table 1.** Model parameters calculated from the offline descent-ascent model of Hofmann et
 994 al. (1992) and the corresponding values used in the online drifter simulations. Parameters
 995 were calculated for three locations and for five months ($N = 15$) for embryos with initial
 996 diameters of 620 μm and 640 μm . The parameters include the mean sinking rate of the
 997 embryos, their period of descent (*hatching time for 620 μm embryos), mean ascent rate
 998 (representing larval swimming for 620 μm embryos and buoyant ascent for 640 μm embryos),
 999 and development time to furcilia I (F1). Variability in offline parameters is reported as the
 1000 mean \pm standard deviation. Circulation model inputs represent generalized values applied
 1001 across all drifters to simulate the descent–ascent cycle and the onset of diel vertical migration
 1002 (DVM). The circulation model hatching depth was not explicitly parameterized.

Range of Tested Parameters	620 μm	640 μm
Mean descent rate (m day^{-1})	132.7 \pm 1.6	82.1 \pm 7.0
Period of descent* (hours)	127.5 \pm 15.6	32.2 \pm 7.5
Mean ascent rate (m day^{-1})	60.1 \pm 6.6	176.2 \pm 6.2
Development time to furcilia I stage (days)	65.5 \pm 13.9	65.5 \pm 13.9
Input parameters for circulation model	620 μm	640 μm
Descent rate (m day^{-1})	145	80
Period of descent (hours)	120	36
Ascent rate (m day^{-1})	68.2	180
Elapsed time from embryo release until start of DVM (days)	63	63

1003
 1004 **Table 2** Larval krill use of nursery grounds by initial embryo size (Em Size), averaged across
 1005 three start years (2016–2018), with annual metrics reported for each nursery ground (NG).
 1006 Larval influx is the mean number of larvae that entered the NG. Retention duration is shown

1007 as the median retention duration and interquartile range (IQR). Larvae retained is the mean
 1008 the number of larvae retained for 21 days or longer. Retention rate is the mean percent of the
 1009 larval influx that was retained. The localized spawning percentage shows the mean
 1010 proportion of retained larvae inferred to have originated (spawned) locally within the
 1011 nursery region. Interannual data are in Supplementary Table 1.

Em Size (μm)	NG	Larval Influx (N)	Median retention duration and IQR (days)	Retained Larvae (N)	Retention Rate %	Localized spawning %
620	BS	316	48 (61)	237	74%	54%
	GS	96	11 (30)	35	37%	NA
	GP	84	20 (33)	39	46%	NA
	MB	75	50 (67)	56	74%	64%
640	BS	1191	41 (61)	833	70%	35 %
	GS	574	25 (42)	299	52%	26%
	GP	1288	44 (80)	895	69%	34%
	MB	1171	58 (71)	960	82%	34%

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