



#### **OPEN ACCESS**

EDITED BY

Michael Lehning, Swiss Federal Institute of Technology Lausanne, Switzerland

REVIEWED BY

Frans Berkhout, King's College London, United Kingdom Alan Robock, Rutgers University, United States

\*CORRESPONDENCE

Martin Siegert

RECEIVED 13 November 2024 ACCEPTED 16 July 2025 PUBLISHED 09 September 2025

#### CITATION

Siegert M, Sevestre H, Bentley MJ, Brigham-Grette J, Burgess H, Buzzard S, Cavitte M, Chown SL, Colleoni F, DeConto RM, Fricker HA, Gasson E, Grant SM, Gulisano AM, Hancock S, Hendry KR, Henley SF, Hock R, Hughes KA, Karentz D, Kirkham JD, Kulessa B, Larter RD, Mackintosh A, Masson-Delmotte V, McCormack FS, Millman H, Mottram R, Moon TA, Naish T, Nath C, Orlove B, Pearson P, Rogelj J, Rumble J, Seabrook S, Silvano A, Sommerkorn M, Stearns LA, Stokes CR, Stroeve J and Truffer M. Safeguarding the polar regions from dangerous geoengineering: a critical assessment of proposed concepts and future prospects. Front Sci (2025) 3:1527393. doi: 10.3389/fsci.2025.1527393

#### COPYRIGHT

© 2025 Siegert, Sevestre, Bentley, Brigham-Grette, Burgess, Buzzard, Cavitte, Chown, Colleoni, DeConto, Fricker, Gasson, Grant, Gulisano, Hancock, Hendry, Henley, Hock, Hughes, Karentz, Kirkham, Kulessa, Larter, Mackintosh, Masson-Delmotte, McCormack, Millman, Mottram, Moon, Naish, Nath, Orlove, Pearson, Rogelj, Rumble, Seabrook, Silvano, Sommerkorn, Stearns, Stokes, Stroeve and Truffer. This is an openaccess article distributed under the terms of the Creative Commons Attribution License (CC BY). The use, distribution or reproduction in other forums is permitted, provided the original author(s) and the copyright owner(s) are credited and that the original publication in this journal is cited, in accordance with accepted academic practice. No use, distribution or reproduction is permitted which does not comply with these terms.

# Safeguarding the polar regions from dangerous geoengineering: a critical assessment of proposed concepts and future prospects

Martin Siegert<sup>1\*</sup>, Heïdi Sevestre<sup>2</sup>, Michael J. Bentley<sup>3</sup>,
Julie Brigham-Grette<sup>4</sup>, Henry Burgess<sup>5</sup>, Sammie Buzzard<sup>6</sup>,
Marie Cavitte<sup>7</sup>, Steven L. Chown<sup>8</sup>, Florence Colleoni<sup>9</sup>,
Robert M. DeConto<sup>4</sup>, Helen Amanda Fricker<sup>10</sup>,
Edward Gasson<sup>1,11</sup>, Susie M. Grant<sup>12</sup>, Adriana Maria Gulisano<sup>13,14,15</sup>,
Susana Hancock<sup>16</sup>, Katharine R. Hendry<sup>12</sup>, Sian F. Henley<sup>17</sup>,
Regine Hock<sup>18,19</sup>, Kevin A. Hughes<sup>12</sup>, Deneb Karentz<sup>20</sup>,
James D. Kirkham<sup>12,16</sup>, Bernd Kulessa<sup>21</sup>, Robert D. Larter<sup>12</sup>,
Andrew Mackintosh<sup>8</sup>, Valérie Masson-Delmotte<sup>22</sup>,
Felicity S. McCormack<sup>8</sup>, Helen Millman<sup>23</sup>, Ruth Mottram<sup>24</sup>,
Twila A. Moon<sup>25</sup>, Tim Naish<sup>26</sup>, Chandrika Nath<sup>27</sup>, Ben Orlove<sup>28,29</sup>,
Pam Pearson<sup>16</sup>, Joeri Rogelj<sup>30,31,32</sup>, Jane Rumble<sup>33</sup>,
Sarah Seabrook<sup>34</sup>, Alessandro Silvano<sup>35</sup>, Martin Sommerkorn<sup>36</sup>,
Leigh A. Stearns<sup>37</sup>, Chris R. Stokes<sup>3</sup>, Julienne Stroeve<sup>38</sup>

<sup>1</sup>University of Exeter, Penryn, United Kingdom, <sup>2</sup>Secretariat, Arctic Monitoring and Assessment Programme, Arctic Council, Tromsø, Norway, <sup>3</sup>Department of Geography, University of Durham, Durham, United Kingdom, <sup>4</sup>Department of Earth, Geographic, and Climate Sciences, College of Natural Sciences, University of Massachusetts-Amherst, Amherst, MA, United States, <sup>5</sup>Arctic Office, Natural Environment Research Council, United Kingdom Research and Innovation (UKRI-NERC), Cambridge, United Kingdom, <sup>6</sup>Department of Geography and Environmental Sciences, Northumbria University, Newcastle upon Tyne, United Kingdom, <sup>7</sup>Department of Water and Climate, Vrije Universiteit Brussel (VUB), Brussels, Belgium, <sup>8</sup>Securing Antarctica's Environmental Future (SAEF), Monash University, Clayton, VIC, Australia, 9Section of Geophysics, National Institute of Oceanography and Applied Geophysics, Trieste, Italy, <sup>10</sup>Scripps Polar Center, Cecil H. and Ida M. Green Institute of Geophysics and Planetary Physics, Scripps Institution of Oceanography, University of California, San Diego, San Diego, CA, United States, <sup>11</sup>Department of Earth and Environmental Science, University of Exeter, Penryn, United Kingdom, <sup>12</sup>British Antarctic Survey, UKRI-NERC, Cambridge, United Kingdom, <sup>13</sup>Department of Atmospheric Sciences, Instituto Antártico Argentino, Buenos Aires, Argentina, 14 CONICET-UBA, Institute for Astronomy and Space Physics, CABA, Argentina, 15 UBA, FCEyN, Physics Department, CABA, Argentina, <sup>16</sup>International Cryosphere Climate Initiative, Stockholm, Sweden, <sup>17</sup>School of Geosciences, University of Edinburgh, Edinburgh, United Kingdom, 18 Department of Geosciences, University of Oslo, Oslo, Norway, <sup>19</sup>Geophysical Institute, University of Alaska Fairbanks, Fairbanks, AK, United States, <sup>20</sup>Departments of Biology and Environmental Science, University of San Francisco, San Francisco, CA, United States, <sup>21</sup>Department of Geography, School of Biosciences, Geography and Physics, Swansea University, Swansea, United Kingdom, <sup>22</sup>Laboratory for Climate and Environmental Sciences (LSCE), Institut Pierre Simon Laplace, Université Paris Saclay, Gif-sur-Yvette, France, 23 Global Systems Institute, Department of Geography, University of Exeter, Exeter, United Kingdom, <sup>24</sup>Danish Meteorological Institute, Copenhagen, Denmark, 25 National Snow and Ice Data Center, Cooperative Institute for Research in Environmental Sciences, University of Colorado Boulder, Boulder, CO, United States, <sup>26</sup>Antarctic Research Centre, Victoria University of Wellington, Wellington, New Zealand, <sup>27</sup>Clare Hall, University of Cambridge, Cambridge, United Kingdom, 28School of International and Public Affairs, Columbia University, New York, NY, United States, <sup>29</sup>International Research Institute for Climate and Society, University of Columbia, New York, NY, United States, 30 Centre for Environmental Policy, Imperial College London, London, United Kingdom, <sup>31</sup>Grantham Institute, Imperial College London, London, United Kingdom, <sup>32</sup>Energy, Climate, and Environment (ECE) Program, International Institute for Applied Systems Analysis, Laxenburg, Austria, 33 Foreign, Commonwealth and Development Office, London, United Kingdom, 34 Earth Sciences New Zealand, Wellington, New Zealand, 35Department of Physical Oceanography, School of Ocean and Earth Science (SOES), University of Southampton, Southampton, United Kingdom, 36WWF Global Arctic Programme, World Wide Fund for Nature, Solna, Sweden, <sup>37</sup>Department of Earth and Environmental Science, University of Pennsylvania, Philadelphia, PA, United States, 38 Centre for Earth Observation Science, University of Manitoba, Winnipeg, MT, Canada

#### **Abstract**

Fossil-fuel burning is heating the planet with catastrophic consequences for its habitability and for the natural world on which our existence depends. Halting global warming requires rapid and deep decarbonization to "net zero" carbon dioxide (CO<sub>2</sub>) emissions, which needs to be achieved by 2050 if warming is to remain within the limits set out by the 2015 Paris Agreement. However, some scientists and engineers claim that a mid-century decarbonization target will not be reached, and they propose that we should focus on technological geoengineering "fixes" or "climate interventions" that could delay or mask some of the impacts of global warming. They often cite the need to slow warming in polar regions because they are experiencing rates of warming higher than the global average, with severe and irreversible projected consequences both locally (e.g., on fragile ecosystems) and globally (e.g., on sea level). Several geoengineering concepts exist for polar regions, but they have not been fully examined by the polar science community, nor integrated with an understanding of polar dynamics and responses. Here, we evaluate five of those polar geoengineering concepts and highlight the significant issues and risks relating to technological availability, logistical feasibility, cost, predictable adverse consequences, environmental damage, scalability (in space and time), governance, and ethics. According to our expert assessment, none of these geoengineering ideas pass scrutiny regarding their use in the coming decades. Instead, we find that the proposed concepts would be environmentally dangerous. It is clear to us that the assessed approaches are not feasible, and that further research into these techniques would not be an effective use of limited time and resources. It is vital that these ideas do not distract from the priority to reduce greenhouse gas (GHG) emissions or from the critical need to conduct fundamental research in the polar regions.

KEYWORDS

global warming, sea level rise, decarbonizing, geoengineering, governance, Arctic, Antarctic, polar

# **Key points**

- Five geoengineering concepts proposed for the polar regions fail to meet the essential criteria required for them to be considered responsible approaches toward limiting the escalation of climate-related risks. These criteria include feasibility and likelihood of success.
- Geoengineering in sensitive polar regions would cause severe environmental damage and comes with the possibility of grave unforeseen consequences.
- Polar regions have complex environmental protection and governance frameworks that would probably reject polar geoengineering fieldwork and large-scale projects.
- Polar geoengineering would require hundreds of billions of dollars in initial costs, plus decades of ongoing maintenance, both of which are presently unavailable and highly unlikely to be secured over necessarily short timescales to address climate change.

- Geoengineering could be used by bad actors as a strategy to create the illusion of a climate solution without committing to decarbonization.
- Minimizing risk and damage from climate change is best achieved by mitigating its causes through immediate, rapid, and deep decarbonization, rather than attempting interventions in fragile polar ecosystems.

# Introduction

The burning of fossil fuels and the resulting global warming are unequivocally damaging the polar regions. Human-caused global warming reached 1.3°C above the preindustrial level in 2023 (1, 2) and, in annual terms, over 1.5°C in 2024 (3). The Arctic is currently warming three times faster than the global average (4), resulting in the unprecedented loss of sea ice extent and volume (5, 6), changes

in snow duration and extent (7), widespread permafrost thaw (8), glacier retreat (9), accelerating Greenland ice loss (10), increasing wildfires (11), changing vegetation distributions (12), and other profound ecosystem changes. These changes are already affecting the daily lives of Indigenous and local communities that rely on local-to-regional resources (13). Across Antarctica, the average atmospheric warming rate is twice the global average (14). Over the past two decades, ice shelves along some of the Antarctic periphery have collapsed (15). The West Antarctic Ice Sheet is experiencing accelerating mass loss (10) and record-high temperatures have been observed in the Antarctic Peninsula (16) and over the East Antarctic plateau (17, 18). Record lows in Antarctic sea ice extent have also been observed in recent years (19, 20). These changes are harming marine and terrestrial species, ecosystems, and ecosystem services across the Antarctic and Southern Ocean, and these effects are projected to worsen (21–23).

Polar environments provide substantial and important services and benefits to humanity (24): changes here have far-reaching implications for the entire planet. Global sea level rise is accelerated by land ice loss from mountain glaciers and the Greenland and Antarctic ice sheets (10, 25). Additionally, the polar regions influence the global climate through extreme weather events, disruptions to the global ocean circulation, carbon cycling, ocean acidification, and other feedback mechanisms (26). The large, bright surfaces of sea ice and ice sheets reflect over 50% of incoming solar radiation back into space, helping to cool the polar regions and the planet overall. The polar regions support a high level of biodiversity, particularly in marine environments, making them some of the most important yet fragile ecosystems on Earth (27, 28). Both the Arctic and the Antarctic play crucial roles in carbon sequestration. In the Arctic, carbon is stored in the permafrost, tundra, and boreal forests (29), and in the Arctic Ocean and the Southern Ocean at least half of all carbon dioxide (CO<sub>2</sub>) uptake by the world's oceans occurs (30-32). This makes the polar regions an important carbon sink but exposes their marine environments to greater acidification, threatening the viability of shelled organisms at the base of polar marine food webs (33, 34).

The polar regions are vital to the overturning circulation of the world's oceans, as they drive the sinking of cold, dense water, which is essential to global ocean currents (35, 36). Direct measurements in the deep Southern Ocean suggest that a ~30% slowdown in the Antarctic overturning has occurred over the past few decades in both the Weddell and Ross Seas, which has been linked to both meltwater and wind changes (37, 38). Model projections suggest that this slowdown will continue for at least the next few decades and that the southern limb of the ocean overturning circulation could collapse within this century (39). The associated freshening of the Southern Ocean is affecting ecosystem functions and loss.

Continuing the current trajectory of fossil fuel consumption and associated global warming is not viable considering the scale and magnitude of the consequences for the polar regions, and their severe regional and global implications (40); under current policies, an intermediate scenario points to an increase in global temperature from pre-industrial levels of approximately 3°C by 2100 (41). By contrast, rapid reductions in carbon emissions, reaching net zero

emissions by mid-century, could stabilize global temperatures and curb polar changes, protecting the world and future generations from the worst effects of high emissions (42, 43).

Some scientists have proposed technological interventions in the climate, oceans, sea ice, or ice sheets as a potential complement to or alternative for reducing emissions (44, 45). Various terms are used to describe these approaches, with different framings resulting in different public perceptions (46). Some of these labels suggest a greater level of control than is possible, such as "climate engineering" and "emergent climate technologies". Other terms, such as "climate repair" and "climate restoration" give the misleading perception of being able to return to a previous climate state. The term "climate intervention" is sometimes preferred, drawing analogies to medicine. However, "intervention" suggests a precise and temporary measure, which does not accurately represent the reality of these approaches. In this article, we use "geoengineering", which encompasses a wider range of activities, such as the large-scale polar engineering ideas that are discussed here. We assess five prominent geoengineering concepts that aim to change heat uptake or distribution around the planet through various means: by increasing solar reflectivity within the atmosphere; diverting key ocean circulation patterns in the polar regions to reduce ice melting; increasing sea ice albedo or thickness; slowing the flow of ice sheets to restrict ice loss to the oceans; and enhancing CO2 uptake by the oceans. We show that all five of these concepts are unlikely to be effective due to a range of factors: the scale and immediacy of the climate problem; harmful side effects within and far beyond the polar regions; failure to address key environmental protection concerns; greater governance and financing challenges than those that already exist in the context of the Paris Agreement; and the risk of distracting from urgent decarbonization efforts, thereby delaying and narrowing remaining and feasible carbon reduction pathways. Our assessment is focused on the polar regions, and is not exhaustive; therefore, it does not include CO2 removal (CDR) technologies and practices, and/or nature-based solutions. Although the development of some CDR technologies, such as direct air capture, is still nascent, they will have to play a role in climate change mitigation (25).

# Part 1: Polar regions' geoengineering assessments

We evaluate five prominent geoengineering concepts: (i) stratospheric aerosol injection (SAI); (ii) sea curtains (or sea walls); (iii) sea ice management (which involves modifying albedo and thickening sea ice); (iv) slowing ice sheet flow through basal water removal; and (v) ocean fertilization. Our evaluation is structured around the following six categories: scope of implementation, which spans both the Arctic and the Antarctic; effectiveness; feasibility; negative consequences; cost; and governance with respect to their deployment at scale (Supplementary Table 1). While these categories are consistent across all case studies, each proposal has unique technological, logistical, and ethical considerations that are discussed in detail in their respective sections.

# Proposed concept 1: Stratospheric aerosol injection

# Background

Solar radiation modification (SRM) refers to the intentional modification of the Earth's radiative budget to offset some of the effects of increasing greenhouse gases (GHGs)—but without reducing the gases themselves (25, 47). SAI is one of the proposed methods, which we discuss below (Figure 1); others include marine cloud brightening, surface albedo enhancement, cirrus cloud thinning, and space-based mirrors.

Budyko (48) published one of the first studies addressing climate change and the possibility of direct climate intervention via aerosol injections. Most of the historical research on aerosol injection stems from the study of the climate effects of major explosive volcanic eruptions. In 1991, Mount Pinatubo in the Philippines erupted, sending an eruption column 40 km

into the atmosphere and creating a massive "cloud" in the middle-to-lower stratosphere. This event injected approximately 17 megatons of sulfur dioxide (SO<sub>2</sub>) into the stratosphere. The resulting aerosol cloud quickly spread globally, achieving full coverage within a year. It significantly reduced the amount of net radiation reaching the Earth's surface, causing observable climate effects. Specifically, there was a surface cooling of up to 0.5°C-0.6°C in the Northern Hemisphere, equivalent to a reduction in net radiation of 4 W/m<sup>2</sup>. Globally, temperatures cooled by approximately 0.5°C for nearly 2 years (49). The event demonstrated that the release of 15-20 megatons of SO<sub>2</sub> into the stratosphere could generate enough aerosols to counteract some global warming. More recent studies show that effective SAI would require lofting hundreds of thousands to millions of tons of material each year to altitudes of up to ~20 km to counter a substantial fraction of warming caused by GHG loading (25, 50, 51).

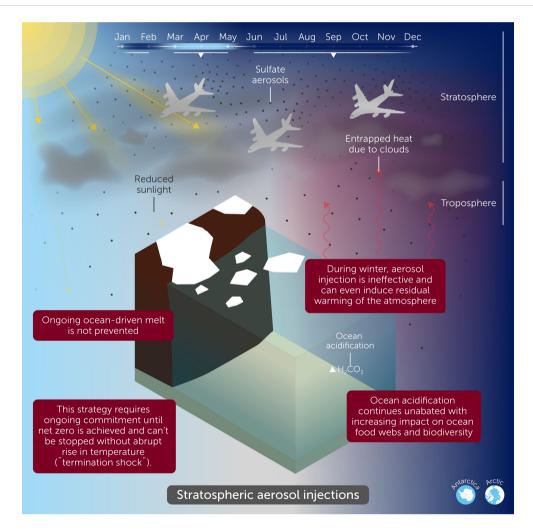


FIGURE 1

Stratospheric aerosol injection (SAI) in polar environments. SAI in polar regions will not be possible year-round, due to winter darkness, and may have unwanted and unintended consequences for regional climates, including those across territorial boundaries. The pale blue shading shows the effective time period for SAI in the Arctic. The Inset icons (bottom right) show whether the option has been proposed for Antarctica, the Arctic, or both.

Current research focuses primarily on the following types of aerosols:

- (i) Sulfur dioxide (SO<sub>2</sub>), which is converted to sulfuric acid aerosols in the stratosphere.
- (ii) Sulfate aerosols, which are similar to those produced by volcanic eruptions. These effectively reflect sunlight and thus could reduce global temperatures but also pose risks, such as potential damage to the ozone layer and the creation of acid rain (52, 53).
- (iii) Titanium dioxide (TiO<sub>2</sub>), which is potentially less chemically reactive than SO<sub>2</sub>, and thus may be less damaging to the ozone layer, but is still in the experimental stage (52).
- (iv) Calcium carbonate (CaCO<sub>3</sub>), which might neutralize some of the acidifying effects of other aerosols and reduce ozone depletion risks but is also still under investigation (52, 53).

#### Effectiveness and feasibility

In the absence of any field assessments and pilot studies, and based on a variety of modeling investigations, the Intergovernmental Panel on Climate Change (IPCC) (25) has noted large uncertainties in the effectiveness of solar geoengineering deployment (51), substantial knowledge gaps regarding its cost-effectiveness, and a range of risks associated with deployment. In the polar regions, the uncertainties are even higher due to seasonally reduced sunlight and aerosol lifetimes.

It has been suggested that aerosols could be injected into the stratosphere using balloons, airships, and/or artillery. However, the most efficient method would probably involve a fleet of custombuilt and/or specially modified aircraft (54, 55). Smith and Wagner (50) estimate that 60,000 flights annually—0.15% of all current flights globally (56)—would be required for a lasting effect. Smith (57) recently highlighted the Boeing 777F as a potential candidate, although it would require substantial modifications, with a fleet of 90 planes focusing on deployment in the polar regions. Deployment strategies would include aerosol injection (i) in the tropics or subtropics (at or equatorward of 30°); (ii) uniformly across all latitudes; and (iii) in the sub-polar or polar regions (at or poleward of 60°). Typically, the tropopause is lowest in altitude in the polar regions, facilitating larger payloads and requiring fewer planes to achieve a given injection mass (54, 57).

SAI would require a highly specific deployment to have a significant impact in the polar regions (58, 59). The radiative forcing from stratospheric aerosols depends on the amount of local incoming solar radiation and top-of-atmosphere albedo. Polar regions are less responsive to aerosol injection during sunlit periods because of their lower insolation and the higher albedo produced by ice and snow (60–62). Furthermore, injections are completely ineffective during the winter months in the polar regions. The Brewer–Dobson circulation, which is characterized by rising air in the tropics and descending air in the mid and high latitudes, affects the distribution and lifetime of stratospheric aerosols. Therefore, aerosols injected at high latitudes have a shorter lifetime and more localized cooling effects owing to their rapid removal by the poleward movement and descent of air (63),

which calls into question the effectiveness of SAI in the Arctic. The effectiveness of SAI in preventing ocean-driven glacier retreat and sea level rise is also likely to be limited (64).

"Termination shock" is the rapid and severe warming that could occur by the unmasking of ongoing GHG emissions if any future large-scale deployment of solar geoengineering were halted (65). The sudden cessation would remove the reflective aerosols from the stratosphere, causing the Earth's climate to adjust abruptly to the conditions dictated by GHG concentrations, resulting in a swift, significant rise in global temperatures. Climate models suggest that following SAI termination, global surface temperatures would increase within a decade or two to values consistent with GHG forcing. This introduces profound systemic uncertainty regarding the long-term viability of SAI and its potential to destabilize the global climate system further. Major water cycle changes would occur, including a rapid increase in global mean precipitation (66). Parker and Irvine (65) claim that timely policies could avert termination shock. However, relying on SAI without GHG removal would probably require prolonged (even permanent) injection, entailing continued financial, ecological and human costs.

The readiness of SAI for implementation varies greatly between polar-specific and global deployments (67, 68). If a launch decision were made by 2030, polar SAI could theoretically be operational regionally by 2040. On a global scale, however, if an international decision to proceed followed a breach of a temperature threshold of 2°C, the decision could be made around 2050 under continual warming from today, meaning that SAI would only be fully operational by 2060 (69). However, many obstacles stand in the way of what is theoretically possible, including financing, governance, and the determination of-and taking responsibility for-negative consequences. The lack of international governance frameworks to oversee SAI deployment compounds these uncertainties, making it even more challenging to predict the longterm impact of this geoengineering option. Overcoming these and other issues will take far longer than the implementation of available actions to strengthen GHG emission reductions and engage transformative adaptation.

# Negative consequences

In the polar regions, atmospheric injections would be ineffective in winter owing to the lack of sunlight. Instead, they could lead to residual warming in winter (due to the blanketing effect of cloud cover), potentially altering the high-latitude seasonal cycle (70, 71), with potential consequences for the water cycle through atmospheric drying. Atmospheric injections could also cause stratospheric heating, which may alter atmospheric circulation patterns, leading to wintertime warming over northern Eurasia, among other unintended consequences (71).

SAI would not address the non-temperature-related effects of GHG emissions, such as ocean acidification. In fact, they could exacerbate acidification, especially if sulfates (the original source of "acid rain" emissions from coal-fired power plants) were used, with consequent negative impacts on ecosystems and economies. The Arctic and the Southern Ocean already show seasonally corrosive conditions for shelled organisms due to current atmospheric CO<sub>2</sub>

levels, and shell damage from increasing acidification was observed from samples taken as early as the 2000s (72). Year-round corrosive conditions are projected to spread throughout much of the polar marine ecosystem with even moderate emissions pathways at  $CO_2$  levels above 450 ppm (26). Ocean acidification is a problem with a very long time frame; it would take 30,000–50,000 years to return to 2012's pH levels (73). Acidification rates are already higher today than at any point in the last 300 million years (74). Exacerbated acidification raises the specter of extinction for many polar marine shelled animals and the economically important species that depend on them, such as krill and cod (34, 75).

Injections of sulfate aerosols can also lead to chemical reactions in the stratosphere that deplete ozone (76). Ozone depletion increases the amount of ultraviolet radiation reaching the Earth's surface, which has harmful effects on human health (see below), ecosystems, and biodiversity as a whole (77, 78).

Using SAI to differentially cool one polar region alone could disrupt global climate patterns. An imbalance in cooling between the hemispheres could shift the Intertropical Convergence Zone, potentially interfering with seasons (63, 79, 80) and reducing rainfall in some regions, thereby affecting water availability and agriculture (81).

Concerns exist regarding the effect of aerosol inhalation on human health. Sulfuric acid aerosols can irritate the respiratory system, causing coughing, airway irritation and inflammation and, in severe cases, respiratory distress. Chronic exposure has been associated with an increased risk of laryngeal cancer owing to long-term irritation and inflammation. This could lead to increased healthcare costs and add a burden on medical infrastructure, with particular risks for vulnerable populations (82).

# Costs

SAI is frequently portrayed as a relatively inexpensive method of climate intervention.

However, Smith et al. (69), looking at a time horizon of 15 years in the future, estimated the direct costs at approximately US\$13.5 billion for acquiring 90 Boeing 777 aircraft, US\$3.2 billion for necessary infrastructure, and approximately US\$1 billion annually for operations. If these expenses were distributed among 30 countries, each would contribute approximately US\$55 million/year. Another recent article has suggested that the annual operating costs, which are larger than the capital costs estimated by Smith et al. (69), would likely grow significantly because of the increase in the amount of material that would be required for injection as GHG concentrations rise. Their study estimates that the injection rate would increase "by a factor of 34" between 2020 and 2080, under a scenario of global warming at Representative Concentration Pathway 4.5 (RCP4.5) (83), so the costs would grow sharply beyond 2040.

However, these calculations assume that the interventions will work as planned, only consider direct costs, and exclude indirect ones such as monitoring and measuring possible impacts, along with the costs of liability related to adverse impacts. Previous studies, such as the one by Smith (84), and more recently by Bronsther and Xu (85), indicate that the perception of SAI as a

low-cost solution may be a misleading oversimplification of the financial realities. Furthermore, such estimates do not include potential unforeseen negative impacts on (i) weather systems, ecosystems, human health, and agriculture; (ii) the effectiveness of renewable energy technologies and the requirements of multicentury commitments (86–88); (iii) the immense potential costs associated with termination shock (see below); (iv) the mistrust and rivalry concerning geopolitical and economic interests among key states (89); and (v) the legal challenges in determining and assigning attribution and liability for such impacts, and the absence of insurance and/or whether insurance is even possible.

#### Governance

The absence of governance for SAI and the absence of deliberation spaces (90) are major challenges (46), with some scholars arguing that deployment may be largely ungovernable (91, 92). In the polar regions, SAI could even be used as a disruptive tool in a wider geopolitical context (89). Unilateral deployment of SAI by one country or a group of countries, but opposed by others, could make international cooperation on addressing climate change harder to achieve and lead to increased geopolitical tensions in general (93, 94). Although political disputes over geoengineering have already been shown to include much deeper issues (95), these tensions are usually ascribed to the fact that SAI has global effects that could vary between countries. While some regions might experience cooling and reduced climate risks, others might face adverse effects such as changes in precipitation patterns and agricultural productivity. This could exacerbate existing inequalities between developed and developing countries. No current international agreements explicitly prevent countries from using geoengineering strategies that could negatively impact other countries (96). In the current geopolitical environment, which is characterized by increasing tension and rivalry, SAI might also prove politically infeasible owing to countries distrusting the intentions of states advocating for research or deployment and may contribute to international tension if subject to disinformation campaigns (89).

No enforceable framework currently exists for international cooperation on SAI, despite some calls for an ethical framework and the recent development of suggested principles (46, 97–99). Efforts to introduce the topic at successive United Nations (UN) Environment Assemblies have failed due to widespread opposition among governments. In 2021, the Swedish Space Agency rejected serving as host for the Harvard-based Stratospheric Controlled Perturbation Experiment (SCoPEx) due to opposition from the Indigenous Arctic population, wider civil society and several Swedish universities; the same project had originally been proposed to take place over Navajo lands in the United States (100).

The European Union has stated that it "...does not consider SRM as a solution, as it does not address the root cause of the problem, which is the increase in greenhouse gases in the atmosphere" (101). In line with this position, the European Commission's Group of Chief Scientific Advisors recently recommended a Europe-wide moratorium on the deployment of

SRM technologies, citing significant risks and uncertainties (102–104). The report also states that "SRM funding should not divert funds and intellectual effort away from climate change mitigation or other related research activities" (102), reinforcing the need to prioritize proven emissions reduction strategies over speculative geoengineering interventions.

In January 2023, Mexico announced that it would ban solar geoengineering experiments following unauthorized tests conducted by the United States-based startup Make Sunsets (105), which released weather balloons containing SO<sub>2</sub> into the atmosphere without governmental authorization. Understanding the power relations and risks inherent in non-state funding of solar geoengineering is a further related challenge (106), particularly as there is a lack of transparency surrounding the funding of geoengineering projects (107).

In January 2022, 400 scientists proposed a non-use agreement on solar geoengineering with five core principles: no public funding, no outdoor experiments, no patents, no deployment, and no support in international institutions; the African Union and the European Parliament have called for support on this agreement (108). Establishing such an agreement should be feasible as there are many precedents in international law that could be used to impose an international ban on the development of solar geoengineering technology, including the Chemical Weapons Convention and the Treaty on the Prohibition of Nuclear Weapons (109).

Another significant concern is that the implementation of SAI would potentially distract and detract from necessary efforts to reduce GHGs (110).

The potential social disruptions would be substantial. There could be significant public resistance due to concerns about the potential side effects and the ethical considerations of deliberately altering the climate. Public protests and legal challenges could arise, complicating the implementation of these programs (111).

# Proposed concept 2: Sea curtains/sea walls

### Background

Wolovick and Moore (112) and Moore et al. (113) proposed installing artificial structures to prevent warm water masses from reaching the buttressing ice shelves and grounding zones of the Antarctic and Greenland ice sheets to reduce ocean warmthinduced ice-sheet melting and mass loss. The initial idea of building artificial sills up to 100 m high, or artificial pinning points up to 300 m high, on which an ice shelf could ground, has now been discarded. Now favored are "seabed curtains" (114, 115) comprising flexible buoyant structures anchored to the seabed at depths between 700 and 1,000 m (114), with 150–500 m high curtains (depending on location) extending upward as a series of thin, flexible overlapping panels (Figure 2).

#### Effectiveness and feasibility

The installation of such structures at depth and on various substrates is extremely challenging; this is compounded further by the harsh environment and the remoteness of the envisaged target areas, especially in the Amundsen Sea region of West Antarctica. The prefabricated foundations would have to be substantial to counteract the buoyancy forces of the curtain, and their necessary size may even exceed the capacity of ships to transport them (115). The proposed "outer" curtain route in the Amundsen Sea is on soft seafloor sediment. The potential for instability and mass movement due to "loading" on the seabed would need to be carefully assessed. In contrast, the "inner" curtain routes would require foundations built at least partly on rough bedrock channels, with steep slopes and the potential for strong currents scouring the topography (116-118). Shaping a prefabricated foundation structure to such a rough surface would be challenging even on land and even more so at depths of several hundred meters in Antarctic waters with icebergs and sea ice present, and especially with the need for precise installation of anchors in both hard and soft surfaces (115).

According to numerical modeling, such structures might successfully divert warm water from one ice shelf system, but their net efficacy may be compromised by the unintended rerouting of this water to other systems (119). Another modeling study concluded that reducing grounding line melt rates by up to 5 m/year below present levels could promote the re-advance of some ice streams (even after extensive prior retreat). However, reversing Antarctica's contribution to sea level rise also requires a substantially increased ice-surface accumulation (120). Feldmann et al. (121) estimate that at least 7 Gt of artificial snow would be required to balance the ice sheet around the Amundsen Sea, but this vast amount of ice would need to be applied over 10 years to be effective and, presumably, sustained at a high level thereafter. These studies call into question the usefulness and practicality of the approach.

Another challenge is that the deployment site is in one of the harshest and most remote environments on Earth, with a transit time of over 1 week from the nearest port and access likely to be possible only for a few months each year owing to polar darkness and extreme ice and weather conditions. Even today, the handful of ice-strengthened research ships that try to undertake short expeditions in this area are hampered by such conditions, often have to change work locations, and frequently fail to get to the region entirely due to impenetrable sea ice. Hazardous sea ice and iceberg collision risk have frequently thwarted attempts to reach the installation sites proposed by Keefer et al. (115) over the 40-year period in which this region has been targeted by ice-breaking vessels for academic research. A literature survey of cruise reports, and interviews with the principal investigators of research expeditions undertaken in the region, has revealed that 56% of cruises experienced at least partial disruption due to sea ice or had significant difficulty entering or exiting the area; 22% were unable to access the region altogether (122, 123). For example, the Inner Bay curtain site proposed for offshore Thwaites Glacier (115) has only been accessed once in the last 40 years despite frequent expeditions targeting the glacier front (117), while several

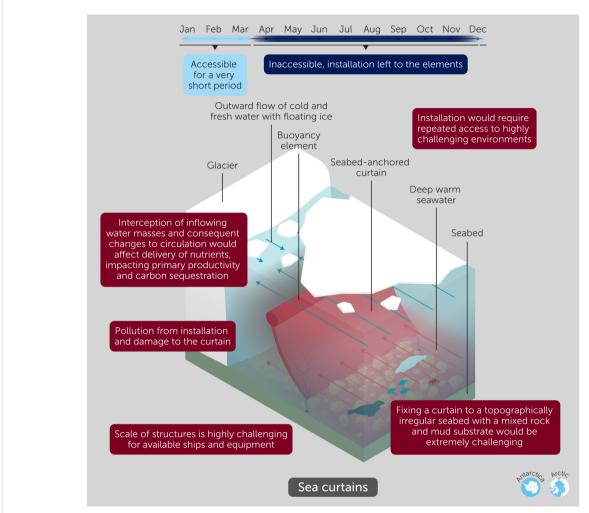


FIGURE 2
Sea curtains to block warm water from flowing towards ice sheet grounding zones. Installing structures spanning many tens of kilometers is a massive technological challenge that will require operations across some of the world's roughest seas and sustained work in ice-covered locations that even modern ice-strengthened vessels cannot always reach. These curtains will probably have unwanted consequences on ocean circulation and ecosystems. The pale blue bar shows the relatively short operational window for ships in the Amundsen Sea, Antarctica. The Inset icons (bottom right) show whether the option has been proposed for Antarctica, the Arctic, or both.

oceanographic moorings remain unrecovered in other parts of the Amundsen Sea continental shelf region despite multiple attempts to retrieve them in different years. Furthermore, even in years when sea ice conditions have been favorable, the risk of iceberg collisions has forced the abandonment of activities. This includes several locations where ships necessarily remain stationary for relatively short periods of time, such as for sediment drilling (124), and also the far greater task of installing complex foundations across rugged topography.

Keefer et al. (115) have argued that all installation vessels would need to be ice-class strengthened to Lloyd's A1 classification and capable of holding position. Currently, very few ships have this capability, with the majority of them committed to Northern Hemisphere operations; new ones would cost on the order of US\$500 million each. These cost estimates are based only on sea curtains in the Amundsen Sea area, but there are other parts of the ice sheet margin that are potentially vulnerable and even more remote from logistical hubs, such as the Cook Ice Shelf, and the Mertz and Ninnis Glaciers in East Antarctica (125). Any installation elsewhere in Antarctica would be scaled in cost accordingly. Similar ideas have been proposed for Greenland, where artificial barriers could be built in glacial fjords to prevent warm deep waters from reaching marine-terminating glaciers (113). While building a berm in Greenland's fjords would be less logistically challenging than in Antarctica, the effectiveness of such interventions might be limited since the Greenland Ice Sheet is also highly susceptible to atmospheric-driven melting (126).

# Negative consequences

The installation of sea curtain structures is likely to have farreaching detrimental marine environmental consequences. These include potential effects on oceanic circulation, sea ice, and marine ecosystems from the interception of circulating water masses, the deflection of water masses elsewhere, and the impacts of any water

masses that leak through barriers or "overtop" the uppermost elements (115).

The presence of barriers to water masses will also act as barriers to marine life, such as demersal fish, mobile benthic invertebrates, and the seabirds and marine mammals that feed at depth in these regions (127). Interception of inflowing water masses and the consequent changes in circulation and sea ice will affect the delivery of nutrients—most notably iron (128)—from upwelling and from ice melt in the region, with consequent impacts on primary productivity and carbon sequestration. Installation of towed floating structures, as suggested by Keefer et al. (115), may also introduce non-native species to the seabed and water column around Antarctica with unknown consequences for marine ecosystems (129, 130).

Other potential consequences include pollution from installation and damage to the curtains, additional emissions associated with transportation and installation, and the need for ongoing maintenance over decadal timescales (the planned lifetime of the curtains being 25 years). The material to be used for the curtains is not yet clear. There is the possibility that elements of the curtains, which are hundreds of meters long, could degrade and/or become dislodged over time, posing a severe threat to marine life and shipping.

Installing such a structure would entail considerable emissions. Transiting an icebreaker more than 2,500 km from the nearest major port while towing large floating structures through rough seas would be operationally risky and carbon-intensive. For example, the German research vessel Polarstern consumes approximately 900 tons of fuel/month (equivalent to 2,800 tons of  $CO_2$ ) with a far higher consumption when breaking through sea ice. Towing large structures through sea ice is extremely difficult, and so the working season would probably be <3 months/year. Each transit for a single curtain section would probably take well over a month, with a likely maximum of four to five round trips per year being possible, each ship emitting 10,000-15,000 tons of  $CO_2$ .

The complexity of these logistical challenges amplifies systemic uncertainties about the long-term feasibility of sea curtains. Any structural failure due to an iceberg collision, strong currents, or foundational instability could compromise the entire intervention, negating potential benefits and leading to unforeseen environmental consequences. Such uncertainties, combined with high costs and engineering demands, underscore the speculative nature of this approach.

#### Costs

The costs of installing sea curtains will probably far exceed the maximum of US\$80 billion for an 80 km structure, spread over a decade, as estimated by Keefer et al. (115). Their cost estimate of US\$1 billion/km can be compared with the real-world cost of the Thames Barrier, a 500 m long structure across the River Thames in London, United Kingdom. This was built on the surface in water <10 m deep immediately adjacent to an enormous logistics hub (i.e., with no shipping required). Nevertheless, it took 8 years to build and cost £535 million in 1982 (131), equivalent to £2.4

billion (US\$3.1 billion or US\$6.2 billion/km) at 2024 prices. The Three Gorges Dam in China, which is approximately 2 km wide and 180 m high, cost ~US\$37 billion in 2008 (132), equivalent to US\$54 billion in 2024; approximately 25 times more per kilometer than presently reported for the Amundsen Sea curtain.

#### Governance

Governance of sea curtains in Antarctica would fall under the Antarctic Treaty System. Article 8 of the Protocol on Environmental Protection requires a Comprehensive Environmental Evaluation (CEE) for any activity with more than a "minor or transitory" impact (133). The Commission for the Conservation of Antarctic Marine Living Resources (CCAMLR) may also have concerns regarding the potential consequences for the marine ecosystem. The installation of a sea curtain would be orders of magnitude larger than any construction activity seen previously in Antarctica and would require not only an estimation of all likely impacts but also a rigorous associated program of impact monitoring and mitigation. Obtaining international consensus to intervene in Antarctic processes on such a large scale has never been attempted before, and those advocating for the use of sea curtains have given this necessary requirement scant consideration thus far. Flamm and Shibata (134) have pointed to significant potential problems for such an idea within the Antarctic Treaty System (see Part 3), suggesting that—even if sea curtains could work without negative impacts—they may precipitate political fragmentation and "pose a risk of 'international discord" in Antarctica, through concerns over "authority, sovereignty, and security".

# Proposed concept 3: Sea ice management—modifying albedo and thickening sea ice

#### Sea ice albedo

# Background

Sea ice cools the Earth by reflecting solar energy. Loss of sea ice in recent years has made the Earth's surface less reflective: according to a recent global study, sea ice has lost 13–15% of its radiative cooling effect since the 1980s, leading to increased solar heating (135). Scattering hollow glass beads over first-year Arctic sea ice could increase its albedo, promoting its survival into a second winter and beyond to re-establish multi-year ice (136).

# Negative consequences and feasibility

The use of glass beads raises significant concerns that must urgently be addressed (Figure 3). First, there is the risk of ecotoxicity. Ecotoxicological testing is underway in vertebrates (136), but the impact on invertebrates appears to be unknown. Potential effects on zooplankton feeding behavior are of particular concern, given the size and nature of the beads. Second, the glass beads are metastable and will readily dissolve in the undersaturated seawater (137), potentially reducing their efficacy. Any means used

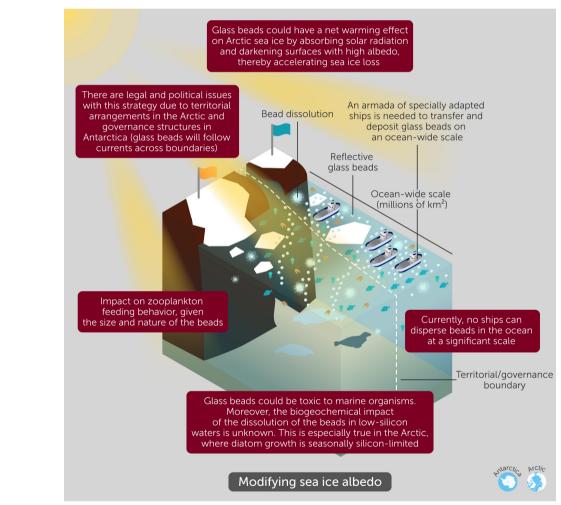
to reduce the solubility of the beads will increase ecotoxicological concerns. The biogeochemical impact of bead dissolution in the low-silicon waters of the Arctic surface ocean, where diatom growth is seasonally silicon-limited (138–140), is unknown. Webster and Warren (141) have shown that massive quantities of glass beads would be required—approximately 360 megatons annually, an amount equivalent to the annual global production of plastic—which poses significant logistical challenges and substantial emissions during production. This scale introduces systemic uncertainties related to supply chains, logistics, and emissions associated with manufacturing, further undermining the feasibility of this approach. Finally, Webster and Warren found that microspheres actually have a net warming effect on Arctic sea ice. They can absorb a significant portion of solar radiation,

darkening surfaces with high albedo and accelerating sea ice loss rather than preserving it.

The Arctic Ice Project, which aimed to use this glass bead technology to slow ice melt, was recently shut down after ecotoxicological tests revealed potential risks to the Arctic food web (https://www.arcticiceproject.org/a-final-chapter/). This underscores the uncertainties and risks of geoengineering interventions in fragile ecosystems.

#### Costs

Beyond the environmental risks, the financial feasibility of deploying glass beads at scale is debatable. In addition to the costs of materials, manufacturing, and logistical operations,



#### FIGURE 3

Using glass beads as a means to reflect sunlight off polar surfaces. Changing albedo by adding particles to the ocean may actually decrease albedo, would require deliberate pollution of ecosystems, and may cross governance boundaries. It is also unlikely to be logistically possible to operate at the scale necessary to make a significant difference. The inset icons (bottom right) show whether the option has been proposed for Antarctica, the Arctic, or both.

research and regulatory approvals would add further expenses, making this approach economically unviable compared to direct GHG mitigation and adaptation strategies.

#### Governance

The use of glass beads also presents important governance challenges. Ocean currents and surface winds carry sea ice across sovereign boundaries in the Arctic Ocean and would do the same for glass beads in surface waters. The international legality of a single nation introducing an environmental disturbance into another would be a serious issue requiring resolution before any deployment. No such agreements exist at present, effectively blocking the roll-out of this technology. Another essential governance factor concerns Indigenous Arctic communities, and their preparedness for and support of these interventions that will affect the natural systems that support their livelihoods and culture (142).

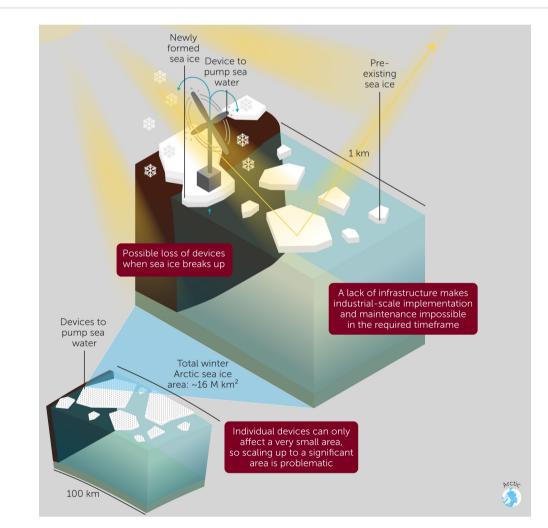
#### Arctic sea ice freezing

#### Background

There is much debate concerning whether Arctic sea ice can be artificially "thickened" to counteract sea ice loss (Figure 4). It has been suggested that sea ice can be thickened by pumping seawater either onto the ice surface, where it will freeze, or into the air, where it will fall as snow onto the ice (143–145). Desch et al. (144) suggest that winter sea ice could be thickened by  $\sim$ 1 m using wind-powered pumps that spray seawater onto the ice surface. They propose that deployment of these pumps across 10% of the Arctic Ocean could "more than reverse current trends of ice loss in the Arctic, using existing industrial capacity".

#### Effectiveness and feasibility

For the technique to be effective on a local scale, Pauling and Bitz (145) suggest that it would have to be deployed "almost



Arctic sea-ice thickening to counteract the loss of ice. Techniques to thicken sea ice would require a very large number of individual devices to be deployed onto the winter sea ice, and it is unlikely to be logistically possible to operate at the scale necessary to make a significant difference. The inset icons (bottom right) show whether the option has been proposed for Antarctica, the Arctic, or both.

immediately". It would not be possible to deploy these pumps at scale within this strict time frame. There are also challenges related to the scale of the deployment area; the long-term average Arctic sea ice extent varies between summer and winter from 5 to 16 million km<sup>2</sup>. According to Desch et al. (144), 10 million pumps would be needed to cover 10% of the Arctic Ocean, and 100 million to cover the entire Arctic. To cover 10% of the Arctic within 10 years, they anticipate that one million pumps would need to be deployed each year, and they concede that this is unlikely to be logistically or financially possible (144). Therefore, sea-ice thickening is simply not feasible for use at a scale and at a rate that would be meaningful for sea ice protection. Furthermore, the scale of implementation of this method would require an unprecedented level of human presence in the High Arctic for initial construction and subsequent maintenance. Other inherent challenges may render the idea infeasible even at a local scale, such as the difficulty of keeping the pumps in place (considering the strong drift of sea ice), the need for year-round maintenance over such a large area, the sinking of the pumps if sea ice melts or breaks up, and governance issues—especially considering that the Arctic Ocean is becoming an increasingly congested space.

In addition to these challenges, it is also unlikely that this approach would be effective as a tool to mitigate global warming. A modeling study by Zampieri and Goessling (146) found that if seawater pumps worked, it could be possible to maintain late-summer sea ice cover at its current extent for the next ~60 years, but maintaining the ice would have a negligible effect on global warming. This contradicts the claim that restoring Arctic sea ice could be a tool to limit global warming (136).

# Costs

Desch et al. (144) estimated production and transport costs of US\$50 billion per year for 10% of the Arctic or US\$500 billion per year for the entire Arctic. Although US\$500 billion only represents ~0.5% of global GDP (~\$100 trillion), it is still an extremely large financial investment for a technology that may not preserve the sea ice and, even if it did, would do little to mitigate strong global warming. Additionally, the long-term financial commitment required for pump maintenance and replacement due to wear and environmental damage adds another layer of uncertainty to the cost-effectiveness of this approach. Redirecting this investment to reduce GHG emissions would provide better value.

# Governance

Any large-scale intervention affecting natural systems in the Arctic must be undertaken with an appreciation of the applicable sovereign and international governance structures (see Part 3). As with the "glass beads" idea, the flow of sea ice across sovereign boundaries means that issues may well arise between nations that agree to the installation of ice-thickening machinery and those that do not, especially for maintenance and removal purposes. In addition, the lack of an international framework to regulate such interventions creates the risk of unilateral action, which could lead to political tensions and disputes over environmental damage. Indigenous Arctic communities must also be included in

governance discussions, as these interventions directly impact the natural systems they depend on for food and cultural practices (142).

# Proposed concept 4: Slowing ice sheet flow through basal water removal

# Background

The rate at which ice is discharged from ice sheets to the ocean is determined by the flow of outlet glaciers and ice streams toward their margins. Increases in flow rates, if not balanced by increased accumulation, contribute to global mean sea level rise. While mountain glaciers dominated the cryospheric contribution to sea level rise over most of the 20th century, ice sheet contributions have risen rapidly since the 1990s, and the combined contribution of both Greenland and Antarctica now exceeds that of mountain glaciers, adding 11.9 mm to global mean sea level between 2006 and 2018 (147). These trends are expected to continue, with the IPCC's Sixth Assessment Report being unable to rule out lowconfidence projections that exceed 15 m of sea level rise by 2300 under a high-emissions scenario (147). It follows, therefore, that reducing the rapid flow of outlet glaciers and ice streams in Greenland and Antarctica could, in principle, mitigate future sealevel rise, especially mass losses from Antarctica, which experiences much less surface melting than Greenland. Building on this logic, some researchers (113, 148) have proposed geoengineering strategies aimed at increasing basal friction to ultimately slow down glacier flow, with a particular focus on removing the lubrication provided by the pressurized subglacial drainage system (113). Here, we present our logic for why we believe that these intervention strategies are scientifically flawed and likely to be logistically impossible.

Ice streams and outlet glaciers exist on a spectrum from those that flow through deep troughs to those that flow over relatively flat beds (149, 150). The former ("topographic ice streams") are common in both Greenland and Antarctica, and the latter ("pure ice streams") are exemplified by those that drain the Siple Coast of West Antarctica but do not occur in Greenland. The limited geoengineering case studies are from the Siple Coast ice streams (148), probably because they dominated much of the earlier literature on ice stream flow mechanisms (150); however, they are not a major contributor to recent sea level rise from Antarctica. Worryingly, the glacier experiencing the most ice loss, Thwaites Glacier, appears to be unconstrained by basal topography, and its deglaciation would probably impact the entire West Antarctic Ice Sheet (151, 152). Irrespective of their topographic setting, the rapid movement of ice streams and outlet glaciers is predominantly enabled by basal sliding-ice slipping along the bed-a phenomenon resulting from a combination of sliding at the icebed interface, and deformation of the bed itself (150, 153). The fastest-flowing ice streams are generally found in regions where water is routed down topographic troughs or accumulates at the base, acting as a lubricant and reducing the effective pressure of the overlying ice (150). Increased water pressure in these areas is

directly linked to increased glacier velocity, resulting in accelerated ice discharge (154).

In Antarctica, subglacial water is generated at the bed primarily by the melting of basal ice due to a combination of frictional heating, geothermal heating, and pressure melting (155). Unlike in Greenland, supraglacial meltwater does not reach the bed in appreciable quantities. It may also originate from Antarctic groundwater; a relatively recent focus of research, which could account for as much as half of the water beneath ice streams (156, 157). While groundwater has been identified at the margins of ice sheets (158), there is currently only limited evidence of groundwater systems beneath Antarctic ice streams (159). However, as there are huge sedimentary basins present beneath the ice (160), it is very probable that water from these basins plays a major role in modulating ice flow (161), especially on decadal timescales (162). Hence, any study looking to extract basal water to reduce water pressure must account for fundamental uncertainties in the role of groundwater.

The rates of basal meltwater production are generally low (mm/year), but over large areas of the ice sheet bed this leads to large volumes of subglacial water available to and transported beneath ice streams (155, 163). Furthermore, the subglacial water system is interconnected and continuously replenished, comprising a network of subglacial streams and lakes (164, 165). In contrast, Greenland's subglacial water is mostly generated at the surface and then delivered to the bed via moulins, providing a direct link between subglacial water and climate change. Importantly, in Antarctica there is no widespread link between basal melting and the atmosphere. However, Kazmierczak et al. (166) demonstrate that the long-term Antarctic response to climate forcing will be modulated to a degree through the flow of basal water.

Geoengineering strategies to reduce ice flow and glacier sliding fall into three main categories (148):

- (i) Drilling/drying: This involves drilling to the subglacial bed and reducing water pressure there by pumping subglacial water to the ice surface. The objective is to increase basal drag by reducing water pressure, thereby slowing glacier flow.
- (ii) Cooling: These approaches aim to reduce glacier flow rates by promoting heat flow from the bed to the surface or by introducing coolants into the bed.
- (iii) Obstructing: This involves constructing obstacles (such as artificial rock outcrops) at the bed to increase basal friction and reduce the dependence of friction on water pressure, thereby impeding glacier flow (167).

## Effectiveness and feasibility

Here, we focus on the scientific and logistical challenges of drilling/drying, but many of these apply equally to cooling and obstructing (Figure 5). We anticipate the effectiveness of drilling/drying to be limited, owing to the highly dynamic spatial and temporal nature of subglacial drainage.

The location, timing, and sphere of influence of each drill hole would be critical. In Antarctica subglacial water production is spatially extensive, but most basal water flows in discrete channels (155) that form intricately branched subglacial hydrologic systems. Therefore, multiple drill holes would be required, each syphoning a different branch—assuming these could be accurately located from surface geophysical surveys. This also assumes that the subglacial drainage system is stable; however, a wealth of evidence shows it to be highly variable, both spatially and temporally (168–170). Furthermore, once the ice stream bed is accessed by a drill, basal sliding will quickly move the drill hole base downstream. Therefore, even if a channel or subglacial lake is connected to the surface by the drill hole, the ice flow above it would act to break the connection within days to weeks. As for groundwater, we presently have no means to consider its extraction.

In Greenland, surface water could be removed or rerouted to prevent it from reaching the bed, but the transient nature of the supraglacial system (169, 171) would necessitate adaptable camps, machinery, and logistics. Maintaining infrastructure on a melting ice surface is a significant challenge requiring almost constant maintenance. Water would have to be removed from the ice sheet and routed elsewhere without it reaching the bed via downstream moulins or crevasses. Additionally, this would have no impact on meltwater generated by frictional heating at the bed of large topographic ice streams (172). Furthermore, channeling large volumes of supraglacial water to the bed would lead to lowpressure channelized drainage systems that would act to reduce annual ice flow velocities, rather than increase them (173). Greenland's subglacial hydrology evolves seasonally; toward the end of the melt season, high volumes of subglacial water evolve into efficient channelized drainage systems that lubricate less of the basal area, thereby decelerating the ice flow (174). Therefore, we consider it highly unlikely that basal water extraction in Greenland, either directly or through supraglacial diversion, is a feasible means to decelerate ice flow, and it may even have the reverse effect.

In Antarctica, targeting the trunk of the ice stream for water removal is unlikely to work as it is the cumulative velocity of the entire ice stream that needs to be reduced. Removing water from upstream and/or in the onset zone of ice streams would be the most effective way to reduce the overall water supply. However, the inland plateau is the most difficult place to access and has the thickest ice, making drilling even more challenging. Moreover, boreholes would have to be kept constantly open because subglacial water production is continuous. Draining the bed in the onset zone or upstream portion of an ice stream would theoretically steepen the ice surface slope and thereby increase driving stresses and pressure melting at the bed, much in the same way that glacier surging is triggered. Indeed, even if drainage slowed ice flow in a particular area, the ice sheet/stream would thicken upstream, and the resulting changes in the ice surface slope would displace basal water into neighboring regions, potentially accelerating flow elsewhere (175). This 'water piracy' mechanism has been attributed to the shutdown of the Kamb Ice Stream (formerly Ice Stream C) (176), which will probably oscillate

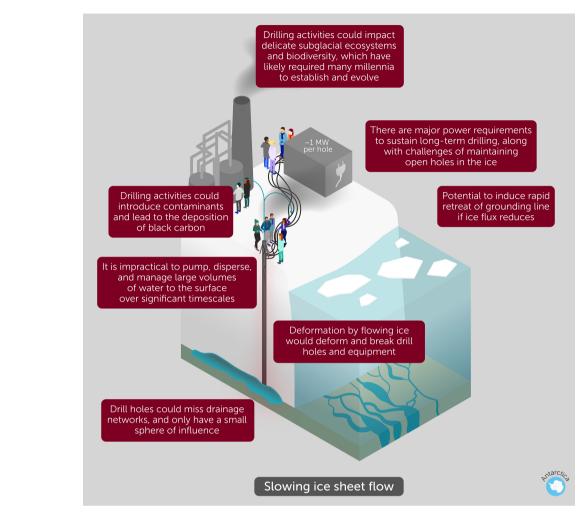


FIGURE 5
Subglacial water removal in ice sheets to slow ice flow to the ocean. Drilling to the bed of thick, flowing ice is highly technologically challenging and has never been undertaken for the sustained period required to maintain the drainage of subglacial water. Subglacial drainage networks are currently pristine and not well mapped, so the introduction of drill holes into the network will probably be highly challenging to achieve reliably and could cause contamination both below and above the ice. The inset icons (bottom right) show whether the option has been proposed for Antarctica, the Arctic, or both.

between periods of fast and slow flow and could be reactivated this century (177).

Both observations and modeling show that the Siple Coast ice streams exhibit significant short-term variability owing to the complex interplay between adjacent ice streams (178). Hence, even if it were theoretically possible to slow down a portion of an ice stream, the resulting thickening would probably reactivate the ice stream, or neighboring ice streams would capture ice and/or meltwater (179). This type of "flow-switching" is preserved in the geological record of paleo-ice stream activity (180–184). This "self-organized" behavior is replicated in numerical ice sheet models (185–187) and during deglaciation of paleo-ice sheets (184). This means that, even if it were possible to slow the flow of one ice stream, there would probably be knock-on effects elsewhere, with other ice streams increasing in flow or reactivating to remove the excess mass.

The rapid flow of major outlet glaciers and ice streams could also be triggered by conditions at the terminus (e.g., major calving events, ice shelf thinning, or grounding line retreat along retrograde bed slopes), which would then propagate inland through increased flow velocity and dynamic thinning (188–190). Hence, even if upstream adjustments to water pressure and flow were successful, continued ice loss at the terminus could lead to modifications to glaciological change at borehole sites, rendering the sites suboptimal. A reduced flux of ice into the grounding zone could even lead to rapid glacier retreat if mass losses at the front (from iceberg calving and submarine melt) do not change. A system that is already close to instability could be "pushed over the edge" if the ice supply were insufficient to keep the grounding zone stable. For example, the grounding zone of the Institute Ice Stream in West Antarctica sits atop a major reverse bed slope that is kept stable by a delicate balance between basal melt rates, ice accumulation and ice flow (191).

Logistically, drilling for subglacial water presents significant challenges and potential consequences. First, drilling to the base of an ice sheet is a monumental undertaking, necessitating advanced equipment and expertise (192). To be most effective, this technique

would require numerous drill holes scattered across the ice sheet, a logistical feat unprecedented in polar fieldwork. The sheer size of an ice sheet means that the endeavor is virtually impossible to execute.

Water under pressure deep beneath the ice surface will not naturally escape in significant volumes: the puncture of Lake Vostok in 2012 created a colossal hydrological shock that bent the drill head and caused massive hydrofractures in the ice base (193). However, only 5 m<sup>3</sup> of drilling fluid escaped to the surface because the water froze in the borehole. The Lake Ellsworth experience in 2012 also points to major challenges in drilling deep holes with water once, let alone repeatedly, and even if successful they will freeze after 24 h unless frequently reamed—requiring substantially more hot surface water (192). Borehole pumping would have to be sustained over long timescales, as the continuously produced basal water would need to be extracted throughout the year, including in winter. Again, because the ice in the interior of Antarctica is extremely cold, the boreholes would freeze rapidly (within hours) without continuous heating. Above the water level, the borehole would close under the pressure of the overburden and require regular reopening or casing. This issue becomes non-linearly more difficult as the water level drops.

Pumping water out from the subglacial system would be a significant task requiring considerable energy resources. Existing hot water drills use pumps that recover water from depths of about 100 m. This corresponds to the water level at overburden pressure in a 1,000 m deep borehole. If lowering the water pressure were successful, then the water would have to be pumped against an everincreasing hydraulic head. Depending on how much water level lowering is needed (which is currently completely unknown), this would dramatically increase the energy requirements and might even exceed the capacity of existing downhole pump technology. Entirely new approaches, such as those used in oil extraction, may be needed. These approaches would require casing of the holes but, in active areas of the ice, significant borehole deformation would make casings vulnerable. Furthermore, successful drying would cause higher basal stresses, which would increase shearing and affect borehole operations.

A cursory power calculation estimates the power requirements to be in the multiple MW range (as exemplified by various deep-drilling projects in Antarctica, e.g., ANDRILL, SWAIS 2C, Subglacial Lake Ellsworth, and IceCube). This is a colossal level of power to accomplish in a remote location, and operating year-round in extreme cold and windy conditions is far beyond the capabilities of existing technologies. Providing such power requirements using nuclear power, for example, would be highly controversial in the context of the Antarctic Treaty. Solar farms at that scale are not realistic given the long polar winter and the formidable challenge of wind drifts. Wind turbines could theoretically work, but existing structures on moving ice are several orders of magnitude smaller than would be necessary.

At present, kerosene is commonly used for power in remote field locations, which poses another environmental risk: the deposition of black carbon on the ice surface. At the scale required for widespread intervention, such deposition would probably be counterproductive, exacerbating rather than mitigating the challenges posed by rapid glacier flow. Additionally, it takes approximately 12 barrels of kerosene to transport 1 barrel of fuel to central Antarctica. Another significant challenge is the management of pumped-out water on the surface. A proposed solution is to turn this into snow, using snowmaking equipment commonly used in ski areas. This is feasible in principle but would produce a large amount of snow and entail additional energy requirements. The necessary management of this snow would require automated procedures to move it away from the snow cannons. The maintenance of surface infrastructure in extreme cold and wind, drifting snow, and riming (freezing water on solarheated surfaces) is also a serious impediment. For example, equipment commonly gets "drifted in" in the Thwaites area, where there is relatively high snow accumulation and wind drift (194). Both the pumping infrastructure and snow-making equipment would require protection from snow drifting.

Many of these challenges require entirely new engineering approaches that would take many years, if not decades, to develop, even with unlimited resources. We emphasize that although they are not considered in detail here, the proposed cooling and obstructing approaches will face similar challenges to drilling/drying, both scientifically and logistically.

### Negative consequences

Despite the scientific flaws of attempting to slow an ice stream, drilling activities could introduce contaminants into pristine subglacial environments, damaging delicate, biodiverse ecosystems that have evolved probably over many millennia. These consequences underscore the importance of thorough risk assessment that considers all environmental impacts. The Scientific Committee on Antarctic Research (SCAR) has agreed to a code of conduct for Antarctic subglacial access (195), endorsed in 2017 by the Antarctic Treaty Consultative Parties (ATCPs). The Antarctic Treaty's Committee for Environmental Protection (CEP) has assessed deep-drilling access to the ice-sheet bed, and following its advice, authorization by ATCPs is only likely to be granted if procedures with proven effectiveness are used to sterilize the drilling water (196). Hot-water drilling with pasteurized water might be possible, but such work is technically challenging (192) and increases energy consumption. The use of ice cores, which use an antifreeze drilling fluid, would probably be incompatible with environmental protection. Drilling could also lead to the unintended alteration of ice structures and crevasse formation (197).

## Costs

The Subglacial Lake Ellsworth program, which in 2012 unsuccessfully aimed to create a 3-km deep borehole for just 24 hours, cost approximately US\$13 million (equivalent to US\$18 million in 2024) (196). The proposed method of water removal would require multiple holes (tens to hundreds), and a continuous connection to the seabed (i.e., 365 days/year for years to come), adding orders of magnitude to the Lake Ellsworth figure, in addition to maintenance and monitoring. It is difficult to imagine such a

large undertaking being possible without the establishment of a new coastal station in the vicinity, to which materials could be delivered by sea. This would serve as a year-round base for operations. The construction of IceCube would not have been possible without the existence of McMurdo, Antarctica's largest base, and the Amundsen–Scott station at the South Pole. Planning, approval, and construction of these types of bases is a decadeslong undertaking.

## Governance

As in the case of sea curtains (proposed concept 2), the implementation of this activity in Antarctica would fall under the Antarctic Treaty System. Article 8 of the Protocol on Environmental Protection requires a CEE for any activity with more than a "minor or transitory" impact (133). Given the high level of uncertainty surrounding both the effectiveness and environmental risks of basal water removal, it is unlikely that such an intervention would meet the Treaty's strict environmental protection standards. This activity would require not only an estimation of all likely impacts but also

a rigorous associated program of monitoring and mitigation of impacts. This governance barrier further highlights the systemic obstacles to implementing this geoengineering concept.

# Proposed concept 5: Ocean fertilization

### Background

Ocean fertilization is a form of geoengineering wherein nutrients whose availability limits phytoplankton growth (e.g., iron) are added to the ocean surface water to promote photosynthesis and stimulate blooms, thereby increasing the biological drawdown of  $\rm CO_2$  from the atmosphere (198) (Figure 6). This enhanced fixation of carbon in surface waters is then expected to sink as particulate organic carbon (POC) to the deep ocean for long-term sequestration. This vertical advection is known as the "biological carbon pump" (BCP), in which a variety of biota from bacteria to zooplankton and fish control the magnitude and efficiency of carbon export to the deep ocean (199).

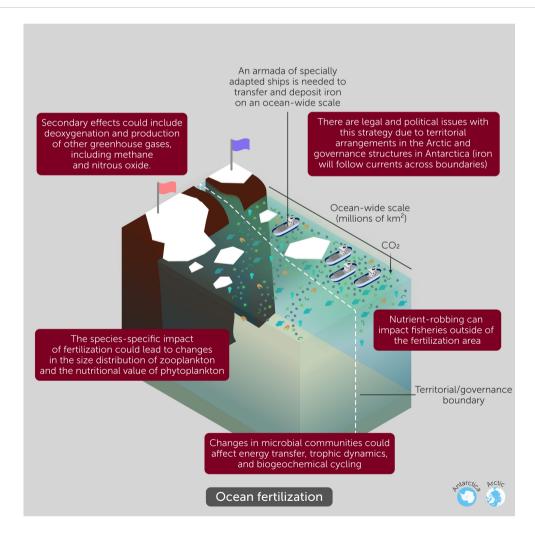


FIGURE 6

Ocean fertilization to "draw down" atmospheric CO<sub>2</sub>. Negative impacts will likely include changes to food web structure, and fertilization could affect nutrients and fisheries elsewhere, including across territorial or governance boundaries. The inset icons (bottom right) show whether the option has been proposed for Antarctica, the Arctic, or both.

Prime areas for ocean fertilization are regions with high nutrient and low chlorophyll (HNLC) concentrations, specifically where macronutrients (e.g., nitrogen, phosphorus, and dissolved silicon) are abundant, but phytoplankton growth is restricted by other factors, such as a lack of micronutrients (e.g., iron). The Southern Ocean, an HNLC environment where productivity is limited by low iron availability, is considered an ideal location for ocean iron fertilization (OIF) (199–201).

# Effectiveness and feasibility

From 1993 to 2011, 19 in situ OIF experiments tested the "iron hypothesis": twelve in the Southern Ocean, three in the Equatorial Pacific Ocean, one in the tropical Atlantic Ocean, and three in the Subarctic (North Pacific Ocean) (202). The results were inconclusive regarding the ability of OIF to increase the vertical export of POC to the deep ocean (203–205). In the Southern Ocean, OIF enhanced photosynthesis; however, an increase in carbon export to depth was documented in only one experiment (202). Export was monitored for only 2 to 6 weeks following iron addition, so it is not known whether longer-term observations would yield more positive results (202, 206, 207).

In terms of feasibility, full-scale deployment would require considerable effort and resources (see later)—involving an armada of ships designed to transfer and deposit iron at considerable oceanwide scale. The means for the creation and supply of suitable material, and docks to support its loading, would also need to be established. None of this is available at present, making the concept's feasibility questionable.

# Negative consequences

Scientists have expressed various concerns about the impacts of OIF on marine communities and ecosystems. While the addition of a limiting nutrient such as iron could enhance phytoplankton growth, there is no control over which species or groups of phytoplankton would be stimulated (205). This introduces significant direct uncertainty about the potential changes in phytoplankton species composition. Phytoplankton communities are typically a diverse mix of taxa. In the Southern Ocean, hundreds of species continually shift in terms of their standing stock and relative abundance. The groups and species that predominate can vary greatly in space and time, and diatoms, dinoflagellates, chrysophytes, and other groups can all be present simultaneously. Natural changes in light, temperature, macro- and micronutrients, and grazing pressure act synergistically to control species succession patterns. The species-specific impact of OIF within this mix of physical and biological controls is unknown and indeed unpredictable. Since phytoplankton cells and colonies range in size from a few microns to millimeters, changes in species composition due to competition and grazing could potentially restructure the size distribution within communities, affecting grazing efficiency among zooplankton and the nutritional value of the phytoplankton. Associated feedback effects on microbial communities in regions of OIF have often been observed (208). These types of changes at the base of a food web and in the microbial loop would affect energy transfer, trophic dynamics, and biogeochemical cycling throughout the Antarctic marine ecosystem and beyond.

Increased carbon export to depth is an intended consequence of OIF. Other (unintended) consequences could result from the trapping of nutrients below fertilized regions, alongside feedback effects, including deoxygenation and the production of other GHGs, e.g., methane and nitrous oxide (N2O). The "nutrient-robbing" impact of OIF has garnered much attention because of its potential broad impacts on net productivity elsewhere in the world. The Southern Ocean plays a significant role in the transport of nutrients to lower latitudes (209), effectively regulating rates of global primary productivity, the efficiency of the BCP, and the transfer of energy to higher trophic levels (including fisheries). While OIF stimulates productivity, it simultaneously depletes stocks of major nutrients, limiting the amount of nutrients transported and consequently decreasing net primary production in regions outside of the fertilization area. In the Southern Ocean, the "nutrient-robbing" signal is exported to lower latitudes by the Subpolar Mode Waters and the Antarctic Intermediate Water, leading to nutrient-depleted surface waters in the tropics (201). Reductions in biological production and organic matter export have been well documented in waters outside the fertilization areas (201, 210). Indeed, recent modeling work has highlighted that when OIF is undertaken in the Southern Ocean, the greater nutrient consumption, coupled with ongoing climate stressors (e.g., increased stratification), leads to a reduction in higher-trophiclevel organisms in the tropics (211). This cascading effect damages important fisheries in the tropical Pacific that are critical to coastal and island communities and for commercial exports—a major social equity concern. The consequences of "nutrient robbing" could be slightly mitigated if OIF were limited to the higherlatitude zones of the Southern Ocean, which are farther from where the intermediate and mode waters originate. However, this approach would provide a far smaller contribution to atmospheric CO<sub>2</sub> removal, while still having negative consequences (201, 211).

The biogeochemical effects of OIF include deoxygenation and the production of  $N_2O$  and methane (201). OIF increases oxygen consumption via the remineralization of the organic matter it produces. This compounds the deoxygenation already caused by global climate change through warmer waters and increased stratification.  $N_2O$  and methane are produced during the remineralization of organic matter produced by OIF, and both have a global warming potential greater than that of  $CO_2$  on a 100-year timescale (300- and 20-fold, respectively). While the production of methane is a minor concern, since its highest production only occurs in anoxic waters,  $N_2O$  generation following OIF ranges from 5% to 12% of the equivalent  $CO_2$  captured during a given OIF event (201).

#### Costs

The practicality of using OIF to mitigate anthropogenic carbon addition to the atmosphere has been questioned due to the difficulty of scaling it up to the magnitude that would be required for commercial application (212). Buesseler and Boyd (213) estimated that the effort needed (relative to the amount of iron used for fertilization) would be six orders of magnitude greater than that of any single experiment conducted in the Southern Ocean, and that the treated area would need to be 10 times larger than the Southern Ocean (which they defined as the area south of 50° latitude, equivalent to the area of Asia). The estimated cost of OIF per ton of CO<sub>2</sub> removed from the atmosphere varies by location, ranging from <US\$100/ton on the shelf to >US\$1,000/ton in offshore areas (214). For reference, global CO<sub>2</sub> emissions from fossil fuels in 2024 were 37.4 billion tons, and 41.6 billion tons when land use (e.g., deforestation) is accounted for (215).

#### Governance

One of the primary challenges of early OIF experiments lies in determining the legality of adding iron (a pollutant) to the ocean (216, 217). The UN addressed this issue in 1999, and OIF is now regulated by the London Protocol to the 1972 Convention on the Prevention of Marine Pollution by Dumping of Wastes and Other Matter (or the "London Convention") (218). A 2013 amendment to the London Convention states that "An ocean fertilization activity may only be considered for a permit if it is assessed as constituting legitimate scientific research taking into account any specific placement assessment framework" (219). Therefore, no large-scale OIF mitigation projects can be implemented. National authorities are responsible for permitting OIF experiments, but several other international agreements also address OIF (220). The UN Convention on Biological Diversity (CBD) does not currently support OIF and the Kyoto Protocol of the UN Framework Convention on Climate Change (UNFCCC) does not allow OIF to be used for carbon credits. OIF is not addressed by the UN Convention on the Law of the Sea (UNCLOS), although it might fall under the UNCLOS definition of marine pollution. The Antarctic Treaty System has not specifically considered OIF; it does have an annex to its Protocol on Environmental Protection on the prevention of marine pollution, but this pertains to waste from ship operations (e.g., food, trash, and oil). The potential impacts of OIF on Southern Ocean ecosystems, food webs, and fisheries would also be of concern to CCAMLR.

OIF is the only geoengineering scheme tested *in situ* in the Southern Ocean, with the last experiment being in 2011 (221). While there has been a long hiatus in OIF experimentation worldwide, there is currently a resurgence of interest in researching geoengineering methods for marine CO<sub>2</sub> removal (mCDR) to alleviate both carbon accumulation in the atmosphere and ocean acidification resulting from the diffusion of CO<sub>2</sub> into the ocean. For example, in 2023, the US National Oceanic and Atmospheric Administration announced US\$24.3 million in funding for such research endeavors (222), and the US Department of Energy (223) has allocated US\$36 million for this joint program. Moreover, new computer technologies (e.g., digital

twins) are being used in marine conservation management, including the planning of OIF strategies (224).

# Part 2: Debunking geoengineering arguments

Geoengineering proposals are often accompanied by a set of recurring arguments intended to justify their exploration or implementation, particularly in the face of growing climate risks and perceived inaction. These arguments typically highlight geoengineering as a necessary, feasible, or even inevitable tool in the broader climate response portfolio. In this section, we examine four of the most common arguments advanced by proponents of geoengineering, evaluating the assumptions they rely on and the implications they carry for science, policy, and governance.

# Argument 1: Mitigation is not happening fast enough

Current policies are insufficient to meet the emissions reductions required to keep warming below 1.5°C, as set out in the Paris Agreement. If countries fully implement their current policies, projections indicate a best-estimate warming of approximately 3°C by 2100 (41, 225, 226). If all current emissions reduction pledges are met, this estimate falls to around 2.5°C. However, countries have also communicated longer-term targets, which often aim to achieve net-zero CO<sub>2</sub> or total GHG emissions (227, 228). Taking these longer-term targets into account reduces the central warming estimate to 2.0°C (229). While this shortfall in climate action is concerning, it does not justify turning to geoengineering as a solution.

Although current policies and pledges do not yet guarantee 1.5°C compliance, they already provide a 1-in-5 chance of limiting long-term warming to 1.5°C, provided all commitments are met (226, 229). More importantly, these policies substantially increase the likelihood of limiting warming to below 2°C, with a 4-in-5 chance and a best estimate of 1.7°C. This underscores the need to accelerate mitigation efforts, rather than resorting to risky geoengineering approaches that fail to address the root causes of climate change.

Climate mitigation efforts are already scaling up, and rapid emissions reductions remain within reach using existing, proven technologies. One of the key flaws in the argument for geoengineering is that it underestimates the accelerating pace of green growth. Many clean technologies, such as renewable energy and electric vehicles, have already entered the rapid S-curve phase of adoption, where deployment scales exponentially as costs fall and efficiency improves (230, 231). Moreover, the Paris Agreement's 5-year ratcheting mechanism is designed to strengthen climate action over time, with countries currently required to submit more ambitious commitments following the first Global Stocktake. These mechanisms are expected to further increase the likelihood of limiting warming to as close to 1.5°C as possible.

Not only do proponents of geoengineering underestimate the potential for accelerated mitigation, but they also risk undermining it. By promoting unproven and potentially dangerous interventions, they distract from politically feasible, technically viable, and economically sound mitigation strategies that require stronger implementation urgently (47, 232).

# Argument 2: It is our moral duty to look at "all the options"

Although there is an intuitive appeal to the notion that "all the options" should be considered in the search for approaches to address anthropogenic warming and its consequences, it is also important to recognize that certain options can carry a significant negative effect, termed here as "moral hazard". Moral hazards are understood as actions or steps that cause individuals, organizations, or societies to increase their exposure to risk because they believe they are not fully responsible for, and/or subject to, the consequences (233). We argue that geoengineering can constitute a moral hazard and, more broadly, constitute "mitigation deterrence". Here, we highlight two moral hazards associated with geoengineering: complacency and predatory delay.

While geoengineering is often framed as one of many tools in a broad climate strategy, its introduction into policy discussions actively competes with mitigation for attention, funding, and legitimacy. Actors, whether individuals or collectives, can focus on only so many priorities at once, and policymakers may be drawn to geoengineering because it appears to offer a quick fix that avoids difficult political trade-offs. Termed "mitigation deterrence" (110), this effect has been discussed by other researchers (234, 235). The problem is not merely complacency but also the misdirection of political and financial resources away from proven, necessary solutions. If geoengineering becomes the central focus, it risks sidelining efforts to phase out fossil fuels, and rather than continuing to explore a wide range of options, policymakers may decide that geoengineering alone can offer sufficient protection against climate risks. They could fall victim to what is termed the "single action bias"—the closing of searches for alternative solutions, once one step has been taken (236). Flood mitigation plans provide examples of the negative consequences of single action bias. The city of New Orleans long relied on levees and flood walls to protect it from storm surges, and did not support other actions, including wetland restoration or expanded evacuation plans; when Hurricane Katrina struck in 2005, the results of the levee failures were catastrophic (237). Similarly, smaller cities along the Missouri River, such as Hamburg, Iowa, Saint Joseph, Missouri and Fremont, Nebraska, also trusted the levee systems built in the past for protection, until the floods in 2019 led to failure or overtopping of the levees; the residents, who would have benefited from additional measures including resilience planning and evacuation drills, faced severe damage (238, 239). Catastrophic flooding in Kolkata, India and Fukushima, Japan, also indicates the failure of relying on dams and embankments alone.

Behavioral studies at the household level have demonstrated this bias in various settings. For example, homeowners in flood-prone areas who are less likely to purchase insurance or elevate their homes if they have simply stocked up on water, flashlights, and batteries (240), or consumers in a cafeteria who undertake fewer behaviors to reduce food waste (e.g., ordering less food) if they are told that uneaten food will be composted (241).

Predatory delay differs from the form of delays discussed above, which are largely unintended consequences of the complacency generated by particular actions. First introduced by Steffen (242), predatory delay refers to deliberate efforts by powerful institutions to slow the implementation of actions that address the root causes of problems to preserve their own financial and political power. Although some individuals within institutions may have genuine motivation to seek alternative actions, the overall goal of the institution is to postpone actions that would be costly to them, even if the actions would have broad benefits to society and the environment. Kramer (243) details the interactions of large firms and lobbyists to influence policymakers and regulatory agencies to create such delays. There is a lack of transparency associated with the funding of some geoengineering projects, which could leave them open to interference by bad or rogue actors, including the fossil fuel industry (107). Muffett and Feit (244) have detailed the funding of geoengineering research by fossil fuel firms, showing how this research is being used to support the ongoing production and utilization of oil, gas, and coal for several decades to come. Almond et al. (245) documented how energy centers at US universities that receive funding from fossil fuel companies promote "climate actiondelaying tactics" that forestall "the transition to current renewable energy sources such as solar and wind" to a significantly higher extent than universities that refuse this type of funding. In this way, geoengineering could be compared to the tobacco industry's proposals for filter cigarettes as a way to reduce the risk of cancer without reducing the consumption of tobacco (246).

There is also a growing body of literature on how geoengineering funding (including from super-rich individuals and foundations), while being based on a perception of "moral duty", may lead to the establishment of undemocratic values and power relations (106), increased unilateralism (247), misconstrued morals and public acceptance (248, 249), and a lack of awareness regarding international conflict versus cooperation (250).

# Argument 3: Geoengineering will "buy us time" to adapt and find other solutions

Geoengineering approaches are frequently presented as temporary measures to delay climate change while societies develop long-term solutions. However, this argument fails to account for the contrasting temporalities between geoengineering deployment and mitigation efforts. Overwhelmingly, geoengineering concepts and methods are in the early-to-middle stages of research and design, meaning that their large-scale implementation remains speculative and probably decades away. Even if they worked as intended (which is highly uncertain), they would still require extensive political deliberation, governance

frameworks, ethical considerations, resource allocation, and infrastructure development before deployment. Therefore, geoengineering will not "buy us time" in the near term; rather, it diverts time and limited resources toward uncertain technological fixes that are unlikely to be implemented soon, if at all. In contrast, many mitigation technologies already exist and are ready to be scaled up locally and globally. Investing in mitigation is a direct response to the urgency of addressing climate change, while geoengineering remains a gamble, with long- and uncertain-time horizons. Therefore, it makes more sense to invest in improving and scaling existing technologies that reduce emissions than to work on speculative technologies that do not address emissions. With an increased focus on the implementation of existing technologies, we can make substantial global progress on reducing emissions, with limited need for additional new technological solutions.

# Argument 4: Geoengineering will prevent tipping points from being crossed

The concept of "tipping points" in the climate system has been a subject of debate for several years. While many studies have emphasized the severe risks associated with crossing these thresholds (147, 232, 251, 252), others have argued that the focus on tipping points can create confusion and, potentially, lead to poor decision-making regarding climate action and environmental protection (253).

The polar regions are home to, or have a direct influence on, four of the five most vulnerable climate tipping points: Greenland Ice Sheet collapse, West Antarctic Ice Sheet collapse, abrupt permafrost thaw, and the collapse of the Labrador–Irminger Subpolar Gyre (42, 251). These Earth system components are at risk of crossing their tipping points with ~1.5°C of warming, and the risk increases as temperatures rise. Some may have already crossed these thresholds (152, 232, 254, 255).

While several studies have framed SRM as a means to delay the crossing of potential tipping points (256–259), others have indicated that it is less effective than GHG mitigation strategies due to the uncertainties and risks involved (260). Moreover, these studies have relied on simplified scenarios that neglect sociopolitical factors that could cause harm or raise ethical dilemmas, which would further restrict the potential of geoengineering as an emergency solution (261–263).

Geoengineering interventions face significant challenges due to the limited timeframe for their deployment. The urgent need for immediate action to prevent the crossing of tipping points necessitates the rapid implementation of strategies. However, the development and application of geoengineering methods would require extensive research, planning, and international cooperation, making it impossible to meet the narrow timeframe available.

Given the substantial uncertainties and challenges associated with geoengineering, it should not be considered a reliable solution for preventing the crossing of Earth system tipping points. Therefore, the focus should remain on aggressive emission reduction strategies and robust adaptation measures. The necessarily limited amount of public research funding available needs to be focused on developing a deeper understanding of Earth system tipping points to minimize associated risks, and thus on the fundamental understanding of the natural systems involved, rather than on researching the impacts of large-scale interventions in these poorly known systems.

# Part 3: International governance and decision-making

In Part 1, we touched on governance matters that would need to be overcome before the deployment of any geoengineering intervention. In this section, we discuss more broadly the existing international governance arrangements in the polar regions and whether or how they are set up to perform an essential role in preventing or setting strict terms for these interventions. Notably, geoengineering presents unique obstacles to effective governance that are not found in other areas of international environmental collaboration (e.g., transnational river basin management or stratospheric ozone management). First, the approaches differ greatly between the technology-centered geoengineering advocates and the institutional and legal frameworks of international regulatory agencies, so they lack common sets of references on which to base planning. Second, stakeholders who seek approval from governance authorities operate on short-term timeframes, while researchers and government regulators adopt medium- and long-term timeframes to address the numerous uncertainties in this area. Finally, geoengineering promoters emphasize trade-offs between risks, which conflict with the precautionary priorities of environmental governance. Together, these barriers severely impede the discussions that are a necessary precursor for effective governance (264).

## Antarctica and the Southern Ocean

Antarctic governance is managed through the Antarctic Treaty System, which uses a consensus-based model for decision-making (265, 266). The Antarctic Treaty, which entered into force in 1961 and currently has 58 signatory states, sets Antarctica aside as a continent for peace and science. The Protocol on Environmental Protection to the Antarctic Treaty (hereafter "the Protocol"), which entered into force in 1998, states that activities in the Antarctic Treaty area (the areas south of latitude 60°S) must be planned and conducted to limit adverse impacts on the Antarctic environment and dependent and associated ecosystems, and to avoid, among other things (i) adverse effects on climate or weather patterns; (ii) significant adverse effects on air or water quality; (iii) significant changes in the atmospheric, terrestrial (including aquatic), glacial or marine environments; and (iv) detrimental changes in the distribution, abundance or productivity of species or populations of species of fauna and flora [Article 3(2a and b) (133)]. The Protocol also states that activities in the Antarctic Treaty area

must be planned and conducted on the basis of information sufficient to allow prior assessments of, and informed judgments about, their possible impacts on the Antarctic environment [Article 3(2c) (133)]. If there is insufficient information on the likely impacts, the use of geoengineering technologies would be contrary to the Protocol and consequently unlikely to receive authorization to proceed.

All activities undertaken within the Antarctic Treaty area must be subject to an Environmental Impact Assessment (EIA), as stipulated in Annex I to the Protocol (267). The level of the EIA prepared by the Party undertaking the activity will depend on whether the anticipated activity is likely to have an impact that is less than, equal to, or greater than "minor or transitory". The Protocol does not define these terms; however, any geoengineering activities would almost certainly trigger the highest level of EIA, known as a Comprehensive Environmental Evaluation (CEE). To date, CEEs have been undertaken for activities such as station construction, but some proposed geoengineering projects are of a scale several orders of magnitude greater. The Protocol also sets out monitoring requirements to ensure that the impacts of the activity are consistent with the Protocol, and to provide information that is useful for minimizing or mitigating impacts, and, where appropriate, information on the need for suspension, cancellation, or modification of the activity (Annex I, Article 5).

International consultation is key to Antarctic governance and, in a system driven by consensus-based decision-making, controversial or unproven geoengineering activities have a very low likelihood of approval. Promoters of geoengineering from the ~139 countries that have not acceded to the Antarctic Treaty (representing ~35% of the global population) are not obligated to comply with any of these requirements. However, non-Treaty parties are unlikely to be able to support independent expeditions to the continent, and the use of large-scale geoengineering methods without sufficient consultation with the Antarctic Treaty Consultative Meeting (ATCM) would probably result in an international outcry (134, 268), as the ATCPs closely guard the status of the Treaty and their rights as Parties.

In practice, no Party project for which a CEE has been presented to the ATCPs has been prevented from going ahead (albeit sometimes with significant delays), including some projects involving what elsewhere might be considered high-impact activities (269). Given the potential global impact of using large-scale geoengineering methods in Antarctica and the international effort needed to deliver them, any concerns may not only be raised in the ATCM system, but may also face challenges outside it (e.g., via the UNFCCC, UN Environment Programme, or UN General Assembly), which the existing ATCPs would want to avoid.

Under Article 7 of the Protocol, which bans mineral resource activities except for scientific research, the ATCM has blocked commercial extraction in Antarctica. Given the large scale of proposed Antarctic geoengineering, the ATCM could take similar action. Almost 20 years after its agreement, Annex VI to the Protocol (Liability arising from Environmental Emergencies; 2005) has yet to enter into force. Once enacted, however, it would impose a liability on non-state operators to pay for a "response"

(clean-up) for environmental damage caused. Even if private, rather than state, operators stepped in to undertake geoengineering activities, the requirement for a CEE and to obtain a permit from the relevant Party would remain the same.

Overall, the prospects of addressing proposals for geoengineering activities in Antarctica depend on the environmental impact assessments conveyed to the Antarctic Treaty System and the political will of its members or parties to the agreements to give effect to the obligations they have set themselves. The Antarctic Treaty System's agreements (including CCAMLR with respect to marine living resources) constitute a set of rigorous governance arrangements that are wholly underestimated by geoengineering proponents.

At the most recent ATCM and CEP meeting (Milan, 23 June - 3 July 2025 (270)), on the basis of a paper by SCAR, the CEP advised the ATCM that a precautionary approach should be taken towards geoengineering activities and that, at this point, geoengineering methods in the Antarctic should not be conducted due to their unknown environmental consequences. The ATCM endorsed the recommendation to adopt a precautionary approach towards all geoengineering activities in the Antarctic area. Most ATCPs supported SCAR's conclusion that, based on the current state of knowledge, none of the options outlined in the papers met essential criteria to be considered safe, responsible or feasible and should not be considered for Antarctica. They also emphasised that geoengineering should not be considered as a solution to climate change in Antarctica.

### The Arctic

Unlike the Antarctic, the Arctic region falls predominantly within national jurisdictions (see: https://www.durham.ac.uk/media/ durham-university/research-/research-centres/ibru-centre-forborders-research/maps-and-databases/arctic-maps-2024-january/ Map-1-IBRU-Arctic-map-04-01-24-(revised-Russia-claimed).pdf). There are four "high seas" areas, as defined by UNCLOS, including the central Arctic Ocean itself; however, this represents a small fraction of the total Arctic Ocean area. While the governance of activities and their impacts at the surface and in the water column within high seas areas resides with a variety of existing international mechanisms, such as the IMO and the Convention for the Protection of the Marine Environment of the North-East Atlantic (OSPAR) and the emerging UN High Seas Treaty, the governance of activities and their impacts at and below the seafloor may change in the future. Many coastal Arctic states are conducting long-term research to establish the northward reach of their continental shelf, thereby seeking to provide evidence to extend their national seabed boundaries, in many cases out into the central Arctic Ocean. This is all in line with established UNCLOS requirements and practices. There is no timetable for the resolution of claims arising from the submission of this evidence.

All land in the Arctic is within nationally regulated and internationally agreed-upon boundaries. Consequently, the dominant legal and regulatory governance mechanisms, including those that would relate to geoengineering, are overwhelmingly those of the eight

Arctic states themselves: the United States, Canada, the Kingdom of Denmark (including Greenland), Iceland, Norway, Sweden, Finland, and Russia. The major international organization in the region is the Arctic Council, created in 1996 to provide a high-level forum for promoting cooperation, coordination, and interaction among the Arctic states and the Permanent Participants (PPs), namely, six representative organizations of Indigenous peoples. State and nonstate Observers also participate in discussions at the discretion of Member States and PPs. The Council offers a space for discussion and mechanisms for addressing issues of common concern, with a special emphasis on protecting the Arctic environment and promoting sustainable development. Established by the Ottawa Declaration, the Council is not a Treaty organization under the terms of the Vienna Convention—unlike the Antarctic Treaty. There are well-established procedures for both state and non-state organizations to apply for Observer status with the Arctic Council, although approval requires consensus from all eight Arctic States. Unlike Antarctica, where no state has legal jurisdiction, there is no opportunity for states to unilaterally join the Council, which by definition is limited to states with territory within the Arctic Circle.

The Arctic Council has strongly addressed governance issues across the Arctic internationally, although rapid climate change and recent geopolitical tensions have placed it under unprecedented pressure (271). Nevertheless, history shows that cooperation on scientific and environmental issues is vulnerable to disruption by overall distrust and geopolitical dynamics between member states (272). Following Russia's full-scale invasion of Ukraine in 2022, the activities of the Arctic Council's six Working Groups halted as all member states except Russia announced their non-participation. Under Norway's current chairship (2023-2025), significant working-level activity has resumed across the Working Groups. However, the organization has yet to resume in-person ministerial and senior officials' meetings involving all eight members. With the new United States administration of 2025, the coming years are likely to remain challenging, with increasing geopolitical uncertainty and persisting tensions with no clear resolution in sight.

The Arctic Council is not a regulatory or law-making body, as legislative authority remains within the national jurisdictions of its member states. The Council has never sought, nor do its members intend for it to have, the power to enforce collective decisions. Other organizations, including the North East Atlantic Fisheries Commission, have the authority to adopt conservation and management measures across portions of the Arctic Ocean, but only within the scope of agreed-upon national jurisdiction. While the Council's work has covered an extremely wide range of issues connected to environmental, ecosystem, and human change in the Arctic—assessing evidence and producing guidelines and best practices—it has chosen not to look specifically at geoengineering.

Under the auspices of the Arctic Council, the eight Arctic states have negotiated the following legally binding international agreements: maritime search and rescue; marine oil pollution prevention and response; and Arctic science cooperation. These are implemented by the states within their national jurisdictions,

not through Council mechanisms. These agreements may have relevance to proposals for geoengineering activities in the Arctic, including their prohibition in the region.

While the Arctic Council has relatively weak oversight of international waters where, for example, any proposed silicate spheres could be located, these efforts almost certainly would require activity in national (or asserted national) waters. Objects that drift from international waters into the territorial waters of a member state (273) come under that state's jurisdiction. Any decision for the Arctic Council to address geoengineering would require both the championing of an incoming chairship (e.g., Denmark in 2025–2027 or Sweden in 2027–2029) and the consensus of all eight Arctic states during the priority-setting process ahead of each 2-year rotating chairship.

As many parties view the Arctic as an opportunity for economic growth and resource extraction, and as a zone of geopolitical tension, there is little prospect of long-term multilateral and low-tension collaboration on geoengineering at present (272). While governance responsibility predominantly lies with the individual Arctic States—either individually or collectively—the effects and impacts of geoengineering cannot be guaranteed to come from or stay within such territorial boundaries.

In contrast to the Antarctic, the Arctic is home to Indigenous communities whose traditional way of life is under increasing pressure from climate change and social and cultural challenges associated with geopolitical and post-colonial development. Strategies that seek to engineer a reduction in the rate of sea level rise via glacier modifications address neither climate impacts nor local cultural challenges. In many cases, these interventions likely worsen local socioeconomic and environmental conditions by potentially putting further pressure on activities such as marine harvesting of fish and mammal species with, for example, increased underwater noise and disturbance of wildlife populations on land, while offering no substantial benefit to local communities. On the other hand, resource extraction of minerals to support the green transition and agriculture, as well as the development of renewable power resources which the Arctic region is rich in, can benefit both local and distal populations and represent a far more sustainable path for climate change mitigation both locally and globally with additional economic benefits for local Arctic populations if implemented sensitively.

It is pertinent to ask: *why* would a nation such as Greenland embrace a geoengineering solution to sea level rise, for example? The sea level is not rising there—rather, the loss of ice sheet mass is causing gravity-induced sea level lowering (274). Moreover, Greenland's contribution to global GHG emissions is small and its local ecosystems would be at risk from geoengineering.

In summary, both the Arctic and the Antarctic lack the effective governance mechanisms needed to offer protections against the significant potential harms of geoengineering. We wish to underscore this commonality, while recognizing the significant differences between the two regions—the greater strength of national institutions in the Arctic; the greater dominance of a single treaty in the Antarctic; the wider range of economic

interests in the Arctic; and the absence of historic Indigenous peoples in the Antarctic, just to mention a few.

# Part 4: Protecting the polar regions without geoengineering

In this section we offer an alternative approach to protect the polar regions that does not rely on physical interventions in natural systems or the intentional use of pollutants. We do not propose these as isolated solutions or quick-fix interventions, but rather as components of a transformation of humanity's relationship to our planet, known as "climate-resilient development". This transformation comprises shifts in major systems, such as energy systems, land use systems, and food systems, which can be best addressed jointly (275). We focus on two components particularly relevant to polar regions: decarbonization and maintenance of protected areas.

#### Decarbonization

Rapid decarbonization to "net-zero" emissions by mid-century, and into "negative emissions" thereafter will lead to a rapid and permanent climate response (the so-called "Zero CO<sub>2</sub> Emissions Commitment" or ZEC) (276). This decarbonization will require significant changes in energy, land use, and food systems. The global warming climate system is likely to stabilize (i.e., global surface temperatures cease increasing) within 20 years of net zero CO<sub>2</sub> (43). This assessment is good news for the polar regions across various systems and realms. Not all changes can be stopped, but their rates of change would be comprehensively reduced if and when net-zero emissions are reached.

Two investigations have underlined the environmental benefits of net-zero CO2 and climate stabilization. Rintoul et al. (21) "hindcasted" two scenarios from the perspective of 2070: one with unabated emissions and a second with ambitious decarbonization leading to net-zero emissions. In the latter scenario, the global temperature stabilized at 0.9°C above pre-industrial levels, as opposed to 2.6°C in the former (similarly for Antarctic air temperature). These scenarios were matched by effective governance of Antarctica and the Southern Ocean under ambitious decarbonization, as opposed to ineffective policies. The relative benefits of decarbonization, based on literature and expert opinion, proved significant in all cases (21). The Antarctic contribution to sea level rise by 2070 was limited to 6 cm, as opposed to 27 cm under high emissions [see also (277)]. Southern Ocean temperatures increased by only 0.7°C as opposed to 1.9°C. Sea ice loss was restricted to 12% of present levels, as opposed to 43%. Ice shelf volume was reduced by only 8% as opposed to 23%. In terms of ocean acidification, decarbonization retained surface super-saturation for aragonite, while continued fossil fuel burning and deforestation led to waters that are corrosive to the aragonite shells of pteropods. Biological invasions were restricted to twice those of today, versus tenfold. Ecosystem failure occurred under both scenarios, but under unabated fossil fuel burning, penguins and krill were adversely affected. With enhanced policy, far fewer people were present in the region, with an accompanying reduction in marine resource extraction. Rintoul et al.'s (21) simple vision for Antarctica is appealing, as it explains how every polar system can be helped by decarbonization, and certainly to the extent that direct intervention (or geoengineering) would be unnecessary.

A similar assessment of the Antarctic Peninsula explained the environmental consequences of limiting global warming to 1.5°C relative to pre-industrial levels (278). Such an outcome, under rapid deep decarbonization to net-zero CO<sub>2</sub> globally by mid-century, would benefit all realms of this environment. This study compared the present situation (1.2°C in 2019) and that at 1.5°C, revealing the magnitude of the "best case" changes. For example, in a 1.5°C scenario the number of days per year when surface air temperatures exceed 0°C would increase from 25–80 to 35–130. While the resulting melting may lead to ponding of water on floating ice shelves, a dramatic loss of ice shelves is not expected. Sea ice would be more limited, but it would remain around the Peninsula, especially in the South. Thus, the sea ice and krill-dependent food web would migrate southward, leaving the northern Peninsula with an increase in fur seals, elephant seals, and gentoo penguins.

While such changes are unwelcome, being an indirect but obvious consequence of fossil fuel burning, they alone would not necessitate an artificial geoengineering intervention at the scales being proposed, and none of the interventions would address these changes. Further warming would probably bring an increase in "extreme events", such as atmospheric and marine heatwaves (15) and continued sharp reductions in sea ice. However, while much of the natural polar environment would be altered or displaced, it would still be present in a manner deserving of environmental protection (e.g., with respect to its marine biodiversity). Indeed, under the ZEC scenario, once net zero has been achieved negative emissions would rapidly (within ~20 years) act to cool the planet, potentially reversing changes under 1.5°C warming. This scenario is already technically, financially, and politically possible (51, 279), and it represents our best means of protecting the Antarctic environment, with major co-benefits for human health and planetary habitability.

The Arctic situation may be profoundly different. Continued global warming will result in enhanced regional warming here, leading to further sea-ice loss, permafrost thawing, and ecosystem damage, all of which will adversely affect the livelihoods, cultures and living environments of Indigenous peoples. Global climate stabilization offers the best chance to protect the Arctic; however, significant further change may already be unavoidable and, as this paper highlights, geoengineering is not a solution.

# Protected areas across the polar regions

Well-managed protected areas, especially those that are highly protected—the equivalent of International Union for Conservation of Nature (IUCN) Categories I or II (280)—significantly benefit biodiversity and ecosystem services. These benefits include higher species richness, abundance, and functional diversity within protected areas and ecological spillover that benefits biodiversity outside the protected area. Substantial evidence exists for these effects on land (281–284). Marine protected areas have similar benefits: evidence-based syntheses and specific studies demonstrate higher biomass/abundances and functional diversity of species within strictly protected areas (compared with outside), along with ecological and fishery spillover (i.e., the outward emigration of individuals that benefits a fishery) (285–288).

Although climate change threatens their effectiveness (289, 290), protected areas provide significant benefits to biodiversity under climate change in both terrestrial and marine systems (291–293). Moreover, protected areas can mitigate the effects of climate change itself, by ensuring the maintenance of ecosystem services such as carbon sequestration (294), and enhancing ecosystem resilience by minimizing other impacts (28). Consequently, protected areas remain a central component of the mitigation response to climate change impacts on biodiversity. Key priorities include various strategies to grow, connect, and manage the network of protected areas, and to improve their overall individual efficacy (295–298).

The climate change mitigation benefits of protected areas have also been recognized in the Antarctic and the Arctic (22, 299); however, it is important to acknowledge the substantial differences between the two regions (300). Although protected areas cannot entirely mitigate impacts such as ocean acidification (301), they can play an important role. Moreover, reducing GHG emissions will mitigate impacts such as ocean acidification, whereas geoengineering activities that fail to reduce atmospheric CO<sub>2</sub> concentrations will not, leaving biodiversity and ecosystem services exposed to a growing threat.

To be effective in a climate change impact mitigation framework, polar protected areas will need broader coverage of biodiversity features and areas, greater connectivity, and more flexibility in their boundaries in response to changing ecological circumstances. Their greatest challenge is ultimately political. For example, the ATCPs and the Members of the CCAMLR have been slow to designate protected areas (302), and there are growing difficulties in protected area discussions in both of these Antarctic Treaty system bodies (303). However, previous difficulties in international environmental agreements have been overcome to great success, such as the 1987 Montreal Protocol to phase out the production of substances responsible for ozone depletion and the 2016 CCAMLR agreement to designate the Ross Sea region as a Marine Protected Area.

# Summary

Over the past three years, polar geoengineering has gained increased media and public attention, with various ideas largely based on computer modeling and limited field studies aiming to increase Arctic sea ice, enhance atmospheric reflectivity, halt the transfer of warmth from the ocean to the grounded ice sheet, and slow the flow of grounded ice into the ocean. In this paper, we have critically assessed five such approaches, identifying significant challenges across six key categories.

# 1 Scientific feasibility

All five proposed concepts are based on the assumption that natural systems can be forcibly controlled through intervention. However, each method has significant oversights that raise serious doubts about whether implementation would achieve the desired outcomes. A key example is Proposed Concept 4—Slowing Ice Sheet Flow—which involves drilling to the base of ice streams to remove basal water and reduce ice flow. Even if water is successfully extracted, it is likely to be replaced by groundwater or redirected from other sources, requiring continuous drilling of new holes. Our limited understanding of Antarctica's subglacial system, and polar systems in general, makes effective planning for such efforts highly uncertain.

#### 2 Environmental risks

One of the central themes throughout this paper is the significant and compounding uncertainties associated with the climate and biogeophysical impacts of the five reviewed geoengineering options. These uncertainties span direct, indirect, and systemic effects and remain vast and largely unquantified. In some cases, the proposed geoengineering intervention uses a direct pollutant, such as in the case of glass microbeads used to increase Arctic sea ice albedo, the SO2 used for atmospheric aerosol injection, and the iron used for ocean fertilization. In other cases, the interventions disrupt weather patterns and the circulation of ocean currents and nutrients. Without rigorous quantification of error bounds and a significant reduction in these uncertainties, it is challenging to foresee a scenario in which decision-makers could proceed with confidence. In fact, the inability to resolve fundamental uncertainties about the impacts may represent one of the most compelling arguments against deploying geoengineering interventions at present.

#### 3 Financial costs

In addition to environmental costs, there are significant and potentially prohibitive financial costs associated with geoengineering. All of the proposed ideas would require funding at a scale of many billions of US dollars, which is presently unavailable through either national or international arrangements. It is unknown where the funds needed to pursue geoengineering interventions would come from, but they would dwarf present scientific budgets. While fundamental research in glaciology contributes to our understanding of climate dynamics, informs mitigation strategies, and supports adaptation planning, large-scale geoengineering research seeks to develop

interventions that carry high levels of uncertainty and risk. If these funds were made available, they would be far more effectively spent on fundamental climate research and efforts to decarbonize rapidly.

# 4 Governance challenges

Currently, no legal or governance framework supports these geoengineering proposals. In the case of Antarctica, such arrangements are offered through the CEP and ATCM. At its most recent meeting the CEP advised that geoengineering methods in the Antarctic should not be conducted. The ATCPs have largely taken this advice noting that such geoengineering has no place in the region, though with some also noting that any such proposals would need formal CEEs before commencing any work. Furthermore, any expansion of governance arrangements in both the Arctic and the Antarctic hinges on the consent of the majority of or all affected states. Current research on the interaction between international politics and geoengineering indicates that the latter will either prove to be politically infeasible or could even contribute to rising geopolitical tensions if deployed unilaterally. Additionally, while reducing carbon emissions is both feasible and undeniably beneficial, international coordination on this effort has been slow and challenging. Given these difficulties, it is even less likely that consensus could be achieved on geoengineering interventions that carry far greater uncertainties and potential risks.

## 5 Scale and time constraints

Limiting the escalation of severe climate-related risks demands that we cut GHG emissions to net zero by mid-century. It has been suggested that geoengineering approaches could be used as a temporary measure to "buy time" while societies develop longterm climate solutions. However, this argument overlooks the stark contrast in timelines between geoengineering and mitigation efforts. Most geoengineering approaches are still in the early stages of research, with large-scale implementation probably decades away, if they prove viable at all. Even assuming they work as intended, their deployment would require extensive political negotiations, governance structures, and infrastructure investments, making them an uncertain and delayed response to climate change. In contrast, many mitigation technologies already exist and can be rapidly scaled up. Investing in proven emissions-reduction strategies directly addresses climate change now, whereas geoengineering diverts attention and resources toward speculative solutions with uncertain outcomes.

# 6 The risk of false hopes

Geoengineering proposals offer false hope that the effects of global warming can be avoided by means other than rapid, deep cuts to GHG emissions. Two key risks are (i) complacency, in which

decision-makers focus on geoengineering at the expense of proven decarbonization strategies, and (ii) predatory delay, in which powerful actors may promote geoengineering to justify continued emissions and preserve their own financial or political interests under the pretense of climate action.

All five proposed geoengineering interventions we have discussed here are extremely unlikely to mitigate the effects of global warming in polar regions and are likely to have serious adverse and unintended consequences. The only realistic and effective approach is rapid and sustained decarbonization to net zero. Given the limited time available, our full attention must remain on proven strategies rather than speculative interventions with unproven benefits and substantial risks.

# Supplementary material

The Supplementary Material for this article can be found online at: https://www.frontiersin.org/articles/10.3389/fsci.2025.1527393/full#supplementary-material

## **Statements**

## **Author contributions**

Writing – original draft: MSi, HS, MJB, JB-G, HB, SB, MC, SLC, FC, RMD, HAF, EG, SMG, AMG, SH, KRH, SFH, RH, KAH, DK, JDK, BK, RDL, AM, VM-D, FSM, HM, RM, TAM, TN, CN, BO, PP, JoR, JaR, SS, AS, MSo, LAS, CRS, JS, MT.

Writing – review & editing: MSi, HS, MJB, JB-G, HB, SB, MC, SLC, FC,RMD, HAF, EG, SMG, AMG, SH, KRH, SFH, RH, KAH, DK, JDK, BK, RDL, AM, VM-D, FSM, HM, RM, TAM, TN, CN, BO, PP, JoR, JaR, SS, AS, MSo, LAS, CRS, JS, MT.

Conceptualization: MSi, HS.

Funding acquisition: MSi.

Investigation: MSi, HS, MJB, JB-G, HB, SB, MC, SLC, FC,RMD, HAF, EG, SMG, AMG, SH, KRH, SFH, RH, KAH, DK, JDK, BK, RDL, AM, VM-D, FSM, HM, RM, TAM, TN, CN, BO, PP, JoR, JaR, SS, AS, MSo, LAS, CRS, JS, MT.

Project administration: MSi.

Validation: MSi, HS, MJB, JB-G, HB, SB, MC, SLC, FC,RMD, HAF, EG, SMG, AMG, SH, KRH, SFH, RH, KAH, DK, JDK, BK, RDL, AM, VM-D, FSM, HM, RM, TAM, TN, CN, BO, PP, JoR, JaR, SS, AS, MSo, LAS, CRS, JS, MT.

Visualization: MJB, MSi.

# Data availability statement

The original contributions presented in the study are included in the article/supplementary material. Further inquiries can be directed to the corresponding author.

# **Funding**

The authors declared financial support was received for this work.

SC, AM, and FM received support from the Australian Research Council (ARC) Special Research Initiative (SRI), Securing Antarctica's Environmental Future (SAEF, no. SR200100005). FM also acknowledges funding from the ARC Discovery Early Career Research Award (no. DE210101433).

MJB's involvement was supported by funding received from the European Research Council (ERC) under the European Union's (EU) Horizon 2020 research and innovation program (grant agreement no. 885205).

JS was supported by the Canada 150 Research Chairs program, C150 grant no. 50296.

AG is a member of the Carrera del Investigador Científico, CONICET and was partially supported by the Argentine grants no. PICT 2019-02754 (FONCyT-ANPCyT) and no. UBACyT-20020190100247BA (UBA).

RH received funding from the Norwegian Research Council, Project no. 324131, ERC-2022-ADG grant no. 01096057 GLACMASS and National Aeronautics and Space Administration (NASA) grant no. 80NSSC20K1296.

TN was funded by the Antarctic Science Platform Contract - ANTA1801.

RM was supported through the Horizon Europe-funded OCEAN: ICE project, which is co-funded by the EU program for research and innovation under grant agreement no. 101060452 and by United Kingdom Research and Innovation (UKRI).

VM-D received a Synergy Grant from the ERC under the EU's Horizon 2020 research and innovation program (AWACA: Atmospheric WAter Cycle over Antarctica: Past, Present and Future, grant agreement no. 951596).

RDL received funding from the US National Science Foundation through a subaward to BAS (No. 1556528) for management of the International Thwaites Glacier Collaboration Science Coordination Office.

# Conflict of interest

JRu is employed by the Polar Regions Department of the United Kingdom Foreign Commonwealth and Development Office (FCDO).

KRH is the Chair of Antarctic Science Ltd. and the Honorary Secretary of the Challenger Society for Marine Science.

SC is a Past President and current Honorary Life Member of the Scientific Committee on Antarctic Research (SCAR).

CN is currently Executive Director of SCAR.

VM-D is a member of the French High Council on Climate, an independent body tasked with issuing advice and recommendations to the French government.

HB is the President of the International Arctic Science Committee 2022–26.

Neither the organizations mentioned in this statement nor any of the article's funders were involved in the study design, collection, analysis, interpretation of data, the writing of this article, or the decision to submit it for publication.

The remaining authors declared that this work was conducted in the absence of any financial relationships that could be construed as potential conflicts of interest.

The authors MS, JR, LS and MT declared that they were an editorial board member of Frontiers at the time of submission. This had no impact on the peer review process and the final decision.

#### Generative AI statement

The authors declared that no generative AI was used in the creation of this manuscript.

Any alternative text (alt text) provided alongside figures in this article has been generated by Frontiers with the support of artificial intelligence and reasonable efforts have been made to ensure accuracy, including review by the authors wherever possible. If you identify any issues, please contact us.

## Publisher's note

All claims expressed in this article are solely those of the authors and do not necessarily represent those of their affiliated organizations, or those of the publisher, the editors and the reviewers. Any product that may be evaluated in this article, or claim that may be made by its manufacturer, is not guaranteed or endorsed by the publisher.

# References

- 1. Betts RA, Belcher SE, Hermanson L, Klein Tank A, Lowe JA, Jones CD, et al. Approaching 1.5°C: how will we know we've reached this crucial warming mark? *Nature* (2023) 624(7990):33–5. doi: 10.1038/d41586-023-03775-z
- 2. Forster PM, Smith C, Walsh T, Lamb WF, Lamboll R, Hall B, et al. Indicators of global climate change 2023: annual update of key indicators of the state of the climate system and human influence. *Earth Syst Sci Data* (2024) 16(6):2625–58. doi: 10.5194/essd-16-2625-2024
- 3. World Meteorological Organization. WMO confirms 2024 as warmest year on record at about 1.55°C above pre-industrial level [online] (2025). Available at: https://wmo.int/news/media-centre/wmo-confirms-2024-warmest-year-record-about-155degc-above-pre-industrial-level
- 4. Rantanen M, Karpechko AY, Lipponen A, Nordling K, Hyvärinen O, Ruosteenoja K, et al. The Arctic has warmed nearly four times faster than the globe since 1979. Commun Earth Environ (2022) 3(1):168. doi: 10.1038/s43247-022-00498-3
- 5. Meier WN, Hovelsrud GK, Van Oort BE, Key JR, Kovacs KM, Michel C, et al. Arctic sea ice in transformation: a review of recent observed changes and impacts on biology and human activity. *Rev Geophys* (2023) 61(1):185–217. doi: 10.1002/2013RG000431
- 6. Notz D, Stroeve J. The trajectory towards a seasonally ice-free Arctic Ocean. Curr Clim Change Rep (2018) 4(4):407–16. doi: 10.1007/s40641-018-0113-2
- 7. Mudryk L, Santolaria-Otín M, Krinner G, Ménégoz M, Derksen C, Brutel-Vuilmet C, et al. Historical Northern Hemisphere snow cover trends and projected

changes in the CMIP6 multi-model ensemble. Cryosphere (2020) 14(7):2495-514. doi: 10.5194/tc-14-2495-2020

- 8. Biskaborn BK, Smith SL, Noetzli J, Matthes H, Vieira G, Streletskiy DA, et al. Permafrost is warming at a global scale. *Nat Commun* (2019) 10(1):264. doi: 10.1038/s41467-018-08240-4
- 9. Kavan J, Szczypińska M, Kochtitzky W, Farquharson L, Bendixen M, Strzelecki MC. New coasts emerging from the retreat of Northern Hemisphere marine-terminating glaciers in the twenty-first century. *Nat Clim Chang* (2025) 15(5):528–37. doi: 10.1038/s41558-025-02282-5
- 10. Otosaka IN, Shepherd A, Ivins ER, Schlegel N-J, Amory C, van den Broeke MR, et al. Mass balance of the Greenland and Antarctic ice sheets from 1992 to 2020. *Earth Syst Sci Data* (2023) 15(4):1597–616. doi: 10.5194/essd-15-1597-2023
- 11. Kim IW, Timmermann A, Kim JE, Rodgers KB, Lee SS, Lee H, et al. Abrupt increase in Arctic-subarctic wildfires caused by future permafrost thaw. *Nat Commun* (2024) 15(1):7868. doi: 10.1038/s41467-024-51471-x
- 12. Myers-Smith IH, Kerby JT, Phoenix GK, Bjerke JW, Epstein HE, Assmann JJ, et al. Complexity revealed in the greening of the Arctic. *Nat Clim Chang* (2020) 10(2):106–17. doi: 10.1038/s41558-019-0688-1
- 13. Nuttall M. An environment at risk: Arctic indigenous peoples, local livelihoods and climate change. In: Ørbæk JB, Kallenborn R, Tombre I, Hegseth EN, Falk-Petersen S, Hoel AH, editors. *Arctic alpine ecosystems and people in a changing environment*. Berlin: Springer-Verlag (2007) 19–35. doi: 10.1007/978-3-540-48514-8\_2
- 14. Casado M, Hébert R, Faranda D, Landais A. The quandary of detecting the signature of climate change in Antarctica. *Nat Clim Chang* (2023) 13(10):1082–8. doi: 10.1038/s41558-023-01791-5
- 15. Siegert MJ, Bentley MJ, Atkinson A, Bracegirdle TJ, Convey P, Davies B, et al. Antarctic extreme events. Front Environ Sci (2023) 11:1229283. doi: 10.3389/fenvs.2023.1229283
- 16. Gorodetskaya IV, Durán-Alarcón C, González-Herrero S, Clem KR, Zou X, Rowe P, et al. Record-high Antarctic Peninsula temperatures and surface melt in February 2022: a compound event with an intense atmospheric river. *NPJ Clim Atmos Sci* (2023) 6(1):202. doi: 10.1038/s41612-023-00529-6
- 17. Wille JD, Alexander SP, Amory C, Baiman R, Barthélemy L, Bergstrom DM, et al. The extraordinary March 2022 East Antarctica "heat" wave. Part I: observations and meteorological drivers. *J Clim* (2024) 37(3):757–78. doi: 10.1175/JCLI-D-23-0175.1
- 18. Wille JD, Alexander SP, Amory C, Baiman R, Barthélemy L, Bergstrom DM, et al. The extraordinary March 2022 East Antarctica heat wave. Part II: impacts on the Antarctic ice sheet. *J Clim* (2024) 37(3):779–99. doi: 10.1175/JCLI-D-23-0176.1
- 19. Parkinson CL. A 40-year record reveals gradual Antarctic sea ice increases followed by decreases at rates far exceeding the rates seen in the Arctic. *Proc Natl Acad Sci USA* (2019) 116(29):14414–23. doi: 10.1073/pnas.1906556116
- 20. Gilbert E, Holmes C. 2023's Antarctic sea ice extent is the lowest on record. Weather (2024) 79(2):46–51. doi: 10.1002/wea.4518
- 21. Rintoul SR, Chown SL, DeConto RM, England MH, Fricker HA, Masson-Delmotte V, et al. Choosing the future of Antarctica. *Nature* (2018) 558(7709):233–41. doi: 10.1038/s41586-018-0173-4
- 22. Brooks CM, Ainley DG, Jacquet J, Chown SL, Pertierra LR, Francis E, et al. Protect global values of the Southern Ocean ecosystem. *Science* (2022) 378(6619):477–9. doi: 10.1126/science.add9480
- 23. Chown SL, Leihy RI, Naish TR, Brooks CM, Convey P, Henley BJ, et al., editors Antarctic climate change and the environment: a decadal synopsis and recommendations for action. Cambridge: Scientific Committee on Antarctic Research (2022). Available at: https://documents.ats.aq/atcm44/att/atcm44\_att111\_e.pdf
- 24. Stoeckl N, Adams V, Baird R, Boothroyd A, Costanza R, Hatton MacDonald D, et al. The value of Antarctic and Southern Ocean ecosystem services. *Nat Rev Earth Environ* (2024) 5(3):153–5. doi: 10.1038/s43017-024-00523-3
- 25. Lee H, Calvin K, Dasgupta D, Krinner G, Mukherji A, Thorne P, et al., editors. Climate change 2023: synthesis report. Contribution of Working Groups I, II and III to the sixth assessment report of the Intergovernmental Panel on Climate Change. Geneva, Switzerland: Intergovernmental Panel on Climate Change (2023) 35–115. doi: 10.59327/IPCC/AR6-9789291691647
- 26. Pörtner H-O, Roberts DC, Masson-Delmotte V, Zhai P, Tignor M, Poloczanska E, editors. *Special report on the ocean and cryosphere in a changing climate*. Cambridge: Cambridge University Press (2019). doi: 10.1017/9781009157964
- 27. Post E, Bhatt US, Bitz CM, Brodie JF, Fulton TL, Hebblewhite M, et al. Ecological consequences of sea-ice decline. *Science* (2013) 341(6145):519–24. doi: 10.1126/science.1235225
- 28. Constable AJ, Melbourne-Thomas J, Muelbert MMC, McCormack S, Brasier M, Caccavo JA, et al. Marine ecosystem assessment for the southern ocean: summary for policymakers. Cambridge: Scott Polar Research Institute (2023). doi: 10.5281/zenodo.8359585
- 29. Hugelius G, Loisel J, Chadburn S, Jackson RB, Jones M, MacDonald G, et al. Large stocks of peatland carbon and nitrogen are vulnerable to permafrost thaw. *Proc Natl Acad Sci USA* (2020) 117(34):20438–46. doi: 10.1073/pnas.1916387117
- 30. Sabine CL, Feely RA, Gruber N, Key RM, Lee K, Bullister JL, et al. The oceanic sink for anthropogenic CO2. *Science* (2004) 305(5682):367–71. doi: 10.1126/science.1097403

- 31. Williams RG, Ceppi P, Roussenov V, Katavouta A, Meijers AJS. The role of the Southern Ocean in the global climate response to carbon emissions. *Phil Trans R Soc A* (2023) 381(2249):20220062. doi: 10.1098/rsta.2022.0062
- 32. Long MC, Stephens BB, McKain K, Sweeney C, Keeling RF, Kort EA, et al. Strong Southern Ocean carbon uptake evident in airborne observations. *Science* (2021) 374 (6572):1275–80. doi: 10.1126/science.abi4355
- 33. Dupont S, Pörtner HO. A snapshot of ocean acidification research. *Mar Biol* (2013) 160(8):1765–71. doi: 10.1007/s00227-013-2282-9
- 34. Arctic Monitoring and Assessment Programme. AMAP assessment 2018: Arctic Ocean acidification. Tromsø: AMAP (2018). Available at: https://www.amap.no/documents/doc/amap-assessment-2018-arctic-ocean-acidification/1659
- 35. Marshall J, Speer K. Closure of the meridional overturning circulation through Southern Ocean upwelling. *Nat Geosci* (2012) 5(3):171–80. doi: 10.1038/ngeo1391
- 36. Lee SK, Lumpkin R, Gomez F, Yeager S, Lopez H, Takglis F, et al. Human-induced changes in the global meridional overturning circulation are emerging from the Southern Ocean. *Commun Earth Environ* (2023) 4(1):69. doi: 10.1038/s43247-023-00727-3
- 37. Gunn KL, Rintoul SR, England MH, Bowen MM. Recent reduced abyssal overturning and ventilation in the Australian Antarctic Basin. *Nat Clim Chang* (2023) 13(6):537–44. doi: 10.1038/s41558-023-01667-8
- 38. Zhou S, Meijers AJS, Meredith MP, Abrahamsen EP, Holland PR, Silvano A, et al. Slowdown of Antarctic Bottom Water export driven by climatic wind and sea-ice changes. *Nat Clim Chang* (2023) 13(7):701–9. doi: 10.1038/s41558-023-01695-4
- 39. Li Q, England MH, Hogg AM, Rintoul SR, Morrison AK. Abyssal ocean overturning slowdown and warming driven by Antarctic meltwater. *Nature* (2023) 615(7954):841–7. doi: 10.1038/s41586-023-05762-w
- 40. Constable AJ, Harper S, Dawson J, Holsman K, Mustonen T, Piepenburg D, et al. Cross-chapter paper 6: polar regions. Climate change 2022: impacts, adaptation and vulnerability. contribution of the WGII to the 6th assessment report of the intergovernmental panel on climate change. Cambridge: Cambridge University Press (2022) 2319–68. doi: 10.1017/9781009325844.023
- 41. United Nations Environment Programme. Emissions gap report 2023: broken record temperatures hit new highs, yet world fails to cut emissions (again). Nairobi: UNEP (2023). doi: 10.59117/20.500.11822/43922
- 42. Masson-Delmotte V, Zhai P, Pirani A, Connors SL, Péan C, Berger S, et al. editors. Climate change: the physical science basis. Contribution of Working Group I to the sixth assessment report of the Intergovernmental Panel on Climate Change. Cambridge: Cambridge University Press (2021). doi: 10.1017/9781009157896
- 43. Palazzo Corner S, Siegert M, Ceppi P, Fox-Kemper B, Frölicher TL, Gallego-Sala A, et al. The Zero Emissions Commitment and climate stabilization. *Front Sci* (2023) 1:1170744. doi: 10.3389/fsci.2023.1170744
- 44. Shepherd JG. Geoengineering the climate: science, governance and uncertainty. The Royal Society (2009). Available at: https://royalsociety.org/topics-policy/publications/2009/geoengineering-climate/
- 45. MacAyeal DR, Mankoff K, Minchew B, Moore J, Wolovick M. *Glacial climate intervention: a research vision*. United States Antarctic Program Data Center (2024). doi: 10.15784/601797
- 46. National Academies of Sciences, Engineering, and Medicine. Reflecting sunlight: recommendations for solar geoengineering research and research governance. Washington, DC: The National Academies Press (2021). doi: 10.17226/25762
- 47. Shukla PR, Skea J, Slade R, Al Khourdajie A, van Diemen R, McCollum D, et al., editors. Climate change 2022: mitigation of climate change. Working Group III contribution to the sixth assessment report of the Intergovernmental Panel on Climate Change. Cambridge: Cambridge University Press (2022). doi: 10.1017/9781009157926
- 48. Budyko MI. On present-day climatic changes. Tellus (1977) 29(3):193–204. doi: 10.1111/j.2153-3490.1977.tb00725.x
- 49. McCormick MP, Thomason LW, Trepte CR. Atmospheric effects of the Mt Pinatubo eruption. Nature~(1995)~373(6513):399-404.~doi: 10.1038/373399a0
- 50. Smith W, Wagner G. Stratospheric aerosol injection tactics and costs in the first 15 years of deployment. *Environ Res Lett* (2018) 13(12):124001. doi: 10.1088/1748-0236/aca004
- 51. Masson-Delmotte V, Zhai P, Pörtner H-O, Roberts D, Skea J, Shukla PR, et al., editors. Global warming of 1.5°c: an IPCC Special Report on impacts of global warming of 1.5°c above pre-industrial levels and related global greenhouse gas emission pathways, in context of strengthening response to climate change, sustainable development, and efforts to eradicate poverty. Cambridge: Cambridge University Press (2018). doi: 10.1017/9781009157940
- 52. Huynh HN, McNeill VF. The potential environmental and climate impacts of stratospheric aerosol injection: a review. *Environ Sci Atmos* (2024) 4(2):114–43. doi: 10.1039/D3EA00134B
- 53. Sofieva VF, Rozanov A, Szelag M, Burrows JP, Retscher C, Damadeo R, et al. CREST: a climate data record of stratospheric aerosols. *Earth Syst Sci Data* (2024) 16(11):5227–41. doi: 10.5194/essd-16-5227-2024
- 54. Moriyama R, Sugiyama M, Kurosawa A, Masuda K, Tsuzuki K, Ishimoto Y. The cost of stratospheric climate engineering revisited. *Environ Res Lett* (2017) 12(3):034019. doi: 10.1088/1748-9326/aa6cc6

- 55. Kravitz B, Robock A, Oman L, Stenchikov G, Marquardt AB. Sulfuric acid deposition from stratospheric geoengineering with sulfate aerosols. *J Geophys Res* (2009) 114(D14):D14109. doi: 10.1029/2009JD011918
- 56. Official Aviation Guide. Airline frequency and capacity statistics [online] (2021). Available at: https://www.oag.com/airline-frequency-and-capacity-statistics
- 57. Smith W. Considerations for implementing stratospheric aerosol injection in polar regions. *Geoeng J* (2020) 14(2):112–30. doi: 10.1088/1748-9326/aae98d
- 58. Goddard PB, Kravitz B, MacMartin DG, Visioni D, Bednarz EM, Lee WR. Stratospheric aerosol injection can reduce risks to Antarctic ice loss depending on injection location and amount. *J Geophys Res Atmos* (2023) 128(22):e2023JD039434. doi: 10.1029/2023JD039434
- 59. Lee WR, MacMartin DG, Visioni D, Kravitz B, Chen Y, Moore JC, et al. High-latitude stratospheric aerosol injection to preserve the Arctic. *Earths Future* (2023) 11(1):e2022EF003052. doi: 10.1029/2022EF003052
- 60. Govindasamy B, Caldeira K. Geoengineering Earth's radiation balance to mitigate climate change from a quadrupling of CO2. *Glob Planet Change* (2000) 24(1):1–14. doi: 10.1016/S0921-8181(99)00093-7
- 61. Ridley JK, Blockley EW. Brief communication: solar radiation management not as effective as CO2 mitigation for Arctic sea ice loss in hitting the 1.5 and 2°C COP climate targets. *Cryosphere* (2018) 12(10):3355–60. doi: 10.5194/tc-12-3355-2018
- 62. Tilmes S, Richter JH, Kravitz B, MacMartin DG, Mills MJ, Simpson IR, et al. CESM1(WACCM) Stratospheric aerosol geoengineering large ensemble project. *Bull Am Meteorol Soc* (2018) 99(11):2361–71. doi: 10.1175/BAMS-D-17-0267.1
- 63. Kravitz B, MacMartin DG, Visioni D, Boucher O, Cole JNS, Haywood J, et al. Comparing different generations of idealized solar geoengineering simulations in the Geoengineering Model Intercomparison Project (GeoMIP). *Atmos Chem Phys* (2021) 21(6):4231–47. doi: 10.5194/acp-21-4231-2021
- 64. Irvine PJ, Keith DW, Moore J. Brief communication: understanding solar geoengineering's potential to limit sea level rise requires attention from cryosphere experts. *Cryosphere* (2018) 12(7):2501–13. doi: 10.5194/tc-12-2501-2018
- 65. Parker A, Irvine PJ. The risk of termination shock from solar geo-engineering. Earths Future (2018) 6(3):456–67. doi: 10.1002/2017EF000735
- 66. Douville H, Raghavan K, Renwick J, Allan RP, Arias PA, Barlow M, et al. Water cycle changes. *Climate change*. Cambridge, United Kingdom: Cambridge University Press (2023) 1055–210. doi: 10.1017/9781009157896.010
- 67. Diamond MS, Wanser K, Boucher O. Cooling credits are not a viable climate solution. Clim Change (2023) 176(7):96. doi: 10.1007/s10584-023-03561-w
- 68. Duffey A, Irvine P, Tsamados M, Stroeve J. Solar geoengineering in the polar regions: a review. Earths Future (2023) 11(6):e2023EF003679. doi: 10.1029/2023EF003679
- 69. Smith W, Bartels MF, Boers JG, Rice CV. On thin ice: solar geoengineering to manage tipping element risks in the cryosphere by 2040. *Earths Future* (2024) 12(8): e2024EF004797. doi: 10.1029/2024EF004797
- 70. Jiang J, Cao L, MacMartin DG, Simpson IR, Kravitz B, Cheng W, et al. Stratospheric sulfate aerosol geoengineering could alter the high-latitude seasonal cycle. *Geophys Res Lett* (2019) 46(23):14153–63. doi: 10.1029/2019GL085758
- 71. Banerjee A, Butler AH, Polvani LM, Robock A, Simpson IR, Sun L. Robust winter warming over Eurasia under stratospheric sulfate geoengineering the role of stratospheric dynamics. *Atmos Chem Phys* (2021) 21(9):6985–97. doi: 10.5194/acp-21-6985-2021
- 72. Bednaršek N, Naish K-A, Feely RA, Hauri C, Kimoto K, Hermann AJ, et al. Integrated assessment of ocean acidification risks to pteropods in the northern high latitudes: regional comparison of exposure, sensitivity and adaptive capacity. *Front Mar Sci* (2021) 8:671497. doi: 10.3389/fmars.2021.671497
- 73. Hönisch B, Ridgwell A, Schmidt DN, Thomas E, Gibbs SJ, Sluijs A, et al. The geological record of ocean acidification. *Science* (2012) 335(6072):1058–63. doi: 10.1126/science.1208277
- 74. Pelejero C, Calvo E, Hoegh-Guldberg O. Paleo-perspectives on ocean acidification. *Trends Ecol Evol* (2010) 25(6):332–44. doi: 10.1016/j.tree.2010.02.002
- 75. Green HL, Findlay HS, Shutler JD, Land PE, Bellerby RGJ. Satellite observations are needed to understand ocean acidification and multi-stressor impacts on fish stocks in a changing Arctic Ocean. Front Mar Sci (2021) 8:635797. doi: 10.3389/fmars.2021.635797
- 76. Tilmes S, Müller R, Salawitch R. The sensitivity of polar ozone depletion to proposed geoengineering schemes. *Science* (2008) 320(5880):1201–4. doi: 10.1126/science.1153966
- 77. Moch JM, Mickley LJ, Eastham SD, Lundgren EW, Shah V, Buonocore JJ, et al. Overlooked long-term atmospheric chemical feedbacks alter the impact of solar geoengineering: implications for tropospheric oxidative capacity.  $AGU\ Adv\ (2023)\ 4(5):e2023AV000911$ . doi: 10.1029/2023AV000911
- 78. Zarnetske PL, Gurevitch J, Franklin J, Groffman PM, Harrison CS, Hellmann JJ, et al. Potential ecological impacts of climate intervention by reflecting sunlight to cool Earth. *Proc Natl Acad Sci USA* (2021) 118(15):e1921854118. doi: 10.1073/pnas.1921854118
- 79. Haywood JM, Jones A, Bellouin N, Stephenson D. Asymmetric forcing from stratospheric aerosols impacts Sahelian rainfall. *Nat Clim Change* (2013) 3(7):660–5. doi: 10.1038/nclimate1857

- 80. Määttänen A, Lameille T, Kloeck C, Boucher O, Ravetta F. Uncertainties and confidence in stratospheric aerosol injection modelling: a systematic literature review. Oxford Open Climate Change (2024) 4(1):kgae007. doi: 10.1093/oxfclm/kgae007
- 81. Tilmes S, Visioni D, Jones A, Haywood J, Séférian R, Nabat P, et al. Stratospheric ozone response to sulfate aerosol and solar dimming climate interventions based on the G6 Geoengineering Model Intercomparison Project (GeoMIP) simulations. *Atmos Chem Phys* (2022) 22(7):4557–79. doi: 10.5194/acp-22-4557-2022
- 82. Effiong U, Neitzel RL. Assessing the direct occupational and public health impacts of solar radiation management with stratospheric aerosols. *Environ Health* (2016) 15:7. doi: 10.1186/s12940-016-0089-0
- 83. Feinberg A. Annual solar geoengineering: mitigating yearly global warming increases. Climate (2024) 12(2):26. doi: 10.3390/cli12020026
- 84. Smith W. The cost of stratospheric aerosol injection through 2100. Environ Res Lett (2020) 15(11):114004. doi: 10.1088/1748-9326/aba7e7
- 85. Bronsther J, Xu Y. The social costs of solar radiation management. npj Clim. Action (2025) 4:69. doi: 10.1038/s44168-025-00273-y
- 86. Baur S, Nauels A, Nicholls Z, Sanderson BM, Schleussner C-F. The deployment length of solar radiation modification: an interplay of mitigation, net-negative emissions and climate uncertainty. *Earth Syst Dynam* (2023) 14(2):367–81. doi: 10.5194/esd-14-367-2023
- 87. Baur S, Sanderson BM, Séférian R, Terray L. Solar radiation modification challenges decarbonization with renewable solar energy. *Earth Syst Dynam* (2024) 15(2):307–22. doi: 10.5194/esd-15-307-2024
- 88. Baur S, Sanderson BM, Séférian R, Terray L. Change in wind renewable energy potential under stratospheric aerosol injections. *Earths Future* (2024) 12(10): e2024EF004575. doi: 10.1029/2024EF004575
- 89. Corry O, McLaren D, Kornbech N. Scientific models versus power politics: how security expertise reframes solar geoengineering. *Rev Int Stud* (2024), 1–20. doi: 10.1017/S0260210524000482
- 90. Fritz L, Baum CM, Low S, Sovacool BK. Public engagement for inclusive and sustainable governance of climate interventions. *Nat Commun* (2024) 15(1):4168. doi: 10.1038/s41467-024-48510-y
- 91. Lockley A, Mundra I, Smith PT. Legitimacy and justifiability of non-state geoengineering. Futures (2023) 152:103210. doi: 10.1016/j.futures.2023.103210
- 92. Biermann F, Oomen J, Gupta A, Ali SH, Conca K, Hajer MA, et al. Solar geoengineering: the case for an international non-use agreement. *Wiley Interdiscip Rev Clim Change* (2022) 13(3):e754. doi: 10.1002/wcc.754
- 93. Sovacool BK, Baum C, Low S. The next climate war? Statecraft, security, and weaponization in the geopolitics of a low-carbon future. *Energy Strategy Rev* (2023) 45:101031. doi: 10.1016/j.esr.2022.101031
- 94. Sovacool BK, Baum CM, Low S. Reviewing the sociotechnical dynamics of carbon removal. *Joule* (2023) 7(1):57–82. doi: 10.1016/j.joule.2022.11.008
- 95. McLaren D, Corry O. Clash of geofutures and the remaking of planetary order: faultlines underlying conflicts over geoengineering governance. *Glob Policy* (2021) 12(S1):20–33. doi: 10.1111/1758-5899.12863
- 96. Brent KA, McGee J. The regulation of geoengineering: a gathering storm for international climate change policy? *Air Qual Clim Change* (2012) 46(4):22–7. doi: 10.3316/informit.961245963007370
- 97. American Geophysical Union. Position statement on climate intervention: the urgent priority of a healthy climate requires enhanced climate intervention research and community engagement. AGU (2022). Available at: https://www.agu.org/science-policy/position-statements/climate-intervention
- 98. United Nations Environment Programme. One atmosphere: an independent expert review on solar radiation modification research and deployment. Nairobi: UNEP (2023). Available at: https://wedocs.unep.org/handle/20.500.11822/41903.
- 99. American Geophysical Union. Ethical framework principles for climate intervention research. Washington, D.C.: AGU ESS Open Archive (2024). doi: 10.22541/essoar.172917365.53105072/v1
- 100. Saami Council. Petition: support the indigenous peoples voices call on Harvard to shut down the SCoPEx project [online] (2021). Available at: https://www.saamicouncil.net/s/Letter-to-Scopex-Advisory-Committee-24-February.pdf.
- 101. Scientific Advice Mechanism, European Commission's Group of Chief Scientific Advisors. Scoping paper: solar radiation modification. Scientific Advice Mechanism to the European Commission. (2023). Available at: https://research-and-innovation.ec.europa.eu/system/files/2023-08/Scoping\_paper\_SRM.pdf.
- 102. European Commission's Group of Chief Scientific Advisors, Directorate-General for Research and Innovation. *Solar radiation modification*. Brussels: Publications Office of the European Union (2024). doi: 10.2777/391614
- 103. European Commission Directorate-General for Research and Innovation. *Opinion on solar radiation modification ethical perspectives.* Brussels: Publications Office of the European Union (2024). doi: 10.2777/951016
- 104. European Commission's Group of Chief Scientific Advisors, Directorate-General for Research and Innovation. *Summary of the scientific opinion on solar radiation modification*. Brussels: Publications Office of the European Union (2025). doi: 10.2777/1707046

- 105. Mexican Secretariat of Environment and Natural Resources, Government of Mexico. Experimentation with solar geoengineering will not be permitted in Mexico [online; Spanish]. (2023). Available at: https://www.gob.mx/semarnat/prensa/la-experimentacion-con-geoingenieria-solar-no-sera-permitida-en-mexico.
- 106. Surprise K, Sapinski J. Whose climate intervention? Solar geoengineering, fractions of capital, and hegemonic strategy. *Cap Class* (2023) 47(4):539–64. doi: 10.1177/03098168221114386
- 107. Talati S, Buck HJ, Kravitz B. How to address solar geoengineering's transparency problem. *Proc Natl Acad Sci USA* (2025) 122(3):e2419587122. doi: 10.1073/pnas.2419587122
- 108. Solar Geoengineering Non-Use Agreement. [website]. Available at: https://www.solargeoeng.org/.
- 109. Gupta A, Biermann F, van Driel E, Bernaz N, Jayaram D, Kim RE, et al. Towards a non-use regime on solar geoengineering: lessons from international law and governance. *Transnat Environ Law* (2024) 13(2):368–99. doi: 10.1017/S2047102524000050
- 110. McLaren D. Mitigation deterrence and the moral hazard of solar radiation management. Earths Future (2016) 4(12):596602. doi: 10.1002/2016EF000445
- 111. Tracy SM, Moch JM, Eastham SD, Buonocore JJ. Stratospheric aerosol injection may impact global systems and human health outcomes. *Elem Sci Anthropocene* (2022) 10(1):47. doi: 10.1525/elementa.2022.00047
- 112. Wolovick MJ, Moore JC. Stopping the flood: could we use targeted geoengineering to mitigate sea level rise? *Cryosphere* (2018) 12(9):2955–67. doi: 10.5194/tc-12-2955-2018
- 113. Moore JC, Gladstone R, Zwinger T, Wolovick MJ. Geoengineer polar glaciers to slow sea-level rise. Nature (2018) 555(7696):303–5. doi: 10.1038/d41586-018-03036-4
- 114. Wolovick M, Moore J, Keefer B. The potential for stabilizing Amundsen Sea glaciers via underwater curtains. *PNAS Nexus* (2023) 2(4):pgad103. doi: 10.1093/pnasnexus/pgad103
- 115. Keefer B, Wolovick M, Moore JC. Feasibility of ice sheet conservation using seabed anchored curtains. *PNAS Nexus* (2023) 2(3):pgad053. doi: 10.1093/pnasnexus/pgad053
- 116. Kirkham JD, Hogan KA, Larter RD, Arnold NS, Nitsche FO, Golledge NR, et al. Past water flow beneath Pine Island and Thwaites glaciers, West Antarctica. *Cryosphere* (2019) 13(7):1959–81. doi: 10.5194/tc-13-1959-2019
- 117. Hogan KA, Larter RD, Graham AGC, Arthern R, Kirkham JD, Totten RL, et al. Revealing the former bed of Thwaites Glacier using sea-floor bathymetry: implications for warm-water routing and bed controls on ice flow and buttressing. *Cryosphere* (2020) 14(9):2883–908. doi: 10.5194/tc-14-2883-2020
- 118. Hogan KA, Arnold NS, Larter RD, Kirkham JD, Noormets R, Ó Cofaigh C, et al. Subglacial water flow over an Antarctic palaeo-ice stream bed? *J Geophys Res Earth Surface* (2022) 127(2):e2021JF006442. doi: 10.1029/2021JF006442
- 119. Gürses Ö, Kolatschek V, Wang Q, Rodehacke CB. Brief communication: a submarine wall protecting the Amundsen Sea intensifies melting of neighboring ice shelves. *Cryosphere* (2019) 13(9):2317–24. doi: 10.5194/tc-13-2317-2019
- 120. Alevropoulos-Borrill A, Golledge NR, Cornford SL, Lowry DP, Krapp M. Sustained ocean cooling insufficient to reverse sea level rise from Antarctica. *Commun Earth Environ* (2024) 5(1):150. doi: 10.1038/s43247-024-01297-8
- 121. Feldmann J, Levermann A, Mengel M. Stabilizing the West Antarctic Ice Sheet by surface mass deposition. Sci~Adv~(2019)~5(7):eaaw4132. doi: 10.1126/sciadv.aaw4132
- 122. Larter RD, Anderson JB, Graham AGC, Gohl K, Hillenbrand C-D, Jakobsson M, et al. Reconstruction of changes in the Amundsen Sea and Bellingshausen Sea sector of the West Antarctic Ice Sheet since the Last Glacial Maximum. *Quat Sci Rev* (2014) 100:55–86. doi: 10.1016/j.quascirev.2013.10.016
- 123. Dorschel B, Hehemann L, Viquerat S, Warnke F, Dreutter S, Tenberge YS, et al. The international bathymetric chart of the Southern Ocean Version 2. *Sci Data* (2022) 9(1):275. doi: 10.1038/s41597-022-01366-7
- 124. Gohl K, Freudenthal T, Hillenbrand C-D, Klages J, Larter R, Bickert T, et al. MeBo70 seabed drilling on a polar continental shelf: operational report and lessons from drilling in the Amundsen Sea embayment of West Antarctica. *Geochem Geophys Geosyst* (2017) 18(11):4235–50. doi: 10.1002/2017GC007081
- 125. Stokes CR, Abram NJ, Bentley MJ, Edwards TL, England MH, Foppert A, et al. Response of the East Antarctic Ice Sheet to past and future climate change. *Nature* (2022) 608(7922):275–86. doi: 10.1038/s41586-022-04946-0
- 126. Hanna E, Topál D, Box JE, Buzzard S, Christie FD, Hvidberg C, et al. Short- and long-term variability of the Antarctic and Greenland ice sheets. *Nat Rev Earth Environ* (2024) 5(3):193–210. doi: 10.1038/s43017-023-00509-7
- 127. Zheng Y, Heywood KJ, Webber BGM, Stevens DP, Biddle LC, Boehme L, et al. Winter seal-based observations reveal glacial meltwater surfacing in the southeastern Amundsen Sea. *Commun Earth Environ* (2021) 2(1):40. doi: 10.1038/s43247-021-00111-z
- 128. Dinniman MS, St-Laurent P, Arrigo KR, Hofmann EE, van Dijken GL. Analysis of iron sources in Antarctic continental shelf waters. *J Geophys Res Oceans* (2020) 125(5):e2019JC015736. doi: 10.1029/2019JC015736
- 129. McCarthy AH, Peck LS, Hughes KA, Aldridge DC. Antarctica: the final frontier for marine biological invasions. *Glob Change Biol* (2019) 25(7):2221–41. doi: 10.1111/gcb.14600

- 130. Hughes KA, et al. Invasive non-native species likely to threaten biodiversity and ecosystems in the Antarctic Peninsula region. *Global Change Biol* (2020) 26:2702–16. doi: 10.1111/gcb.14938
- 131. United Kingdom Environment Agency. The Thames Barrier protecting London and the Thames Estuary for 40 years [online]. (2024). Available at: https://www.gov.uk/government/news/the-thames-barrier-protecting-london-and-the-thames-estuary-for-40-years
- 132. Reuters. China says Three Gorges Dam cost \$37 billion [online]. (2009). Available at: https://www.reuters.com/article/economy/china-says-three-gorges-dam-cost-37-billion-idUSPEK84588/#:~:text=BEIJING%2C%20Sept%2014%20(Reuters), several%20times%20an%20original%20estimate.
- 133. Secretariat of the Antarctic Treaty. Protocol on Environmental Protection to the Antarctic Treaty (1991). Available at: https://www.ats.aq/e/protocol.html.
- 134. Flamm P, Shibata A. 'Ice Sheet Conservation' and international discord: governing (potential) glacial geoengineering in Antarctica. *Int Aff* (2025) 101(1):309–20. doi: 10.1093/ia/iiae281
- 135. Duspayev A, Flanner MG, Riihelä A. Earth's sea ice radiative effect from 1980 to 2023. Geophys Res Lett (2024) 51(14):e2024GL109608. doi: 10.1029/2024GL109608
- 136. Field L, Ivanova D, Bhattacharyya S, Mlaker V, Sholtz A, Decca R, et al. Increasing Arctic sea ice albedo using localized reversible geoengineering. *Earths Future* (2018) 6(6):882–901. doi: 10.1029/2018EF000820
- 137. Rickert D, Schlüter M, Wallmann K. Dissolution kinetics of biogenic silica from the water column to the sediments. *Geochim Cosmochim Acta* (2002) 66(3):439–55. doi: 10.1016/S0016-7037(01)00757-8
- 138. Krause JW, Duarte CM, Marquez IA, Assmy P, Fernández-Méndez M, Wiedmann I, et al. Biogenic silica production and diatom dynamics in the Svalbard region during spring. *Biogeosciences* (2018) 15(21):6503–17. doi: 10.5194/bg-15-6503-2018
- 139. Krause JW, Schulz IK, Rowe KA, Dobbins W, Winding MHS, Sejr MK, et al. Silicic acid limitation drives bloom termination and potential carbon sequestration in an Arctic bloom. *Sci Rep* (2019) 9(1):8149. doi: 10.1038/s41598-019-44587-4
- 140. Giesbrecht KE, Varela DE. Summertime biogenic silica production and silicon limitation in the Pacific Arctic region from 2006 to 2016. *Glob Biogeochem Cycles* (2021) 35(1):e2020GB006629. doi: 10.1029/2020GB006629
- 141. Webster MA, Warren SG. Regional geoengineering using tiny glass bubbles would accelerate the loss of Arctic sea ice. *Earths Future* (2022) 10(10):e2022EF002815. doi: 10.1029/2022EF002815
- 142. Chuffart R, Cooper AM, Wood-Donnelly C, Seddon L. Old Sea, New Ice: sea ice geoengineering and indigenous rights in Arctic Ocean governance. *Polar J* (2023) 13(2):195–215. doi: 10.1080/2154896X.2023.2269688
- 143. Flannery BP, Kheshgi H, Marland G, MacCracken MC. Geoengineering climate. In: Watts RG, editor. *Engineering response to global climate change: planning a research and development agenda*. Boca Raton: CRC Press (1997).
- 144. Desch SJ, Smith N, Groppi C, Vargas P, Jackson R, Kalyaan A, et al. Arctic ice management. *Earths Future* (2017) 5(1):107–27. doi: 10.1002/2016EF000410
- 145. Pauling AG, Bitz CM. Arctic sea ice response to flooding of the snow layer in future warming scenarios. *Earths Future* (2021) 9(10):e2021EF002136. doi: 10.1029/2021EF002136
- 146. Zampieri L, Goessling HF. Sea ice targeted geoengineering can delay Arctic sea ice decline but not global warming.  $Earths\ Future\ (2019)\ 7(12):1296-306$ . doi: 10.1029/2019EF001230
- 147. Fox-Kemper B, Hewitt HT, Xiao C, Aðalgeirsdóttir G, Drijfhout SS, Edwards TL, et al. Ocean, cryosphere and sea level change. In: Masson-Delmotte V, Zhai P, Pirani A, Connors SL, Péan C, Berger S, et al., editors. Climate change: the physical science basis. Contribution of Working Group I to the sixth assessment report of the Intergovernmental Panel on Climate Change. Cambridge: Cambridge University Press (2021) 1211–1362. doi: 10.1017/9781009157896.011 Available at: https://www.ipcc.ch/report/ar6/wg1/downloads/report/IPCC\_AR6\_WGI\_Chapter09.pdf
- 148. Lockley A, Wolovick M, Keefer B, Gladstone R, Zhao L-Y, Moore JC. Glacier geoengineering to address sea-level rise: a geotechnical approach. Adv Clim Change Res (2020) 11(4):401–14. doi: 10.1016/j.accre.2020.11.008
- 149. Truffer M, Echelmeyer KA. Of isbrae and ice streams. *Ann Glaciol* (2003) 36:66–72. doi: 10.3189/172756403781816347
- 150. Stokes CR. Geomorphology under ice streams: moving from form to process. Earth Surf Process Landf (2018) 43(1):85–123. doi: 10.1002/esp.4259
- 151. Holt JW, Blankenship DD, Morse DL, Young DA, Peters ME, Kempf SD, et al. New boundary conditions for the West Antarctic Ice Sheet: subglacial topography of the Thwaites and Smith glacier catchments. *Geophys Res Lett* (2006) 33(9):L09501. doi: 10.1029/2005GL025588
- 152. Reese R, Garbe J, Hill EA, Urruty B, Naughten KA, Gagliardini O, et al. The stability of present-day Antarctic grounding lines part 2: onset of irreversible retreat of Amundsen Sea glaciers under current climate on centennial timescales cannot be excluded. *Cryosphere* (2023) 17(9):3761–83. doi: 10.5194/tc-17-3761-2023
- 153. Cuffey KM, Paterson WSB. The physics of glaciers. (4th edition). London: Butterworth (2010).

- 154. Schoof C. Ice-sheet acceleration driven by melt supply variability. Nature (2010) 468(7325):803–6. doi: 10.1038/nature09618
- 155. Dow CF, Ross N, Jeofry H, Siu K, Siegert MJ. Antarctic basal environment shaped by high-pressure flow through a subglacial river system. *Nat Geosci* (2022) 15 (11):892–8. doi: 10.1038/s41561-022-01059-1
- 156. Christoffersen P, Bougamont M, Carter SP, Fricker HA, Tulaczyk S. Significant groundwater contribution to Antarctic ice streams hydrologic budget. *Geophys Res Lett* (2014) 41(6):2003–10. doi: 10.1002/2014GL059250
- 157. Siegert MJ, Kulessa B, Bougamont M, Christoffersen P, Key K, Andersen KR, et al. Antarctic subglacial groundwater: a concept paper on its measurement and potential influence on ice flow. In: Siegert MJ, Jamieson SSR, White DA, editors. Exploration of subsurface Antarctica: uncovering past changes and modern processes. London: Geological Society of London (2018) 197–213. doi: 10.1144/SP461.8
- 158. Mikucki JA, Auken E, Tulaczyk S, Virginia RA, Schamper C, Sørensen KI, et al. Deep groundwater and potential subsurface habitats beneath an Antarctic dry valley. *Nat Commun* (2015) 6:6831. doi: 10.1038/ncomms7831
- 159. Gustafson CD, Key K, Siegfried MR, Winberry JP, Fricker HA, Venturelli RA, et al. A dynamic saline groundwater system mapped beneath an Antarctic ice stream. *Science* (2022) 376(6593):640–4. doi: 10.1126/science.abm3301
- 160. Aitken ARA, Kulessa B, Schroeder D, Jordan T, Whittaker J, Anandakrishnan S, et al. Antarctic sedimentary basins and their influence on ice-sheet dynamics. *Rev Geophys* (2023) 61(3):e2021RG000767. doi: 10.1029/2021RG000767
- 161. Li L, Aitken AR, Lindsay MD, Kulessa B. Sedimentary basins reduce stability of Antarctic ice streams through groundwater feedbacks. *Nat Geosci* (2022) 15(8):645–50. doi: 10.1038/s41561-022-00992-5
- 162. Robel AA, Sim SJ, Meyer C, Siegfried MR, Gustafson CD. Contemporary ice sheet thinning drives subglacial groundwater exfiltration with potential feedbacks on glacier flow. *Sci Adv* (2023) 9(33):eadh3693. doi: 10.1126/sciadv.adh3693
- 163. Pattyn F. Antarctic subglacial conditions inferred from a hybrid ice sheet/ice stream model. *Earth Planet Sci Lett* (2010) 295(3-4):451-61. doi: 10.1016/j.epsl.2010.04.025
- 164. Fricker HA, Scambos T, BindsChadler R, Padman L. An active subglacial water system in West Antarctica mapped from space. *Science* (2007) 315(5818):1544–8. doi: 10.1126/science.1136897
- 165. Livingstone SJ, Li Y, Rutishauser A, Sanderson RJ, Winter K, Mikucki JA, et al. Subglacial lakes and their changing role in a warming climate. *Nat Rev Earth Environ* (2022) 3(2):106–24. doi: 10.1038/s43017-021-00246-9
- 166. Kazmierczak E, Sun S, Coulon V, Pattyn F. Subglacial hydrology modulates basal sliding response of the Antarctic ice sheet to climate forcing. *Cryosphere* (2022) 16(10):4537–52. doi: 10.5194/tc-16-4537-2022
- 167. Schoof C. The effect of cavitation on glacier sliding. Proc R Soc A Math Phys Eng Sci (2005) 461(2055):609–27. doi: 10.1098/rspa.2004.1350
- 168. Bell RE. The role of subglacial water in ice-sheet mass balance. Nat Geosci (2008) 1(5):297–304. doi:  $10.1038/\rm{ngeo186}$
- 169. Bartholomew I, Nienow P, Sole A, Mair D, Cowton T, King MA. Short-term variability in Greenland Ice Sheet motion forced by time-varying meltwater drainage: implications for the relationship between subglacial drainage system behavior and ice velocity. *J Geophys Res* (2012) 117(F3):F03002. doi: 10.1029/2011JF002220
- 170. Siegfried MR, Fricker HA. Illuminating active subglacial lake processes with ICESat-2 laser altimetry. *Geophys Res Lett* (2021) 48(14):GL091089. doi: 10.1029/2020GL091089
- 171. Das SB, Joughin I, Behn MD, Howat IM, King MA, Lizarralde D, et al. Fracture propagation to the base of the Greenland ice sheet during supraglacial lake drainage. *Science* (2008) 320(5877):778–81. doi: 10.1126/science.1153360
- 172. Echelmeyer K, Harrison WD. Jakobshavns Isbræ, West Greenland: seasonal variations in velocity or lack thereof. J Glaciol (1990) 36(122):82-8. doi: 10.3189/50022143000005591
- 173. Davison BJ, Sole AJ, Livingstone SJ, Cowton TR, Nienow PW. The influence of hydrology on the dynamics of land-terminating sectors of the Greenland ice sheet. *Front Earth Sci* (2019) 7:10. doi: 10.3389/feart.2019.00010
- 174. Sundal AV, Shepherd A, Nienow P, Hanna E, Palmer S, Huybrechts P. Meltinduced speed-up of Greenland ice sheet offset by efficient subglacial drainage. *Nature* (2011) 469(7331):521–4. doi: 10.1038/nature09740
- 175. Wright AP, Siegert MJ, Le Brocq AM, Gore DB. High sensitivity of subglacial hydrological pathways in Antarctica to small ice-sheet changes. *Geophys Res Lett* (2008) 35(17):L17504. doi: 10.1029/2008GL034937
- 176. Anandakrishnan S, Alley RB. Stagnation of Ice Stream C, West Antarctica by water piracy. *Geophys Res Lett* (1997) 24(3):265–8. doi: 10.1029/96GL04016
- 177. van der Wel N, Christoffersen P, Bougamont M. The influence of subglacial hydrology on the flow of Kamb Ice Stream, West Antarctica. *J Geophys Res Earth Surf* (2013) 118(1):97–110. doi: 10.1029/2012JF002570
- 178. Catania G, Hulbe C, Conway H, Scambos TA, Raymond CF. Variability in the mass flux of the Ross Ice Streams, West Antarctica, over the last millennium. *J Glaciol* (2012) 58(210):741–52. doi: 10.3189/2012JoG11J219

179. Conway H, Catania G, Raymond CF, Gades AM, Scambos TA, Engelhardt H. Switch of flow direction in an Antarctic ice stream. *Nature* (2002) 419(6906):465–7. doi: 10.1038/nature01081

- 180. Dowdeswell JA, Ottesen D, Rise L. Flow switching and large-scale deposition by ice streams draining former ice sheets. *Geology* (2006) 34(4):313–6. doi: 10.1130/G22253.1
- 181. Cofaigh CO, Evans DJA, Smith IR. Large-scale reorganization and sedimentation of terrestrial ice streams during Late Wisconsinan Laurentide Ice Sheet deglaciation. *Geol Soc Am Bull* (2010) 122(5-6):743-56. doi: 10.1130/B26476.1
- 182. Stokes CR, Clark CD, Storrar R. Major changes in ice stream dynamics during deglaciation of the north-western margin of the Laurentide Ice Sheet. *Quat Sci Rev* (2009) 28(7–8):721–38. doi: 10.1016/j.quascirev.2008.07.019
- 183. Winsborrow MCM, Stokes CR, Andreassen K. Ice stream flow switching during deglaciation of the southwestern Barents Sea. Geol Soc Am Bull (2012) 124(3-4):275-90. doi: 10.1130/B30416.1
- 184. Stokes CR, Margold M, Clark CD, Tarasov L. Ice stream activity scaled to ice sheet volume during Laurentide Ice Sheet deglaciation. *Nature* (2016) 530(7590):322–6. doi: 10.1038/nature16947
- 185. Payne AJ, Dongelmans PW. Self-organization in the thermomechanical flow of ice sheets. *J Geophys Res* (1997) 102(B6):12219–33. doi: 10.1029/97JB00513
- 186. Robel AA, DeGiuli E, Schoof C, Tziperman E. Dynamics of ice stream temporal variability: modes, scales, and hysteresis. *J Geophys Res Earth Surf* (2013) 118:925–36. doi: 10.1002/jgrf.20072
- 187. Kyrke-Smith TM, Katz RF, Fowler AC. Subglacial hydrology as a control on emergence, scale, and spacing of ice streams. *J Geophys Res Earth Surf* (2015) 120(8):1501-14. doi: 10.1002/2015JF003505
- 188. Pfeffer WT. A simple mechanism for irreversible tidewater glacier retreat. *J Geophys Res Earth Surf* (2007) 112(F3):F03S25. doi: 10.1029/2006JF000590
- 189. Nick FM, Vieli A, Howat IM, Joughin IM. Large-scale changes in Greenland outlet glacier dynamics triggered at the terminus. *Nat Geosci* (2009) 2(2):110–4. doi: 10.1038/ngeo394
- 190. Gudmundsson GH, Paolo FS, Adusumilli S, Fricker HA. Instantaneous Antarctic ice sheet mass loss driven by thinning ice shelves. *Geophys Res Lett* (2019) 46(13):13903–9. doi: 10.1029/2019GL085027
- 191. Ross N, Bingham RG, Corr HFJ, Ferraccioli F, Jordan TA, Le Brocq A, et al. Steep reverse bed slope at the grounding line of the Weddell Sea sector in West Antarctica. *Nat Geosci* (2012) 5(6):393–6. doi: 10.1038/ngeo1468
- 192. Siegert MJ, Makinson K, Blake D, Mowlem M, Ross N. An assessment of deephot-water drilling as a means to undertake direct measurement and sampling of Antarctic subglacial lakes: experience and lessons learned from the Lake Ellsworth field season 2012-13. *Ann Glaciol* (2014) 55(65):59–73. doi: 10.3189/2014AoG65A008
- 193. Siegert MJ. A 60-year international history of Antarctic subglacial lake exploration. In: Siegert MJ, Jamieson SSR, White DA. *Exploration of subsurface Antarctica: uncovering past changes and modern processes.* London: Geological Society (2018). doi: 10.1144/sp461.5
- 194. Johnson JS, O'Donnell JP, Thomas ER. *In situ* measurements of snow accumulation in the Amundsen Sea Embayment during 2016. *Antarct Sci* (2018) 30(3):197–203. doi: 10.1017/S0954102018000068
- 195. Siegert MJ, Kennicutt MC. Governance of the exploration of subglacial Antarctica. Front Environ Sci (2018) 6:103. doi: 10.3389/fenvs.2018.00103
- 196. Siegert MJ, Clarke RJ, Mowlem M, Ross N, Hill CS, Tait A, et al. Clean access, measurement and sampling of Ellsworth subglacial Lake: a method to explore deep Antarctic subglacial lake environments. *Rev Geophys* (2012) 50(1):RG1003. doi: 10.1029/2011RG000361
- 197. Scott JBT, Smith AM, Bingham RG, Vaughan DG. Crevasses triggered on Pine Island Glacier, West Antarctica, by drilling through an exceptional melt layer. *Ann Glaciol* (2010) 51(55):65–70. doi: 10.3189/172756410791392763
- 198. Arrigo KR, Tagliabue A. Iron in the Ross Sea: 2. Impact of discrete iron addition strategies. J Geophys Res (2005) 110(C3):C03010. doi: 10.1029/2004JC002568
- 199. Jiang H-B, Hutchins DA, Zhang H-R, Feng Y-Y, Zhang R-F, Sun W-W, et al. Complexities of regulating climate by promoting marine primary production with ocean iron fertilization. *Earth Sci Rev* (2024) 249:104675. doi: 10.1016/j.earscirev.2024.104675
- 200. Martin JH, Gordon RM, Fitzwater SE. Iron in antarctic waters. *Nature* (1990) 345(6271):156–8. doi: 10.1038/345156a0
- 201. Oschlies A, Koeve W, Rickels W, Rehdanz K. Side effects and accounting aspects of hypothetical large-scale Southern Ocean iron fertilization. Biogeosciences (2010) 7(12):4017–35. doi: 10.5194/bg-7-4017-2010
- 202. Yoon J-E, Yoo K-C, Macdonald AM, Yoon H-I, Park K-T, Yang EJ, et al. Reviews and syntheses: ocean iron fertilization experiments past, present, and future looking to a future Korean Iron Fertilization Experiment in the Southern Ocean (KIFES) project. *Biogeosciences* (2018) 15(19):5847–89. doi: 10.5194/bg-15-5847-2018
- 203. De Baar HJW, Boyd PW, Coale KH, Landry MR, Tsuda A, Assmy P, et al. Synthesis of iron fertilization experiments: from the Iron Age in the Age of Enlightenment. *J Geophys Res* (2005) 110(C9):C09S16. doi: 10.1029/2004JC002601

- 204. Aumont O, Bopp L. Globalizing results from ocean in situ iron fertilization studies. Glob Biogeochem Cycles (2006) 20(2):GB2017. doi: 10.1029/2005GB002591
- 205. Quéguiner B. Iron fertilization and the structure of planktonic communities in high nutrient regions of the Southern Ocean.  $Deep\ Sea\ Res\ II\ (2013)\ 90:43-54.$  doi: 10.1016/j.dsr2.2012.07.024
- 206. Boyd PW, Law CS. The Southern Ocean iron release experiment (SOIREE) introduction and summary. Deep Sea Res II (2001) 48(11-12):2425-38. doi: 10.1016/80967-0645(01)00002-9
- 207. Buesseler KO, Andrews JE, Pike SM, Charette MA. The effects of iron fertilization on carbon sequestration in the Southern Ocean. *Science* (2004) 304 (5669):414–17. doi: 10.1126/science.1086895
- 208. West NJ, Obernosterer I, Zemb O, Lebaron P. Major differences of bacterial diversity and activity inside and outside of a natural iron-fertilized phytoplankton bloom in the Southern Ocean. *Environ Microbiol* (2008) 10(3):738–56. doi: 10.1111/j.1462-2920.2007.01497.x
- 209. Sarmiento JL, Gruber N, Brzezinski MA, Dunne JP. High-latitude controls of thermocline nutrients and low latitude biological productivity. *Nature* (2004) 427 (6969):56–60. doi: 10.1038/nature02127
- 210. Sarmiento JL, Orr JC. Three-dimensional simulations of the impact of Southern Ocean nutrient depletion on atmospheric CO2 and ocean chemistry. *Limnol Oceanogr* (1991) 36(8):1928–50. doi: 10.4319/lo.1991.36.8.1928
- 211. Tagliabue A, Twining BS, Barrier N, Maury O, Berger M, Bopp L. Ocean iron fertilization may amplify climate change pressures on marine animal biomass for limited climate benefit. *Glob Change Biol* (2023) 29(18):5250–60. doi: 10.1111/gcb.16854
- 212. Rohr T. Southern Ocean iron fertilization: an argument against commercialization but for continued research amidst lingering uncertainty. *J Sci Policy Gov* (2019) 15:5847. doi: 10.38126/JSPG150114
- 213. Buesseler KO, Boyd PW. Climate change. Will ocean fertilization work? Science (2003) 300(5616):67-8. doi: 10.1126/science.1082959
- 214. Bach LT, Tamsitt V, Baldry K, McGee J, Laurenceau-Cornec EC, Strzepek RF, et al. Identifying the most (cost-)efficient regions for CO2 removal with iron fertilization in the Southern Ocean. *Glob Biogeochem Cycles* (2023) 37(11): e2023GB007754. doi: 10.1029/2023GB007754
- 215. Friedlingstein P, O'Sullivan M, Jones MW, Andrew RM, Hauck J, Landschützer P, et al. Global carbon budget 2024. Earth Syst Sci Data (2025) 17 (3):965–1039. doi: 10.5194/essd-17-965-2025
- 216. Freestone D, Rayfuse R. Ocean iron fertilization and international law. *Mar Ecol Prog Ser* (2008) 364:227–33. doi: 10.3354/meps07543
- 217. Eick M. A navigational system for uncharted waters: the London Convention and London Protocol's assessment framework on ocean iron fertilization. *Tulsa Law Rev* (2013) 46(2):351–78. Available at: https://digitalcommons.law.utulsa.edu/tlr/vol46/iss2/6
- 218. International Maritime Organization. Ocean fertilization under the LC/LP [online]. Available at: https://www.imo.org/en/ourwork/environment/pages/oceanfertilization-default.aspx.
- 219. International Maritime Organization. On the amendment to the London Protocol to regulate the placement of matter for ocean fertilization and other marine geoengineering activities [resolution LP.4(8)]. Eighth meeting of contracting parties to the 1996 Protocol. (2013). Available at: https://www.cdn.imo.org/localresources/en/KnowledgeCentre/IndexoflMOResolutions/LCLPDocuments/LP.4(8).pdf.
- 220. Keating-Bitonti C. Geoengineering: ocean iron fertilization. Congressional Research Service (2022). Available at: https://sgp.fas.org/crs/misc/R47172.pdf
- 221. Bowie AR, van der Merwe P, Quéroué F, Trull T, Fourquez M, Planchon F, et al. Iron budgets for three distinct biogeochemical sites around the Kerguelen Archipelago (Southern Ocean) during the natural fertilisation study, KEOPS-2. *Biogeosciences* (2015) 12(14):4421–45. doi: 10.5194/bg-12-4421-2015
- 222. National Oceanic and Atmospheric Administration Ocean Acidification Program. Announcing \$24.3M investment advancing marine carbon dioxide removal research [online] (2023). Available at: https://oceanacidification.noaa.gov/fy23-noppmcdr-awards/
- 223. United States Department of Energy. DOE announces \$36 million to advance marine carbon dioxide removal techniques and slash harmful greenhouse gas pollution [online] (2023). Available at: https://www.energy.gov/articles/doe-announces-36-million-advance-marine-carbon-dioxide-removal-techniques-and-slash
- 224. Tzachor A, Hendel O, Richards CE. Digital twins: a stepping stone to achieve ocean sustainability? NPJ Ocean Sustain (2023) 2(1):16. doi: 10.1038/s44183-023-00023-9
- 225. Rogelj J, Den Elzen MGJ, Portugal Pereira J. Chapter 4. The emissions gap in 2030 and beyond. The UNEP emissions gap report 2023: broken record temperatures hit new highs, yet world fails to cut emissions (again). Nairobi: UNEP (2023). Available at: https://www.unep.org/interactives/emissions-gap-report/2023/section\_-1
- 226. Rogelj J, Den Elzen MGJ, Portugal Pereira J. No more hot air ... Please! With a massive gap between rhetoric and reality, countries draft new climate commitments. Nairobi: UNEP (2024). Available at: https://www.undp.org/sites/g/files/zskgke326/files/2025-01/egr2024.pdf
- 227. Höhne N, Gidden MJ, den Elzen M, Hans F, Fyson C, Geiges A, et al. Wave of net zero emission targets opens window to meeting the Paris Agreement. Nat Clim Chang (2021) 11(10):820–2. doi: 10.1038/s41558-021-01142-2

- 228. Rogelj J, Geden O, Cowie A, Reisinger A. Net-zero emissions targets are vague: three ways to fix. Nature~(2021)~591(7850):365-8.~doi:~10.1038/d41586-021-00662-3
- 229. Rogelj J, Fransen T, den Elzen MGJ, Lamboll RD, Schumer C, Kuramochi T, et al. Credibility gap in net-zero climate targets leaves world at high risk. *Science* (2023) 380(6649):1014–6. doi: 10.1126/science.adg6248
- 230. Roser M. Why did renewables become so cheap so fast? Our World in Data [online] (2020). Available at: https://ourworldindata.org/cheap-renewables-growth? qls=QMM\_12345678.0123456789
- 231. Jung F, Jung F, Schröder M, Timme M. Exponential adoption of battery electric cars. *PloS One* (2023) 18(12):e0295692. doi: 10.1371/journal.pone.0295692
- 232. Lenton TM, McKay DI, Loriani S, Abrams JF, Lade SJ, Donges JF, et al. *The global tipping points report 2023* [thesis]. Exeter: University of Exeter (2023). Available at: https://hal.umontpellier.fr/hal-04548845
- 233. Rowell D, Connelly LB. A history of the term moral hazard. *J Risk Ins* (2012) 79 (4):1051–75. doi: 10.1111/j.1539-6975.2011.01448.x
- 234. Mahajan A, Tingley D, Wagner G. Fast, cheap, and imperfect? US public opinion about solar geoengineering. *Environ Polit* (2019) 28(3):523–43. doi: 10.1080/09644016.2018.1479101
- 235. Stephens JC, Kashwan P, McLaren D, Surprise K. The dangers of mainstreaming solar geoengineering: a critique of the national academies report. *Environ Polit* (2023) 32(1):157–66. doi: 10.1080/09644016.2021.1989214
- 236. Weber EU. Understanding public risk perception and responses to changes in perceived risk. In: Balleisen EJ, Bennear LS, Krawiec KD, Wiener JB, editors. *Policy shock: recalibrating risk and regulation after oil spills, nuclear accidents and financial crises*. Cambridge: Cambridge University Press (2017) 58–81. doi: 10.1017/9781316492635.004
- 237. Van Heerden II. The failure of the New Orleans levee system following Hurricane Katrina and the pathway forward. *Public Adm Rev* (2007) 67:24-35. doi: 10.1111/j.1540-6210.2007.00810.x
- 238. Smith M. An Iowa Town fought and failed to save a levee. Then came the flood. New York Times (2019). Available at: https://www.nytimes.com/2019/03/20/us/flooding-iowa-hamburg.html
- 239. Lieb D. States boost flood prevention as damage costs soar. The Chicago Tribune. (2019). Available at: https://www.chicagotribune.com/2019/08/12/states-boost-flood-prevention-as-damage-costs-soar/
- 240. Buchanan MK, Oppenheimer M, Parris A. Values, bias, and stressors affect intentions to adapt to coastal flood risk: a case study from New York City. Weather Clim Soc (2019) 11(4):809–21. doi: 10.1175/WCAS-D-18-0082.1
- 241. Qi D, Roe BE. Foodservice composting crowds out consumer food waste reduction behavior in a dining experiment. Am J Agric Econ (2017) 99(5):1159–71. doi: 10.1093/ajae/aax050
- 242. Steffen A. Predatory delay and the rights of future generations. Medium (2016). Available at: https://medium.com/@AlexSteffen/predatory-delay-and-the-rights-of-future-generations-69b06094a16
- 243. Kramer RC. The politics of predatory delay: climate crimes of political omission and socially organized denial. *Carbon criminals, climate crimes*. New Brunswick: Rutgers University Press (2020) 84–124. Available at: https://chooser.crossref.org/?doi=10.2307%2Fj.ctvxw3p51.7
- 244. Muffett C, Feit S. Fuel to the fire: how geoengineering threatens to entrench fossil fuels and accelerate the climate crisis. Washington, DC: Center for International Environmental Law (2019). Available at: https://www.ciel.org/wp-content/uploads/2019/02/CIEL\_FUEL-TO-THE-FIRE\_How-Geoengineering-Threatens-to-Entrench-Fossil-Fuels-and-Accelerate-the-Climate-Crisis\_February-2019.pdf
- 245. Almond D, Du X, Papp A. Favourability towards natural gas relates to funding source of university energy centres. Nat Clim Chang (2022) 12(12):1122-8. doi: 10.1038/s41558-022-01521-3
- 246. Harris B. The intractable cigarette 'filter problem'.  $Tob\ Control\ (2011)\ 20(1):$  i10–6. doi: 10.1136/tc.2010.040113
- 247. Stephens JC, Kashwan P, McLaren D, Surprise K. Toward dangerous US unilateralism on solar geoengineering. *Environ Polit* (2023) 32(1):171-3. doi: 10.1080/09644016.2022.2156182
- 248. Cherry TL, Kroll S, McEvoy DM, Campoverde D, Moreno-Cruz J. Climate cooperation in the shadow of solar geoengineering: an experimental investigation of the moral hazard conjecture. *Environ Polit* (2023) 32(2):362–70. doi: 10.1080/09644016.2022.2066285
- 249. Fritz L, Baum CM, Brutschin E, Low S, Sovacool BK. Climate beliefs, climate technologies and transformation pathways: contextualizing public perceptions in 22 countries. *Glob Environ Change* (2024) 87:102880. doi: 10.1016/j.gloenvcha.2024.102880
- 250. Buck HJ. Environmental peacebuilding and solar geoengineering. Front Clim (2022) 4:869774. doi: 10.3389/fclim.2022.869774
- 251. Armstrong McKay DI, Staal A, Abrams JF, Winkelmann R, Sakschewski B, Loriani S, et al. Exceeding 1.5°C global warming could trigger multiple climate tipping points. *Science* (2022) 377(6611):eabn7950. doi: 10.1126/science.abn7950
- 252. Lenton TM, Rockström J, Gaffney O, Rahmstorf S, Richardson K, Steffen W, et al. Climate tipping points too risky to bet against. Nature~(2019)~575(7784):592-5. doi: 10.1038/d41586-019-03595-0

- 253. Kopp RE, Gilmore EA, Shwom RL, Adams H, Adler C, Oppenheimer M, et al. 'Tipping points' confuse and can distract from urgent climate action. *Nat Clim Chang* (2025) 15(1):29–36. doi: 10.1038/s41558-024-02196-8
- 254. Rosier SHR, Reese R, Donges JF, De Rydt J, Gudmundsson GH, Winkelmann R. The tipping points and early warning indicators for Pine Island Glacier, West Antarctica. *Cryosphere* (2021) 15(3):1501–16. doi: 10.5194/tc-15-1501-2021
- 255. Klose AK, Coulon V, Pattyn F, Winkelmann R. The long-term sea-level commitment from Antarctica. *Cryosphere* (2024) 18(9):4463–92. doi: 10.5194/tc-18-4463-2024
- 256. Xie M, Moore JC, Zhao L, Wolovick M, Muri H. Impacts of three types of solar geoengineering on the Atlantic Meridional Overturning Circulation. *Atmos Chem Phys* (2022) 22(7):4581–97. doi: 10.5194/acp-22-4581-2022
- 257. Moore JC, Yue C, Zhao L, Guo X, Watanabe S, Ji D. Greenland ice sheet response to stratospheric aerosol injection geoengineering. *Earths Future* (2019) 7 (12):1451–63. doi: 10.1029/2019EF001393
- 258. Sutter J, Jones A, Frölicher TL, Wirths C, Stocker TF. Climate intervention on a high-emissions pathway could delay but not prevent West Antarctic Ice Sheet demise. *Nat Clim Change* (2023) 13(9):951–60. doi: 10.1038/s41558-023-01738-w
- 259. Chen Y, Ji D, Zhang Q, Moore JC, Boucher O, Jones A, et al. Northern-high-latitude permafrost and terrestrial carbon response to two solar geoengineering scenarios. *Earth Syst Dynam* (2023) 14(1):55–79. doi: 10.5194/esd-14-55-2023
- 260. Schleussner CF, Ganti G, Lejeune Q, Zhu B, Pfleiderer P, Prütz R, et al. Overconfidence in climate overshoot. *Nature* (2024) 634(8033):366–73. doi: 10.1038/s41586-024-08020-9
- 261. Sillmann J, Lenton TM, Levermann A, Ott K, Hulme M, Benduhn F, et al. Climate emergencies do not justify engineering the climate. *Nat Clim Change* (2015) 5(4):290–2. doi: 10.1038/nclimate2539
- 262. Corry O. The international politics of geoengineering: the feasibility of Plan B for tackling climate change. Sec Dialogue (2017) 48(4):297–315. doi: 10.1177/0967010617704142
- 263. McLaren DP. Whose climate and whose ethics? Conceptions of justice in solar geoengineering modelling. Energy Res Soc Sci (2018) 44:209–21. doi: 10.1016/j.erss.2018.05.021
- 264. Möller I. Political perspectives on geoengineering: navigating problem definition and institutional fit. *Glob Environ Polit* (2020) 20(2):57–82. doi: 10.1162/glep\_a\_00547
- 265. Triggs G. The Antarctic treaty system: a model of legal creativity and cooperation. In: Berkman PA, Lang MA, Walton DWH, Young OR, editors. *Science diplomacy: Antarctica, science, and the governance of international spaces*. Washington, DC: Smithsonian Institution Scholarly Press (2011) 39–50
- 266. Hughes KA, Lowther A, Gilbert N, Waluda CM, Lee JR. Communicating the best available science to inform Antarctic policy and management: a practical introduction for researchers. *Antarct Sci* (2023) 35(6):438–72. doi: 10.1017/S095410202300024X
- 267. Bastmeijer CJ, Roura R. Environmental impact assessment in Antarctica. In: Bastmeijer CJ, Koivurova T, editors. *Theory and practice of transboundary environmental impact assessment.* Leiden: Martinus Nijhoff (2007) 175–219
- 268. Dodds K, et al. *Handbook on the Politics of Antarctica*. Edward Elgar (2017). Available at: https://www.e-elgar.com/shop/gbp/handbook-on-the-politics-of-antarctica-9781784717674.html.
- 269. Hemmings AD, Kriwoken LK. High level Antarctic EIA under the Madrid Protocol: state practice and the effectiveness of the comprehensive environmental evaluation process. *Int Environ Agreements* (2010) 10(3):187–208. doi: 10.1007/s10784-010-9119-5
- 270. Antarctic Treaty Consultative Meeting. Final report of the forty-seventh Antarctic Treaty Consultative Meeting (preliminary version). Milan, Italy. 23 June 3 July 2025. Volume I. Buenos Aires: Secretariat of the Antarctic Treaty. (2025). Available at: https://documents.ats.aq/ATCM47/fr/ATCM47\_fr011\_e.pdf.
- 271. Smieszek M. Do the cures match the problem? Reforming the Arctic Council. Polar Rec (2019) 55(3):121-31. doi: 10.1017/S0032247419000263
- 272. Kornbech N, Corry O, McLaren D. Securing the 'great white shield'? Climate change, Arctic security and the geopolitics of solar geoengineering. *Coop Confl* (2025) 60(2):286–307. doi: 10.1177/00108367241269629
- 273. Landriault M, Chater A, Rowe EW, Lackenbauer PW. Governing complexity in the Arctic region. Abingdon: Routledge (2020).
- 274. Kurtze DA. Gravitational effects of ice sheets on sea level. Am J Phys (2022) 90(5):351–8. doi: 10.1119/5.0067924
- 275. Reckien D, Magnan AK, Singh C, Lukas-Sithole M, Orlove B, Schipper ELF, et al. Navigating the continuum between adaptation and maladaptation. *Nat Clim Change* (2023) 13(9):907–18. doi: 10.1038/s41558-023-01774-6
- 276. Jones CD, Frölicher TL, Koven C, MacDougall AH, Matthews HD, Zickfeld K, et al. The Zero Emissions Commitment Model Intercomparison Project (ZECMIP) contribution to C4MIP: quantifying committed climate changes following zero carbon emissions. *Geosci Model Dev* (2019) 12(10):4375–85. doi: 10.5194/gmd-12-4375-2019
- 277. Jourdain NC, Amory C, Kittel C, Durand G. Changes in Antarctic surface conditions and potential for ice shelf hydrofracturing from 1850 to 2200. *Cryosphere* (2025) 19(4):1641–74. doi: 10.5194/tc-19-1641-2025
- 278. Siegert MJ, Atkinson A, Banwell A, Brandon M, Convey P, Davies B, et al. The Antarctic Peninsula under a 1.5°C global warming scenario. *Front Environ Sci* (2019) 7:102. doi: 10.3389/fenvs.2019.00102

279. Mann ME. Warming ends when carbon pollution stops. Front Sci (2023) 1:1256273. doi: 10.3389/fsci.2023.1256273

- 280. Dudley N, editor. *Guidelines for applying protected area management categories*. Gland: International Union for Conservation of Nature (2008). Available at: https://portals.iucn.org/library/sites/library/files/documents/pag-021.pdf
- 281. Coetzee BWT, Gaston KJ, Chown SL. Local scale comparisons of biodiversity as a test for global protected area ecological performance: a meta-analysis. *PloS One* (2014) 9(8):e105824. doi: 10.1371/journal.pone.0105824
- 282. Gray CL, Hill SLL, Newbold T, Hudson LN, Börger L, Contu S, et al. Local biodiversity is higher inside than outside terrestrial protected areas worldwide. *Nat Commun* (2016) 7:12306. doi: 10.1038/ncomms12306
- 283. Barnes AE, Davies JG, Martay B, Boersch-Supan PH, Harris SJ, Noble DG, et al. Rare and declining bird species benefit most from designating protected areas for conservation in the UK. *Nat Ecol Evol* (2023) 7(1):92–101. doi: 10.1038/s41559-022-01927-4
- 284. Brodie JF, Mohd-Azlan J, Chen C, Wearn OR, Deith MCM, Ball JGC, et al. Landscape-scale benefits of protected areas for tropical biodiversity. *Nature* (2023) 620 (7975):807–12. doi: 10.1038/s41586-023-06410-z
- 285. Soykan CU, Lewison RL. Using community-level metrics to monitor the effects of marine protected areas on biodiversity. *Conserv Biol* (2015) 29(3):775–83. doi: 10.1111/cobi.12445
- 286. Gill DA, Mascia MB, Ahmadia GN, Glew L, Lester SE, Barnes M, et al. Capacity shortfalls hinder the performance of marine protected areas globally. *Nature* (2017) 543 (7647):665–9. doi: 10.1038/nature21708
- 287. Cheng BS, Altieri AH, Torchin ME, Ruiz GM. Can marine reserves restore lost ecosystem functioning? A global synthesis. *Ecology* (2019) 100(4):e02617. doi: 10.1002/ecv.2617
- 288. Medoff S, Lynham J, Raynor J. Spillover benefits from the world's largest fully protected MPA. *Science* (2022) 378(6617):313–6. doi: 10.1126/science.abn0098
- 289. Bruno JF, Bates AE, Cacciapaglia C, Pike EP, Amstrup SC, Van Hooidonk R, et al. Climate change threatens the world's marine protected areas. *Nat Clim Change* (2018) 8(6):499–503. doi: 10.1038/s41558-018-0149-2
- 290. ASamoah EF, Beaumont LJ, Maina JM. Climate and land-use changes reduce the benefits of terrestrial protected areas. *Nat Clim Chang* (2021) 11(12):1105–10. doi: 10.1038/s41558-021-01223-2
- 291. Jacquemont J, Blasiak R, Le Cam C, Le Gouellec M, Claudet J. Ocean conservation boosts climate change mitigation and adaptation. *One Earth* (2022) 5 (10):1126–38. doi: 10.1016/j.oneear.2022.09.002
- 292. Xu X, Huang A, Belle E, De Frenne P, Jia G. Protected areas provide thermal buffer against climate change. Sci~Adv~(2022)~8(44):eabo0119. doi: 10.1126/sciadv.abo0119
- 293. Mi C, Ma L, Yang M, Li X, Meiri S, Roll U, et al. Global protected areas as refuges for amphibians and reptiles under climate change. *Nat Commun* (2023) 14(1):1389. doi: 10.1038/s41467-023-36987-v
- 294. Duncanson L, Liang M, Leitold V, Armston J, Krishna Moorthy SM, Dubayah R, et al. The effectiveness of global protected areas for climate change mitigation. *Nat Commun* (2023) 14(1):2908. doi: 10.1038/s41467-023-38073-9
- 295. Dinerstein E, Joshi AR, Vynne C, Lee ATL, Pharand-Deschênes F, França M, et al. A "Global Safety Net" to reverse biodiversity loss and stabilize Earth's climate. *Sci Adv* (2020) 6(36):eabb2824. doi: 10.1126/sciadv.abb2824
- 296. Elsen PR, Monahan WB, Dougherty ER, Merenlender AM. Keeping pace with climate change in global terrestrial protected areas. *Sci Adv* (2020) 6(25):eaay0814. doi: 10.1126/sciadv.aay0814
- 297. Pinsky ML, Rogers LA, Morley JW, Frölicher TL. Ocean planning for species on the move provides substantial benefits and requires few trade-offs.  $Sci\ Adv\ (2020)\ 6(50)$ : eabb8428. doi: 10.1126/sciadv.abb8428
- 298. Ranius T, Widenfalk LA, Seedre M, Lindman L, Felton A, Hämäläinen A, et al. Protected area designation and management in a world of climate change: a review of recommendations. *Ambio* (2023) 52(1):68–80. doi: 10.1007/s13280-022-01779-z
- 299. Hughes KA, Convey P, Turner J. Developing resilience to climate change impacts in Antarctica: an evaluation of Antarctic Treaty System protected area policy. *Environ Sci Policy* (2021) 124:12–22. doi: 10.1016/j.envsci.2021.05.023
- 300. Bennett JR, Shaw JD, Terauds A, Smol JP, Aerts R, Bergstrom DM, et al. Polar lessons learned: long-term management based on shared threats in Arctic and Antarctic environments. *Frontiers Ecol Environ*. (2015) 13:316–3324. doi: 10.1890/140315
- 301. Nissen C, Lovenduski NS, Brooks CM, Hoppema M, Timmermann R, Hauck J. Severe 21st-century ocean acidification in Antarctic Marine Protected Areas. *Nat Commun* (2024) 15(1):259. doi: 10.1038/s41467-023-44438-x
- 302. Chown SL, Brooks CM, Terauds A, Le Bohec C, van Klaveren-Impagliazzo C, Whittington JD, et al. Antarctica and the strategic plan for biodiversity. *PloS Biol* (2017) 15(3):e2001656. doi: 10.1371/journal.pbio.2001656
- 303. Brooks CM, Crowder LB, Curran LM, Dunbar RB, Ainley DG, Dodds KJ, et al. Science-based management in decline in the Southern Ocean. *Science*. (2016) 354(6309):185–7. doi: 10.1126/science.aah4119