

Contribution to the Sheet Description of theWells-next-thesea district (Sheet 130) -

Upper Cretaceous: Chalk Group

Integrated Geosurveys (Southern England) Programme Internal Report IR/04/112

BRITISH GEOLOGICAL SURVEY

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Contribution to the Sheet Description of the Wells-nextthe-sea district (Sheet 130) -Upper Cretaceous: Chalk Group

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Contents

Contentsi				
Sui	nmar	۲۷	ii	
1	Intro	Introduction1		
2			2	
	2.1	Grey Chalk Subgroup	2	
	2.2	White Chalk subgroup		
Re	feren	۔ Ces	6	

FIGURES

- Figure 1. Correlation and stratigraphical interpretation of key boreholes in the Chalk Group of west Norfolk. Interpretation of North Pickenham Borehole follows Mortimore & Wood (1986).
- Figure 2. Occurrence of Chalk Group biozones in the Wells district. Based on data from N B Peake (BGS archives) and I P Wilkinson (see text).

TABLES

Table 1.The stratigraphy of the Chalk Group of the Wells district, and its relationship to
lithostratigraphical subdivisions in southern England and north-east England.

Summary

This report describes the stratigraphy of the Chalk Group in the Wells-next-the-Sea district (Sheet 130), and forms a contribution to the Sheet Description and Sheet Explanation.

1 Introduction

Chalk is typically a very fine-grained, white limestone, predominantly composed of the disaggregated skeletal remains (coccoliths) of tiny planktonic algae that flourished in the seas of the Late Cretaceous. The Chalk Group is composed of almost pure calcium carbonate in the form of low magnesian calcite, except the lower part, which contains up to 30% clay. Flints, clay-rich horizons (marls), beds of indurated, mineralised chalk (hardgrounds), and coarsely bioclastic chalk horizons also occur, and some of these are geographically extensive marker-horizons.

The Chalk of north Norfolk, including the Wells district, in part belongs to the 'Northern Province' of Mortimore *et al.* (2001). This designation recognises differences in lithology and fauna of the Chalk of Yorkshire, Lincolnshire and northern East Anglia compared to southern England ('Southern Province'); differences that begin to appear in the Chalk extending north-eastwards from the Marlborough and Berkshire Downs into southern and central East Anglia ('Transitional Province'). In general, the Chalk of the 'Northern Province' is more indurated; a consequence of the secondary infilling of pore spaces by clacite, although the mechanism by which this has occurred is still unclear (Mortimore *et al.*, 2001). Stylolites are also a feature of the Northern Province Chalk, as seen in the core of the Bircham Borehole at the southern margin of the Wells district (Fig. 1), and demonstrating that locally extensive pressure solution has occurred. Flints occur at lower stratigraphical levels in 'Northern Province' Chalk compared to southern England, and are more frequently of tabular form (Mortimore *et al.*, 2001). However, not all of the north Norfolk Chalk succession is of Northern Province type, and it seems likely that the boundary between the Northern Province and Transitional Province fluctuated with time (Mortimore *et al.*, 2001).

Although a separate lithostratigraphical classification is applied to the Chalk of the 'Northern Province' compared to southern England (Rawson *et al.*, 2001; Table 1), as yet there is no agreed Chalk lithostratigraphy for the 'Transitional Province'. In the Wells district, an extensive cover of Quaternary sediments has prevented detailed mapping of the Chalk Group, the only mapped subdivision being the boundary between the Grey and White Chalk subgroups. However, sporadic outcrop and borehole data show that it is possible to locally recognise greater variation in the Chalk Group than is capable of being mapped (Fig. 1). Locally, a feature occurs above the base of the White Chalk that may equate with a level at or slightly below the 'Chalk Rock', and thus approximate to the base of the traditional 'Upper Chalk' (Table 1).

In the absence of outcrop data, biostratigraphy has historically been the key to understanding the Chalk of north Norfolk. The work of Peake and Hancock (1970) showed that macrofossil collections from more than 200 localities could be used to construct a biozonal map of the Chalk for the northern part of East Anglia. This work provides a more detailed picture of stratigraphical variation in the Chalk than can generally be achieved from feature mapping, and is a guide to correlation with successions to the north and south, and with the BGS Trunch Borehole [TG 2933 3455], a key reference borehole for the Chalk of north Norfolk (Wood et al., 1994). Figure 2 shows the occurrence and biozonal classification of localities in the Wells district, including those used in the construction of Peake & Hancock's biozonal map. The biozonal classification used by Peake & Hancock (1970), and in part employed on Fig. 2, differs from the standard macrofossil scheme; its relationship to current nomenclature is shown on Table 1. Microfossil samples from some of the localities on Figure 2 (numbered 1 to 10) were examined by I P Wilkinson for biostratigraphically significant foraminifera (Wilkinson, 2004). The results generally confirm previous biozonal assignments by Peake & Hancock (1970). An exception is the Titchwell Parish Pit [TF 762 433] (Locality 3 of Figure 2), assigned by Peake & Hancock (1970) to the plana Zone, but actually belonging to the M. coranguinum Zone on the basis of microfossil evidence (Wilkinson, 2004) and macrofossil data (Mortimore et al., 2001). Localities near Stiffkey and Warham show that the youngest Chalk in the Wells district equates with the

basal part of the Portsdown Chalk of southern England (Wilkinson, 2004), an interval in which marl seams typically occur. Marl seams have been recorded in the slightly younger *B. mucronata* Zone equivalent of the Portsdown Chalk in the Trunch Borehole (Wood *et al.*, 1994), but there is no lithological data for the Wells district.

2 Stratigraphy

Boreholes at South Creake [TF 8573 3401] and North Creake [TF 8566 3863], in the southern and central parts of the Wells district, show the Chalk Group to be between 140 and 160 m thick thereabouts. Younger Chalk Group strata occur further east, making the maximum thickness of the Group about 300 m based on biozonal thickness data (Peake & Hancock 1970; Wood *et al.*, 1994). Following Rawson *et al.* (2001), the Chalk of the Wells district has been subdivided into Grey Chalk Subgroup and White Chalk Subgroup, but extensive cover of Quaternary strata has prevented the mapping of formational boundaries. Nevertheless, a more detailed understanding of the succession can be gleaned from sporadic outcrops, boreholes and archival data.

2.1 GREY CHALK SUBGROUP

In the Bircham Borehole, just to the south of the Wells district, the Grey Chalk Subgroup is about 13 m thick (Fig. 1), compared to about 50 m in southern East Anglia (Bristow, 1990), and more than 70 m in southern England (Mortimore *et al.*, 2001). A similarly thin Grey Chalk Subgroup occurs in the Trunch Borehole, and represents a regional trend of increasing condensation and induration of this interval in north Norfolk, perhaps related to the presence of a large igneous intrusion at depth (Wood *et al.*, 1994; Mortimore *et al.*, 2001, fig. 1.15). The Subgroup comprises hard, marly chalk with stylolites, hardgrounds and shell-rich intervals. Its lithology is typical of the Grey Chalk Subgroup in the 'Northern Province' (= Ferriby Chalk Formation; Table 1), and several regionally developed marker beds can be recognised.

The basal part of the Grey Chalk Subgroup in the core of the Bircham Borehole is represented by fragments of very hard, nodular and glauconitised chalk. This probably represents the Paradoxica Bed, named for the abundance of burrows of *Thalassinoides paradoxica* (Woodward) and seen at the base of the Chalk Group in the cliffs at Hunstanton, just west of the Wells district (Gallois, 1994). Above this, the Bircham Borehole core contains a few metres of shell-rich chalk that probably represent the Lower and Upper Inoceramus beds, seen at Hunstanton and traceable into the Lincolnshire succession north of The Wash (Gaunt *et al.*, 1992; Gallois, 1994). These bioclastic beds are largely formed or the fragmental remains of the bivalve *Inoceramus* ex gr. *crippsi* Mantell, but may include other forms.

The Totternhoe Stone is a widespread marker in the Grey Chalk Subgroup of northern, eastern and parts of southern England. In the Bircham Borehole it appears to be represented by 0.3 m of grey, marly and silty chalk with glauconitised pebbles overlying a glauconitised hardground. At Hunstanton the Totternhoe Stone is 0.7 m thick, and just over a metre thick in the Trunch Borehole (Gallois, 1994; Wood *et al.*, 1994). The base of the Totternhoe Stone equates with the junction of the West Melbury Marly Chalk and Zig Zag Chalk formations in southern England.

The higher part of the Grey Chalk Subgroup contains the Nettleton Stone, an indurated unit of slightly gritty chalk with an oyster-rich marl at its base. In the Bircham Borehole, a 0.2 m thick marl, containing oyster remains, occurs less than a metre above the inferred Totternhoe Stone, and might represent the base of the Nettleton Stone. The Nettleton Stone has been described at several localities in the adjacent King's Lynn district, including Dersingham [TF 7013 3092] and Heacham [TF 688 368], just to the west of the Wells district (Gallois, 1994), and it is widely developed in the Chalk north of The Wash (Whitham, 1991; Gaunt *et al.*, 1994).

2.2 WHITE CHALK SUBGROUP

The White Chalk Subgroup probably has a maximum thickness of about 285 m in the Wells district, based on borehole and biozonal data. Although poorly exposed in the district, data from abandoned chalk pits and borehole records from adjacent areas provide some insight into the detailed stratigraphy.

The base of the Subgroup is the base of the Plenus Marls Member. In north Norfolk, the member is much attenuated from the several metres of clay-rich chalk that is typical of the unit in southern England. In the Kings Lynn district, the Plenus Marls are just a few centimetres thick, and have been recorded at Barrett Ringstead [TF 6787 3979] and in the former pit at Heacham [TF 688 368], just to the west of Sheet 130 (Gallois, 1994). In the Bircham Borehole, the Plenus Marls are not clearly visible in the core samples, however, their horizon is inferred from a small inflection on the resistivity log, coincident with a downward change from shelly, coarse-textured chalk to smooth-textured chalk (Fig. 1).

Hard, nodular, shell-rich chalk, equating with Hólywell Nodular Chalk Formation of the 'Southern Province' succession, occurs above the inferred Plenus Marls in the Bircham Borehole. This interval is just over 20 m thick in the Bircham succession, contains sporadic nodular flints, and is clearly represented by a distinctive high resistivity interval on resistivity logs for this borehole and the nearby Manor Farm Borehole at Docking (Fig. 1). Very strongly indurated chalk, represented by a high resistivity spike on borehole resistivity logs (Fig. 1), occurs immediately above the Plenus Marls and probably represents the Melbourn Rock. Flint is not typical of the Holywell Nodular Chalk in southern England, but does occur at this level in parts of East Anglia (Wood, 1996). The shelly chalk is formed of the broken-up remains of the inoceramid bivalve *Mytiloides*, and is also present at the base of the Welton Chalk Formation in north-east England (Whitham, 1991). In north-east Norfolk, the Holywell Nodular Chalk is only c. 7 m thick in the Trunch Borehole, and there is some evidence for thinning of the lower part of the White Chalk Subgroup north of Sedgeford and Docking (see *P. (S.) plana* Zone below).

The higher part of the core in the Bircham Borehole is poorly bioclastic. It contains a thick (c. 50 mm) grey marl seam about 10 m above the top of the Holywell Nodular Chalk, and an old core log for the highest part of the borehole (for which no core survives) records a thick (c. 0.3 m) flint interval. This succession is probably similar to that described in the Swaffham and Thetford districts by Mortimore & Wood (1986). These authors recorded a succession of named marker-marls (based on the work of Ward et al., 1968) between the top of the Holywell Nodular Chalk and the base of the 'Chalk Rock' (a diachronous indurated unit marking the base of the traditional 'Upper Chalk'; see below), with thick flints developed in the interval below the 'Chalk Rock'. There are four principal marl seams, comprising (in ascending stratigraphical order): the Pilgrims Walk Marl, the Mount Ephraim Marl, the Twin Marl and the Grimes Graves Marl (Fig. 1). The lowest two of these probably occur in the Manor Farm and Bircham boreholes, in the south-western part of the Wells district, and most have been identified in the Trunch Borehole. The marls can be recognised by distinctive inflections on borehole resistivity and gamma logs and form an easily correlatable framework (Fig. 1). The succession is very similar to the higher part of the Welton Chalk Formation and basal Burnham Chalk Formation described by Whitham (1991) in north-east England, and marl seam correlations have been established with this and the southern England succession (Mortimore & Wood, 1986; Fig. 1). The thick flint in the higher part of the Bircham core is probably analogous to the flint-rich interval that occurs below the 'Chalk Rock' in the Thetford area, named the Brandon Flint Series (Mortimore & Wood, 1986), and at the junction of the Welton and Burnham Chalk formations described by Whitham (1991).

Figure 1 shows that, based on marl seam correlations, strata equivalent to the New Pit Chalk Formation are likely to be less than 15 m thick in the Wells district, considerably condensed compared to most southern England localities (Table 1, Fig. 1). In contrast, strata equivalent to the lower part of the Lewes Nodular Chalk Formation are relatively expanded, even compared to basinal areas in southern England (Table 1, Fig. 1). Lithologically, the whole of the succession above the base of the Plenus Marls in the Bircham Borehole, can most easily be regarded as Welton Chalk Formation.

The remainder of the White Chalk mostly equates with the Upper Chalk of traditional classification; an interval that is largely masked by Quaternary deposits in the Wells district. Historically, recognition of the base of the Upper Chalk in north Norfolk has been problematic because of poor surface expression of the indurated horizon ('Chalk Rock') traditionally used to map this boundary (Whitaker *et al*, 1893; Whitaker & Jukes-Browne, 1899). Because of these difficulties, biostratigraphical subdivisions have formed the basis for previous lithological discussion of strata equivalent to the middle and higher parts of the White Chalk Subgroup within the district (Jukes-Browne & Hill, 1904; Peake & Hancock, 1970; Wood *et al.*, 1994), and in the absence of any current means of lithological subdivision, this method is adopted herein. Thickness estimates for zones are based on data from Peake and Hancock (1970) for the whole of north Norfolk and the Trunch Borehole (Wood *et al.*, 1994).

Plesiocorys (Sternotaxis) plana Zone (25 to c 37 m thick): Exposures of this interval occur at several localities in the eastern part of the Wells district (Fig. 2). At Bircham Newton [TF 7695 3390] Jukes-Browne & Hill (1904) recorded hard, white chalk with grey, nodular and tabular flints at regular intervals and larger, Paramoudra-like flints. In the Trunch Borehole, the plana Zone succession compares closely with the Burnham Chalk of northeast England and many of the key flints and marls can be correlated between the successions. Locally, in the Wells district, the *P*. (*S.*) plana Zone is coincident with a feature, possibly formed by the 'Chalk Rock' and therefore marking the base of the traditional 'Upper Chalk'. North and north-east of Sedgeford, this feature is only 20 - 30 m above the base of the White Chalk Subgroup, but at least 40-50 m of White Chalk Subgroup probably occur below the *P*. (*S.*) plana Zone in the Bircham and Docking boreholes (Fig. 1). Evidence from the Trunch Borehole suggests that this discrepancy might possibly be explained by rapid northward thinning in the lower part of the White Chalk Subgroup (see Holywell Nodular Chalk above).

Micraster cortestudinarium Zone (c. 20 m thick): There are no known exposures of this zone in the Wells district, but the fauna from this interval in the Trunch Borehole shares features with the succession in north-east England (Wood *et al.*, 1994).

Micraster coranguinum Zone (51 to c. 76 m thick): Mortimore *et al.* (2001) recorded that the chalk of this interval resembled the Seaford Chalk Formation of southern England, and that flinty chalk with marl seams, equivalent to the basal part of the Seaford Chalk, occurred in the Titchwell Parish Pit [TF 762 433] in the north eastern part of the Wells district. Chalk belonging to the *coranguinum* Zone mostly occurs in the central part of the Wells district (Fig. 2), and at South Creake [TF 8553 3625] is 'rather hard chalk with layers of flint' (Jukes-Browne & Hill, 1904). The moderately high diversity of the fauna recorded at South Creake (Woods, 1998) might suggest assignment to the higher (Santonian) part of the *coranguinum* Zone. In the Trunch Borehole there are possible correlatives of named flint marker beds in southern England. This record of flint in the higher part of the *coranguinum* Zone, and its known occurrence in superjacent zones in north Norfolk, contrasts with the complete absence of flint in the partly coeval Flamborough Chalk Formation in north-east England (Whitham, 1993).

Uintacrinus socialis Zone (see *M. testudinarius* Zone for thickness): In north Norfolk, this interval comprises soft chalk, mainly lacking flint (Peake & Hancock, 1970, fig. 3). There is a record of the occurrence of this zone near East Barsham (Fig. 2).

Marsupites testudinarius Zone (28 to 42 m thick, including U. socialis Zone): There are at least four localities in the eastern part of the Wells district where chalk belonging to this zone has been recorded (Fig. 2). The chalk is apparently similar to that belonging to the socialis Zone (Peake & Hancock, 1970), but Jukes-Browne & Hill (1904) noted rather hard testudinarius Zone chalk at Houghton St. Giles [TF 9235 3542]. Faunas from Houghton St

Giles include common oysters, especially *Pseudoperna boucheroni* (Woods, 1998), suggesting that the oyster-rich bioclastic facies that overlies the *testudinarius* Zone in the Trunch Borehole might here occur within this zone.

Uintacrinus anglicus Zone: There is no record of this zone in north Norfolk, although it probably occurs in the southern part of East Anglia and definitely occurs in north-east England (Mitchell, 1995). Any strata that do equate with this zone presumably form the basal part of the overlying 'Zone of *Gonioteuthis*' (sensu Peake & Hancock, 1970).

'Zone of Gonioteuthis' (78 to c. 97 m thick): The chalk assigned to this interval by Peake & Hancock (1970) equates with the (?U. anglicus Zone), O. pilula Zone, and the greater part of the G. quadrata Zone of southern England (Table 1). The highest part of the quadrata Zone, assigned to the 'Basal mucronata Chalk' by Peake & Hancock, 1970 (Peake & Hancock, 2000), occurs outside the Wells district. In the eastern part of the Wells district, 14 localities are assigned to the 'Zone of Gonioteuthis' (Fig. 2), the best exposure being the former quarry at Wells [TF 928 429]. Here, nearly 40 m of succession was visible, equating with the upper O. pilula Zone and lower G. quadrata Zone of southern England. The succession contained flint and marl seams, suggesting similarity to the Newhaven Chalk Formation of southern England (Peake & Hancock, 1970, 2000; Mortimore et al., 2001). A strongly developed marl in the Wells succession has previously been correlated with the Old Nore Marl in the Newhaven Chalk of southern England (Wood et al., 1994), but may in fact be younger, and closer to the top of the *pilula* Zone (Mortimore et al., 2001). In the Trunch Borehole, coarse-grained, bioclastic chalk occurs in the lower part of the *pilula* Zone, and this facies, characterised by common oyster remains, appears to occur at several localities near Wells and also at Great and Little Walsingham (Woods, 1998). Microfossil data from localities near Warham and Stiffkey show that the chalk thereabouts equates with the basal part of the Portsdown Chalk of southern England (Wilkinson, 2004; Table 1).

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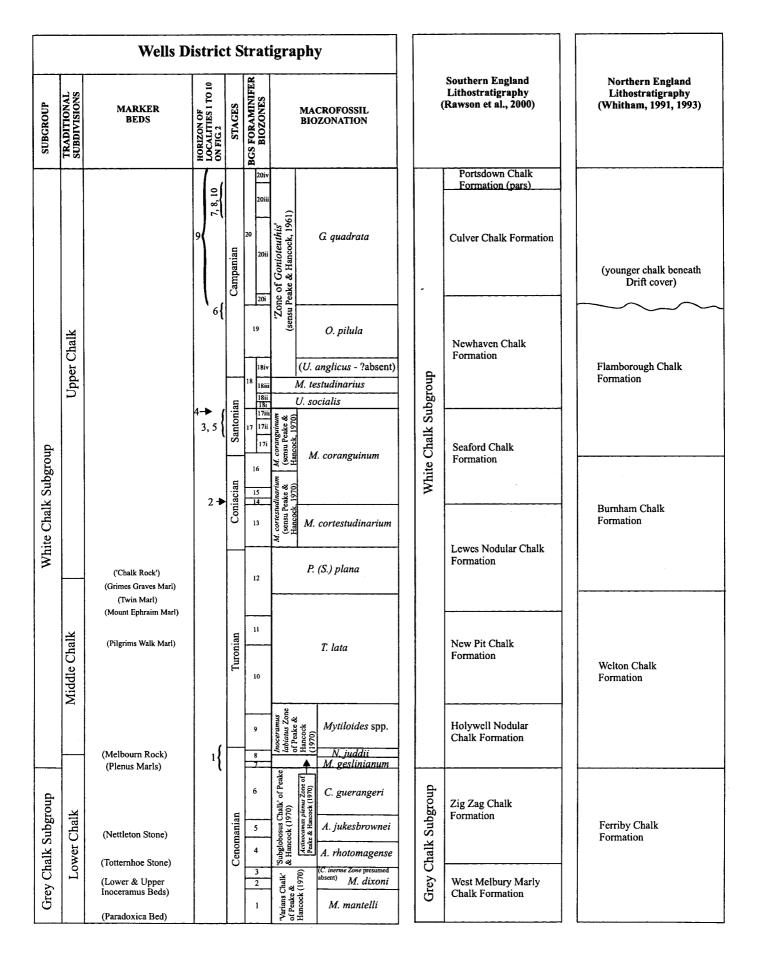
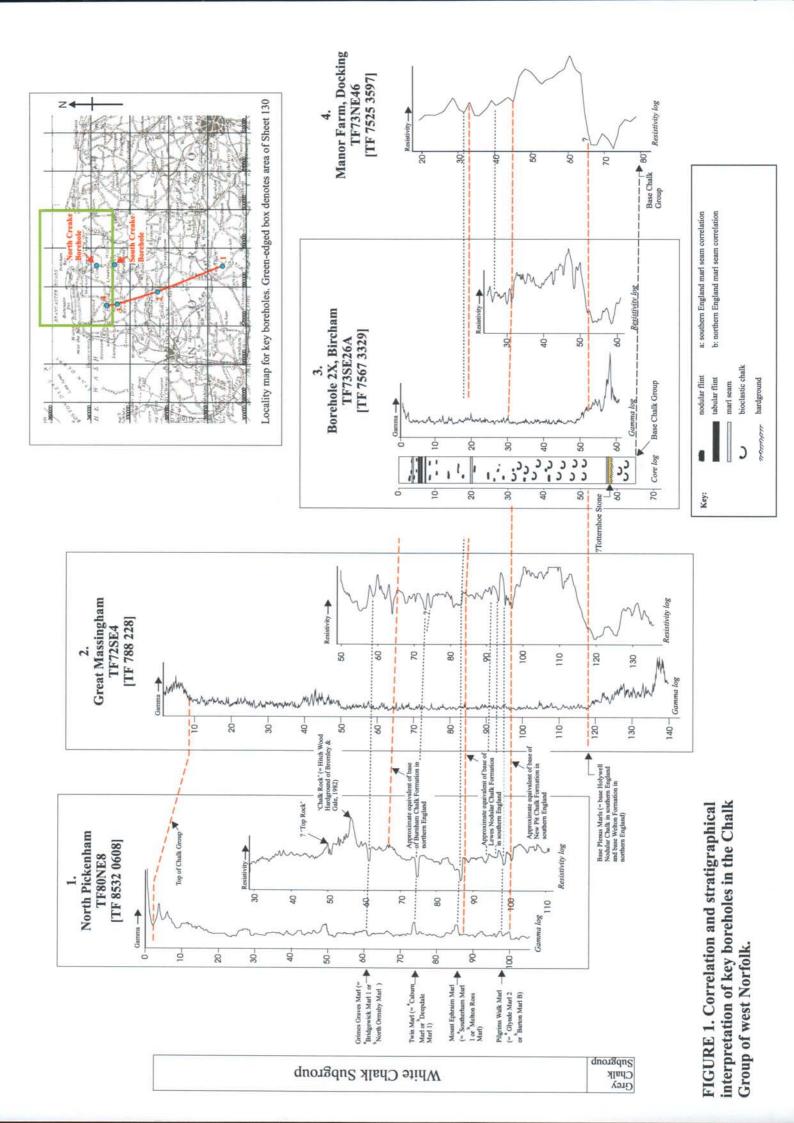


TABLE 1.The stratigraphy of the Chalk Group in the Wells district, and its relationship to
lithostratigraphical subdivisions in southern England and north-east England.



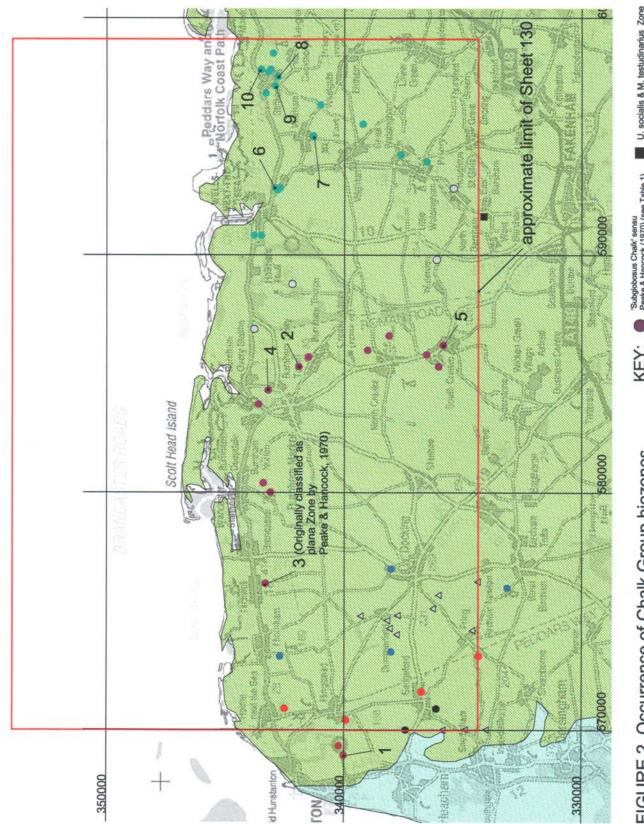




FIGURE 2. Occurrence of Chalk Group biozones in the Wells district. Based on data from N B Peake (BGS archives) and I P Wilkinson (see text).