

1 **Characterising groundwater-surface water connectivity in lower Gandak catchment, a barrage**
2 **regulated biodiversity hotspot in the mid-Gangetic basin**

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14 **Highlights**

- 15 • Diffuse rainfall dominates groundwater recharge
- 16 • Groundwater baseflow is important for river ecology
- 17 • Evidence for localised groundwater recharge from canals and rivers
- 18 • Adequate flushing sustains low groundwater salinity
- 19 • Local redox conditions control groundwater As, Fe, Mn, NO₃ and U

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23

24 **Abstract**

25 The alluvial aquifer system of the Indo-Gangetic Basin (IGB) is one of the world's most important
26 freshwater resources, sustaining humans and river ecosystems. Understanding groundwater
27 recharge processes and connections to meteoric and surface water is necessary for effective water
28 resource management for human and wider ecological requirements. Parts of the mid-Gangetic
29 Basin, across eastern Uttar Pradesh and Bihar, are characterised by stable long-term groundwater
30 levels, high annual rainfall, and limited historical groundwater use compared to parts of Northwest
31 India for example. In this paper we use a combination of environmental tracers and hydrograph
32 observations to characterise sources of recharge and groundwater-surface water interaction using a
33 transect approach across the catchment of the River Gandak, a major barrage-regulated tributary of
34 the River Ganga. Stable isotope results show that the dominant source of groundwater recharge, in
35 the shallow (0-40 m bgl) Holocene and underlying Pleistocene aquifer system (>40 m bgl), is local
36 rainfall. The shallow Holocene aquifer is also supplemented by local recharge from river and canal
37 seepage and irrigation return flow in the upper and mid parts of the catchment. These observations
38 are corroborated by evidence from detailed groundwater hydrographs and salinity observations,
39 indicating localised canal, river and lake connectivity to groundwater. In the middle and lower
40 catchment, river discharge is dominated by groundwater baseflow during the peak dry season when
41 barrage gates are closed, which contributes to ecological flows for endangered river dolphins and
42 gharial crocodiles. Groundwater residence time tracers indicate active modern recharge in the
43 shallow alluvial aquifer system across the catchment. In the shallow Holocene aquifer elevated
44 arsenic (As), iron (Fe), and manganese (Mn) exceeded WHO drinking water guidelines in a minority
45 of sites, and uranium (U) and fluoride (F) concentrations approach but do not exceed the WHO
46 guideline values. These observations varied across the catchment with higher As, Fe and Mn in the
47 upper and mid catchments and higher U in the lower catchment. Groundwater salinity was typically
48 between 500 and 1000 $\mu\text{S}/\text{cm}$, and isolated higher salinity was due to recharge from flood-plain
49 wetlands and lakes impacted by evaporation. At present, the Gandak catchment has relatively high

50 rainfall and low abstraction, which maintains stable groundwater levels and thus baseflow to the
51 river in the dry season. Potential future threats to groundwater resources, and therefore river
52 ecology due to the sensitivity to changes in baseflow in the catchment, would likely be driven by
53 reductions in local monsoon rainfall, changes in water management practices and increased
54 groundwater use.

55 **Keywords:** Groundwater, Indo-Gangetic Basin, recharge, surface-groundwater interactions, canal,
56 River Gandak

57 **1. Introduction**

58 Groundwater is a critical resource for hundreds of millions of people in the Indo-Gangetic Basin (IGB)
59 who rely on it for drinking water, agriculture and industry (MacDonald et al., 2016). Groundwater
60 also sustains river flows and ecology through baseflow (Mukherjee et al. 2018). India's annual rate of
61 groundwater abstraction is the highest globally, now exceeding $200 \text{ km}^3 \text{ yr}^{-1}$, and continues to
62 increase (Wada et al., 2010). Shallow groundwater has been used for centuries in the IGB, but over
63 the last 40 to 50 years, and most markedly in the last 20 years, there has been a transition in food
64 production from a reliance on surface water and rain-fed agriculture towards increased dependence
65 on groundwater irrigation (Shah 2009). This has its roots in the agrarian economy of the region and
66 the erratic rainfall patterns of the Indian monsoon (Gadgil et al., 1999, Sinha et al., 2007).

67 Groundwater recharge processes and connectivity with surface water varies across the IGB.

68 Managing water resources effectively for human uses as well as river-floodplain ecosystems relies on
69 a good spatiotemporal understanding of connectivity processes. Previous work in the IGB using
70 environmental tracers and groundwater level time-series observations have illustrated the impact of
71 excessive groundwater pumping for irrigation in depleting groundwater levels, and also the
72 important role of surface irrigation in controlling groundwater recharge (Kumar et al., 2011, Sharma
73 et al., 2014, Lapworth et al., 2015; 2018; MacDonald et al., 2016, Joshi et al., 2018, O'Keeffe et al.,
74 2020). Several of these studies (e.g. Lapworth et al., 2015, Joshi et al., 2018) have explored the

75 relative importance of current rainfall versus river and canal water for groundwater recharge using
76 water stable isotopes supported by a range of residence time tracers. These studies have tended to
77 focus on areas of the IGB that have experienced the largest depletion in groundwater levels (e.g.
78 Lapworth et al., 2015; Joshi et al., 2018; van Dijk et al., 2019). However, there are many parts of the
79 basin where long-term groundwater levels are shown to be stable over the last few decades
80 (MacDonald et al., 2016), as a result of limited groundwater exploitation in relation to the role of
81 surface water sources in maintaining groundwater levels.

82 The mid-Ganga region, specifically the lower portion of the Gandak catchment overlying the IGB
83 aquifer in India, is one such area that has received less attention to date, and is the focus of this
84 study (Figure 1). It was specifically chosen for its suitability for exploring groundwater interactions
85 with river and canal water in an area of stable groundwater levels within the IGB.

86 The Gandak basin sustains thriving agriculture, capture fisheries, riverine forests and grasslands with
87 large mega-herbivores and predators including the Tiger (*Panthera Tigris*). The river holds a globally
88 significant breeding population of the critically endangered Gharial crocodile (*Gavialis gangeticus*),
89 besides a population of the Ganges river dolphin *Platanista gangetica* (Choudhary et al. 2012). The
90 Valmiki Tiger Reserve and a proposed Conservation Reserve on the downstream stretch of the river
91 are part of the Gandak biodiversity conservation regimes. In 1960 a barrage was built across the
92 river at the foot of the Shivalik-Himalayan hills on the India-Nepal border, which was followed by the
93 expansion of a large canal network. This canal network supports agricultural production mostly in
94 the dry-season, but suffers from major inefficiencies and has contributed to waterlogging in the
95 command area (Chowdary et al., 2008, Khan et al., 2014). Compared to other regions within the IGB
96 in Northwest India, groundwater pumping for irrigation and drinking water supply in the Gandak
97 region is less intense and mainly restricted to the shallow part of the aquifer system (MacDonald et
98 al., 2016, Bonsor et al., 2017).

99 In this paper, we examine evidence of whether local rainfall and surface water sources (*i.e.* leakage
100 or seepage from canals, the River Gandak and other surface water sources) maintain groundwater
101 recharge and groundwater discharge as baseflow to the river, in the Gandak catchment. We then
102 extend this to discuss how anthropogenic hydrological interventions, especially flow regulation by a
103 barrage and abstraction through canals, have altered groundwater recharge and flow processes
104 within this sub-catchment of the IGB where rainfall is higher (> 1000 mm) and groundwater
105 abstraction is not excessive. Our primary working hypothesis is that canal leakage/seepage and
106 other surface water sources recharge the alluvial aquifer system, but this is largely restricted to the
107 shallow upper part of the aquifer, and that the influence of canal leakage/seepage as a recharge
108 source reduces as you move down catchment and away from the upper canal command area. We
109 test the hypothesis using a suite of environmental tracers to characterise recharge sources and
110 processes along a sampling transect from the upper to lower Gandak catchment (Figure 1) within the
111 non-mountainous (plains) zone of the basin, and supplement this with new targeted high frequency
112 groundwater level observations and spot river flow measurements.

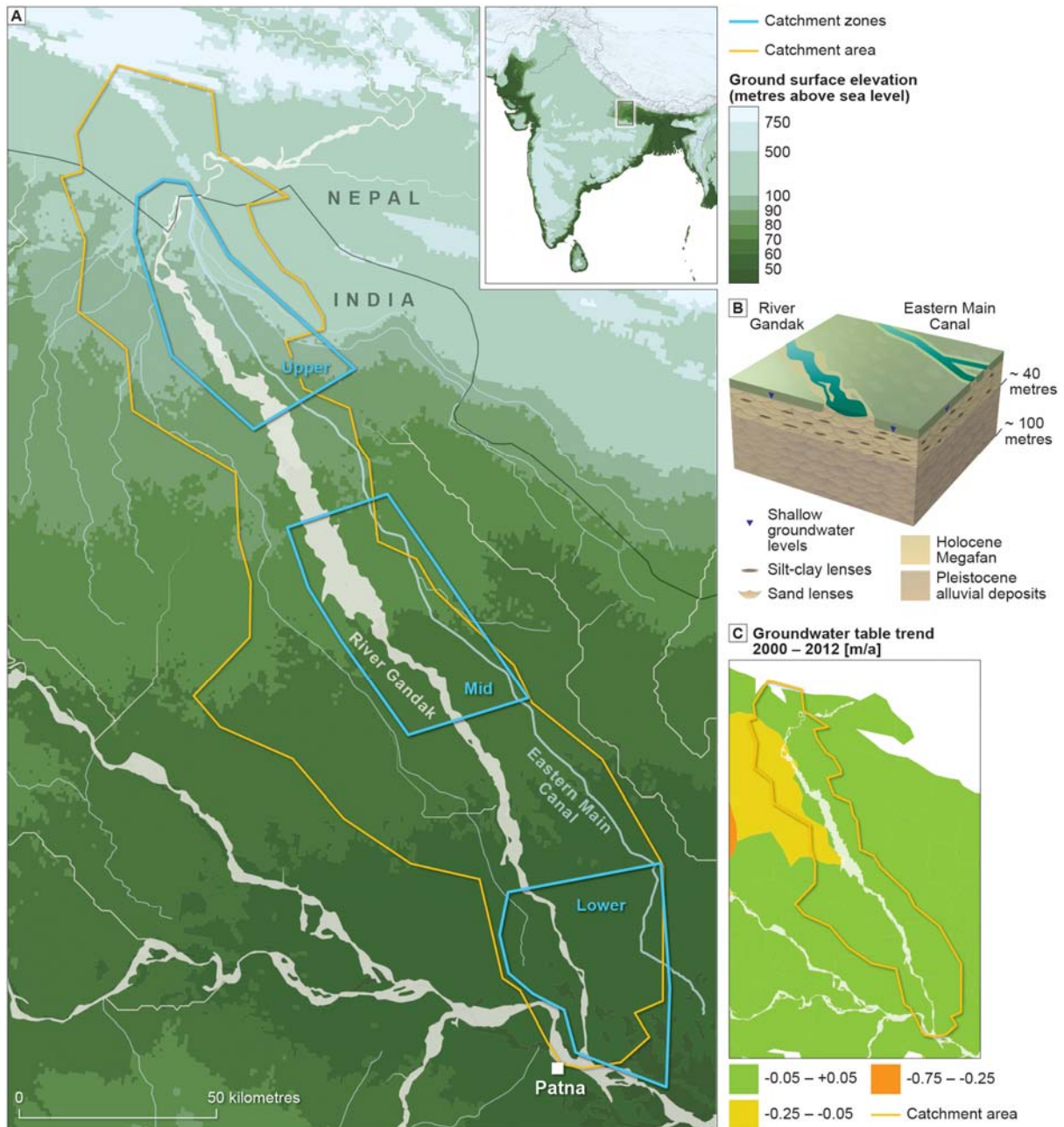
113 **2. Methods**

114 **2.1 Study area**

115 The River Gandak rises on the Tibetan plateau and flows through Nepal before entering India in
116 north-west Bihar where the Shivalik mountains meet the alluvial plains. From there it flows
117 southeast for 335 km before joining the River Ganga close to Patna (Figure 1a). The Gandak
118 catchment area is 46,000 km², but this study is focussed on the downstream section of the Gandak
119 catchment in India, which covers an area of 7600 km², and is henceforth referred to in this paper as
120 the 'catchment'. The River Gandak has the third highest contributing flow of any tributary to the
121 Ganga (MacDonald et al., 2015), and is prone to periodic flooding (Chowdary et al., 2008). Long-term
122 annual average rainfall in the catchment is c.1200 mm (Singh et al., 2008). The lower section of the
123 catchment receives 800-1200 mm of rainfall annually and the upper section receives 1200-2000 mm

124 annually (Sinha and Friend, 1994; Pai et al., 2014), with high inter-annual variability. Annual average
125 discharge for the Gandak recorded at Dumariaghat is $1555 \text{ m}^3 \text{ s}^{-1}$. The flow in the River Gandak has
126 been shown to gain after the Barrage, reflecting in-catchment contributions to flow (Sinha and
127 Friend, 1994).

128 The flow in the River Gandak is regulated by the Gandak barrage on the India-Nepal border, from
129 where three main canal branches distribute water from the river: the Nepalese; western; and
130 eastern canal command branches (Jha and Prasad, 2002). The eastern canal branch supplies Bihar
131 and is within the current study area. Large scale irrigation using canal water began in the early 1970s
132 (Jha and Prasad, 2002). The canal system is clay lined and leakage losses of up to 50% are reported
133 increasing groundwater recharge, particularly in the uppermost section of the Eastern branch
134 (Bonsor et al., 2017). There is intense use of canal water for irrigation in the northern section of the
135 catchment and extensive use of shallow groundwater for irrigation across the whole catchment
136 (Government of Bihar 2014; O'Keeffe et al., 2020). For much of the dry season, canal flows do not
137 reach the southern section of the catchment and there was no change, between 2000-2012, in
138 groundwater levels across the catchment in a recent study by MacDonald et al. (2016).



139

140 **Figure 1.** Study area location, sedimentology and long-term groundwater trends a) Gandak
 141 catchment and study zones in the upper, mid and lower Gandak, b) schematic sedimentary
 142 architecture of the Gandak megafan (adapted from Pati et al., 2019), c) long-term 2000-2012
 143 groundwater trends for the Gandak and surrounding area (source MacDonald et al., 2016).

144 **2.2 Hydrogeology and sedimentology of Gandak catchment**

145 The IGB is the largest area of modern alluvial sedimentation in the world, with sedimentary deposits
 146 up to several km deep in places, (Sinha and Sarkar, 2009; Densmore et al., 2016). Holocene and

147 Pleistocene alluvium form two distinct units within the overall IGB aquifer system, from which all
148 groundwater in the basin is abstracted. At local scales, these sedimentary units are composed of
149 alternating gravels, sands, silts and occasional clays, in sequences of channel and inter-channel
150 deposits forming laterally discontinuous aquifer units, typically less than a few kilometres across
151 (Samadder et al., 2011; Bonsor et al., 2017). At a basin scale, vertical connectivity reduces from
152 proximal to distal regions as low permeability layers become more dominant (MacDonald et al.,
153 2016, Bonsor et al., 2017).

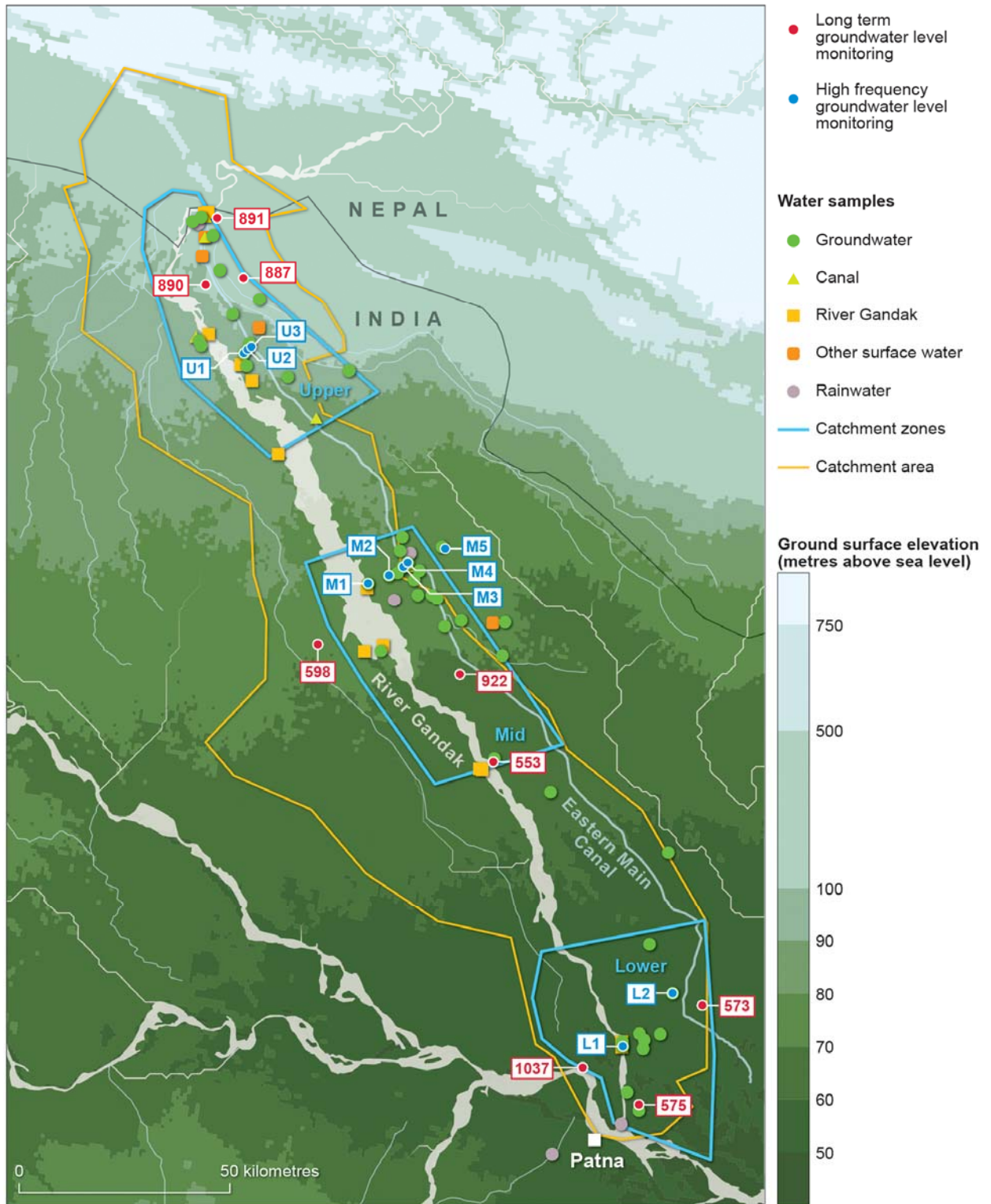
154 The Gandak megafan directly underlies the study catchment, composed of Holocene alluvial
155 sediments deposited by the River Gandak system to produce an alluvial plain with low relief (Sinha
156 and Friend., 1994; Pati et al., 2019). The Holocene megafan sediments were deposited as the River
157 Gandak migrated eastward, forming up to 40 m of mainly sand deposits with minor clay units, with
158 the sands fining, and clay increasing, towards the distal parts of the megafan (Pati et al., 2019).
159 There are very few gravel deposits in the megafan due to the trapping, abrasion and reworking of
160 gravel within the streambed in Nepal (Sinha and Sarkar 2009; Dingle et al., 2017). The megafan
161 deposits can form significant high permeability aquifers (Bonsor et al., 2017). Beneath the Holocene
162 mega fan is a thick sequence of Pleistocene aged alluvium. There are very few boreholes that
163 directly characterise these deposits in the Gandak, but the few that do indicate significant and
164 extensive sand (Sinha et al., 2014). They are most likely similar to Pleistocene deposits elsewhere in
165 the Ganga basin composed of alternating coarse and fine sands, silts and occasionally clays,
166 deposited within sequences of channel and inter-channel deposits, with each package rarely more
167 than a few kilometres across, and any one individual unit is generally less than 50 m thick (Bonsor et
168 al., 2017) (Figure 1b).

169 Mohindra et al (1992) provide a comprehensive assessment of the pedology of the Gandak megafan,
170 and report a distinct increase in soil clay accumulation (illuviated clay) away from the Young Gandak
171 Plain, close to the river, towards the Old Gandak Plain (see DEM on Figure 1a). This is accompanied

172 by a significant increase in soil acidity moving from the Young Plain soils (pH 7.8-8.8) to the Old Plain
173 soils (pH 5.8-6.2), related to the decalcification of parent material (Mohindra et al., 1992). Both the
174 sedimentary architecture and these changes in soil characteristics within the catchment may
175 influence localised recharge processes and shallow groundwater chemistry.

176 ***2.3 Water sampling and analysis***

177 A sampling transect down the catchment was established on the eastern side of the River Gandak, in
178 the eastern canal command area. Samples were collected for different water sources (groundwater,
179 surface water and rainfall) to characterise changes in groundwater and surface water down the
180 transect. Sampling was focussed on three zones within the Gandak: the (i) upper (ii) mid and (iii)
181 lower sections of the catchment (Figure 2). This enabled groundwater recharge processes to be
182 evaluated in different parts of the catchment where the use of surface water for irrigation varies -
183 from the intensive use of canal water in the upper and mid sections to lower use in the lower
184 section. Boreholes were selected in consultation with local farmers to enable groundwater sampling.



185

186 **Figure 2.** Sample sites and groundwater level monitoring points

187 Groundwater chemistry was sampled under pre-monsoon conditions following purging of boreholes

188 and achieving stable readings for field parameters. A total of 61 groundwater samples was collected

189 from existing boreholes from a depth range between <10-70 m below ground level (bgl), and a single
190 sample from a deeper borehole at 150 m bgl. Eighteen River Gandak samples, 4 other surface water
191 samples, 7 Gandak canal water samples and 18 rainfall samples were also collected. Some of the 61
192 groundwater samples, including handpumps, were taken from boreholes that were in regular use
193 and therefore did not require extensive purging prior to sampling, other sites not in regular use were
194 purged prior to sampling.

195 In total, 109 water samples were collected for stable isotope analysis; 13 samples for groundwater
196 residence time analysis and 45 samples for inorganic chemistry analysis (O Dochartaigh et al.,
197 2020b). A total of 55 of these surface and groundwater samples were measured for specific
198 electrical conductivity (SEC) in the field, for the other grab samples where field measurements were
199 not possible and laboratory SEC measurements were undertaken. Rainfall samples were collected
200 using a totaliser with a tube and 'dip-in' design suitable for the climate of India, which minimises
201 evaporative effects (IAEA/GNIP 2014). The presence of *kankar* carbonate deposits make the use of
202 radio-carbon dating of groundwater problematic in this region, and elsewhere in NW India, and
203 therefore other groundwater residence time tracers were used. Groundwater residence time tracers
204 samples for chlorofluorocarbon (CFC)-11, 12 and sulphur hexafluoride (SF₆) were collected
205 unfiltered, and without atmospheric contact, in sealed containers by the displacement method of
206 Oster et al. (1996). A recharge temperature of 28 °C was assumed for calculating tracer
207 concentrations. SF₆ data were corrected for excess air at 3 cc/L. All chemical analysis was undertaken
208 at BGS geochemistry and groundwater tracer laboratories in the UK. Stable isotope results are
209 reported as a deviation from Vienna Standard Mean Ocean Water (vs. VSMOW) in per mil (‰)
210 difference using delta (δ) notation.

211 **2.4 Ground-water level monitoring**

212 Groundwater levels measured in the Gandak catchment over the period 1983 to 2013 were obtained
213 from the Central Groundwater Board (CGWB) monitoring network (Macdonald et al., 2016). These

214 time-series typically consist of 2-6 manual dips per year. Nine monitoring points were selected to
215 assess long-term trends in groundwater levels across the three catchment zones (upper, mid and
216 lower): seven to the east of the Gandak and two to the west.

217 Additionally, automatic groundwater level loggers, recording at 15-minute intervals, were installed
218 in 10 boreholes in farmers' fields along three transects, one in each of the three zones
219 (O Dochartaigh et al., 2020a). Six sites (U1-3 and M1-3) were newly constructed piezometers to a
220 depth of 20 m. These were used to quantify detailed seasonal changes in groundwater levels and
221 characterise responses along transects from the River Gandak to, and beyond, the Eastern Main
222 Canal, as well as the possible effect of shallow groundwater pumping for irrigation. The observations
223 span the period 1 April 2017 to 31 January 2019, but the exact period of monitoring was different
224 between the boreholes, due to the different dates of their construction, or in one case (M5) the loss
225 of the logger. The locations of the monitored boreholes in the three transects are shown in Figure 2.

226 ***2.5 Seasonal monitoring of river discharge and rainfall data***

227 River discharge was measured at reference sites in the three zones, at approximate distances of 0,
228 44, 88, 140, 190, and 295 km from the Gandak barrage in the: i) post-monsoon season (November
229 2017, when all barrage gates were open), and ii) summer (March 2018, when barrage gates were
230 closed). River discharge measurements were undertaken using the velocity-area method and
231 Manning's equation (from river energy slope and roughness coefficients). A decrease in discharge
232 with increasing distance was regarded as contributions to recharge, and an increase was regarded as
233 gains in discharge or base-flows contributed by the aquifer.

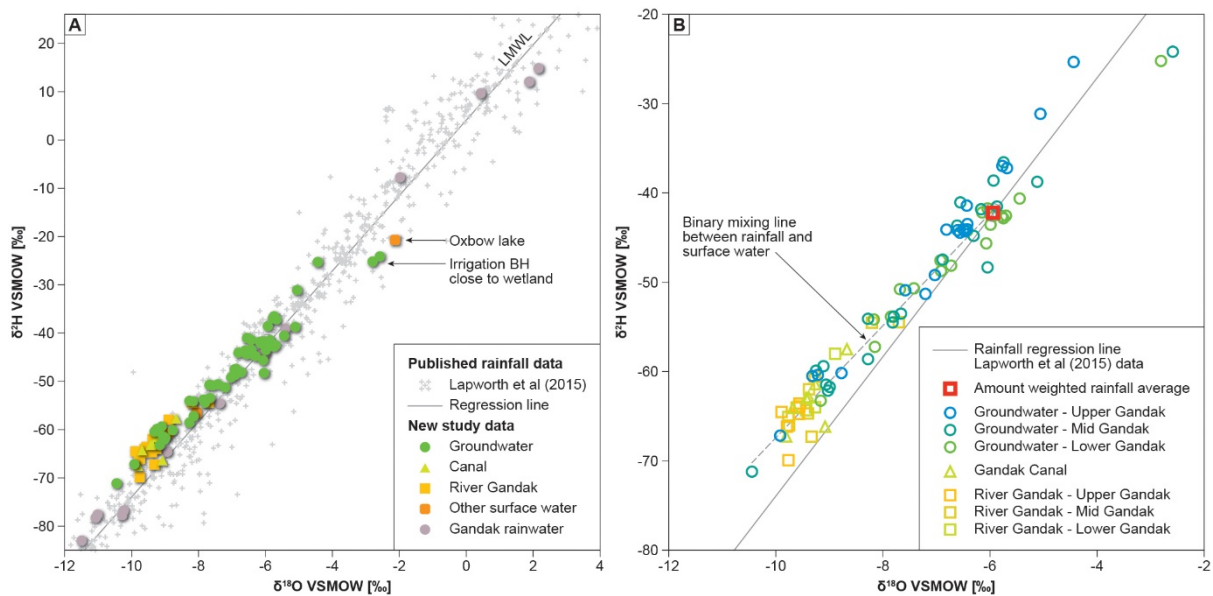
234 Estimates of daily rainfall for the three catchment zones were derived from the Indian
235 Meteorological Department's 0.25° gridded dataset (Pai et al., 2014).

236

237 **3. Results**

238 **3.1 Water stable isotopes**

239 Figure 3 shows a summary of the water stable isotope results as a cross-plot of $\delta^{18}\text{O}$ vs $\delta^2\text{H}$ for the
240 full range of water types collected as part of this study. Groundwater and surface water results are
241 differentiated by catchment zone (Figure 3b). An indicative binary mixing line between groundwater
242 and surface water (Figure 3b) illustrates surface water-groundwater interactions and mixing
243 between water from meteoric rainfall and surface water sources. The local meteoric water line and
244 amount-weighted rainfall mean values are shown for comparison. A small sub-set of sites (n=3)
245 showing significant enrichment (Figure 3a), from an oxbow lake sample and two shallow boreholes
246 close to wetlands, are indicative of evapotranspiration processes prior to recharge.



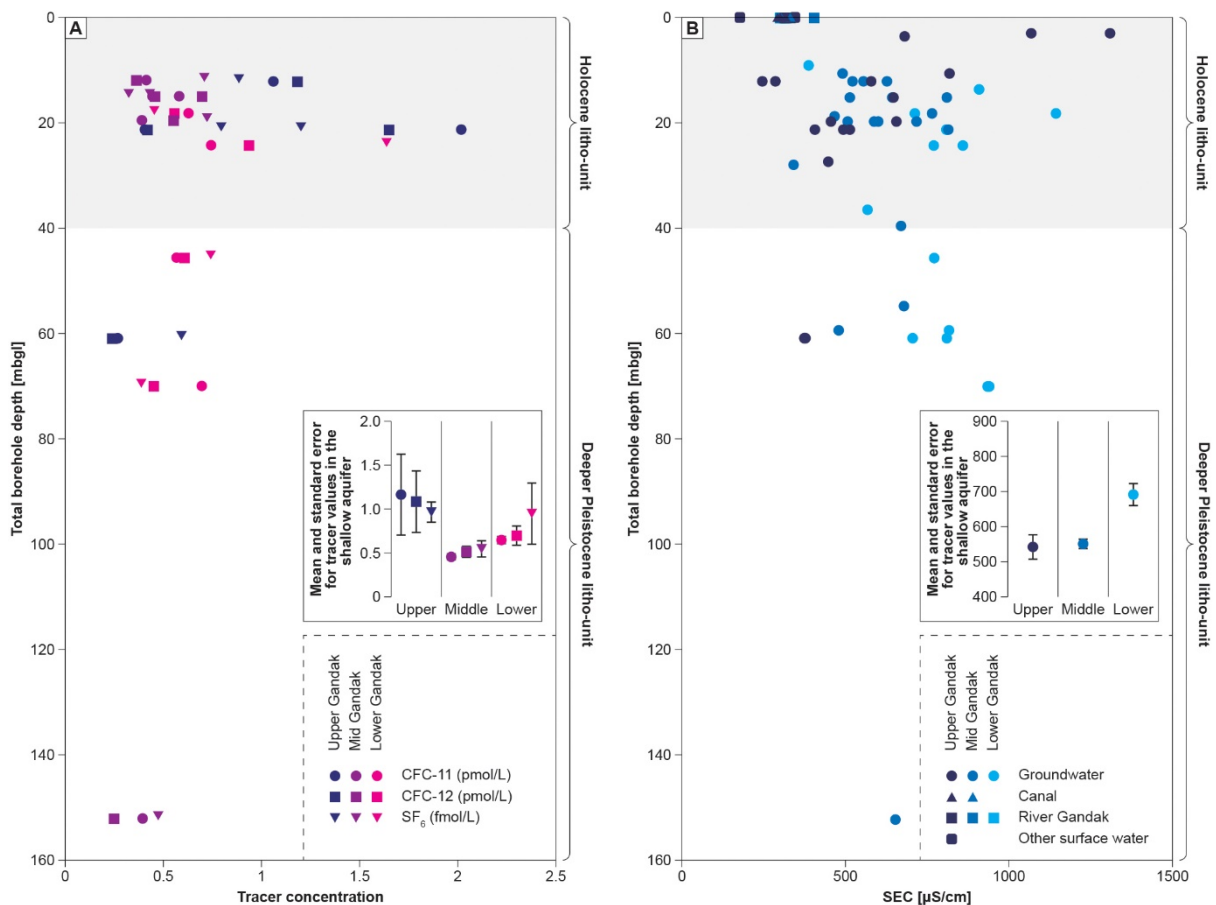
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248 **Figure 3.** Water stable isotope results: a) all data by water type b) groundwater and surface water
249 results by catchment zone (upper, mid and lower catchment). Local meteoric water line (LMWL) and
250 amount-weighted rainfall mean from Lapworth et al. (2015) using data for NW India.

251 **3.2 Groundwater residence time tracers**

252 The variations in groundwater residence time tracer concentration (CFC-12, CFC-11, SF_6) and salinity
253 with total borehole depth are shown in Figure 4. There is an overall decrease in tracer
254 concentrations in groundwater with depth (Figure 4a), but there is considerable variability within the

255 shallow (<40 m bgl) groundwater system, within the Holocene aquifer, and there are few
 256 observations from the deeper (>40 m bgl) Pleistocene aquifer. Within the shallow aquifer, c. 30% of
 257 the observed tracer concentrations are indicative of dominantly recent recharge. The remaining 70%
 258 of samples from the shallow aquifer, and all from the deeper aquifer, all contain some modern
 259 tracers but are all are predominantly older water. The upper catchment zone had the largest
 260 proportion of shallow (<40 m bgl) samples with high concentrations of modern tracers, for CFC
 261 tracers these were significantly higher than the other zones, however this was not found to be the
 262 case for SF₆ (Figure 4a). In the mid catchment, groundwater showed no evidence of changes in tracer
 263 concentration with depth, and shallow groundwaters had lower tracer concentrations compared to
 264 the other two zones (Figure 4a). There was also no marked change in salinity with depth (Figure 4b),
 265 but surface waters showed lower salinity and lower variability than groundwaters. The salinity of
 266 samples in the lower catchment zone are significantly higher for the Holocene aquifer (Figure 4b),
 267 compared to the upper and mid zones which have comparable SEC values (Figure 4b).

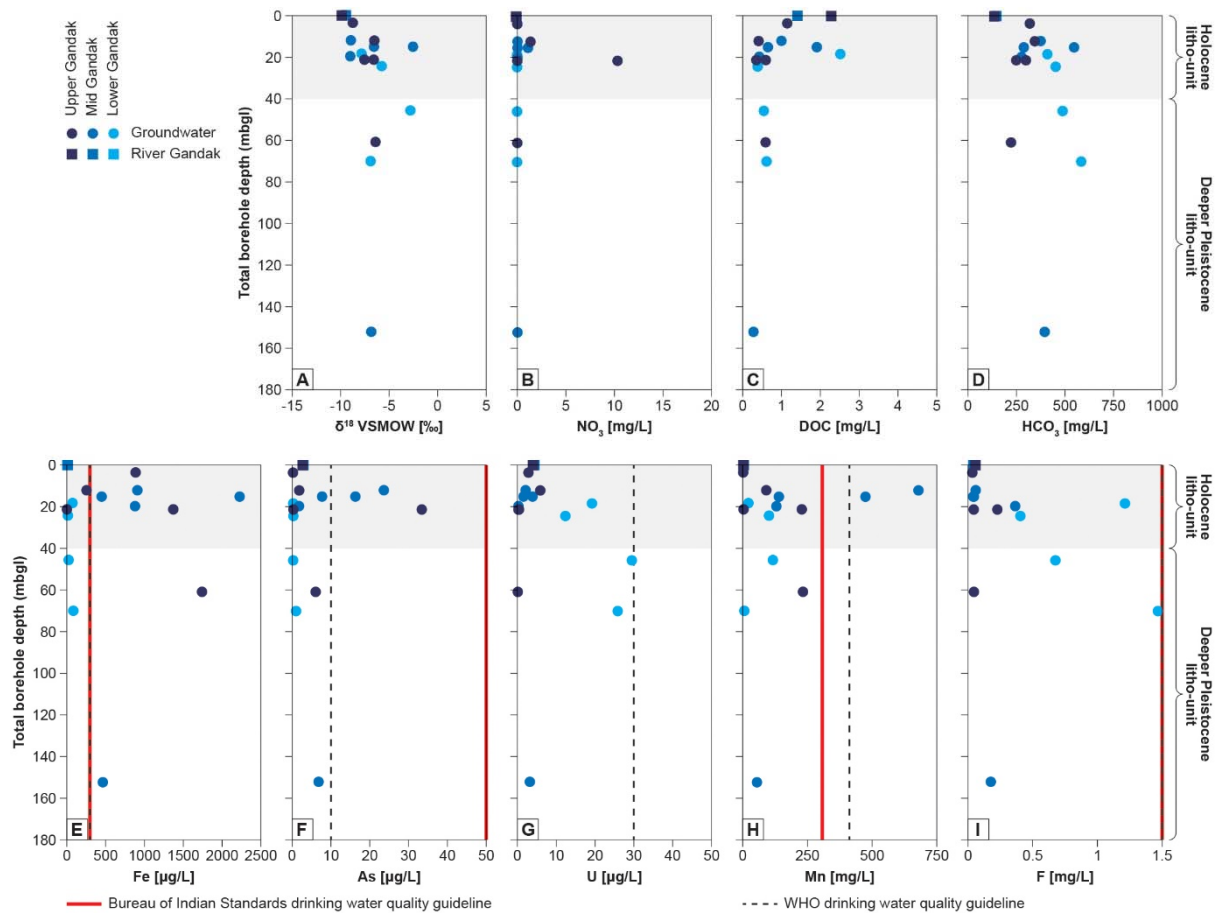


268

269 **Figure 4.** Groundwater residence time and surface and groundwater salinity results for upper, mid
270 and lower catchment zones, as depth profiles. a) Tracer concentrations, b) Salinity plotted as SEC
271 ($\mu\text{S}/\text{cm}$). Data is plotted using total borehole depth: casing depth information was not available.

272 **3.3 Inorganic hydrochemistry**

273 Depth variations for selected hydrochemical parameters in groundwater are shown with canal and
274 river water data for comparison (Figure 5a-i). With the exception of dissolved organic carbon (DOC),
275 surface waters have lower variability and lower concentrations for all parameters compared to
276 groundwater, and $\delta^{18}\text{O}$ values are depleted and tightly clustered relative to groundwater. Several
277 parameters for groundwater samples (e.g. stable isotopes, SEC, DOC, and some redox sensitive
278 parameters, such as nitrate (NO_3), arsenic (As) and manganese (Mn), are characterised by higher
279 variability in the shallow Holocene than in the deeper Pleistocene aquifer. Other parameters,
280 including U, F and Fe, are variable in both shallow and deeper aquifers (Figure 5). The majority of
281 groundwater samples have low nitrate ($<5 \text{ mg}/\text{L NO}_3$) and As ($<10 \mu\text{g}/\text{L}$) concentrations. In a few
282 groundwater samples ($n=5$) from the Holocene aquifer, As concentrations exceed the WHO drinking
283 water guideline value of $10 \mu\text{g}/\text{L}$ (WHO 2017), but none exceed the Bureau of Indian Standards (BIS
284 2012) of $50 \mu\text{g}/\text{L}$. Fluoride concentrations are all below $1.5 \text{ mg}/\text{L}$ (BIS and WHO drinking water
285 guideline value). For U no sites exceed Indian drinking water limit of $60 \mu\text{g}/\text{L}$ or WHO limit of 30
286 $\mu\text{g}/\text{L}$, but the highest U value ($29.5 \mu\text{g}/\text{L}$) approaches the WHO limit. Nine samples exceed the BIS
287 limit of $0.3 \text{ mg}/\text{L}$ for iron, and two samples exceed the BIS of $300 \mu\text{g}/\text{L}$ for manganese. DOC
288 concentrations are all below $3 \text{ mg}/\text{L}$, but are higher in the shallow Holocene aquifer.



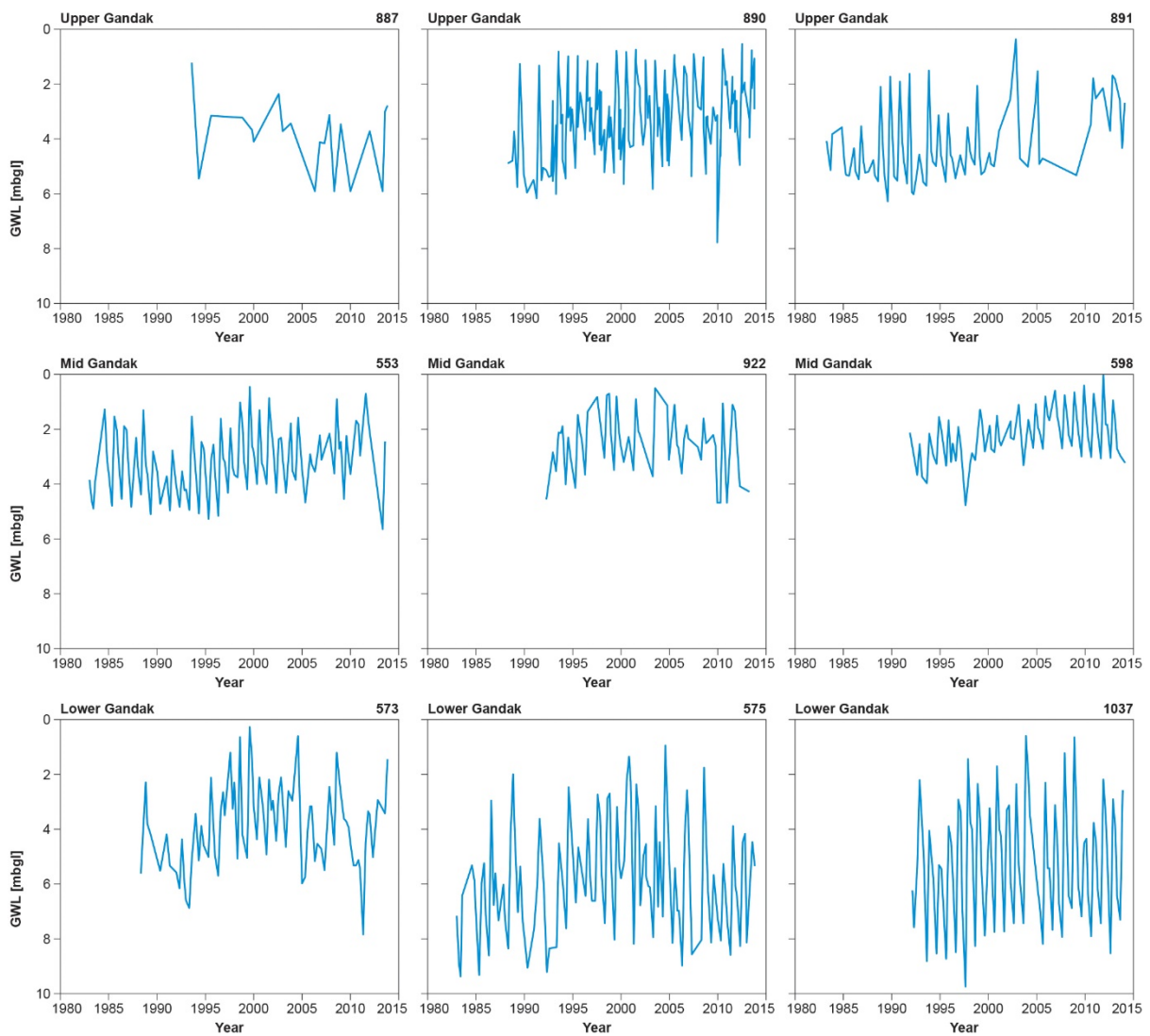
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290 **Figure 5.** Inorganic hydrochemistry depth profiles. a) $\delta^{18}\text{O}$ VSMOW (‰), b) NO_3 (mg/L), c) DOC
 291 (mg/L), d) HCO_3 (mg/L), e) Fe ($\mu\text{g/L}$), f) As ($\mu\text{g/L}$), g) U ($\mu\text{g/L}$), h) Mn ($\mu\text{g/L}$), i) F (mg/L). e-i: India BIS
 292 drinking water guideline value shown by solid vertical line, WHO drinking water guideline value is
 293 shown by the dashed vertical line (WHO 2017).

294 **3.4 Groundwater level monitoring**

295 Long-term (>20 years) groundwater hydrographs for nine monitoring sites (three sites in each of the
 296 upper, mid and lower catchment zones) show shallow groundwater levels (< 10 m bgl) across the
 297 Gandak catchment (Figure 6). A distinct unimodal monsoon signal is seen at all sites, with no
 298 evidence of long-term changes in groundwater levels at any of the monitoring sites. The amplitude
 299 of seasonal groundwater level changes are comparable for the upper and mid catchment zones
 300 (typically <4 m), but larger for the lower catchment zone (typically >6 m). Groundwater levels are
 301 also deeper in the lower zone than the upper and mid zones: mean and standard deviations for

302 monitored groundwater levels for the upper, mid and lower catchment zones were 3.5 ± 1.5 , 2.6 ± 1.1
 303 and 5 ± 2.1 m bgl respectively.



304

305 **Figure 6.** Long-term groundwater level records from the Gandak catchment (1983-2013), source
 306 MacDonald et al (2016). Borehole IDs 598 and 1037 are from the west of the River Gandak; all other
 307 hydrographs are from the east of the River Gandak. For borehole locations refer to Figure 2.

308 The groundwater levels time-series for the 10 shallow (<20 m) piezometers installed along the three
 309 transects are shown in Figure 7. Time-series of daily rainfall are also shown, which on average
 310 declines from north to south; mean rainfall was 5.2, 3.4 and 2.3 mm/d, for the upper mid and lower
 311 catchments, respectively, over this 22-month period. As would be expected, all groundwater levels

312 responded to seasonal monsoon rainfall, peaking between August and September in 2017 and 2018.

313 However, other responses can also be identified which are described below.

314 Considering the mid-catchment first, the groundwater level in (M4) located 480 m to the east of the

315 Eastern Main Canal rose (by 90 cm) and then fell between January and April 2019. Given that there

316 was no rainfall during this period, we attribute this to recharge from the canal. The level in

317 piezometer M3, 780 m to the west of the Eastern Main Canal, between March and April 2018, was

318 similar to that of M4, and therefore also appears to be influenced by canal recharge. The recession

319 curves for piezometers M1 and M2, which are between 4.8 and 10 km from the Eastern Main Canal,

320 are smoother than those near to the canal (M3, M4).

321 In the upper catchment, the pattern of groundwater level fluctuations in piezometers U2 and U3

322 between mid-March and the end of April 2018 was similar to that in boreholes M3 and M4 (Figure

323 7). Also, groundwater levels rise in December 2018 – January 2019, indicating that these

324 piezometers are also responding to canal recharge or seepage. Piezometer U1 is close to the River

325 Gandak but further from the canal; outside of the monsoon season, the groundwater levels fall

326 rapidly and there is little evidence of a link to the canal network. This is likely due to the fact that the

327 majority of river water is being diverted into the canal network outside of the monsoon season

328 (Neupane et al., 2010; Choudhary et al 2012; Dixit and Shukla 2017).

329 In the lower catchment, the two groundwater level hydrographs follow a similar pattern to each

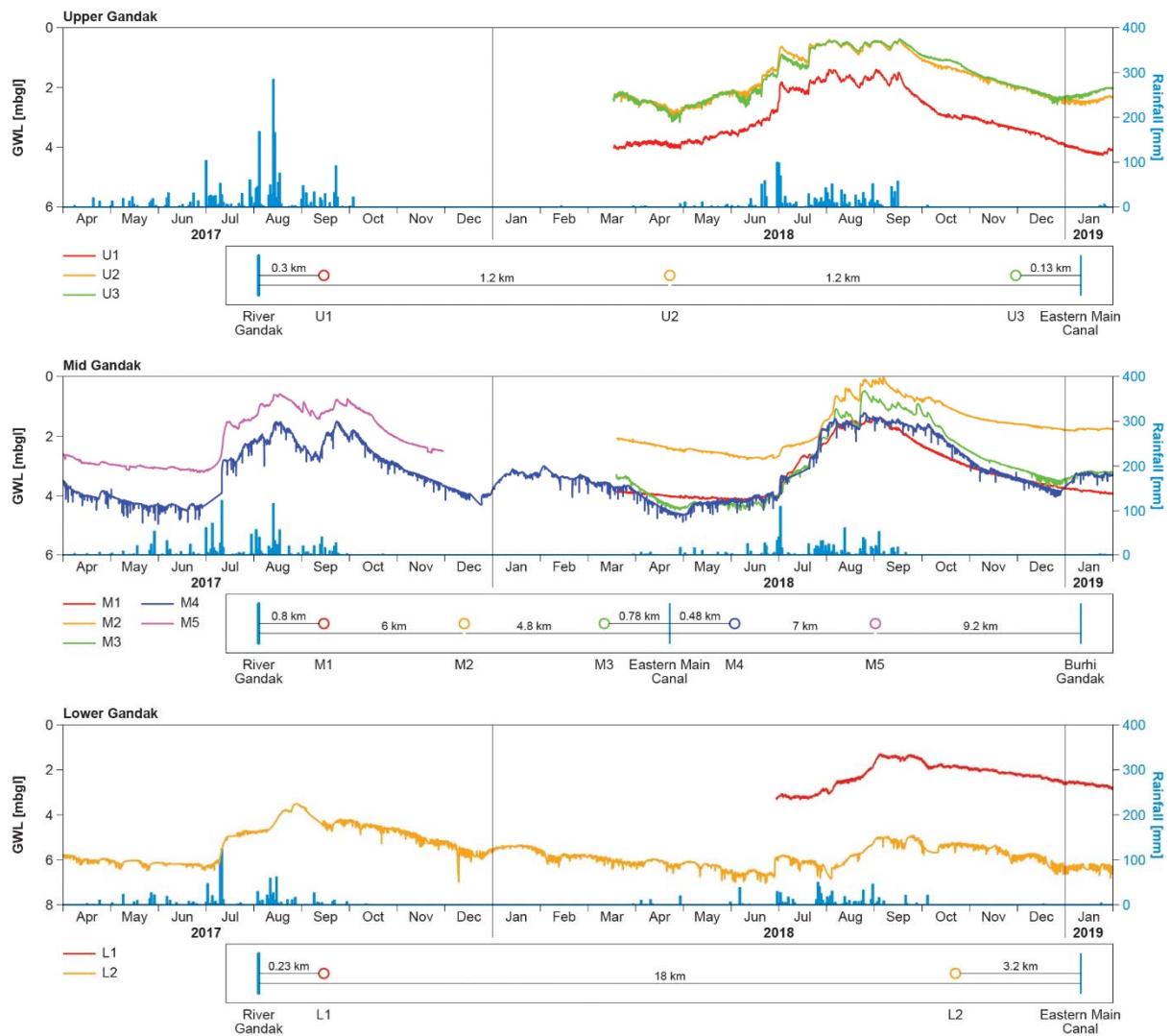
330 other, with relatively linear recession curves. Recessions continue almost unbroken between the

331 monsoons apart from a rise in the groundwater level in borehole L2, occurring over approximately

332 one month, started just before the end of 2018. This is coincident with the rise in borehole M4, and

333 may also be indicative of a response due to canal recharge, though this is again uncertain.

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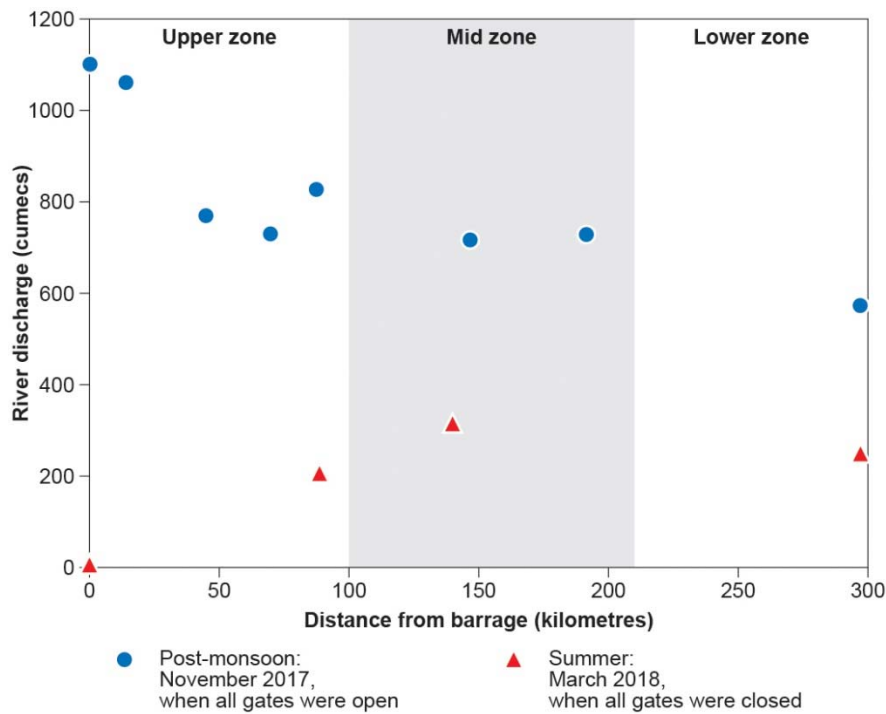


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336 **Figure 7.** Transects of high frequency groundwater observations (2017) close to the eastern Gandak
 337 canal in the upper, mid and lower catchment zones.

338 **3.5. Seasonal changes in river discharge**

339 River discharge decreased , at $1.85 \text{ m}^3 \cdot \text{s}^{-1} \cdot \text{km}^{-1}$ with downstream distance from the barrage in the
 340 post-monsoon period (November 2017), when the barrage gates were open, indicating that the river
 341 was recharging the aquifer at this time. In March 2018 when all barrage gates were closed, river
 342 discharge gradually increased at approximately $0.75 \text{ m}^3 \cdot \text{s}^{-1} \cdot \text{km}^{-1}$ moving downstream from zone 1 to
 343 2, beyond which it was approximately stable (Figure 8). This indicated contribution of groundwater
 344 from the aquifer to the river.



345

346

Figure 8. River flows downstream of Triveni barrage, during the post-monsoon (blue points:

347

November 2017, when all gates were open) and during summer (red triangles: March 2018 when all

348

gates were closed).

349

4. Discussion

350

4.1 Evolution in groundwater recharge and discharge across the Gandak catchment

351

Evidence from stable isotope and residence time tracers can delineate changes in groundwater

352

recharge sources across the Gandak catchment, particularly with reference to surface water and

353

rainfall stable isotopes signatures. As with other studies (*e.g.* Lapworth et al 2015; Joshi et al., 2018)

354

distinct isotopic signatures can often be used to differentiate between surface water and diffuse

355

meteoric recharge sources in groundwater. Rainfall samples collected as part of this study show

356

close agreement with data from Lapworth et al. (2015) (Figure 3a), despite being some 1000 km

357

apart, and delineate a consistent MWL across this region of northern India (*e.g.* Krishan et al., 2014).

358

The surface water (canal and River Gandak) samples in this study are significantly depleted with

359

respect to the amount weighted rainfall value for the catchment, indicating that they originate from

360 a higher elevation (*i.e.* the Himalayas). Canal water cannot be distinguished isotopically from the
361 river water due to their common source in the River Gandak at the Gandak barrage (Figure 3b). The
362 majority of groundwater samples have water stable isotope signatures that are comparable with
363 local meteoric rainfall (Figure 3). Overall, the isotopic evidence suggests that rainfall recharge
364 dominates across the Gandak. However, there is isotopic evidence for surface water recharge
365 sources in shallow groundwater samples mainly from the upper and middle of the catchment. This is
366 corroborated by the piezometer hydrographs, which show some, but limited, groundwater
367 interaction with canals and rivers (Figure 7), and is consistent with the river discharge measurements
368 post monsoon which indicate the river loses water downstream (Figure 8).

369 Baseflow to the River Gandak as the dry season progresses is shown by an increase in river discharge
370 downstream in zones 2 and 3. With no other inflow to the Gandak or from rainfall, the increase in
371 discharge downstream must be from groundwater. The baseflow to the river thus has a vital role in
372 sustaining habitat for endangered gharial crocodiles (*Gavialis gangeticus*), Ganges river dolphins
373 (*Platanista gangetica*), fresh-water turtles, fish, and other river fauna (Choudhary et al. 2012, Sinha
374 2018). River discharge measurements do not indicate significant baseflow in the lower catchment in
375 zone 3. Previous modelling studies of the lower catchment confirm that the alluvial aquifer adjacent
376 to the river is recharged by the River Gandak between May-October and January-February, but also
377 suggest that the process is reversed during other months (Singh et al., 2018). Overall, this zone
378 represents a system less affected by recharge from surface water due to intermittent canal flow. The
379 deeper groundwater levels recorded in this zone (Figure 6, 7) could help account for the reduction in
380 baseflow and will increase unsaturated zone travel times during recharge. This is reflected in the
381 overall lower residence time tracer concentrations in shallow groundwaters in the mid and lower
382 catchment (Figure 4).

383 **4.2 Changes in salinity across the catchment**

384 Most groundwater samples had higher SEC (typically > 500 $\mu\text{S}/\text{cm}$) compared to surface waters
385 (<500 $\mu\text{S}/\text{cm}$). Salinity depth profiles (Figure 4b) show highest variability in very shallow (0-20 m bgl)
386 groundwater samples. In contrast, shallow (0-50 m bgl) SEC concentrations were less than those
387 recorded in Punjab in northwest India, which is likely to be due to a combination of higher
388 evaporation rates, more intensive groundwater pumping regimes, and potentially also greater
389 historical use of fertilisers in Northwest India compared to the Gandak (Lapworth et al., 2017; Foster
390 et al., 2018).

391 Compared to other arid settings, such as northwest India and Pakistan, the higher precipitation,
392 lower evaporation and supply of low SEC water from surface water sources (200-300 $\mu\text{S}/\text{cm}$) in the
393 Gandak catchment is currently adequate to maintain relatively low SEC conditions (500-1000 $\mu\text{S}/\text{cm}$)
394 in groundwater system. Apart from a few isolated samples where there is evidence of local higher
395 salinity surface water bodies (e.g. wetlands and oxbow lakes) impacting on groundwater, there is
396 compelling evidence that canal leakage in the Gandak catchment results in lower SEC groundwaters
397 close to canals compared to more distal areas where the influence of canals is limited. However,
398 there is evidence of soil sodicity and salinity problems in the upper Gandak as a result of
399 waterlogging (Singh and Khan 2002).

400 The differences in SEC depth profiles across the catchment (Figure 4) are interpreted as greater
401 contributions of low residence time, low SEC, water from surface water sources in the upper and
402 middle catchment compared to the lower catchment. Furthermore, the overall low SEC found across
403 the whole Gandak catchment (typically <1000 $\mu\text{S}/\text{cm}$) contrasts with the high SEC observed with
404 other parts of the IGB where canal irrigation is widespread but which have significantly more arid
405 conditions than the Gandak (MacDonald et al., 2016).

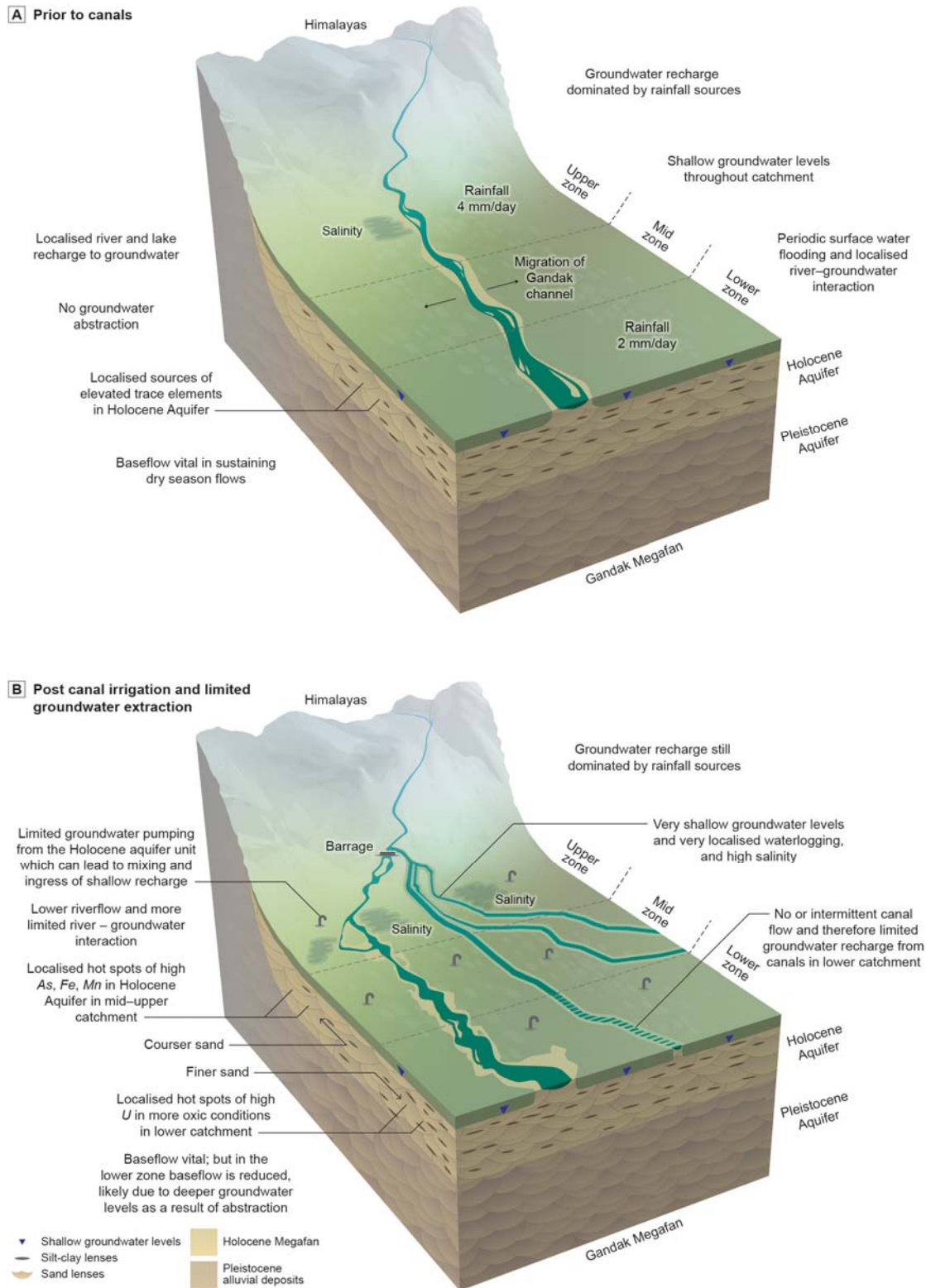
406 ***4.3 Connection between shallow Holocene aquifer and deeper Pleistocene aquifer***

407 The tracer evidence (e.g. Figure 4 and 5) shows that the contribution of surface water (river or canal)
408 sources to shallow groundwater recharge is largely restricted to the shallow Holocene aquifer (0-40

409 m bgl). At greater depths, in the Pleistocene aquifer, groundwater is largely consistent with more
410 diffuse meteoric rainfall recharge, and therefore appears to be less connected hydraulically to
411 surface water recharge sources. This is also observed in many other parts of the IGB (e.g. Joshi et al.,
412 2018, Lapworth et al., 2015) and is consistent with the idea that the River Gandak was probably not
413 the major source of groundwater recharge during the Holocene. The fact that this older meteoric
414 rainfall signature in the shallow Holocene aquifer has not been overprinted in the lower catchment
415 zone is likely to be due to relatively low groundwater abstraction as well as the limited impact of
416 recharge from canals in this zone compared to the upper and mid catchment.

417 ***4.4 Conceptual Model for the Gandak River in India***

418 In summary, three lines of hydrochemical evidence (stable isotopes, residence time indicators and
419 salinity) and the observations of groundwater level and river discharge from upstream to
420 downstream confirm the hypothesis that canal leakage and river sources recharge the alluvial
421 aquifer system, but this is largely restricted to the upper Holocene aquifer, and that the influence of
422 canal leakage as a recharge source reduces down the catchment where canal flows are lower. The
423 observations of higher residence time tracer concentrations, lower SEC, and more depleted water
424 stable isotope signatures in shallow groundwater in the upper catchment compared to the mid and
425 lower catchment (Figures 3-5), and higher SEC in both shallow and deeper aquifer layers in the lower
426 catchment, all confirms this hypothesis. However, overall the isotopic evidence (Figure 3b) shows
427 that meteoric rainfall sources are the dominant source of recharge across the catchment, even in the
428 upper parts of the catchment where canal flows are higher. Dry-season groundwater baseflow helps
429 to sustain ecological flow regimes needed for endangered species, but is reducing in the lower
430 section of the Gandak (zone 3). Figure 9 summarises key results from this study and shows a
431 schematic conceptual model highlighting the dominance of local meteoric recharge sources across
432 the Gandak catchment and the evolution of recharge processes locally within the Gandak catchment
433 due to the influence of canals and more recent groundwater pumping.



438 **4.5 Groundwater quality constraints and controls**

439 Elevated arsenic, iron and manganese, above WHO drinking water guideline values, were found for a
440 small proportion of groundwaters within the Gandak catchment (Figure 5). These exceedances were
441 almost exclusively for groundwaters in the Holocene aquifer (<40 m bgl) with one occurrence of high
442 iron concentration (>300 µg/L) in the deeper Pleistocene aquifer (Figure 5e). Exceedances for arsenic
443 (n=3, 20%) were below the BIS of 50 µg/L in all cases; for manganese (n=2, 13%), the maximum
444 concentration was 680 µg/L; and for iron (n=6, 40%), the maximum concentration was 2225 µg/L.
445 These exceedances were only found in the upper and mid catchment, and are linked to reducing
446 conditions in the shallow aquifer, which facilitate the reductive dissolution of iron oxyhydroxides and
447 the release of bound arsenic (e.g. Nickson et al., 2000; BGS/DHPE 2001; Richards et al., 2020).
448 Shallow DOC concentrations were comparable for all three catchment zones (range 0.5-2.5 mg/L),
449 but lower concentrations (c.0.5 mg/L) are found at depth in the Pleistocene deposits across the
450 catchment.

451 In contrast, uranium concentrations approaching the WHO limit of 30 µg/L were only found for
452 groundwaters in the lower catchment, with a maximum concentration of 29.5 µg/L, not exceeding
453 the WHO guideline value of 30 µg/L (Figure 5). The deeper groundwater levels and higher amplitude
454 of seasonal groundwater level fluctuations in the lower catchment are likely to facilitate more oxic
455 groundwater conditions compared to the upper and mid catchment, where groundwater levels are
456 typically shallower with smaller seasonal fluctuations. The presence of geogenic sources of U (e.g. U
457 minerals and iron oxide coatings on aquifer grains) is likely to be similar across the catchment.
458 However, the prevailing oxic and alkaline conditions and carbonate complexation with U (VI) in the
459 lower catchment zone can facilitate U mobilisation (Barnett et al., 2002; Kumar et al., 2011), and
460 results in significantly higher U concentrations in the lower catchment compared to the other two
461 catchment zones. These concentrations are comparable with concentrations found by Richards et al
462 (2020), but are not as high as those detected by Lapworth et al (2017) and Coyte et al (2019) in

463 north-west India, but aquifer redox conditions are comparable and concentrations are approaching
464 the WHO guideline value of 30 µg/L.

465 Outside the Gandak catchment, high arsenic groundwater concentrations (> 50 µg/L) have been
466 extensively reported in the mid Gangetic Basin, largely restricted to districts along the main channel
467 of the River Ganga (Chakraborti et al., 2003, 2016; Saha 2008, 2009; Saha et al., 2010; Singh 2015;
468 Kumar et al., 2016; Richards et al., 2020). Reductive dissolution mechanisms are inferred (Nickson et
469 al., 2000, BGS/DPHE, 2001) from As and Fe correlations, although there is almost no reporting of
470 DOC groundwater concentrations from this region.

471 Our results for spatial and depth As and U distributions within the Gandak corroborate recently
472 published results by Richards et al (2020). Singh (2015) reports parts of three districts within the
473 Gandak (West Champaran in the mid Gandak catchment zone and Saran and Vaishali in the lower
474 catchment) as having arsenic concentrations > 50 µg/L, in each case at locations close to the River
475 Gandak. Neither the results from the current study nor the majority of previously published studies,
476 with the notable exception of a study by Jangle et al. (2016) for the Vaishali district, suggest that
477 there is a widespread issue of high arsenic in the shallow Holocene aquifer across large parts of the
478 Gandak catchment. Rather, the current evidence suggests that groundwater with elevated arsenic
479 concentrations in the Gandak may be a localised phenomenon, linked to isolated hot-spots where As
480 release from sediments within the Holocene aquifer is facilitated through reductive dissolution.

481 Nitrate concentrations were found to be low throughout the study catchment, most likely driven by
482 the prevailing low oxygen groundwater conditions in the upper and mid catchment, which are likely
483 due to the shallow groundwater levels and a limited supply of shallow sources of nitrate during
484 recharge (e.g. Seitzinger et al., 2006). However, very low nitrate concentrations are also found in the
485 lower catchment, where unsaturated zones are somewhat deeper, suggesting that contamination
486 from the use of N fertilisers is less prevalent across the Gandak compared to other regions of the IGB
487 (Agrawal et al., 1999; Lapworth et al., 2017). Groundwater salinity values found in the Gandak were

488 low overall (typically 500-1000 $\mu\text{S}/\text{cm}$), with only isolated higher SEC due to recharge from wetlands
489 impacted by evaporation. This reflects the relatively humid climate of the Gandak compared to
490 elsewhere in the IGB, where shallow groundwater levels and high rates of evapotranspiration can
491 lead to increasing groundwater salinity, which can constrain water use (MacDonald et al., 2016).

492 ***4.6 Potential threats to groundwater resources and endangered species***

493 Low groundwater abstraction, return flows, seepage from canals and high meteoric recharge in
494 Gandak have sustained a high-quality groundwater resource throughout much of the Gandak
495 catchment where it overlies the IGB aquifer within India. As a consequence, groundwater baseflow
496 contributes to sustaining habitat for the endangered species in spite of flow regulation. Increased
497 groundwater abstraction may be possible however this should proceed cautiously as recharge
498 depends mostly on the reliability of the monsoon, and declining groundwater levels will impact
499 environmental flows within the River Gandak. Increased abstraction may also lead to zones where
500 the water quality is degraded, due to ingress of surface contaminants or mobilisation of in-situ
501 contaminants and this would need to be monitored. Current groundwater salinity values across the
502 catchment suggest that the water is suitable for irrigation. However, shallow groundwater levels are
503 leading to localised salinity in groundwater and soil through phreatic salinization. Greater
504 abstraction in the more waterlogged sections of the upper Gandak may be beneficial for reducing
505 the impact of soil sodicity. However, increase groundwater use for irrigation coupled with reducing
506 rainfall and surface water irrigation increase the likelihood of increases in groundwater salinity due
507 to irrigation returns (Foster et al., 2018).

508 Increased abstraction from the deeper Pleistocene sediments is likely to lead to a mixing of shallow
509 and deep groundwater due to the sediment architecture of the fan deposits and the lack of
510 extensive low K sequences to block downward migration of groundwater. Therefore, should
511 pumping from the deeper parts of the aquifer system increase significantly, contaminants may
512 migrate to depth within the aquifer system.

513 Groundwater recharge is dominated by meteoric sources, so the largest direct threat to
514 groundwater in the Gandak is the changing pattern of the Indian monsoon, which may impact on
515 local rainfall and recharge. Long-term climate records show highly variability in monsoon rainfall
516 totals (Lacombe and McCartney 2014) and future climate changes are highly uncertain. For
517 example, Warwade et al. (2018) report between 15-20% coefficient of variation between 1901-2002
518 for monsoon and annual rainfall across Bihar. Additionally, studies have noted increasing glacial melt
519 water inflows into the River Gandak, as for many north-south flowing tributaries of the Ganga
520 (Anand et al. 2018). However, while there have been historically no significant downward trends for
521 rainfall totals for the lower Gandak catchment, parts of the Himalayas to the north of the study area
522 show significant long-term reductions in precipitation (Lacombe and McCartney 2014), which may
523 impact on surface water flows reaching the lower Gandak catchment. Increased intra-seasonal
524 variability of the Indian monsoon has also been reported (Befort et al., 2016; Vinnarasi and Dhanya
525 2016). There is an indirect threat from increased abstraction, although direct recharge from
526 irrigation canals is currently only important locally. If river flows and canal flows reduce, an increase
527 in groundwater abstraction to meet demand would be anticipated.

528 Groundwater abstraction for irrigation across the Gandak catchment is largely restricted to the
529 shallow Holocene aquifer. It is likely that this, combined with adequate recharge from canal and
530 river water, is hydraulically restricting the downward movement of shallow recharge to greater
531 depths within the Pleistocene aquifer, and maintaining groundwater levels. This contrasts with
532 observations in many other parts of the IGB (Mukherjee et al 2015; Lapworth et al., 2015;
533 MacDonald et al., 2016, Bhanja et al 2017; van Dijk et al., 2019). In addition, the availability of canal
534 flows for irrigation for much of the year in the upper and mid Gandak catchment zones may
535 continue to reduce the need for shallow groundwater abstraction for irrigation in these zones.
536 Shallow groundwater levels and smaller amplitude of seasonal groundwater level fluctuations in the
537 upper and mid zones relative to the lower catchment (Figure 6) reflect this. By contrast, in the distal
538 lower catchment the lack of canal flows for most months of the year could lead not only to reduced

539 recharge from canal returns, but also to increased demand for groundwater abstraction for
540 irrigation. The lower groundwater levels in this zone (Figure 6) likely reflect the lower rainfall here
541 than in the upper catchment and lower recharge from surface water sources, particularly canals; and
542 perhaps a greater reliance on groundwater for irrigation.

543 There are no significant downward trends in groundwater levels in the Gandak catchment (Figure 6),
544 indicating that there is adequate replenishment to the shallow groundwater system to sustain
545 current levels of groundwater abstraction. Residence time tracer results indicate that in all parts of
546 the catchment there is modern recharge to the Holocene aquifer, sourced from meteoric rainfall
547 across the catchment and additionally from canal water returns in the upper and mid catchment
548 zones, which is sufficient to sustain current groundwater demands. This is in stark contrast with the
549 situation in other parts of the basin in Northwest India and Northern Pakistan (Lapworth et al., 2015;
550 MacDonald et al 2016), where considerably higher historical rates of groundwater abstraction and
551 lower average rainfall occur. Another important factor is the current more limited use of deeper
552 groundwater resources in the Gandak compared to other parts of the IGB (Bonsor et al., 2017).

553 Conjunctive use of groundwater and surface-water irrigation has been a recurring emphasis of water
554 management policymaking in the IGB aquifer (National Institute of Hydrology, 2000, Khan et al.
555 2014) due to problems of canal irrigation, especially with waterlogging and inefficient/erratic water
556 supply. For the Gandak this could be encouraging farmers to shift towards more groundwater use in
557 farm-level irrigation. However, although there is capacity for increased abstraction, groundwater
558 and river levels and quality will need to be carefully monitored along with changes in precipitation to
559 minimise degradation of the groundwater resource and the fragile river ecosystem which depend on
560 baseflow. This may, over time, accelerate the use of groundwater resources in this region in the
561 future, as has been observed elsewhere across the IGB, which may put further pressure on
562 groundwater resources and river ecology.

563

564 **5. Conclusions and future outlook**

565 Quantifying groundwater recharge sources and processes in the mid IGB and understanding ground-
566 surface water interactions is fundamental to effective management of water resources for people
567 and nature, and for forecasting future changes. This study contributes to ongoing debate over
568 groundwater status across the IGB and the different anthropogenic impacts on groundwater
569 recharge and differences in groundwater level responses observed across the IGB. This is the first
570 assessment of groundwater residence times in the Gandak alluvial aquifer, and the first study to
571 integrate a range of environmental tracers and groundwater level monitoring and spatially explicit
572 river discharge measurements to conceptualise groundwater recharge process in the mid IGB. When
573 compared with findings from other recent studies in the IGB this study highlights the heterogeneity
574 of groundwater recharge process and groundwater level responses across the IGB. Key conclusions
575 for the Gandak catchment are:

- 576 • Rainfall recharge sources dominate across the catchment, but stable isotope and
577 residence time data show that groundwater recharge from canal leakage has been
578 happening for decades, but is confined to the shallow aquifer (<40 m bgl). The shallow
579 aquifer contains tracer concentrations indicative of active modern recharge.
- 580 • Groundwater baseflow to the River Gandak during the dry season plays a vital role in
581 sustaining river flows, habitats, and vital aquatic ecosystems, especially for endangered
582 species such as the gharial crocodile (*Gavialis gangeticus*) and Ganges river dolphins
583 (*Platanista gangetica*) besides fresh-water turtles and fisheries.
- 584 • There is little evidence that the River Gandak had a major role in recharging the
585 groundwater system throughout the Holocene aquifer system; while this process may be
586 important locally, the dominant recharge mechanism appears to be rainfall.
- 587 • In the Holocene part of the aquifer system, groundwater chemistry is highly variable,
588 especially for DOC, As, Mn, NO₃, U and other redox sensitive constituents

- 589
- There is some localised contamination from shallow geogenic sources of As in the
590 Holocene aquifer, exceeding 10 µg/L in a few cases, but no samples in this study
591 exceeded the BIS guideline value of 50 µg/L. Some groundwater samples showed
592 elevated U (max = 29.5 µg/L), approaching the WHO drinking water guideline value, all
593 from the lower catchment, where groundwater levels are generally lower.
 - There is no evidence of widespread salinity build up in the areas served by irrigation
594 canals, unlike in other more arid settings in the IGB where this is a widespread issue due
595 to high evapotranspiration rates and shallow groundwater levels. Although there is
596 some evidence of higher salinity in the lower Gandak catchment, where canal flows are
597 less reliable, by contrast there is also evidence that canal water returns may reduce
598 salinity in parts of the upper catchment.
 - In the deeper Pleistocene aquifer system (>40 m bgl), groundwater chemistry shows less
600 variation and stable isotope signatures are consistent with recharge from meteoric
601 rainfall only, not canal or river water. Residence time tracers and stable isotope results
602 indicate a very low proportion of modern recharge entering this system due to the fact
603 there is currently limited groundwater abstraction from the deeper part of the aquifer.
 - Our study shows that in the upper and mid catchment there is evidence of surface water
605 recharge to the Holocene aquifer, but further away from the main canals in all
606 catchment zones the influence of the canals reduces. There is less impact on
607 groundwater recharge from canals in the lower catchment.
- 608

609

610

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