

Laboratory measurements of high-frequency, acoustic broadband backscattering from sea ice and crude oil

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Abstract: Recent decreases in summer sea ice cover are spurring interest in hydrocarbon extraction and shipping in Arctic waters, increasing the risk of an oil spill in ice covered waters. With advances in unmanned vehicle operation, there is an interest in identifying techniques for remote, underwater detection of oil spills from below. High-frequency (200–565 kHz), broadband acoustic scattering data demonstrate that oil can be detected and quantified under laboratory grown sea ice and may be of use in natural settings. A simple scattering model based on the reflection coefficients from the interfaces agrees well with the data.

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1. Introduction

It is estimated that as much as 13% of the world's undiscovered oil is located in the Arctic, mostly in offshore waters less than 500 m deep.¹ Because of technological advances, the retreat of Arctic sea ice extent, and other economic and political factors, further hydrocarbon extraction in Arctic waters is increasingly likely. With areas of significant new exploration offshore likely on the Alaskan outer continental shelf² there is the potential for an oil spill in areas that are seasonally covered by drifting pack ice. Airborne techniques for detection of oil trapped beneath pack ice have not yet proven effective in complex ice conditions, and on-ice techniques are logistically difficult in such an environment.² However, improvements in under ice operation of autonomous vehicles³ have opened up the possibility of oil spill detection from below a floating ice cover. To develop this capability, analysis of the efficacy of sensors for oil detection in a variety of ice conditions is required.

Congelation ice, also referred to as columnar ice, forms during calm sea conditions. Under turbulent conditions initial ice formation can take a different form (e.g., frazil ice). After the frazil layer consolidates additional thickening occurs in the form of congelation ice, which constitutes the majority of the Arctic ice pack.⁴ Congelation ice forms through the downward vertical growth of crystals into the ocean below. However, in the presence of saltwater, as is typical in freezing of binary mixtures, a pure solid is not formed, but rather a complex mixture of ice and liquid brine. As the ice continues to grow the brine layers begin to freeze as they continue to cool, pinching off into pockets of dense brine and a series of vertical channels that allow brine expelled during freezing to drain into the sea water. At the base of the columnar

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crystals near the water/ice interface there are narrow, planar dendrites interspersed with brine inclusions. The columnar crystals are typically less than a few centimeters in diameter while the spacing of the dendrites is 1 mm or less. The volume of brine within the bottom portion of the ice is determined by the amount of trapped salt and the ice temperature, which affect the connectivity of the brine. In this zone near the base of the ice, referred to as the skeletal layer, the bulk properties vary between those of sea water and pure ice.

This work focuses on the use of acoustic backscattering techniques for the detection and characterization of crude oil under a growing sheet of congelation ice. The physical scales associated with the complex structure of the skeletal layer suggest that high-frequency (>10 kHz) acoustic scattering techniques are well suited to studies aimed at remotely characterizing sea ice.^{5–9} The results presented here demonstrate that high-frequency, broadband techniques can be applied in both the temporal and frequency domains to identify a thin layer of crude oil, approximately 4 cm thick, beneath laboratory-grown congelation ice. Additionally, the measurements performed here suggest that the structure of the skeletal layer is significantly modified by the presence of even a very thin layer of oil, thinner than the oil layer thickness resolvable given the available bandwidth. The corresponding changes observed in the acoustic backscattering may provide another potential avenue of information for detecting and characterizing even very thin oil layers under congelation ice.

2. Methods

Measurements of high-frequency broadband acoustic scattering of sea ice and crude oil under sea ice were conducted in a 2000 L (1.37 m deep, 1.4 m in diameter) cylindrical tank in a cold room at Hamburgische Schiffbau-Versuchsanstalt GmbH (HSVA) in Hamburg, Germany from December 12–20, 2013. Throughout the experiment the room temperature was maintained at -18°C . A few hours before ice growth began on December 12, a uniform salinity profile ($>99.8\%$ NaCl) of approximately 32 PSU was measured in the tank using a RBR*concerto* conductivity, temperature, and depth (CTD) logger. The range to the surface prior to ice growth was 90 cm. Ice growth proceeded uninterrupted until the afternoon of December 16, at which point the acoustically inferred range to the water/ice interface was 78 cm, suggesting an ice thickness of 12 cm [Fig. 1(a)]. At that point 50 L of North Sea medium crude oil (API gravity = 30.75) was deployed through a pipe frozen into the ice using hydrostatic pressure; care was taken to avoid introducing air.

The pulse-echo system (Ref. 10, Fig. 2) was controlled by a National Instruments (NI) data acquisition system (PXI-1000B) with an embedded computer controller (PXI-8175, Windows 2000) and a custom LabView script. Transmitted signals were generated by a 12-bit, 40 MHz Arbitrary Waveform Generator (PXI-5411) and amplified (ENI 2100L: 50 dB gain). A signal sampler (Ritec SS-40, -40 dB) immediately following the amplifier was used to record the transmitted signal. Received signals were filtered and preamplified (RITEC Inc., Model BR-640 A). The transmitted and received signals were sampled at 4 MHz.

The acoustic backscatter measurements were obtained using multiple pairs of identical transducers mounted near the bottom of the tank. After the oil deployment, the range to the water/oil interface was 74 cm [Fig. 1(b)]. The transducers in each pair were closely spaced (<1 cm separation distance between transducers) to approximate monostatic measurements. The transducer pairs included 250 kHz and 500 kHz center frequency piston-like transducers (see Ref. 10, Table I, for details). The transmitted signals were 250 μs linearly modulated chirps from 200–300 kHz and 350–565 kHz.

The backscatter data analyzed here were obtained every morning throughout the experiment with the exception of the weekend. When time permitted additional data were obtained in the afternoons. A third data set was also obtained on December 16 immediately following the release of the oil. During sampling periods each transducer pair pinged a minimum of 30 times at 1 Hz. The observed backscattering from the interfaces did not vary over such short time periods so the pings were ensemble

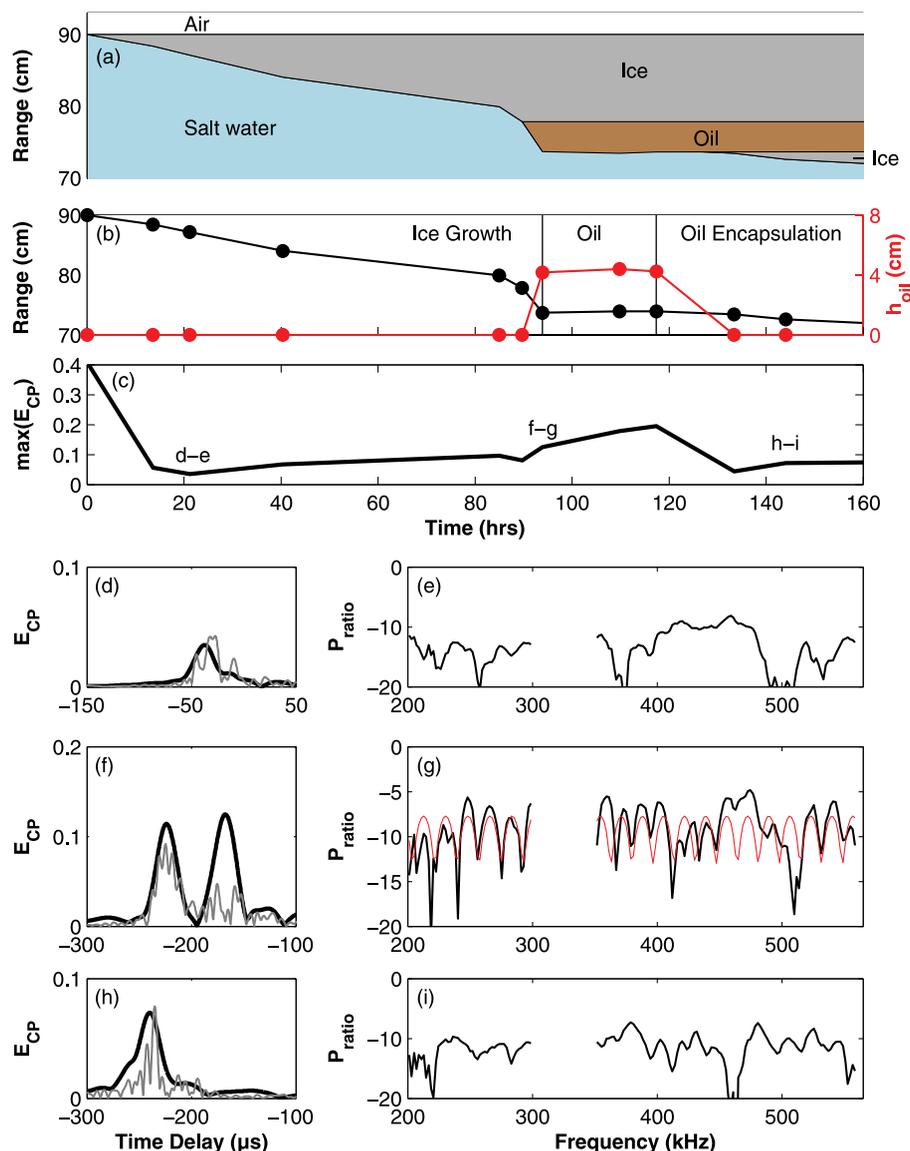


Fig. 1. (Color online) (a) Time series, linearly interpolated from the data in (b), showing the stages of the experiment. (b) Time series of inferred range to the bottom interface (ice or oil) based on the change in time delay from the pressure release surface and the acoustically inferred oil thickness. (c) Time series of the maximum value of E_{CP} . E_{CP} (black 200–300 kHz; gray 350–565 kHz) and P_{ratio} during each of the stages of the experiment—ice growth [(d)–(e)], oil under ice [(f)–(g)], and oil encapsulation [(h)–(i)]. The model is compared to P_{ratio} (black) in (g).

averaged. The data in Fig. 1(b) show the elapsed time, from the beginning of the experiment, at which the data were obtained.

The water/air pressure release surface was used to calibrate the system. Prior to ice growth, the tank was covered and allowed to settle overnight to minimize the surface disturbances. The transmitted and surface-reflected signals were recorded and processed for calibration. Assuming the surface is acoustically smooth, the received calibration voltage signals, $v_{cal}^R(t)$, are phase-shifted replicas of the transmitted signals.

Backscattering measurements were analyzed using pulse compression techniques, which increase signal-to-noise ratios and temporal resolution.¹¹ The compressed

pulse (CP) output is obtained by cross-correlating the received voltage time series with the received calibration time series

$$\text{CP}(t) = k_{\text{CP}}^{-1} v^R(t) \otimes v_{\text{cal}}^R(t), \quad (1)$$

where $v^R(t)$ is the received voltage time series, \otimes is the cross-correlation, and k_{CP} is the autocorrelation function of the received calibration time series at zero time lag. The temporal data are presented in terms of the envelope of the CP output, E_{CP} .

By assuming that the largest peak in the CP output is associated with the mean range to the bottom interface, the range is given by the following: $r = (c_w/2)t$, where c_w is the sound speed of the water ($c_w = 1440$ m/s) and t is the delay time to the first interface peak in E_{CP} . Once introduced, the buoyant oil forms a layer under the ice and the pulse compression output contains peaks from both interfaces. The thickness of the oil is then given by the following: $h_{\text{oil}} = (c_{\text{oil}}/2)(t_{\text{oil,ice}} - t_{\text{w,oil}})$, where c_{oil} is the sound speed of the oil, and $t_{\text{oil,ice}}$ and $t_{\text{w,oil}}$ are the time delays that correspond to the peaks in E_{CP} associated with the oil/ice and water/oil interfaces, respectively.

The frequency-dependent scattering is presented in the form

$$P_{\text{ratio}}(\omega) = \frac{P_{\text{scat}}(\omega)}{P_{\text{inc}}(\omega)} = G(\omega) \frac{V^R(\omega) r_{\text{scat}}}{V_{\text{cal}}^R(\omega) r_{\text{cal}}}, \quad (2)$$

where ω is the angular frequency, P_{scat} is the amplitude of the scattered pressure, P_{inc} is the amplitude of the incident pressure at the scattering interface, r_{scat} is the range to the scattering interface, and r_{cal} is the range to the pressure release surface during calibration. $V^R(\omega)$ and $V_{\text{cal}}^R(\omega)$ are the absolute values of the Fourier transforms of the received voltage time series and the received calibration voltage time series, respectively. $G(\omega) = V_{\text{cal}}^T(\omega)/V^T(\omega)$ is the ratio of the absolute value of the Fourier transform of the transmitted calibration voltage time series to the absolute value of the transmitted voltage time series. This formulation of the scattering data, used previously in studies of scattering from cylinders¹² and double-diffusive microstructure,¹³ accounts for the range dependence of the incident pressure, but not the range dependence of the scattered pressure. Therefore, directly comparing the results at different ranges may be inappropriate. If the scattering is assumed to be consistent with that from a smooth interface, then P_{ratio} reduces to the reflection coefficient. Attenuation in the water is neglected in this analysis due to the short ranges ($r_{\text{scat}} < 1$ m).

3. Results

The acoustically inferred range to the bottom interface throughout the experiment are included in Figs. 1(a) and 1(b). The maximum of E_{CP} , which is an approximation of the reflection coefficient, is plotted in Fig. 1(c). When the ice is thin, the reflection coefficient is high (>0.3) but decreases rapidly with ice growth. E_{CP} then remains less than 0.1 until the oil is introduced. The introduction of the oil increases E_{CP} until encapsulation occurs and E_{CP} decreases again. After the oil was introduced the acoustically inferred range to the water/oil interface remained constant until oil encapsulation began and ice growth continued. The acoustically inferred oil thickness from E_{CP} (200–300 kHz) was 4.2 cm prior to encapsulation. Following the encapsulation of the oil, the received signals no longer contained clear peaks from multiple interfaces.

Examples of the high-frequency, broadband scattering from growing sea ice and crude oil under sea ice are included in Figs. 1(d)–1(i) in both the temporal (left-hand column) and frequency (right-hand column) domains. As shown in pre- and post-oil measurements, the scattered signals using the 200–300 kHz transducers are easier to interpret than the higher frequency transducers because there is a single peak associated with each interface. The envelopes from the 350–565 kHz transducers are more difficult to interpret due to the presence of additional peaks that are not resolved in the 200–300 kHz data [Figs. 1(d), 1(f), and 1(h)]. This complexity is attributed to a combination of the increased temporal resolution and roughness of the basal structure relative to the acoustic wavelength.

4. Comparison of results to a simple 1-D weak-scattering model for the oil layer

To model the frequency spectra following the introduction of the oil, the water/oil and oil/ice interfaces are assumed to be smooth and infinite. Weak scattering is assumed so only the first reflection is considered. Previous work has shown high-frequency attenuation in the ice to be significant⁵ so reflections from the ice/air interface are neglected. Attenuation in the oil is also neglected. These assumptions result in a frequency-dependent total reflection coefficient with interference patterns related to the sound speed, oil thickness, and reflection coefficients at the interfaces. This term is described by

$$\mathcal{R}_{\text{model}}(k) = \frac{\mathcal{R}_{\text{w,oil}} + \mathcal{R}_{\text{oil,ice}} e^{2jk h_{\text{oil}}}}{1 + \mathcal{R}_{\text{w,oil}} \mathcal{R}_{\text{oil,ice}} e^{2jk h_{\text{oil}}}}, \quad (3)$$

where \mathcal{R} is a reflection coefficient calculated for each interface, k is the vertical wave number in the oil, h_{oil} is the thickness of the crude oil layer, and the subscript w refers to the water. The reflection coefficients are related to the mismatch in the acoustic impedance at each interface. The water properties ($c_{\text{w}} = 1440$ m/s, $\rho_{\text{w}} = 1025$ kg/m³) are based on the CTD measurements. The skeletal layer and bulk properties of the ice vary based on their structure and brine content, so sea ice can be modeled acoustically using sound speed profiles dependent on the range from the water/ice interface.^{7–9,14} Because of the high frequencies used in this work, reflection from the sea ice is modeled based on the impedance mismatch at the interface where the ice properties are $c_{\text{ice}} = 1700$ m/s and $\rho_{\text{ice}} = 960$ kg/m³, which are similar to values previously used to model reflections from ice.^{7–9,14} The properties of oil, $c_{\text{oil}} = 1475$ m/s and $\rho_{\text{oil}} = 885$ kg/m³, are based on empirical formulations for the properties of crude oil using the API gravity at 0°. ¹⁵ At the time of deployment the crude oil was approximately 5°C; no temperature measurements of the oil are available post-deployment. Given these sound speeds and densities, the reflection coefficients are $\mathcal{R}_{\text{w,oil}} = 0.06$ and $\mathcal{R}_{\text{oil,ice}} = 0.11$. The density value used for the ice represents a mixture of brine and ice and is difficult to measure in practice. However, the sound speed contrast at the oil/ice interface is much greater than the density contrast so the model is relatively insensitive to a realistic range of densities. In the absence of oil, the reflection coefficient from the water/ice interface is $\mathcal{R}_{\text{w,ice}} = 0.05$, which is consistent with previously reported values for sea ice at high-frequencies^{5–8,14} and the data in Fig. 1(c) prior to the introduction of oil.

The spectra and model comparisons included in Fig. 1(f) show that the model captures the peaks and nulls of the spectra with better performance at lower frequencies due to the previously described temporal resolution and relative roughness of the interfaces. The acoustically inferred oil thickness, h_{oil} , from the 200–300 kHz E_{CP} is 4.2 cm. $\mathcal{R}_{\text{model}}$ for the 200–300 kHz and 350–565 kHz data were determined by minimizing the rms error of the model with different oil thicknesses bounded by the h_{oil} obtained in the temporal domain. The modeled data in Fig. 1(g) used oil thickness of 4.2 cm (200–300 kHz) and 4.4 cm (350–565 kHz). Figures 1(h) and 1(i) show the E_{CP} and P_{ratio} following oil encapsulation, which do not clearly indicate the presence of an oil layer. Post-experiment ice cores showed that the encapsulation of oil produced a distinct, contiguous layer of oil separating the two crystallographically unconnected layers of ice with only small amounts of percolation into the brine channels near the oil/ice interface. This is consistent with previous studies that showed that, with the exception of melting periods when brine channels open, enabling vertical migration, oil entrainment is limited to low concentrations (<5% by vol.) within the relatively high-porosity skeletal layer and discrete brine channels near the oil/ice interface.^{16,17}

5. Discussion

The results presented in Sec. 4 demonstrate that high-frequency, broadband acoustic scattering techniques can be used to identify crude oil spills under congelation ice in a controlled, laboratory environment. Detection of crude oil spills using active acoustics requires balancing two competing considerations regarding the range resolution of

transducers: (1) resolution must be sufficient to resolve both the water/oil and oil/ice interfaces and (2) increases in frequency and range resolution complicate the interpretation of scattered signals due to the roughness of the interfaces. Reported equilibrium thicknesses of crude oil under “smooth” ice range from 0.5–1 cm to 2 cm or more.^{18–20} However, in a natural environment, oil is expected to spread in a thin lens following underwater topography with deeper pools in under-ice depressions. This would result in oil being present under only a fraction of the ice but with thickness varying from those expected under smooth ice to tens of centimeters or more.²¹

The application of high-frequency, broadband techniques to oil spill detection requires further studies to inform interpretation of the scattered signals. In this experiment, the data were easiest to interpret at the lowest frequencies, 200–300 kHz, with a corresponding range resolution of roughly 1 cm. In future work, spanning a lower frequency range (e.g., 80–500 kHz) could provide a balance between signal complexity and resolution across the range of potential oil thicknesses. The roughness and correlation lengths of both sea ice and crude oil under sea ice will affect scattering and are needed for the further development of models. There is also a need to study the impact of the beamwidth and the corresponding acoustic footprint, which varies by frequency for broadband transducers. Proper quantification of scattering from sea ice and oil under ice would also benefit from studies of the range dependence of scattering from these interfaces. Previous studies indicate that for the frequencies discussed here there is a transition between different statistical distributions,^{5,6} suggesting roughnesses between the asymptotic scattering regimes of smooth interfaces and point scatterers. Finally, future studies should consider the impact of oil on the structure of the skeletal layer and the implications for acoustic scattering.

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