

1 **Nitrate drawdown during a shelf sea spring bloom revealed using a novel**
2 **microfluidic *in situ* chemical sensor deployed within an autonomous underwater**
3 **glider**

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17

18 **Abstract**

19 Here we describe, for the first time, the use of a miniaturized Lab-on-Chip
20 (LoC) nutrient sensor deployed within an autonomous underwater vehicle (AUV;
21 Kongsberg Seaglider) to collect high-resolution nitrate (nitrate + nitrite) data in a highly
22 dynamic shelf sea environment. Seasonally stratified temperate shelf seas act as
23 important carbon sinks, where primary production is controlled by the availability of
24 nutrients such as nitrate. Spring phytoplankton blooms and sporadic mixing events can
25 drastically modify the availability of nitrate on temporal scales from hours to days.

26 Traditional sampling methods are unable to capture high frequency events that can be
27 clearly observed using a wet-chemical microfluidic system deployed within a glider.
28 We highlight firstly, an excellent agreement between the LoC and shipboard nitrate +
29 nitrite measurements ($r^2 = 0.98$ $n = 11$) during the onset of a spring bloom. Secondly,
30 the LoC was able to observe a decrease in nitrate within the surface mixed layer from
31 $5.74 \mu\text{M}$ (4th) to $1.42 \mu\text{M}$ (25th), whilst bottom layer concentrations remained constant
32 ($6.86 \pm 0.16 \mu\text{M}$), with an estimated analytical uncertainty of $< 0.2 \mu\text{M}$. Thirdly, the
33 ability of an LoC sensor deployed within an AUV to accurately capture simultaneous
34 biogeochemical and physical parameters at an enhanced resolution, on both spatial and
35 temporal scales, improves our understanding of biogeochemical cycles within the
36 dynamic temperate shelf sea environments.

37

38 **1. Introduction**

39 Seasonally stratified, temperate shelf seas act as important global carbon sinks through
40 the continental shelf pump mechanism (Tsunogai, Watanabe and Sato, 1999; Thomas
41 *et al.*, 2004; Takahashi *et al.*, 2009). Despite the shelf seas' relatively small size (8% of
42 global ocean area), they actually account for 15–30% of total oceanic primary
43 production (Tweddle *et al.*, 2013) and have an average carbon fixation rate per unit area
44 ~2.5 times greater than the open ocean (Simpson and Sharples, 2012). The interplay
45 between light, nutrients and mixing are key drivers of primary production. A well-
46 studied temperate shelf system, the Celtic Sea, shows a clear seasonal cycle whereby
47 initial light and nutrient conditions are ideally suited to support the onset of the spring
48 bloom (Fasham, Holligan and Pugh, 1983). This results in the rapid drawdown of nitrate
49 (NO_3^-) from ~ 6-8 μM to below the limit of detection ($< 0.1 \mu\text{M}$) (e.g. Hickman *et al.*,
50 2012; Davidson *et al.*, 2013). Post-bloom, new primary production is limited to the sub-

51 surface chlorophyll maximum (SCM) where fluxes of NO_3^- into the thermocline fuels
52 new production (Hickman *et al.*, 2009).

53 In the central Celtic Sea, shear generated turbulence (Sharples *et al.*, 2001, 2007;
54 Rippeth *et al.*, 2009) and wind-driven oscillations (Tom P. Rippeth, 2005) are the
55 central mechanisms in driving diapycnal mixing of NO_3^- up into the nutrient deplete
56 surface mixed layer. Wind-driven shear occurs in episodic short-lived spikes (0.5-1 hr)
57 and has the potential to have a large impact on NO_3^- fluxes, with observations indicating
58 that this flux can be up to 4 times greater than when no shear was observed (Sharples
59 *et al.*, 2001; Palmer, Rippeth and Simpson, 2008; Rippeth *et al.*, 2009).

60 At present, it is difficult to effectively sample nitrate at the high resolution required to
61 observe these key short-term mixing events (Sharples *et al.*, 2007; Charlotte Williams
62 *et al.*, 2013). Combined with limited winter data, this can lead to incorrect seasonal
63 estimates of NO_3^- fluxes that are key to primary production and carbon fixation.

64 Chemical *in-situ* sensors can provide high-resolution data necessary to resolve
65 biogeochemical processes occurring in shelf seas (Prien, 2007; Adornato *et al.*, 2010).

66 Wet-chemical analysers, centred on microfluidic Lab-on-Chip (LoC) technology, are
67 at the leading edge of advancements for chemical *in-situ* nutrient measurements
68 (Nightingale, Beaton and Mowlem, 2015). Due to their compactness, low resource use
69 and analytical performance comparable to laboratory-based methods, LoC nutrient
70 sensors are well suited to high-resolution float, glider and mooring deployments.

71 Autonomous underwater vehicles, such as gliders, have already been shown to provide
72 an economic and efficient observation platform to resolve mesoscale and submesoscale
73 structures, allowing for high-resolution sampling of biogeochemical parameters, such
74 as NO_3^- using an *in situ* ultraviolet spectrophotometer (Johnson *et al.*, 2009,2017;
75 Rudnick, 2016).

76 Here we demonstrate the power of coupling the LoC nutrient sensor into a Seaglider to
77 obtain continuous *in-situ* high temporal and spatial resolution nitrite + nitrate (NO_3^-
78 $+\text{NO}_2^-$; hereafter defined as ΣNO_x) measurements over the duration of the spring bloom
79 in April 2015. This enabled short-term mixing events key to establishing the spring
80 bloom and its subsequent decline to be observed.

81

82 **2. Experimental**

83 A single LoC ΣNO_x sensor (Ocean Technology and Engineering Group,
84 National Oceanography Center, Southampton, UK) was integrated within the science
85 bay of a Kongsberg Seaglider (Ogive fairing) and deployed from the *R.S.S Discovery*
86 in the Celtic Sea, as part of the NERC funded Shelf Sea Biogeochemistry program in
87 April 2015. In addition, a second LoC ΣNO_x sensor was mounted on a Conductivity-
88 Temperature-Depth (CTD; Seabird 911 plus) rosette to enable direct comparison of its
89 measurements to concentrations of ΣNO_x from seawater samples collected *in situ* at the
90 same depth and time.

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97 measurements to concentrations of ΣNO_x from seawater samples collected *in-situ* at
98 the same depth and time in July 2015.

99 Both LoC deployments by Seaglider and discrete ship-based CTD rosette
100 samples were collected from the Central Celtic Sea site (CCS; Fig. 1) located ~137

101 miles off the Cornish coast, UK (49°24.134'N, 8°36.248'W), with a water column depth
102 of ~ 145 m. The Seaglider with integrated LoC ΣNO_x sensor was deployed for 21 days
103 (4th-25th April 2015) and completed 776 dives within 10 km of the CCS station. A
104 total of 24 CTD casts within < 22 km of CCS were also conducted, allowing for
105 comparison between discrete water samples and the LoC ΣNO_x sensor. The rosette
106 package consisted of a Seabird 911 plus CTD and 24 bottle Niskin system, which was
107 used to collect discrete seawater samples. Chlorophyll *a* (hereafter Chl-*a*) was
108 measured on a pre-calibrated (spinach extract, Sigma Aldrich) fluorometer (Turner
109 Design Trilogy). Water samples collected from the Niskin bottles were analysed
110 onboard for $\text{NO}_3^- + \text{NO}_2^-$ using a segmented-flow autoanalyzer (Bran & Luebbe)
111 following the colorimetric procedures of Woodward and Rees (2001).

112 The LoC ΣNO_x sensor is composed of a three-layer PMMA chip which contains
113 precision-milled microchannels (150 μm wide, 300 μm deep), mixers and optical
114 components consisting of LEDs (525 nm, Avago Technologies, USA) and photodiodes.
115 Syringe pump, valves and electronics are mounted on the chip, which was encased in a
116 mineral oil-filled housing (PVC, 12 cm diameter, 30 cm height) with an internally fitted
117 pressure-compensating bladder. The LoC ΣNO_x sensor uses colourimetric detection,
118 using the Griess assay (Grasshoff et al., 2009) where NO_3^- is reduced to NO_2^- using an
119 off-chip copper activated cadmium column to enable ΣNO_x to be determined. A
120 detection limit of 20 nM and linear range of up to 350 μM have been demonstrated in
121 laboratory settings (Beaton et al., 2012). The LoC ΣNO_x sensor relies on a standard
122 measurement and a blank measurement to determine the concentration of the sample.
123 All reagents (Griess, imidazole buffer), standard and blank solutions were stored in
124 externally attached gas impermeable 150 mL Flexboy bags (Sartorius, UK) and the
125 waste was collected into a 500 ml Flexboy bag. The LoC ΣNO_x sensor, reagents and

126 standards used in this study have previously been described in detail by Beaton *et al.*,
127 (2012 & 2011) where it was deployed in a dynamic estuarine environment. More
128 recently, the LoC ΣNO_x sensor was deployed on a benthic lander in the Mauritian
129 oxygen minimum zone to examine cross-shelf transport of NO_3^- rich waters (Yücel et
130 al., 2015), and in glacial meltwaters rivers draining the Greenland Ice Sheet (Beaton et
131 al., 2017).

132 The LoC ΣNO_x sensor was integrated into the payload bay of the Seaglider
133 (Fig. 2) and connected by cable directly into one of the glider serial ports. The inlet tube
134 with filter is located on the surface of the payload bay cover ~30 cm behind the CT sail
135 (Seabird Electronics). The Seaglider software uses a CNF file that contains the
136 configuration for each on-board instrument and a CMD file that provides mission
137 parameters. The CNF file enables communication between the Seaglider and the LoC
138 ΣNO_x sensor. The LoC ΣNO_x sensor is set to 'logger' in the CNF file, which enables
139 the glider to send a number of commands. These commands allow the Seaglider to send
140 and receive data to and from the sensor. Some of the key commands are 'clock-set',
141 used only at the start of each dive but enables the sensor to store any time offset between
142 glider and LoC ΣNO_x sensor, 'status' which sends the sensor depth every 5 seconds
143 along with 3 trigger values and 'download' sent at the end of each dive requesting the
144 sensor to send both ascent and descent data files of processed ΣNO_x values. During
145 deployments the CMD file is typically transmitted to the Seaglider by satellite and
146 includes three trigger values that can be passed to the LoC ΣNO_x sensor using the
147 'status' command. These triggers are used to modify sensor behaviour at different
148 depths. Primarily these are used to ensure the LoC ΣNO_x sensor does not take samples
149 on the surface and risk the intake of air, and to aid with additional blank and standard
150 measurements.

151 For deployments on the CTD rosette, the LoC ΣNO_x sensor was programmed
152 in a continuous mode and performed a repeating measurement sequence of artificial
153 seawater blank, sample, and NO_3^- standard (3 μM), until the CTD rosette was
154 recovered to 5 m where the power was turned off. The operation of the LoC sensor was
155 as follows; The seawater sample was drawn into the sensor through a 0.45 μm MILLEX-
156 HP filter unit (Millipore). Both the filtered seawater sample and imidazole buffer were
157 pushed simultaneously through a serpentine mixer (used to aid mixing) before moving
158 through an off-chip cadmium column. Griess reagent was added to the resultant
159 buffered reduced seawater sample and then pushed through the 25mm absorption
160 measurement cell, where the pink-coloured azo dye developed. Absorbance was
161 calculated by comparing the optical intensity measured by the photodiode after a 100s
162 reaction wait time for each measurement. Each step involved 7 flushes prior to the
163 measurement to minimize sample carryover. Photodiode data was recorded at 1 Hz and
164 the average of the last 10 readings of each wait stage was used to calculate absorbance,
165 according to the Beer-Lambert law. Concentrations were calculated by comparing the
166 absorbance of each sample to that of the subsequent standard measurement. This results
167 in one blank-corrected sample measurement and standard every 17min during CTD
168 deployments.

169 For the Seaglider deployments, the LoC ΣNO_x sensor was programmed to
170 obtain a minimum of one artificial seawater blank and a standard measurement (6.5 μM
171 NO_3^-) at the beginning and end of each dive. After the first measurement of blank and
172 standard, continuous sample measurements took place on both the descent and ascent
173 until the glider was at 10 m. Discrete seawater samples were drawn through a 0.45 μm
174 MILLEX-HP filter unit (Millipore) into the sensor within the payload bay. Over the
175 period of sampling (21 days) this resulted in 312 and 199 artificial seawater blanks and

176 standards, which was sufficient to determine both the ΣNO_x concentration and any drift
177 associated with either the artificial seawater blank and/or the NO_3^- standard.

178 The Seaglider is a buoyancy driven autonomous underwater vehicle capable of
179 multi-month deployments collecting high-resolution profiles of physical and
180 biogeochemical parameters to 1000 m with a maximum travel range of 4,600 km
181 (Rudnick et al., 2004). Bilateral communication between the Seaglider and base station,
182 through an Iridium satellite connection, allowed dive configurations to be modified
183 once deployed. Data was transmitted back to shore during deployments to assess the
184 performance of the LoC ΣNO_x sensor. In addition to the LoC ΣNO_x sensor within the
185 science bay, the Seaglider measured conductivity & temperature (non-pumped Sea-
186 Bird SBE13 CT Sail, Seabird Electronics), pressure (Pain Electronics) and
187 fluorescence, turbidity and optical backscatter (Triplet Ecopuck, Wet Labs).
188 Conductivity, temperature and pressure were collected at a frequency of 1 Hz during
189 deployment with all dive profiles lying within 4 km of CCS (Fig. 1). Temperature and
190 conductivity were extracted and processed using the UEA Glider Toolbox (Queste,
191 2013). These routines apply manufacturer calibrations, correct for thermal inertia
192 following the methods of Garau *et al.*, (2011) remove spikes and anomalous data, and
193 draw upon a flight model similar to that described by Frajka-Williams *et al.*, (2011).
194 Four CTD casts, taken within 1.6 km of the glider, were used to calibrate the
195 temperature and salinity. Manufacturer calibrations were initially applied to data from
196 the Wetlabs Triplet for coloured dissolved organic matter (CDOM), backscatter &
197 fluorescence by subtracting the instrument blank and applying a scaling factor.
198 Calibration to convert fluorescence to Chl-*a* is based on the sensor's response to a
199 cultured diatom, *Thalassiosira weissflogii* at a known Chl-*a* concentration (Hemsley et
200 al., 2015; Sea-Bird Scientific, 2017).

201

202 3. Results and Discussion

203 The ability of the LoC ΣNO_x sensor to accurately determine ΣNO_x was
204 assessed in two ways: (i) by comparison of the LoC ΣNO_x sensor mounted on the
205 stainless-steel rosette with discrete water samples collected using Niskin bottles at the
206 same time and (ii) by comparing profiles obtained from the LoC ΣNO_x sensor deployed
207 within the glider with discrete water samples collected from the CTD on the same day
208 at CCS. The LoC ΣNO_x sensor, along with battery and reagents, was a similar size to
209 the 20L Niskin bottle and was mounted in place of a single Niskin bottle on the frame.
210 Once below 5 meters, a pressure sensor on the battery activated the power to the LoC
211 ΣNO_x sensor. The CTD rosette was held at three depths (45, 50 & 90 meters) for at
212 least 90 minutes to allow for triplicate LoC ΣNO_x sensor measurements at the
213 prescribed depth, each bracketed by a blank and standard measurement. During this
214 time period, two Niskin bottles were fired, one within 5 minutes of the first LoC ΣNO_x
215 sensor measurement and the second within 5 minutes of the last LoC ΣNO_x sensor
216 measurement. Water was collected from these discrete bottle firings for determination
217 of $\text{NO}_3^- + \text{NO}_2^-$ concentrations using the segmented flow autoanalyzer in the ship-based
218 laboratory. Figure 3a indicates the excellent agreement between sensor measurements
219 and analysis of discrete water samples with a correlation of $r^2 = >0.99$ ($n = 9$; $p =$
220 <0.001). The estimated analytical uncertainty of the LoC ΣNO_x sensor was calculated
221 from two times the standard deviation of the absorbance value of the deployed standard
222 over the three deployments and was $0.14 \mu\text{M}$ ($n = 10$). This analytical uncertainty is
223 higher than reported values for traditional segmented flow autoanalyzer analysis of
224 ΣNO_x ($0.03\text{-}0.07 \mu\text{M}$; Dafner, 2015) but lower than previous LoC ΣNO_x deployments
225 ($0.4\text{-}1 \mu\text{M}$; Yücel *et al.*, 2015).

226 LoC ΣNO_x data, collected from within the glider, was compared with $\text{NO}_3^- +$
227 NO_2^- values from 24 CTD profiles collected at CCS throughout the 21 day deployment.
228 Unlike the previous assessment of analytical uncertainty, data from the CTD profiles
229 were not collected at the same time (1–10 h window) or depth range (± 3 m) and all
230 samples were collected in a dynamic shelf sea system. Fig. 3b however, shows the
231 excellent agreement the LoC ΣNO_x data and the discrete water sample measurements
232 over the 21day period, with a correlation of $r^2 \geq 0.98$ ($n = 51$; $p \leq 0.001$). The average
233 estimated analytical un- certainty for the LoC ΣNO_x sensor during this period was
234 $0.19\mu\text{M}$ ($n = 142$) similar to the $0.14 \mu\text{M}$ ($n = 10$) for the CTD NO_3^- profiles.

235 Initial Seaglider dives (4th to 9th of April) were configured in a standard flight
236 mode, whereby the Seaglider adjusts its pitch and buoyancy to maintain a uniform glide
237 slope and descent and ascent speed (Eriksen et al., 2001). The LoC ΣNO_x sensor was
238 switched on at the beginning of each dive and completed a blank and standard
239 measurement followed by continuous measurements. The LoC ΣNO_x sensor acquired
240 depth information directly from the Seaglider, and using a depth trigger at 10 m the
241 sensor recognized the Seaglider was diving and after completing its current blank or
242 standard measurement would undertake continuous sample measurements. Triggers
243 were also used to take advantage of extra time at the beginning, apogee and end of dives
244 to undertake extra blank and standard measurements.

245 Fig. 4a shows that when the Seaglider was operated in the standard dive mode,
246 the LoC ΣNO_x sensor carried out 5 sample measurements per 120 m dive with a total
247 dive time of 30 ± 8 min. Moreover, these measurements were always in the same depth
248 ranges within the water column due to the relatively shallow water column ($\sim 145\text{m}$)
249 and sensor operation timings. Fig. 4b shows the excellent agreement be- tween the LoC
250 ΣNO_x sensor and traditional CTD segmented flow autoanalyzer measurements, for one

251 single CTD cast at 02:06 am and sixteen dives by the Seaglider from 00:18 to 09:25 am
252 on the 6th of April 2015. However, the temporal and spatial resolution of ΣNO_x data
253 (over the period of the day within the surface layer and across boundaries such as the
254 nitricline) would not be sufficient to investigate the depletion of ΣNO_x as the spring
255 bloom develops. As the glider can be controlled remotely, to increase the distribution
256 of measurements by the LoC ΣNO_x sensor throughout the water column, (in particular
257 across the nitricline), a second dive methodology - termed a 'loiter' dive - was
258 employed. After the Seaglider has reached its maximum depth for that particular dive
259 and started its ascent, for 30 mins the ascent angle was lowered and the glider 'loitered',
260 thus increasing the resolution of measurements within the water column (Fig. 4a).
261 Selected maximum target depths (90, 60, 40 and 25 m) were used to control the
262 maximum dive depth and ensured a higher number of measurements in areas of interest
263 (Fig. 4c). Loiter dives to 90 m doubled the amount of measurements made compared to
264 the previous standard dives. Fig. 4c shows a comparison between ΣNO_x concentrations
265 from two CTD casts at 02:06 and 08:22 am and LoC ΣNO_x concentrations from seven
266 'loiter' dives from 11:19 am to 16:25 pm on the 15th of April 2015. Once more, good
267 agreement between the LoC ΣNO_x sensor and traditional segmented flow autoanalyzer
268 measurements of CTD discrete samples was observed throughout the whole water
269 column during this 14-h period. To ascertain the ability of the LoC ΣNO_x sensor to
270 make comparable measurements to the segmented flow autoanalyzer, over an extended
271 time period (4th to 25th April 2015), we compared the measurements within the bottom
272 layer at 60–120 m where little changes in ΣNO_x were observed. Excellent agreement
273 between both the segmented flow autoanalyzer ($6.86 \pm 0.09 \mu\text{M}$; $n = 22$) and LoC ΣNO_x
274 sensor ($6.86 \pm 0.16 \mu\text{M}$; $n = 120$) was observed. We have demonstrated the ΣNO_x
275 concentrations measured from the LoC ΣNO_x sensor are com- parable to those of the

276 shipboard measurements analyzed on a segment flow autoanalyzer. Moreover, this
277 shows that accurate measurements can be obtained from the LoC ΣNO_x wet chemical
278 sensor over 21 days in a dynamic shelf environment.

279 On-board calibration with artificial seawater blanks and NO_3^- standard for each
280 dive enables both the monitoring of instrument performance and stability over long-
281 term deployments. To ensure maximum efficiency of the cadmium column, where
282 NO_3^- is reduced to NO_2^- a flow rate of 150 $\mu\text{l}/\text{min}$ was chosen (Beaton et al., 2012).
283 However, it is important to monitor any drift in the efficiency of the cadmium column
284 over time as this may impact on the ΣNO_x concentrations. In this study, a decrease in
285 absorbance values was observed, likely due to the gradual reduction in efficiency of the
286 cadmium column over time. As sample concentrations are calculated from their
287 associated blank and standard measurements, where the ratio of the absorbance of the
288 sample and standard are determined, any drift caused by the decreasing reduction
289 efficiency of the cadmium column is compensated for. Our results demonstrate that any
290 decrease in absorbance values observed did not impact on the accurate determination
291 of ΣNO_x from the LoC as shown by the excellent agreement with traditional
292 autoanalyzer ΣNO_x method from discrete water samples collected throughout the 21-
293 day deployment.

294 The data set presented here was collected during the spring phytoplankton
295 bloom, a period during which integrated net productivity becomes greater than
296 integrated losses and phytoplankton biomass accumulates in surface waters (Sverdrup,
297 1953). Over the 21-day deployment the LoC ΣNO_x sensor was able to accurately
298 capture the large drawdown of ΣNO_x within the surface layer due to the onset of the
299 spring bloom (Fig. 5). Concentrations decreased from 5.74 μM (4th) to 1.42 μM (25th),

300 whilst bottom layer NO_3^- concentrations remained constant ($6.86 \pm 0.16 \mu\text{M}$), as
301 observed in previous studies within the Celtic Sea (Tweedle, 2007; Williams, 2013).

302 At the start of the deployment (4th – 6th April 2015), a small $0.8 \mu\text{M}$ difference
303 between near surface (20–40 m) and bottom water (60–80 m) ΣNO_x concentrations was
304 observed. During this time, surface Chl-a concentrations were relatively low (1.8 mg
305 m^{-3}), but much higher (by 1.5 mg m^{-3}) than those typically observed during the winter
306 in the area (Pingree et al., 1976; Smyth et al., 2010). This suggests that some
307 phytoplankton growth had already occurred prior to deployment of the glider.

308 Between the 4th and 6th April the ΣNO_x concentration between 20 and 40 m
309 remained constant. Any changes in the near surface ($< 20 \text{ m}$) water however, where
310 you might expect the largest drawdown in NO_3^- , were not resolved since the standard
311 dive pattern used during this early period did not result in near surface LoC ΣNO_x
312 measurements being made. Nevertheless, it is clear from Fig. 5 that a large drawdown
313 of $3 \mu\text{M} \Sigma\text{NO}_x$ occurred between the 4th and 11th of April 2015 ($< 40 \text{ m}$) as stratification
314 of the water column initialized. This can be backed up with the observed increase in
315 surface water Chl-a, from 1.8 mg m^{-3} to 3.7 mg m^{-3} , indicative of phytoplankton
316 growth, and thus nutrient drawdown, during the onset of the spring bloom. Changing
317 the dive configuration to ‘loiter’ dives on the 11th of April in- creased the resolution of
318 ΣNO_x surface data.

319 Just as the temperature sensor resolves the gradual deepening and warming of
320 the surface mixed layer, the LoC ΣNO_x sensor resolves the coincident deepening of the
321 nitricline and drawdown of ΣNO_x above it. Between the 4th and 25th April surface
322 waters warmed by $> 1 \text{ }^\circ\text{C}$ and a 40 m deep thermocline is established. During this time

323 there is a 4.2 μM drawdown of ΣNO_x and an increase in Chl-a from a background of
324 1.8 mg m^{-3} to 4–6.8 mg m^{-3} .

325 By the end of the Seaglider deployment (25th) a two-layer water column had
326 developed with a warm, nutrient depleted, 40 m surface layer overlying colder, nutrient
327 rich bottom waters. Previous studies suggest that phytoplankton growth starts to
328 become ΣNO_x limited when concentrations fall below 1 μM (Eppley et al., 1969). The
329 low surface water ΣNO_x concentrations (1.4 μM) measured during the end of the
330 deployment suggest that this was close to happening. This is supported by a coinciding
331 decrease in Chl-a (to $< 2 \text{ mg m}^{-3}$) towards the end of the deployment. Ship-based
332 observations show that it was not until the 28th that surface water concentration were
333 below the limit of detection of 0.1 μM (Birchill et al., 2017).

334 This study has demonstrated for the first time that it is possible to accurately
335 measure ΣNO_x over long-term deployments using a wet chemical nutrient sensor
336 deployed within a glider. Our study provides a novel methodology, of differing
337 operational characteristics to current high-resolution capable methodology (Johnson et
338 al., 2013), enabling an increase in observations of ΣNO_x dynamics in temperate shelf
339 seas during key transitional events (e.g. the onset of stratification and the spring bloom,
340 convective overturning and the autumn bloom) and across fine-scale vertical and
341 horizontal features (e.g. tidal mixing fronts, sub-surface chlorophyll maximum). At
342 present, wet chemical biogeochemical sensors do not have the vertical resolution
343 capabilities of the commonly used physical and optical sensors such as temperature,
344 fluorescence or established ultraviolet spectrophotometer nitrate sensors (Johnson and
345 Coletti, 2002). This resolution could be further improved by decreasing the time taken
346 between measurements. By increasing flow rate, decreasing colour development time,
347 decreasing the number of flushes and increasing N.E.D (naphthylethylenediamine di-

348 hydrochloride) concentration and reaction temperature, an increased measurement
349 resolution could be achieved, but at the expense of measurement sensitivity (Beaton et
350 al., 2012; Pai et al., 1990).

351 For long-term deployment of wet chemical sensors, reduced re- source
352 consumption (power and reagents) and compact size are the main advantages of
353 microfluidic systems. Deployment of sensors on moorings, where solar and wind power
354 are available, negates power constraints. However, for autonomous underwater
355 vehicles power be- comes the greatest limiting resource (Dickey et al., 2008). During
356 the 21-day deployment, the LoC ΣNO_x sensor (version 3.2) had a low power
357 consumption of 1.5 W. This was only marginally higher than the other standard sensor
358 packages on the Seaglider (0.9 W and 0.25 W for the Wetlabs ECO Triplet and SBE
359 pumped payload CTD respectively), but lower than other wet chemical and UV
360 absorption systems (e.g. ISUS V3, Satlantic, USA; NitraVis, YSI, USA; SubChemPak,
361 SubChem Systems, USA). A single dive of the LoC ΣNO_x sensor consumed 2.5 mL of
362 Griess reagent, 2.5 mL of buffer solution and 0.21 mL of standard and blank solution
363 achieving 1 blank and standard measurement and ~ 10 samples on a dive to 120 m. This
364 would enable a total of 400 dives to be made and 4000 sample measurements. During
365 this study, the LoC was only active during periods when the R.S.S Discovery was also
366 sampling at CCS to enable a direct comparison between the two types of measurements
367 over a long-term deployment of the LoC. This resulted in accurate LoC ΣNO_x
368 determined over 21 days. With a sampling strategy focused on achieving the maximum
369 amount of measurements, a profile consisting of 10 sample measurements could be
370 undertaken every ~ 40 min, allowing for increased observations of episodic and
371 transient events unable to be observed by discrete sampling.

372

373 **4. Conclusion**

374 The temporal and spatial variability of biogeochemical processes has been
375 successfully measured through the use of remote sensing, time series moorings and
376 ship-based methods, but these, to an extent, fall short in resolving the dynamic temporal
377 and spatial elements with long- term endurance in a low cost package. From this, there
378 has been a strong call for the development of biogeochemical sensors to be deployed,
379 on stationary and mobile platforms, to provide in situ measurements as part of sensor
380 networks aimed at providing long term monitoring within a low cost package (Adornato
381 et al., 2010; Johnson et al., 2009). The LoC ΣNO_x sensor deployed within the Seaglider
382 in this study clearly demonstrates that nitrate & nitrite can be accurately determined
383 over monthly timescales due to the sensor's low resource use, small size and in situ
384 calibration abilities. Moreover, deployed within a glider with other physical and
385 biochemical data (e.g. CTD and the Wetlabs Triplet sensor measurements), it provided
386 a powerful tool for resolving dynamic biogeochemical processes within a dynamic shelf
387 system.

388

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401

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559 **Figures**

560 **Captions**

561 **Figure 1.** (a) Location of the Central Celtic Sea (CCS) sampling site (black cross) and
562 shelf and shelf edge bathymetry. (b) Location of the Celtic sea on the North West
563 European Shelf. (c) Seaglider tracks (black lines) and surfacing points (colour
564 dots) in relation to CCS from the 4th to 25th of April 2015.

565 **Figure 2.** (a) Diagram of the locations of sensors deployed on the Seaglider
566 (Kongsberg; Not to scale): (1) CT sail, (2) LoC, and (3) EcoPuck. (b) A LoC
567 ΣNO_x Sensor deployed with housing within the sensor payload bay of a
568 Seaglider (Ogive Profile). (c) A LoC ΣNO_x sensor consisting of (1) microfluidic
569 chip, (2) custom electronics, and (3) syringe pump assembly.

570 **Figure 3.** Comparison of autoanalyzer (AA) and LoC ΣNO_x measurements. (a) AA
571 and *in situ* LoC ΣNO_x measurements obtained from CTD casts. (b) AA and LoC
572 ΣNO_x measurements obtained from separate CTD casts (3rd to 6th, 11th, 12th, 15th,
573 16th, 20th, 21st and 25th) and *in situ* LoC glider profiles at CCS.

574 **Figure 4.** (a) Comparison of the two different dives used in this study. Blue and red
575 diamonds indicate where ΣNO_x measurements were obtained. (b) ΣNO_x
576 measurements from one CTD cast (02:06 am; dashed line) and sixteen standard
577 dives (00:18 to 09:25 am; diamonds) from the 6th April 2015. (c) ΣNO_x
578 measurements from two CTD cast (02:06 and 08:22 am; dashed lines) and seven
579 ‘loiter’ dives (11:19 to 16:25 pm; diamonds) from the 16th April 2015.

580 **Figure 5.** (a) temperature ($^{\circ}\text{C}$), (b) LoC ΣNO_z (μM) and (c) mean surface (< 20 m)
581 Chlorophyll-*a* measurements obtained from one Seaglider deployment from the

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4th to 25th April 2015.

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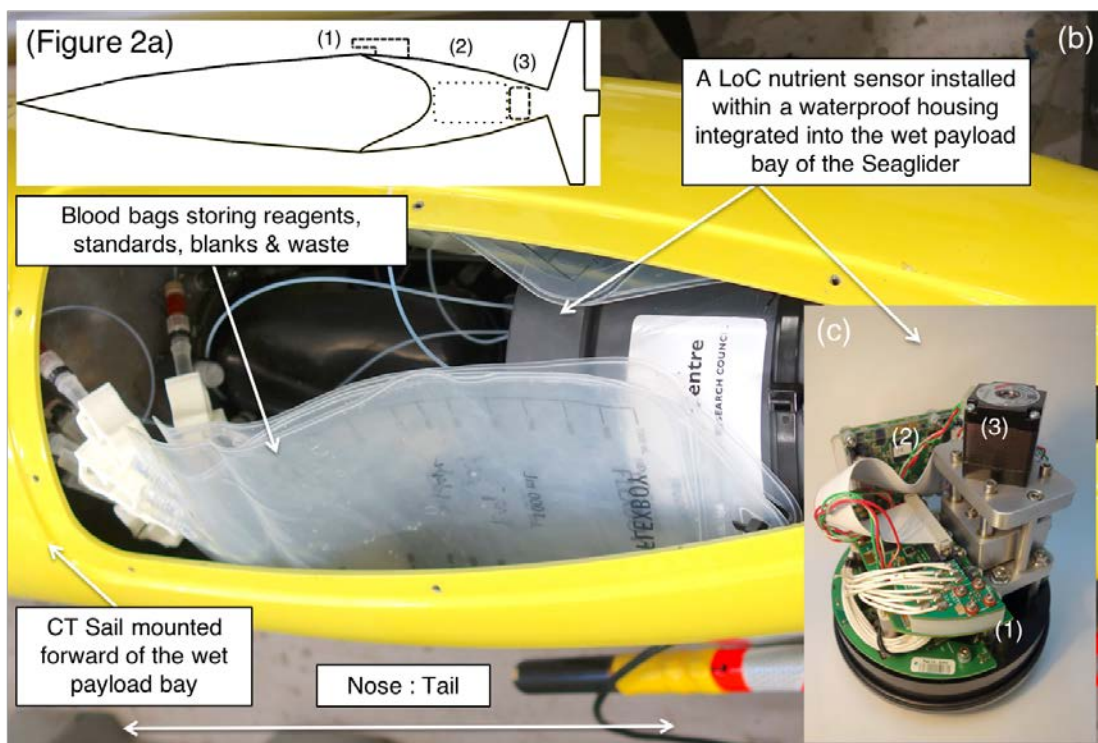
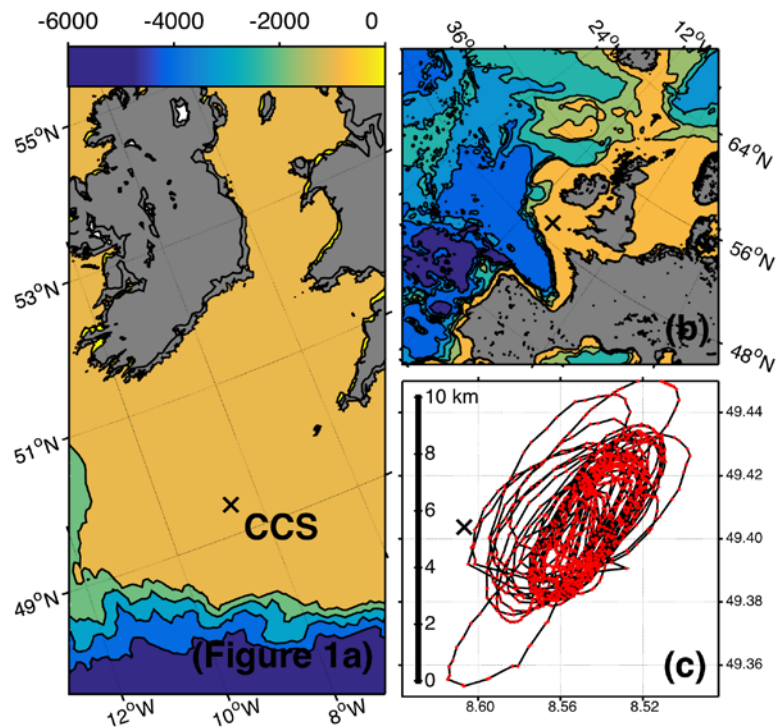
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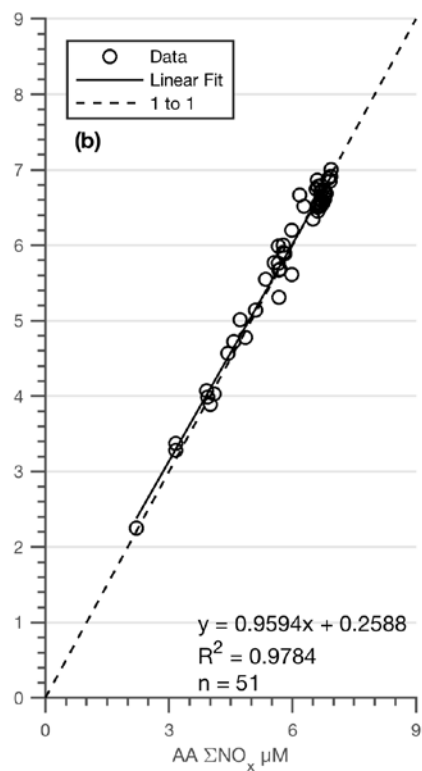
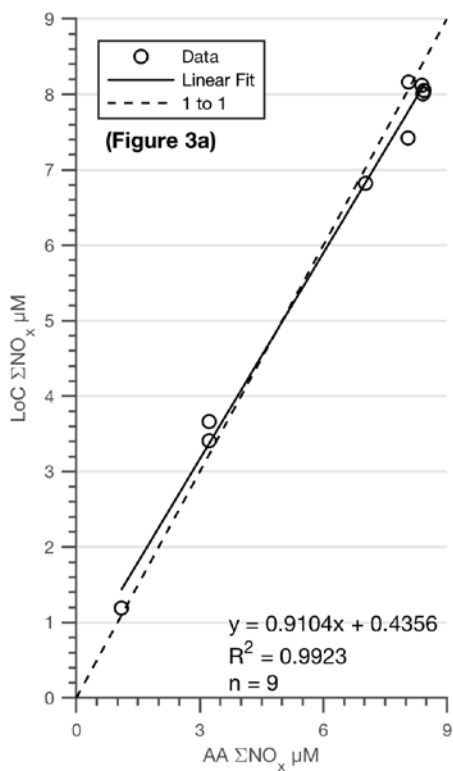
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