

28 The movement of juveniles into deeper waters post-metamorphosis were further confirmed
29 by significant shifts in the element:⁴²Ca ratios (for elements ²⁴Mg, ⁵⁵Mn, ⁸⁸Sr and ¹³⁸Ba).
30 Elemental ‘fingerprints’ for each of the three otolith zones were further found to be distinct
31 through high classification accuracies of the random Forest model ($81.1 \pm 0.1\%$). Holistic
32 approach investigating both otolith microstructure and elemental composition provides
33 further evidence that ontogeny could be a confounding factor in natal (geographic) origin
34 studies if (1) otolith ‘core’ samples are taken from different ontogenetic otolith zones and (2)
35 if elements predominantly controlled by physiology are included in the respective
36 classification analyses.

37 **1. Introduction**

38 The southern blue whiting, *Micromesistius australis* (Norman, 1937), is a mesopelagic
39 schooling fish that occupies waters around New Zealand and the tip of South America
40 (Patagonia). The Patagonian sub-species *Micromesistius australis australis* once contributed
41 to the largest fishery in the South Atlantic. However, following decades of exploitation and
42 concerns over fluctuating landings the spawning stock biomass of *M. a. australis* is currently
43 considered to be critically low (Laptikhovsky et al., 2013). The fishery has mainly targeted
44 spawning aggregations of *M. a. australis* to the south-southwest of the Falkland Islands
45 where spawning occurs between August – October (Macchi et al., 2005; Macchi and Pajaro,
46 1999; Pajaro and Macchi, 2001). Both egg and newly-hatched larval stages have previously
47 been described for *M. a. australis* (Laptikhovsky and Brickle, 2009; Weiss, 1974), as well as
48 embryogenesis (Laptikhovsky and Brickle, 2009). However, very little is known about
49 subsequent stages of early ontogeny including the timing of metamorphosis and the juvenile
50 stage.

51 Metamorphosis occurs when the larvae undergoes dramatic morphological and physiological
52 changes during its development into a juvenile fish (Christensen and Korsgaard, 1999; Sæle
53 et al., 2004). For many mesopelagic and benthic species the end of metamorphosis is marked
54 by the transition from the pelagic larval habitat to the juvenile habitat in deeper waters, or
55 settlement to demersal or benthic habitats (Biagi et al., 1998). Similar changes were
56 previously observed in North Sea plaice (*Pleuronectes platessa*) by hatching larvae in
57 laboratory conditions and monitoring their development and behaviour (Modin et al., 1996;
58 Morales-Nin et al., 2005). In many gadoids the process of metamorphosis results in
59 significant change to the otolith microstructure through the formation of accessory growth
60 centres (AGCs) (Morales-Nin, 2000; Morales-Nin et al., 2005). As growth increments form
61 and expand from the AGCs, they ultimately amalgamate to form a visible ‘prism’ around the
62 otolith core (Morales-Nin et al., 2005). Combined observations of larval development and
63 otolith growth revealed that the formation of the first and last formed AGC are associated
64 with initiation and completion of metamorphosis respectively (Modin et al., 1996; Morales-
65 Nin et al., 2005). The change in habitat between the larval and juvenile stages is visible
66 within the otolith microstructure by means of a check mark resulting from a change in diet
67 related to the change in habitat (Brothers, Edward et al., 1982; Victor, 1982). Although *M. a.*
68 *australis* does not undergo settlement to the benthic habitat, a ‘settlement mark’ has
69 previously been reported in *M. a. australis* otoliths (Barrera-Oro and Tomo, 1988). This mark
70 in *M. a. australis* possibly corresponds to the transition of an individual from the pelagic
71 larval habitat to a deeper, juvenile habitat. The timing at which these features are formed can
72 be estimated through the usage of daily growth increment counts observed within the otolith
73 microstructure (Arai et al., 2000; Buratti and Santos, 2010).

74 Besides structural changes, the otolith microchemistry is also thought to change throughout
75 the early stages of ontogeny (Martin and Thorrold, 2005). For example, Brophy et al. (2004)

76 determined a Mn-rich core in Atlantic herring (*Clupea harengus*) otoliths. Additionally, Arai
77 et al. (2000) reported significant shifts in $^{88}\text{Sr}:^{42}\text{Ca}$ ratios coinciding with the initiation and
78 completion of metamorphosis. Longmore et al. (2011) further observed significant
79 differences in various element: ^{42}Ca ratios between samples taken from parts of the otolith
80 corresponding to different stages of ontogeny. Multi-elemental signatures, or elemental
81 ‘fingerprints’, are frequently used to classify between geographical locations to infer the
82 stock structure or to reconstruct migration patterns of fishes (Campana, 1999; Campana and
83 Thorrold, 2001). These studies assume differential rates of elemental uptake into the otolith
84 microstructure based on the physical or chemical properties of the contemporaneous ambient
85 environmental conditions which vary between geographical regions (Campana and Thorrold,
86 2001; Elsdon et al., 2008). The elemental shifts found in relation to ontogeny could confound
87 the interpretations of these geographical origin studies. Macdonald et al. (2008), for example,
88 revealed increased classification accuracies in the natal origin of three geographically isolated
89 groups of Australian smelt (*Retropinna semoni*) when excluding the ^{55}Mn -rich otolith core
90 (corresponding to the first period of life). Therefore, further elucidating on the relationship
91 between ontogeny and otolith microchemistry is still required.

92 Here, we aim to estimate the timing and duration of the early stages of ontogeny in *M. a.*
93 *australis* using the otolith microstructure and to investigate if the elemental ‘fingerprints’
94 corresponding to these putative stages are (1) distinct and (2) consistent with interpretations
95 of a transition to a new life history stage.

96 **2. Materials and Methods**

97 *2.1 Sample collection*

98 Juvenile *M. a. australis* were sampled within the Falkland Interim Conservation and
99 Management Zone (FICZ) between January and November 2017 (Figure 1).

100 *INSERT FIGURE 1*

101 Samples from January and February were collected during research cruises conducted by the
102 Falkland Islands Fisheries Department (FIFD) aboard *RV Castelo* (ZDLT1), whereas samples
103 from March to November were collected by FIFD fisheries observers from commercial
104 trawls. Research and commercial trawls were conducted using bottom trawl nets with
105 respective mesh sizes of 90 mm (including a 10-15 mm cod end liner) and 110 mm (no liner).
106 Trawl depth ranged between 69 – 275 m and 120 - 287 m for research and commercial trawls
107 respectively. Samples from the research cruises were frozen on board and later processed in
108 the laboratory at the FIFD. Samples from the commercial trawls were processed at sea. Total
109 length (cm) and total weight (g) were recorded for all samples, and otoliths were removed
110 and stored in 95% ethanol. All individuals were considered juveniles (FIFD maturity scale) as
111 gonads were too small to be classified correctly, with length-frequency analyses used to
112 ensure samples belonged to the same cohort before further processing.

113 *2.2 Sample preparation and interpretation*

114 In the laboratory otoliths were mounted on slides and ground down manually on both sides to
115 reveal the nucleus and daily growth increments following methods by Arkhipkin and
116 Shcherbich (2012). However, a modification was required as the processing time of larger
117 otoliths became impractical. For these samples initial grinding was performed using silicon
118 carbide paper at differing coarse grades, e.g. P180 followed by P320/400, to reduce the
119 thickness of the sample before polishing with P600 and P2000. It also became increasingly
120 difficult to reveal clear growth increments from the nucleus to the otolith margin in the larger
121 otoliths. Too thick a section left the primordium unclear, too thin a section and the otolith
122 margin became unreadable (as described in Arkhipkin and Shcherbich (2012)). As a
123 precautionary approach, counts within the primordium of larger otoliths were extrapolated. If

124 a preparation was uninterpretable, the second otolith was processed. In cases where
125 interpretation of growth increments along a single radius was impossible, prominent growth
126 increments or microstructural features were used to change the reading orientation (Morales-
127 Nin and Aldebert, 1997) (Figure 2). Otoliths were initially read at 200x magnification
128 (Olympus BX51) using transmitted light, but this was increased to 400x magnification
129 towards the margin of larger otoliths where increments became narrower.

130 *2.3 Timing of early life history stages*

131 Initial increment counts were conducted from the nucleus (N) to the margin for all samples
132 (Figure 2). The nucleus was hereby defined as the area of the otolith contained within the first
133 discontinuous unit (Dunkelberger et al., 1980).

134 *INSERT FIGURE 2*

135 These ‘total increment counts’ were plotted against date of capture to confirm daily formation
136 of the growth increments. Total increment counts were performed twice by Reader A (TB)
137 and a subsample of 5 otoliths were read by Reader B (AA) to confirm reading precision. The
138 otolith microstructure further revealed a number of features including a ‘prism’ shape with
139 accessory growth centres (AGCs), and one or two distinct hyaline growth zones (the latter
140 only visible in samples caught after July 2017). Increment counts were conducted up to the
141 first formed AGC (AGC1), the last formed AGC (AGC2) and the first hyaline zone (HZ1) to
142 estimate the age at which these features were formed. Three zones within the otolith sections
143 were finally defined by these features according to their corresponding early life history
144 stage: (A) larval (pre-metamorphosis) (N–AGC1), (B) metamorphosis (AGC1–AGC2) and
145 (C) juvenile (AGC2–margin) (Figure 3).

146 *2.4 Trace elemental analyses*

147 A subsample of otoliths was further used to investigate the elemental composition of the
148 otolith throughout the early life history stages of *M. a. australis* by means of laser-ablation
149 inductively coupled plasma mass-spectrometry (LA-ICP-MS) at the British Geological
150 Survey, Keyworth, Nottingham, UK. Samples were re-mounted onto slides (dimensions 40 x
151 28 mm), decontaminated using 1% HNO₃/0.5% HCl solution and subsequently rinsed with
152 Milli-Q before placement into the chamber of the laser ablation system (NewWave UP-193).
153 Reference glasses NIST-610 and NIST-612 were used as external standards. To account for
154 mechanical drift, reference materials were ablated before and after samples, and reference
155 material NIST-610 was further ablated between individuals. For each otolith a line of spots
156 (spot size 35 μm, distance between spots 50 μm, dwell time = 30 s, wash-out time = 30 s)
157 was analysed from the nucleus to the anterior margin (Figure 3).

158 *INSERT FIGURE 3*

159 Irradiance (GW/cm²) and fluence (J/cm²) were recorded for each run. Raw counts (cps) of
160 trace elements (¹³⁸Ba, ⁴²Ca, ⁷Li, ⁵⁵Mg, ²⁴Mn, ²³Na, ⁸⁸Sr) within the ablated material were
161 determined by ICP-MS (Agilent 7500c). Raw counts were converted to element
162 concentrations using Igor Pro 6.34 (Wavemetrics) and extension package Iolite v 2.5.
163 Elements for which counts were within detection limits (two SE of the background signal)
164 were removed from further analyses. Increment number (days after first increment formation)
165 and life history stage were determined for each spot. If a spot covered multiple growth
166 increments, the mean increment number was used.

167 2.5 Statistical analyses

168 All statistical analyses were performed in R V.3.5.1 (R Core Team, 2018). Data exploration
169 was conducted according to methods by Zuur et al. (2010) as summarised below.

170 2.5.1 Age validation

171 Average percent error (APE) (equation 1) was determined between total counts by Reader 1
172 (Comparison 1), and between the subsample read by Reader 1 and Reader 2 (Comparison 2)
173 to confirm reading accuracy using R package ‘FSA’ (Ogle et al., 2018).

$$174 \quad APE = \frac{\sum_{j=1}^n APE_j}{n} \quad \text{where} \quad APE_j = 100 \times \frac{\sum_{i=1}^R (x_{ij} - \bar{x}_j)}{R \bar{x}_j} \quad (1)$$

175 where APE_j is the average percent error for the j th fish, x_{ij} is the i th age estimate on the j th
176 fish, \bar{x}_j is the mean age estimate for the j th fish, R is the number of times that each fish was
177 aged, and n is the number of aged fish in the sample (Beamish and Fournier, 1981).
178 Systematic differences in age readings were determined using methods by Evans and Hoenig
179 (1998). Samples were removed from further analyses if readings 1 and 2 by Reader A
180 differed by >10%. Total number of growth increments were plotted against date of capture to
181 confirm daily formation of growth increments. The data were fitted with a linear regression
182 model of which the regression slope (m) was compared to 1 ($m = 1$ suggests: 1 growth
183 increment = 1 day in age) to validate interpretation of the number of growth increments as
184 age (days).

185 2.5.2 Age vs Length

186 The relationship between total length and age was visualised using a scatterplot. By
187 increasing model complexity and checking model residuals a Gaussian GAM (family =
188 quasipoisson), was produced for which the residual variation was homogenous and normally
189 distributed. The ‘predict’ function in the ‘Deriv’-package (Clausen and Sokol, 2018) was
190 finally used to estimate the length (\pm SE) at Age = 365 days (1 year).

191 2.5.3 Trace elemental analyses

192 Elemental concentrations of the elements ^{138}Ba , ^{55}Mg , ^{24}Mn , ^{23}Na , ^{88}Sr were expressed as a
193 ratio to ^{42}Ca . Elements ratios were used as explanatory variables with ‘Zone’ being the

194 response variable. Outliers in element:⁴²Ca were investigated. A distinction was made
195 between outliers that were considered singular and outliers that were part of a ‘peak’ in
196 element:⁴²Ca concentrations along individual profiles. Outliers applying to the former were
197 removed from further analyses. Collinearity was observed between ⁸⁸Sr and ²³Na (0.83), of
198 which ²³Na was removed upon inspection of the variance inflation factor values (VIF ²³Na =
199 3.54, VIF ⁸⁸Sr = 3.52). To account for the 2-dimensionality estimates of AGC1, AGC2, and
200 HZ1, spots occurring within two SE of their respective mean estimates were removed from
201 further analyses. Visualisation of the data included initial smoothed (method = LOESS, span
202 = 0.8) element:⁴²Ca profiles against age, followed by boxplots depicting element:⁴²Ca ratio
203 distribution against zone. Non-parametric Kruskal-Wallis tests were performed to compare
204 median element:⁴²Ca ratios between life history stages with Wilcoxon rank post-hoc tests
205 where pertinent. A random forest model was finally used to investigate multivariate
206 differences in elemental composition of the otolith microstructure for the different early life
207 history stages. The random forest model is a classification method which uses a subset
208 (training subset) of the data to build classification trees (Breiman, 2001). The remaining data
209 (prediction subset) is subsequently used to determine the prediction ability of those trees by
210 trying a set number of variables, or elements in this study, at each node of the trees. See
211 (Mercier et al., 2011) for more details on this model. Analyses were performed using R
212 package ‘randomForest’ (Liaw and Wiener, 2002). In this study, the training:prediction data
213 set ratio was 70:30, the number of classification trees computed was 500 and the number of
214 variables tried at each split was 4. This procedure was repeated 1000 times. Classification
215 results were converted to percentages (%), ranked, and the 50th, 500th and 950th observations
216 for each value reported. Finally, the order of element contribution to classification accuracy
217 was determined through a stepwise procedure using the Gini index (Breiman, 2001).

218 **3. Results**

219 *3.1 Increment interpretation and age validation*

220 100 individuals (TL range: 6 – 21 cm) were sampled between January and November 2017.
221 Total length was shown to increase with age (Figure 4). GAM regression parameters for
222 $s(\text{Age})$ were found to be significant ($edf = 4.43$, $F = 227.8$, $p = <0.001$) and the deviance
223 explained by the model was 92.9%. A reduction in growth rate (not calculated here) is visible
224 between 175-275 days, which corresponded to the austral winter period. Finally, TL at age 1
225 year (365 days) was estimated at 20.31 ± 0.48 cm.

226 *INSERT FIGURE 4*

227 GI counts were conducted using the methodology described in Section 2.2 (Figure 2). For 19
228 (22%) individuals the first otolith preparation was unreadable resulting in the second otolith
229 being utilised. APE for readings by Reader 1 (Comparison 1), and for the subsample read by
230 Reader 1 and Reader 2 (Comparison 2) were 1.7 and 1.1% respectively. No systematic
231 differences were further determined between readings ($p = 0.40$, $p = 0.17$ for Comparisons 1
232 and 2 respectively). Total increment counts ranged between 87-384 and increased with date
233 of capture (Figure 5).

234 *INSERT FIGURE 5*

235 The slope of the fitted linear model ($m = 0.96$) was compared against 1 (where 1 GI = 1 day)
236 and was found not to be significantly different (t-ratio = 0.571, $p = 0.5693$). Hereafter,
237 number of GIs was therefore interpreted as number of days. The linear model was further
238 used to estimate a mean date of first increment formation for the 2016-cohort as 03-11-2016.

239 *3.2 Timing of early life history stages*

240 Counts from the nucleus (N) to the first accessory growth centre (AGC1) revealed that the
241 initiation of metamorphosis occurred 26-34 days after first increment formation (mean $30.5 \pm$

242 0.3). In the larger otoliths for which reading between N - AGC1 was impossible, we therefore
243 extrapolated this count at 31 increments. Counts further revealed that AGC2 formed 48-75
244 days after first increment formation (mean 60.7 ± 1.8). The average duration of
245 metamorphosis was calculated by subtracting the mean number of increments at AGC2 from
246 the mean number of increments for AGC1 ($60.7 - 30.5$) equating to 30.2 days. The next
247 formed structural feature was the check mark. Counts indicated that the mark formed 60-91
248 days after first increment formation (mean 76.5 ± 2.0 inc.). These results suggest that the
249 movement of juvenile *M. a. australis* into deeper waters occurs roughly 15.8 days after
250 metamorphosis is completed (i.e. $76.5 - 60.7$ days). The results further confirmed that this
251 check mark formed during austral summer and should therefore not be interpreted as a winter
252 annulus. The true winter annulus became visible in otoliths collected after July 2017 (austral
253 winter).

254 3.3 Trace elemental analyses

255 15 Otoliths were successfully ablated and analysed from the core to the margin using LA-
256 ICP-MS. Varying trends in element:⁴²Ca ratios were observed for the four elements ²⁴Mg,
257 ⁵⁵Mn, ⁸⁸Sr and ¹³⁸Ba (Figure 6) and comparisons between ontogenetic zones revealed
258 significant differences in element:⁴²Ca ratios for all elements (Table1).

259 *INSERT FIGURE 6*

260 ²⁴Mg:⁴²Ca ratios ranged between $5.01e^{-5} - 2.57e^{-4}$ $\mu\text{mol}\cdot\text{mol}^{-1}$ across the three otolith zones.
261 ²⁴Mg:⁴²Ca ratios were found to be significantly lower during metamorphosis compared to the
262 larval (pre-metamorphosis) and juvenile otolith zones ($p = 0.01$ and $p = 1.90e^{-3}$ respectively).
263 The opposite trend was observed for ⁵⁵Mn:⁴²Ca ratios which ranged between $7.94e^{-08} - 1.42e^{-$
264 ⁰⁵ $\mu\text{mol}\cdot\text{mol}^{-1}$. Here, the onset of metamorphosis resulted in a significant increase in the
265 ⁵⁵Mn:⁴²Ca ratios ($p = 6.30e^{-11}$). This decreased slightly after metamorphosis, but remained

266 significantly higher compared to the larval (pre-metamorphosis) zone ($p = 1.78e^{-03}$). $^{88}\text{Sr}:^{42}\text{Ca}$
267 ratios ranged between $1.97e^{-3} - 1.1e^{-2} \text{ mmol}\cdot\text{mol}^{-1}$ and were found to be significantly higher
268 during the larval (pre-metamorphosis) phase compared to the metamorphosis and juvenile
269 phases ($p = 7.20e^{-4}$ and $p = 8.10e^{-4}$ respectively). No statistical differences were found in
270 $^{88}\text{Sr}:^{42}\text{Ca}$ ratios between the metamorphosis and juvenile otolith zones. $^{138}\text{Ba}:^{42}\text{Ca}$ ranged
271 between $1.60e^{-06} - 1.92e^{-08} \mu\text{mol}\cdot\text{mol}^{-1}$. Unlike the other three elements included in this
272 study, no shift was observed in $^{138}\text{Ba}:^{42}\text{Ca}$ ratios at the onset of metamorphosis. $^{138}\text{Ba}:^{42}\text{Ca}$
273 ratios were shown to remain stable throughout the entire larval period (including
274 metamorphosis). However, $^{138}\text{Ba}:^{42}\text{Ca}$ ratios were found to be significantly higher in the
275 juvenile zone compared to the larval (pre-metamorphosis) and metamorphosis zone ($p =$
276 $0.91e^{-3}$ and $p = 0.10e^{-3}$ respectively) possibly indicating a movement into deeper waters.

277 *INSERT TABLE 1*

278 Next, the single element: ^{42}Ca data were combined and analysed using a multivariate random
279 forest model to determine whether elemental fingerprints could be distinguished between
280 different ontogenetic zones of the otolith. The model was iterated 1000 times and model
281 classification accuracies ranged between 62.2 – 92.2 % (mean 81.1 ± 0.1 %). To further
282 identify where correct and misclassification was occurring the results for the three zones were
283 separated (Figure 7). Correct classification is indicated by filled squares (Figure 7). Distinct
284 elemental fingerprints were determined for two out of three otolith zones (B and C). For zone
285 A correct classification of spots ranged between 6.25 – 84.62 % (mean 52.3 ± 0.4 %). Spots
286 from zone (A) had the highest misclassification rates at 37.0 ± 0.4 % and 10.7 ± 0.3 % for
287 zones B and C respectively. This indicates the highest similarity in composition of the larval
288 (pre-metamorphosis) zone elemental fingerprint with the metamorphosis zone fingerprint.
289 Spots from the metamorphosis zone revealed the highest classification accuracy in this study
290 at 74.4 – 100 % (mean 91.2 ± 0.1 %). Misclassification of spots from zone B for zone A and

291 C was relatively low at 6.0 ± 0.1 and 2.8 ± 0.1 % respectively. This suggests that the shifts
292 found in single element:⁴²Ca ratios during metamorphosis all contribute to create a distinct
293 elemental fingerprint for this otolith zone. Classification accuracies for spots from the
294 juvenile zone (C) ranged between 50.0 – 100 % (mean 77.9 ± 0.3 %). Classification error of
295 spots from zone C was relatively low at 7.3 ± 0.2 and 14.8 ± 0.2 % for zones A and B
296 respectively. The order of contribution to correct classification for the four elements was
297 further determined as ⁵⁵Mn, ⁸⁸Sr, ²⁴Mg and finally ¹³⁸Ba.

298 *INSERT FIGURE 7*

299 **4. Discussion**

300 *4.1 Age validation*

301 A linear relationship was determined between date of capture and number of GIs within the
302 otoliths of juvenile *M. a. australis*. This relationship validated the daily formation of the
303 observed increments and revealed the mean date of first increment formation as 03-11-2016.
304 The interpretation of, and age at, first increment formation in *M. a. australis* is currently
305 unknown. In the European hake, *Merluccius merluccius*, age at first increment formation was
306 determined at 8 days (Morales-Nin et al., 2005), corresponding to first exogenous feeding.
307 Using this as a proxy for *M. a. australis* combined with a known embryogenic development
308 time of up to ~8 days (Laptikhovsky and Brickle, 2009), reveals a mean spawning date of 18-
309 10-2016. This falls within the estimated spawning period (Sep-Oct) observed for *M. a.*
310 *australis* (Schubnikov *et al.* (1969) in (Barrera-Oro and Tomo, 1988; Pajaro and Macchi,
311 2001), further supporting our validation. TL was shown to increase throughout the first year
312 of life, with a reduction in growth rate observed during the austral winter months (roughly
313 between 175-275 days in age). Our GAM-predicted TL at age 1 year (365 days) of 20.31 cm
314 was similar to the range (19.75 – 24.86 cm TL) reported in previous age and growth papers

315 (Zukowski & Liwoch (1977) in Barrera-Oro and Tomo (1988); Cassia, 2000). However, this
316 study is the first to present validated age estimates for *M. a. australis*.

317 4.2 Timing of early life history events

318 The validated daily growth increments were further used to elucidate the timing and duration
319 of early ontogenetic stages in *M. a. australis*. The order of formation of these microstructural
320 features correspond to the order of the early life history events. Victor (1982) demonstrated
321 through combined field observations and observations on the otolith microstructure that
322 settlement in the wrasse *Halicoeres bivittatus* occurs before metamorphosis. In this study
323 however, the otolith microstructure revealed that accessory growth centres formed before the
324 check mark. From this we determined that metamorphosis occurred before the movement of
325 juvenile *M. a. australis* into deeper waters. The initiation and completion of metamorphosis
326 was estimated to occur $30.5 (\pm 0.3)$ and $60.7 (\pm 1.8)$ days respectively after first increment
327 formation. Similar timings have been reported for the completion of metamorphosis in two
328 separate Argentine hake (*Merluccius hubbsi*) populations at 64.1 - 66.7 days (Buratti and
329 Santos, 2010), and for European hake (*Merluccius merluccius*) at 64 days after first
330 increment formation (Morales-Nin et al., 2005). The check mark (HZ1) was further estimated
331 to form $76.5 (\pm 2.0)$ days after first increment formation. Buratti and Santos (2010) estimated
332 that in *M. hubbsi* the transition from pelagic to demersal habitat was completed ~80 days
333 after first increment formation. These results suggest that the check mark in *M. a. australis*
334 otoliths could therefore indicate the timing of movement to deeper waters in juvenile *M. a.*
335 *australis*. Additionally, Barrera-Oro and Tomo (1988) were the first to suggest that the first
336 check mark observed in *M. a. australis* otoliths was a settlement mark and proposed that this
337 mark corresponded to a shift in habitat (and subsequently diet) rather than a period of reduced
338 growth expected during the winter. Our results support this by showing that the check mark is
339 formed 1) in austral summer and 2) shortly before the first juveniles appear as bycatch in the

340 bottom-trawl fishery. However, the term settlement mark is confusing for *M. a. australis* as
341 this species does not undergo settlement. The term ‘juvenile transition-mark’ (JT-mark) is
342 therefore proposed to describe the movement of juveniles from the pelagic larval habitat to
343 the deeper, juvenile pelagic habitat. The second check mark was only observed in otoliths of
344 older juveniles caught after the austral winter period which coincided with an observed
345 reduction in growth. Using the otolith microstructure to infer the timing of shifts between
346 early stages of ontogeny in *M. a. australis* could further be improved through tank-rearing
347 experiments, where the formation of microstructural features can directly be linked to
348 observed changes in behaviour and morphology (Modin et al., 1996).

349 4.3 Ontogenetic shifts in single element:⁴²Ca ratios

350 Mean element:⁴²Ca ratios reported in this study were similar to those previously reported for
351 *M. a. australis* by Arkhipkin et al. (2009) and Niklitschek et al. (2010), except for ⁸⁸Sr (Table
352 2). Variation in these ratios is likely due to the sampling method used combined with the
353 otolith area (and thus life history stage) studied (Table 2).

354 *INSERT TABLE 2*

355 Significant differences in element:⁴²Ca ratios between otolith zones were found for all
356 elements studied. A significant reduction in ²⁴Mg:⁴²Ca ratios was observed during the
357 metamorphosis period. ²⁴Mg incorporation into the otolith has been shown to be independent
358 of temperature and salinity (Martin and Thorrold, 2005), indicating that ²⁴Mg is not a useful
359 element to infer changes in habitat during early ontogeny. However, ²⁴Mg has been suggested
360 to be under physiological control (Woodcock et al., 2012), with incorporation rates being
361 negatively correlated with otolith precipitation rates and somatic growth (Martin and
362 Thorrold, 2005). Our results, coupled with observations of increased daily increment width (a

363 proxy for otolith precipitation rates) in *M. a. australis* otoliths during metamorphosis, support
364 this.

365 The initiation of metamorphosis was shown to result in a significant increase in $^{55}\text{Mn}:^{42}\text{Ca}$
366 ratios in the juvenile *M. a. australis* otoliths. Similar findings have been reported for juvenile
367 river herring (Alewife *Alosa pseudoharengus*) and blueback herring (*Alosa aestivalis*)
368 otoliths (Turner and Limburg, 2015). This increase was further positively correlated with
369 daily increment width (Turner and Limburg, 2015). Daily increments comprise of two units:
370 (1) a protein-rich layer and (2) a discontinuous, (typically) aragonite crystal layer (Morales-
371 Nin, 2000; Pannella, 1971). Areas of high precipitation rates are characterised by a relative
372 increase of protein-rich layer compared to the aragonite crystal layer (Morales-Nin, 2000),
373 and Izzo et al. (2016) revealed that a relatively large proportion of ^{55}Mn (up to 55.96%) is
374 incorporated into the protein-rich layer of the otolith (compared to 0.04 and 2.24 % for ^{88}Sr
375 and ^{138}Ba respectively). This was supported by Thomas et al. (2019) who suggested that ^{55}Mn
376 (II) ions are co-factors for the kinase *FAM20C* which regulates the biomineralisation of
377 otolith proteins. The increase in otolith-precipitation rates during metamorphosis could
378 therefore explain the shift in $^{55}\text{Mn}:^{42}\text{Ca}$ ratios.

379 $^{88}\text{Sr}:^{42}\text{Ca}$ ratios were shown to decrease throughout the early stages of ontogeny. Similar
380 findings have been reported for *Solea solea* (Pontual et al., 2003) and various *Anguilla sp.*
381 (Arai et al., 2000, 1999a, 1999b, 1999c, 1997). Incorporation rates of ^{88}Sr into the otolith
382 matrix has been widely studied, and are generally negatively correlated with ambient water
383 temperature (see Campana (1999)). The observed decrease in $^{88}\text{Sr}:^{42}\text{Ca}$ ratios within the
384 otoliths of juvenile *M. a. australis* therefore implies an increase in ambient water
385 temperature. Arkhipkin et al. (2004) observed a general warming trend within the first 100-
386 150m of the water column around the Falkland Islands during the austral summer months,
387 which could explain the decrease in $^{88}\text{Sr}:^{42}\text{Ca}$ ratios found in this study.

388 $^{138}\text{Ba}:$ ^{42}Ca ratios have been shown to be positively correlated with dissolved organic matter,
389 which generally increases with depth (Ashford et al., 2005). No shift in $^{138}\text{Ba}:$ ^{42}Ca ratios was
390 observed during metamorphosis in this study. However, an increase was observed for the
391 juvenile phase. This suggests the movement of *M. a. australis* into deeper, more productive
392 upwelling waters after metamorphosis has been completed. This shift could be the start of
393 their transition to the late juvenile/adult habitat depth range of 130-800 m (Niklitschek et al.,
394 2010).

395 *4.4 Classification between early life history stages using elemental 'fingerprints'*

396 This study revealed significant shifts in single element: ^{42}Ca ratios between otolith material
397 laid down during different ontogenetic stages for all elements studied. Resulting classification
398 accuracies were consequently high for two out of three ontogenetic zones (91.2 – 77.9%
399 accuracy for zones B and C respectively). Misclassification of spots from zone A was high
400 ($47.7 \pm 0.3\%$) and was possibly due to the large variability found in the elemental signatures
401 in this zone, or by having the lowest number of spots available to train the random forest
402 model. However, a mean classification accuracy of $81.1 \pm 0.01\%$ suggests that the random
403 forest model was an appropriate method for our classification analyses. The relative
404 simplicity in implementing the random forest model (i.e. no assumed relationship between
405 variables and raw input data) further advocates the use of this classification method for
406 otolith chemistry data. Classification accuracies can be improved by selecting for the optimal
407 number of elements, where the inclusion of more elements does not necessarily improve
408 classification (Mercier et al. (2011); however, see Marriott et al. (2016)). In this study, a high
409 classification accuracy was determined for spots taken from different ontogenetic zones, with
410 ^{55}Mn being the highest contributor to classification accuracy. It could therefore be argued that
411 when using elemental 'fingerprints' to investigate natal origin, for example, elements such as
412 ^{55}Mn should be excluded to avoid confounding ontogenetic signals. This is supported by

413 Macdonald et al. (2008) who revealed increased classification accuracies in the natal origin of
414 three geographically isolated groups of Australian smelt (*Retropinna semoni*) when excluding
415 the ⁵⁵Mn-rich primordium of the otolith.

416 This study revealed that changes to the otolith microstructure of gadoid otoliths can be
417 correlated with changes in microchemistry. We further showed that ontogenetic signals
418 complied with traditional interpretations of certain elements in otolith chemistry data. For
419 future studies element selection must therefore be carefully tailored to address the aims of the
420 study. Element affinity to either the protein- or aragonite layer should also be taken into
421 account when selecting for elements. A study comparing classification accuracies using
422 ‘protein-elements’ vs ‘aragonite-elements’ could provide further insights into optimal
423 element selection in future otolith chemistry research.

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587

Table 1. Summary table for median element/Ca ratios per zone (All = all zones, A = larval, B = metamorphosis, C = juvenile) in *M. a. australis* otoliths. N = number of laser ablation spots available per zone. Summary statistics for elemental comparisons between all zones (Kruskal-Wallis rank sum test) and pairwise post-hoc (Wilcoxon rank sum test) comparisons between zones (A = zone A, B = zone B, and C = zone C) are also shown. MAD equals Median Absolute Deviation.

| Element | Zone | N | Median | MAD | Kruskal-Wallis comparisons | | | Wilcoxon post-hoc comparisons | |
|-------------------|------|-----|----------|----------|----------------------------|------------------|-----------|-------------------------------|-----------|
| | | | | | df | Chi ² | <i>p</i> | A | B |
| ²⁴ Mg | All | 359 | 9.77e-05 | 4.02e-05 | 2 | 14.695 | 6.44e-04* | | |
| | A | 69 | 1.04e-04 | 4.51e-05 | | | | | |
| | B | 200 | 8.81e-05 | 3.39e-05 | | | | 0.01* | |
| | C | 90 | 1.06e-04 | 3.60e-05 | | | | 0.76 | 1.90e-3* |
| ⁵⁵ Mn | All | 359 | 5.89e-06 | 4.15e-06 | 2 | 38.72 | 3.91e-09* | | |
| | A | 69 | 3.04e-06 | 1.68e-06 | | | | | |
| | B | 200 | 6.91e-06 | 3.23e-06 | | | | 6.30e-11* | |
| | C | 90 | 5.00e-06 | 5.00e-06 | | | | 1.78e-03* | 0.05 |
| ⁸⁸ Sr | All | 359 | 4.65e-03 | 1.71e-03 | 2 | 9.2276 | 9.91e-03* | | |
| | A | 69 | 5.25e-03 | 1.20e-03 | | | | | |
| | B | 200 | 4.36e-03 | 1.52e-03 | | | | 7.20e-04* | |
| | C | 90 | 4.14e-03 | 1.47e-03 | | | | 8.10e-04* | 0.36 |
| ¹³⁸ Ba | All | 359 | 4.62e-06 | 2.13e-06 | 2 | 19.133 | 7.01e-05* | | |
| | A | 69 | 4.43e-06 | 1.51e-06 | | | | | |
| | B | 200 | 4.33e-06 | 2.11e-06 | | | | 0.95 | |
| | C | 90 | 5.32e-06 | 1.82e-06 | | | | 0.91e-03* | 0.10e-03* |

* indicates significance

Table 2. Mean element:⁴²Ca ratios determined in this study, Arkhipkin et al. (2009) and Niklitschek et al. (2010) for *M. a. australis* otoliths.

| | Methodology | Sample region | ²⁴Mg | ⁵⁵Mn | ⁸⁸Sr | ¹³⁸Ba |
|----------------------------------|------------------------------|-----------------------|------------------------|------------------------|------------------------|-------------------------|
| This study | spot analyses | juvenile (< 4 months) | 108 | 6 | 4730 | 5 |
| Arkhipkin <i>et al.</i> (2009) | spot analyses (otolith core) | Juvenile (< 4 months) | 18 | 2 | 1909 | 3 |
| Niklitschek <i>et al.</i> (2010) | dissolved otolith fragments | juvenile (< 1 year) | 300 | 18 | 2340 | 11 |

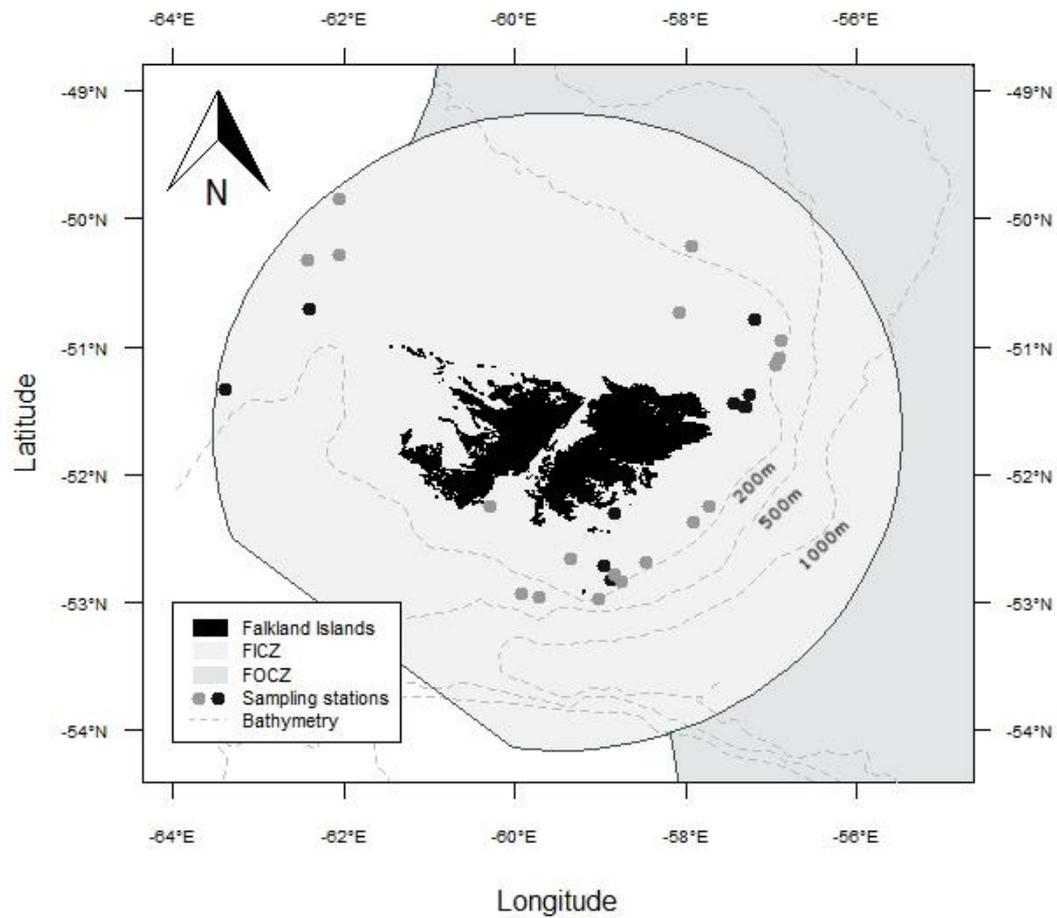


Figure 1. Map of sampling stations within the Falkland Interim Conservation and Management Zone (FICZ). Samples from all stations were used for age validation, and the dark grey stations indicate which samples were further subsampled for trace elemental analyses. FOCZ represents the Falklands Outer Conservation Zone.

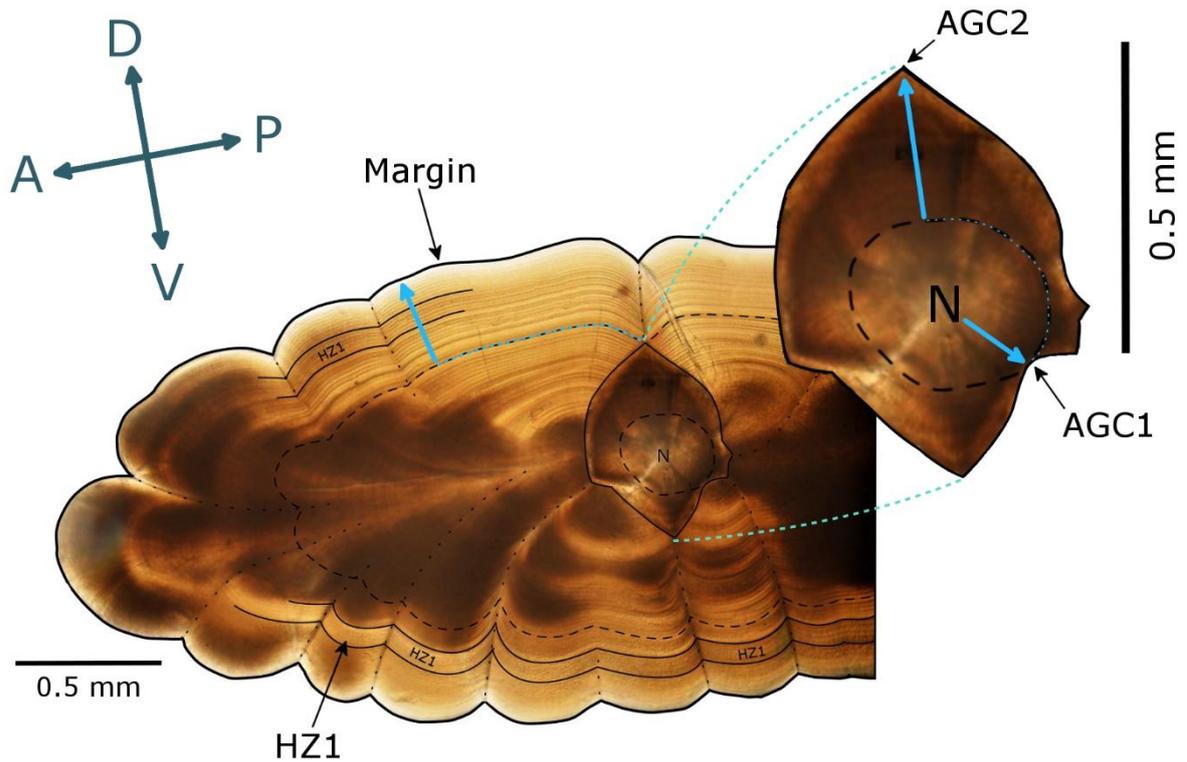


Figure 2. Juvenile *M. a. australis* (TL = 8 cm, 86 days) otolith section ground down to reveal the nucleus (N), prism shape (inset), accessory growth centres (i.e. AGC1, AGC2) and the first hyaline zone (HZ1). Blue arrows indicate direction of reading, blue dashed lines indicate change of reading orientation. Black dashed lines indicate shape of otolith at AGC1 and AGC2 respectively. Top left compass indicates anterior (A), posterior (P), dorsal (D) and ventral (V) orientation of otolith. Scale bars shown for otolith and inset respectively.

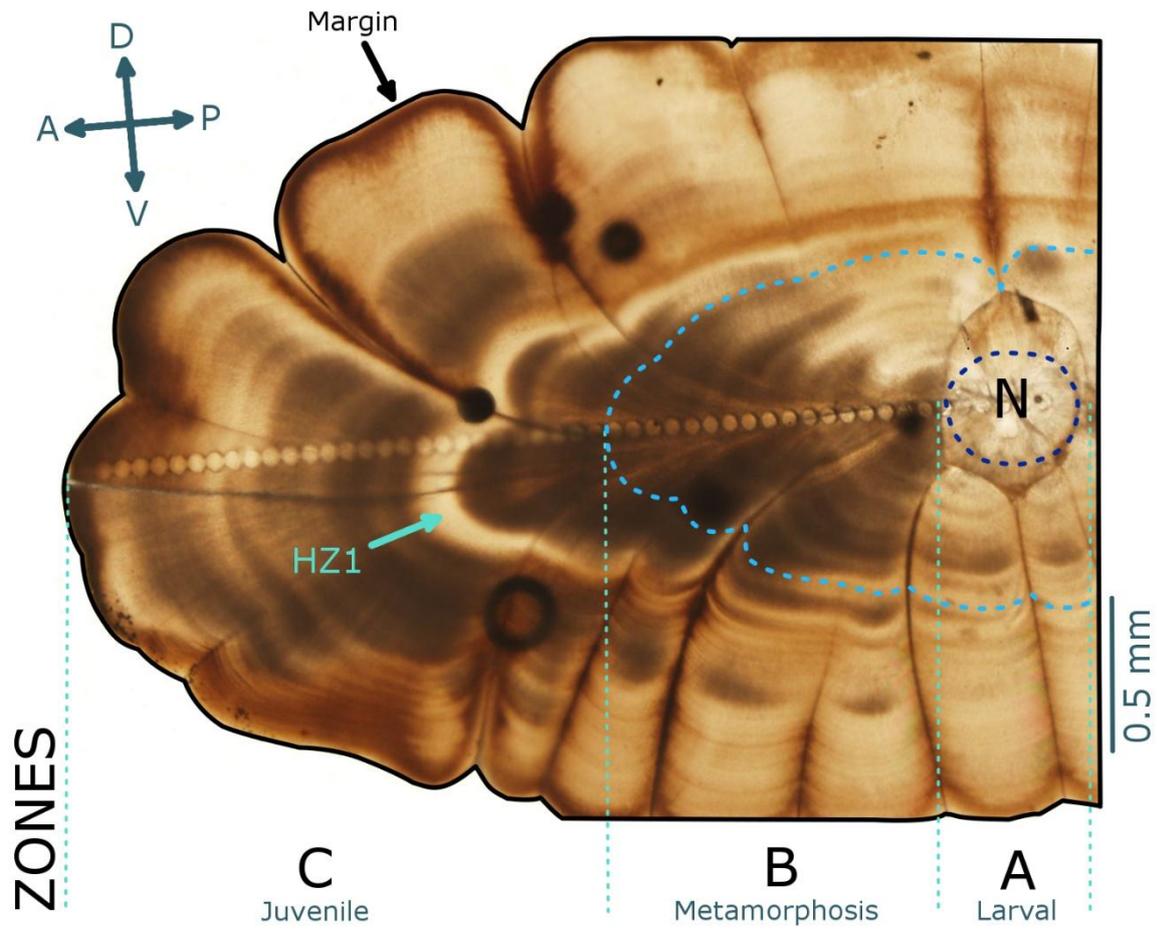


Figure 3. Juvenile *M. a. australis* (12 cm TL, 142 days) otolith section showing line of ablation spots analysed from the nucleus (N) to the margin, covering the first formed hyaline zone (HZ1). Larval and subsequent metamorphosis zones are outlined (dashed lines) to indicate significant change in otolith shape during metamorphosis.

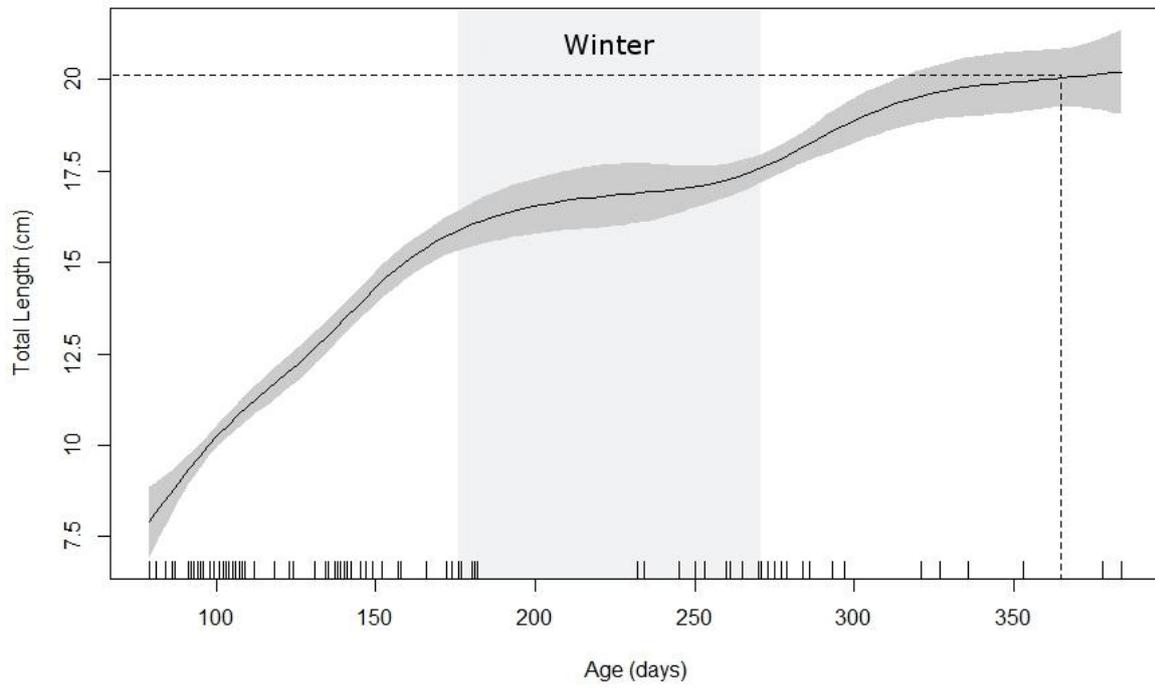


Figure 4. Relationship between total length (cm) and (smoothed) age (days) in juvenile *M. a. australis* of the 2016-cohort. 95% confidence intervals are depicted as shaded bands. Dashed lines indicate total length at 1 year of age. Rug depicts distribution of the data. Austral winter period is indicated by shaded area.

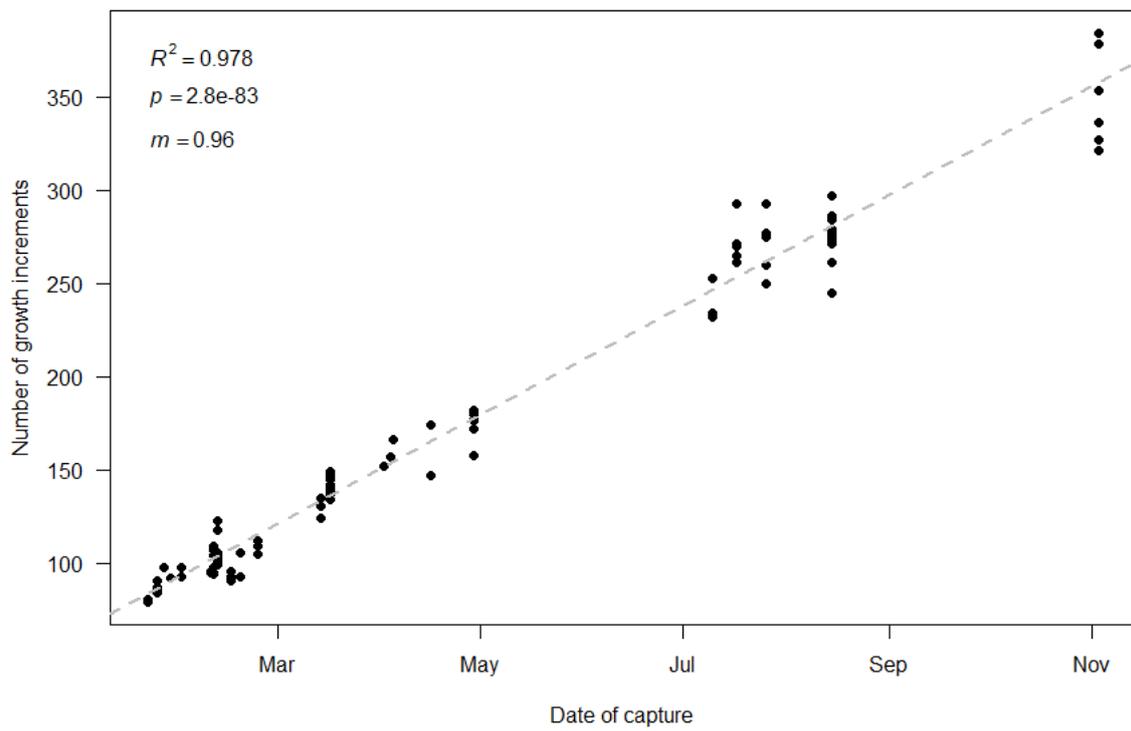


Figure 5. Number of growth increments within juvenile *M. a. australis* otoliths plotted against date of capture. Data was fitted with a (dashed) trend line for which the regression parameters R-squared, significance and slope are shown in figure.

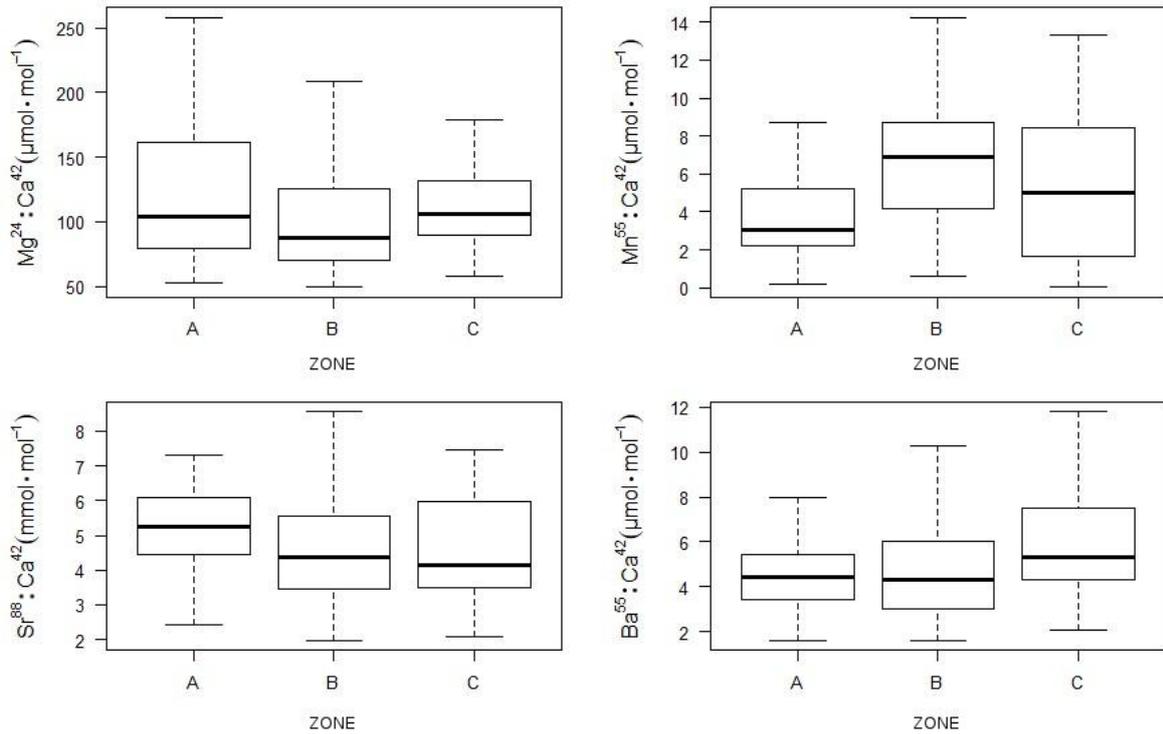


Figure 6. Boxplots depicting median element:⁴²Ca ratios found in different zones (A=larval (pre-metamorphosis), B=metamorphosis, C=juvenile) of *M. a. australis* otoliths. All observations are contained between the lower and upper whiskers (1st and 100th percentile respectively). The lower and upper quartiles represent the 25th and 75% percentile respectively.

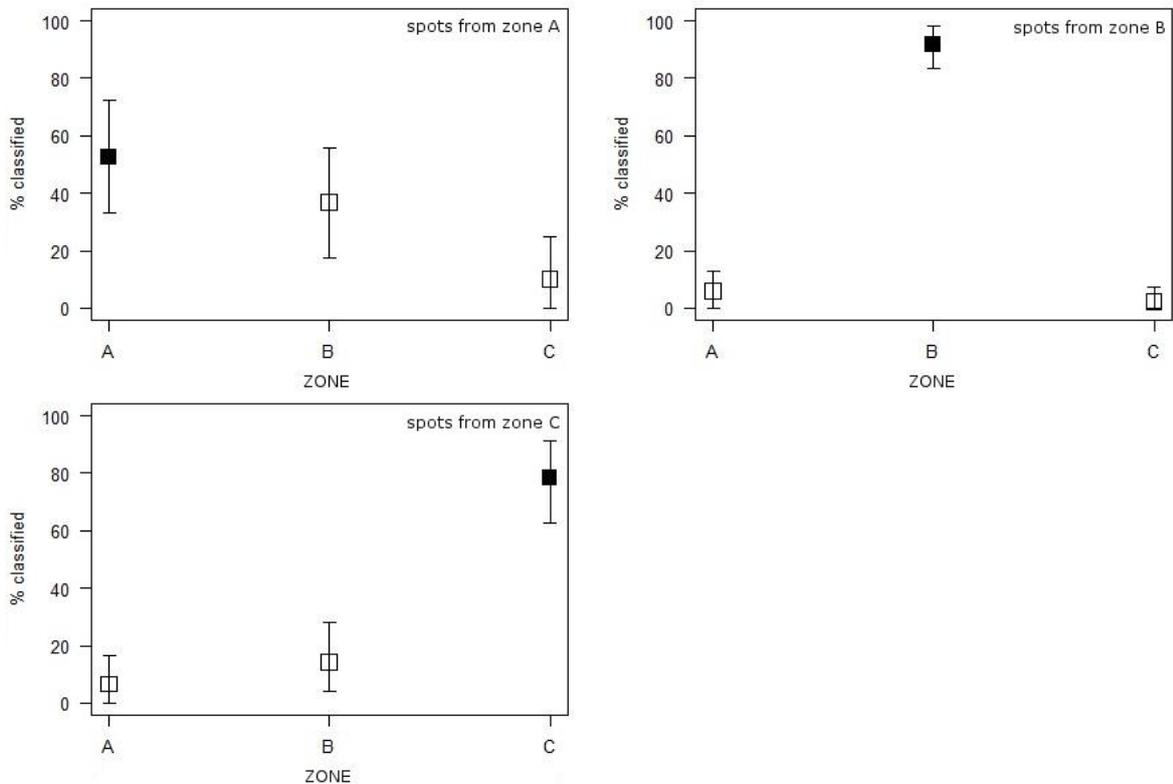


Figure 7. Percentage of spots assigned to each zone (x-axes: A = larval (pre-metamorphosis), B = metamorphosis, C = juvenile) in the random forest model. Model was repeated 1000 times and data points (squares) represent the median

**(500th) observed values with top and lower error bars representing the 950th and 50th observed values respectively.
Filled squares = correct classification, empty squares = classification error.**