

1 MULTIDISCIPLINARY CHARACTERISATION AND MODELLING OF A SMALL  
2 UPLAND CATCHMENT IN SCOTLAND

3  
4 Andreas Scheib<sup>a</sup>, Sarah Arkley<sup>b</sup>, Clive Auton<sup>b</sup>, David Boon<sup>a</sup>, Jeremy Everest<sup>b</sup>, Oliver Kuras<sup>a</sup>,  
5 Stephen Pearson<sup>a</sup>, Michael Raines<sup>a</sup>, John Williams<sup>a</sup>

6  
7 <sup>a</sup> British Geological Survey, Kingsley Dunham Centre, Keyworth, Nottingham, NG12 5GG,  
8 UK

9 <sup>b</sup> British Geological Survey, Murchison House, West Mains Road, Edinburgh, EH9 3LA, UK

10  
11 Manuscript received...

12 Revised version ...

13 Scheib A.J., Arkley S., Auton C., Boon D., Everest J., Kuras O., Pearson S., Raines M.,  
14 Williams J., Multidisciplinary characterisation and modelling of a small upland catchment in  
15 Scotland. *Questiones Geographicae*---, Adam Mickiewicz Press, Poznan 2008 pp. ---.

16 **ABSTRACT**

17 A combination of conventional surveying and non-invasive techniques have been applied to  
18 characterising the geomorphology, soils and shallow substrates of a typical small catchment  
19 in the Southern Uplands in Scotland, in three dimensions. Integration of geospatial,  
20 geophysical and geotechnical data, in the resulting digital 3D model, enable the nature and  
21 extent of individual components of the landscape to be measured and their relationships at

22 depth to interpreted and visualised. This type of baseline data is fundamental to understanding  
23 past, and monitoring and measuring the impacts of future environmental changes in  
24 environmentally sensitive areas.

25

26 Keywords: Quaternary landscape evolution, 3D modelling, non-invasive investigation  
27 techniques, shallow geophysics, geomorphology

28 \*Andreas Scheib, British Geological Survey, Kingsley Dunham Centre, Keyworth,  
29 Nottingham, NG12 5GG, UK, email: [ascheib@bgs.ac.uk](mailto:ascheib@bgs.ac.uk)

## 30 INTRODUCTION

31 This study reports on an integrated survey used to investigate the geomorphology, soils and  
32 substrates of part of a typical small mountainous catchment within Upper Tweeddale in the  
33 Southern Uplands of Scotland. The principle aims of the investigation were to (i) develop a  
34 combination of techniques for mapping the landscape in three dimensions and (ii) to establish  
35 a baseline from which the paleoenvironmental evolution of the area could be assessed and  
36 against which monitoring of impacts of future environmental changes could be undertaken.  
37 This required a multidisciplinary approach, where conventional surveying methods (mapping,  
38 augering, drilling and pitting) were combined with more advanced techniques (terrestrial  
39 LiDAR scanning, Electrical Resistivity Tomography [ERT], Ground Probing Radar [GPR]  
40 and Panda penetrometer traverses) to characterise the landscape and its components. The data  
41 from this integrated field methodology was combined within 3D modelling software, enabling  
42 holistic representation of Quaternary landforms and sediments from a remote upland area for  
43 which little previous data existed.

44 The site, now established by the British Geological Survey (BGS) as the Talla Earth  
45 Observatory, is on open grazing land in the Tweedsmuir Hills, 1km SE of Talla Linnfoots  
46 Farm [<sup>3</sup>315 <sup>6</sup>204], 35 km south of Edinburgh (Figure 1). It forms part (3.2 km<sup>2</sup>) of the  
47 catchment of the Talla Water, a small headwater tributary of the River Tweed, which drains  
48 north eastwards a distance of 660m from the NE margin of the site to feed into the Talla  
49 Reservoir, hence the site's environmental sensitivity. The area is characterised by rugged  
50 topography (Figure 2) and dominated by the effects of erosion and deposition by glaciers that  
51 disappeared in the last 10 000 to 13,000 years. The mountain tops rise to elevations of 700 to  
52 800 m above sea level and the study site lies largely within a pass at 400 to 450m elevation,  
53 between the Tweed Valley and Moffat Dale.

54 Although a varied assemblage of relict glacial, periglacial and post glacial deposits and  
55 landform overlies the bedrock in the region (Table 1) there are few modern published  
56 descriptions of their nature and distribution. Two recent studies from the surrounding area  
57 (Chiverrell, Harvey, Foster, 2007; Foster, Chiverrell, Harvey, Dearing, Dunsford, 2008) have  
58 focused on dating Holocene alluvial fan sequences and their use as proxies for  
59 palaeoenvironmental change over the last few thousand years. The proxy climate and  
60 vegetation record of the Talla area itself, for the last 5,500 years, was described by Chambers,  
61 Barber, Maddy, Brew (1997) from monolith samples of peat from the NE portion of the Earth  
62 Observatory area (Figure 1). They recognised that the site has been subjected to a markedly  
63 oscillatory climate during the late Holocene, with several specific episodes of exceptional  
64 wetness having occurred during the last 3.5ka.

65 <Figure 1 and 2>

66 <Table 1>

## 67 **GEOLOGICAL SETTING**

68 The bedrock of the site comprises highly resistant, weakly metamorphosed greywacke  
69 sandstones of the Gala Group (Upper Ordovician – Silurian) of the Southern Uplands of  
70 Scotland (Floyd, 2001). The glacial sediments and weathered bedrock both form bouldery and  
71 cobbly substrates that are difficult to penetrate by traditional percussive drilling methods,  
72 while steep slopes and unstable ground conditions make trial pitting difficult. Where the  
73 glacial debris has been reworked by frost action, or by rivers and streams the periglacial and  
74 post glacial sediments formed tend to be largely composed of cobbles and boulders and their  
75 genesis is often only clearly evident from their geomorphological expression and position  
76 within the landscape (Figure 3).

77 The Observatory site contains many, but not all of the features and sediments typical of an  
78 upland valley in southern Scotland (Table 1). It shows evidence of Weichselian and probable  
79 Younger Dryas glaciation with deposition of till and moraines on lower valley slopes and para  
80 or postglacial mass movements leading to the accumulation of tallus and solifluction lobes  
81 locally. Basin peat has formed in the valley bottom and hill peat flanks the mountain tops;  
82 organic and bouldery soils are developed on all but the steepest slopes.

83 <Figure 3>

## 84 **TECHNIQUES USED TO CHARACTERISE THE TALLA EARTH OBSERVATORY** 85 **SITE**

86 Field investigations at the site took place in three main phases (Table 2). These were followed  
87 by the construction of a 3D model of the study area. The initial phase of field work  
88 established the geology and overall morphology of the site. The second phase produced data  
89 on the nature of the shallow subsurface to c. 1 m depth and more rigorous acquisition of  
90 digital elevation and ground morphology data. This enabled the targeting of the more complex  
91 areas with non-invasive ground investigation techniques in the third phase of the work.

92 <Table 2>

### 93 **First phase surveys**

94 The Quaternary sediments and features within the Talla Earth Observatory site were  
95 recognised using geological and geomorphological mapping techniques. This involved  
96 interpretation of Digital Elevation Models (DEM) and hill shaded Digital Terrain Models  
97 (DTM) derived from NEXTMap British Elevation data (©Intermap Technologies) at 1 m  
98 vertical and 5 m horizontal resolution, as well as mapping from stereoscopic aerial  
99 photography. This provided baseline data that was incorporated into a walkover

100 geological/geomorphological survey of the Observatory site and surrounding area, which is  
101 the basis for the Quaternary map shown in Figure 4.

102 ***Summary of the geomorphological features and deposits of the Talla site***

103 *Glacial features:* The mountain sides have been oversteepened by glaciation, largely during  
104 the Weichselian, giving the valleys a typical ‘u-shaped’ cross section. It is possible that some  
105 further erosion of the slopes occurred during the Younger Dryas.

106 The lower valley slopes are mantled to an elevation of c. 550 m by a variable thickness of  
107 glacial till. It is of uniform clayey-gravel composition, and dominated by clasts of locally-  
108 sourced greywacke with some rare further-travelled coarser sandstones and lavas. The matrix  
109 of the deposit is highly consolidated and its morphology suggests subglacial deposition.

110 The upper reaches of the valley of the Talla Water, and its junction with the east-west pass  
111 between the Talla and Megget reservoirs, are dominated by accumulations of morainic debris  
112 (Figure 4). The forms of the moraines indicate that they were deposited by a small glacier that  
113 retreated toward the headwater area of the Talla Water. The moraines within the upper Talla  
114 Water (Figure 5) are ‘classical’ Scottish hummocky moraine comparable to those described  
115 by Benn (1992), they form a continuous spread of ‘up-valley’ arcuate ridges, 5-8 m in height.  
116 Natural exposures in the deposits are poor but they are composed of similar material to the till  
117 but with a slightly sandier matrix.

118 A small flat-topped spread of glaciofluvial outwash sand and gravel occurs on the northern  
119 bank of the Talla Water close to the centre of the site (Figure 4); more extensive spreads and  
120 larger outwash fans are associated with the course of the Megget Water NE of the  
121 Observatory area.

122 *Periglacial features:* Large areas of inactive bouldery and cobbly solifluction deposits are  
123 present on the upper slopes and on the hill summits. Low relief ridges and solifluction lobes  
124 are present on the southern slope of Fans Law (Figure 4). The solifluction deposits are  
125 composed primarily of frost shattered material, and appear to be derived from remobilisation  
126 of till.

127 Some small areas of inactive talus occur below bedrock exposures. The relict deposit is  
128 composed of angular clasts of greywacke up to 10cm in diameter.

129 *Holocene Features:* Small amounts of talus are actively accumulating beneath the steepest  
130 bedrock exposures and as cones at the mouths bedrock gullies. Relict alluvial fans are present  
131 at the mouths of gullies on northern valley sides above the Talla Water (Figure 6). Features  
132 associated with more recent fluvial activity reflect the small size of the Talla Water burn, with  
133 relatively narrow terraces, composed of cobbles and pebbles; the modern floodplain is  
134 similarly floored with coarse material. The burn flows directly on bedrock in the uppermost  
135 and lowermost reaches (Figure 4).

136 The northern slope of Calavin Hill and parts of the summit of Fans Law are mantled with thin  
137 relatively unhumified hill peat (Figure 4). The centre of the site (Figure 6) is covered by a  
138 waterlogged mire, composed of well humified peat, that extends over >30 hectares and  
139 blankets the underlying moundy topography.

140 <Figure 4 and 5>

## 141 **Second phase surveys**

### 142 *Auger surveys*

143 Auger surveys were conducted across the Talla site, supplemented by observations from the  
144 sparse natural exposures and in places augmented by shallow pits dug with a spade. These

145 provided detailed data on the nature of the shallow subsurface at 191 data points, to depths  
146 that locally exceeded 1 m. The results were used to refine the resolution of the Quaternary  
147 map (Figure 4), to identify areas of thin Quaternary cover on bedrock and identify 13 Soil  
148 Series and 4 composite soil units, which form the basis of the 11 soil types shown on the soil  
149 map of the site (Figure 7).

150 Soils: The surveys showed that thick, commonly waterlogged, fibrous and humified organic  
151 soils are developed on the peat mires in centre of the site (Figure 7). Bouldery sandy clay  
152 loams, are locally developed on the mountain top solifluction deposits, on till and bedrock  
153 within the tributary valleys, and on the moraines flanking the upper reaches of the Talla  
154 Water. Bleached stony sandy loams with ferruginous pans, blanket moraines lower in the  
155 valley. Thin clay loams cap bedrock crags while thicker clay soils are also developed on till  
156 locally. Thin waterlogged gravel soils dominate on the floodplain of the Talla Water.

157 <Figures 6 and 7>

### 158 *Terrestrial LiDAR survey and dGPS*

159 A high resolution topographic survey of the Talla catchment was undertaken using terrestrial  
160 LiDAR (Light Distance And Ranging), commonly referred to as laser scanning. The aim of  
161 the LiDAR survey was the creation of a detailed surface DTM to aid geological and  
162 geophysical data interpretation.

163 The terrestrial laser scanner deployed was a Riegl LPM-i800HA (inset, Figure 8), with an  
164 effective range up to 800 m at an accuracy of  $\pm 25$  mm. The scanner measures and records  
165 relative distance, and vertical and horizontal angle between the instrument and the target  
166 surface for each point, with a typical measurement rate of 1000 pts/sec. This scanner also  
167 incorporates a high-resolution digital camera, enabling true-colour information for surface  
168 point data, or calibrated digital photo images to be captured. The laser scan technique relies

169 on the fact that the laser beam reflects from target surfaces. The intensity (strength) of the  
170 returned laser beam is recorded, which varies according to the reflectivity of the surface  
171 material being measured. The method is not effective where the subject is moving (e.g.  
172 water), not visible or where the laser beam is reflected by heavy rain or fog. The principal  
173 advantage of the LiDAR surveying technique is the speed, resolution and accuracy of data  
174 acquisition (Buckley, Howell, Enge, Kurz, 2008). It also enables data capture in terrain that is  
175 difficult to access (e.g. steep valley sides). For the Talla survey, multiple scans were taken  
176 from different aspects of the site in order to ensure almost complete terrain coverage, creating  
177 a higher resolution and more complete DTM for the site in less time than it would be possible  
178 by traditional surveying.

179 Key to the success of laser scanning is the accurate horizontal and vertical positioning of the  
180 instrument and other reference points. This is achieved with a dual frequency GPS, which  
181 enabled accurate orientation of the data to British National Grid coordinates. The altitude and  
182 location of the scan and reference point positions were measured using a Leica SR530  
183 differential Global Positioning System (dGPS) corrected to Ordnance Survey National Grid  
184 and Ordnance Datum (OD) (Newlyn). The dGPS was deployed in Real Time Kinematic  
185 mode, allowing positional corrections to be received in real-time via radio link from a  
186 reference receiver. Using this system, altitude data was produced that is accurate to within  
187 0.03 m, and 0.02 m in plan (The Survey Association, 2008). The terrestrial LiDAR data  
188 consists of 'point-clouds' comprising x, y, z point coordinates and intensity values. The raw  
189 point cloud data were input into surface modelling software to produce a solid high-resolution  
190 3D terrain model of the Observatory site.

191 <Figure 8>

192 Figure 8 shows a 3D multiple-scan point-cloud image, shaded by intensity which enhances  
193 the recognition and interpretation of specific features. The image was taken from the southern  
194 side of the catchment, adjacent to Molls Cleuch Burn, looking north. The Burn can be seen in  
195 the left foreground, as a low intensity (black) feature where the laser signal was not returned  
196 due to running water. Another, low intensity area can be seen in the Talla Moss blanket peat  
197 area in the centre right of the figure, also associated with surface water. The medium intensity  
198 (dark grey) area in the centre images the alluvium adjacent to the main channel of Talla  
199 Water. In contrast, lower intensity (light grey) areas on the distant slopes depict the position  
200 of solifluction lobes on the southern flank of Fans Law.

201 Holes in the data for the LiDAR solid surface model occurred at localities where data were  
202 not captured due to the presence of surface water. In order to create an accurate and complete  
203 DTM, the LiDAR DTM was patched with data from the lower-resolution NEXTMap Digital  
204 Surface Model. It was not possible to simply splice the NEXTMap data into the LiDAR  
205 dataset, as it created too many data artefacts, resulting in a noticeably stepped terrain model.  
206 However, individually warped patches of the NEXTMap model were patched onto the  
207 LiDAR data, by creating a correction algorithm that compared elevations between the models  
208 at the margins of the holes in the LiDAR dataset. Applying the correction to the NEXTMap  
209 data created a best-fit join between the models at the margins of the data holes, with the effect  
210 of the warping decreasing towards their centres. This blended the NEXTMap patches into the  
211 LiDAR DTM effectively, to produce a 'smooth' terrain model. This DTM forms the surface  
212 topography of the 3D model of the Talla Earth Observatory and is also the basis of the DEM  
213 of the Talla Linnfoots area (Figure 2).

### 214 **Third phase surveys**

215 Invasive and non-invasive ground investigation techniques were used to investigate and  
216 characterise the shallow subsurface in detail. The invasive methods employed (trial pitting,  
217 shallow percussive drilling and light-weight cone penetration testing) were integrated with the  
218 results from shallow geophysical surveys. Synthesis of the results allowed visualisation and  
219 quantitative analysis of the physical properties of subsurface materials, while cross-calibration  
220 between different types of dataset increased confidence in the final geological interpretation  
221 and modelling of the subsurface conditions at the site.

### 222 *Near-surface geophysical surveys*

223 Geophysical techniques were used to elucidate the form and distribution of the shallow  
224 subsurface deposits in 3D, to identify locations for more detailed invasive investigations, and  
225 to obtain volumetric information about the nature of strata beyond depths that could be  
226 investigated by pitting and drilling.

227 *Electrical Resistivity Tomography (ERT):* As the vertical resolution of ARP techniques are  
228 limited, 2D ERT was employed to provide detailed geophysical cross-sections at key  
229 locations across the Observatory site. 2D ERT is capable of producing tomographic images of  
230 the spatial distribution of bulk subsurface resistivity (as opposed to apparent resistivity) along  
231 linear survey profiles (e.g., Dahlin, 2001). Modern data processing and inverse modelling  
232 algorithms allow for the integration of steep topography and complex geomorphology  
233 typically encountered on upland sites such as at Talla. ERT surveys were carried out on five  
234 major profiles (TAL-A, B, D, F and G) located across the valley at different elevations in the  
235 catchment (Figure 7). Specifically designed multi-channel measurement sequences based on  
236 the Wenner-Schlumberger array geometry were used for data acquisition (Chambers, Kuras,  
237 Meldrum, Ogilvy, Hollands, 2006), resulting in robust datasets with excellent lateral and  
238 vertical resolution, showing resistivity variations across nearly four orders of magnitude.

239 Figure 10a shows the resulting 2D resistivity model of Profile TAL-D as an example. The  
240 profile crosses peat deposits in the centre of Observatory site. The electrical signature of the  
241 waterlogged peat is a conductive layer (shown in blue) with resistivities of below 500  $\Omega\text{m}$ .  
242 The interpreted form of the base of the morainic deposits, seen between profile distances  
243 170 m and 550 m coincides with resistivities above c 5000  $\Omega\text{m}$ .

244 *Automated resistivity Profiling (ARP)*: Due to its sensitivity to water content and variations in  
245 lithology, electrical resistivity was seen as a key geophysical parameter to differentiate  
246 between various types of Quaternary deposits. To provide baseline resistivity across a large  
247 part of the Observatory site, an ARP survey was undertaken (Besson, Cousin, Dabas, Biozard,  
248 Richard, 2005). A towed survey system capable of making measurements of the apparent soil  
249 resistivity across three different depths of investigation was deployed. An ARP survey  
250 involves towing a mobile electrical sensor array across the site using an all-terrain vehicle  
251 (Figure 9). Dynamic positioning was achieved via integrated RTK-GPS and measurements  
252 were processed on board, with the operator guided by a real-time GIS system. The  
253 methodology allowed the production of preliminary digital maps of apparent resistivity (geo-  
254 referenced bitmaps) for approximately 13 hectares within a few hours, at a lateral resolution  
255 of approximately 15 m for the final maps (Figure 10b).

256 The maps show resistivity distributions at 0 - 0.5 m, 0 - 1.0 m and 0 - 1.7 m depths, that  
257 reflect the mapped geology, in particular the extent of the peat moss (shown by low resistivity  
258 values in blue) and the location of river terrace gravels (high resistivity values in red) at  
259 several locations immediately adjacent to the northern bank of the Talla Water (Figure 10b).  
260 However, much additional detail was obtained that reflects the nature of subsurface features,  
261 notably the form and extent of moraines concealed beneath the peat. They are picked out by  
262 high resistivity values that begin to be apparent in 0 - 0.5 m dataset, but are better defined in

263 the deeper datasets, which show that the peat basin is bisected by a buried N-S trending  
264 morainic ridge.

265 <Figure 9 and Figure 10>

266 *Ground Penetrating Radar (GPR)*: GPR is used to investigate the subsurface by penetration  
267 and reflection of high-frequency electromagnetic waves in the ground (Davis and Annan,  
268 1989). In near-surface investigations GPR can provide a high degree of vertical and horizontal  
269 resolution, ideally complementing the ERT technique (Slater and Reeve, 2002), as they are  
270 each controlled by contrasts in different electrical parameters (resistivity and permittivity,  
271 respectively).

272 The suitability of GPR surveying in peatlands was initially reported Theimer, Nobes, Warner,  
273 (1994). Their study found that the principal peatland interfaces detected were the near-surface  
274 aerobic to anaerobic transition within the peat and the peat to mineral basement contact. More  
275 recently, Comas, Slater, Reeve (2004, 2005) have used GPR to determine peat basin  
276 morphology and the stratigraphic controls on pool formation in a domed bog, whilst Sauer  
277 and Felix-Henningsen (2004) used it to measure the thickness of Pleistocene periglacial slope  
278 deposits. At the Talla site, GPR has been used in conjunction with the ARP and ERT  
279 techniques, mentioned above, to characterise the vertical and lateral extent of peat and the  
280 underlying glacial sediments.

281 A Pulse Ekko IV system (Sensors and Software Ltd) with 100 MHz antennae was used to  
282 map the base of peat over traverses TAL-A and TAL-D. It accurately delineated peat  
283 thickness variations and added near-surface detail enhancing interpretation of 2D ERT  
284 sections. The GPR traverses were later calibrated by Panda penetrometer profiles (see below).

285 ***Trial pitting and drilling***

286 Eleven trial pits were excavated using a tracked, low ground-pressure excavator, capable of  
287 traversing both steeply inclined ground and areas of unstable water-saturated peat. Pit sites  
288 (Figure 7) were concentrated at the ends of 3 of the 5 ERT traverses, as well as on the valley  
289 sides and the margins of the blanket mire (where the geological and geophysical surveys  
290 indicated that the surface peat was generally < 0.5 m thick). In all instances the pits enabled a  
291 direct evaluation of the nature of Quaternary sequence to 2-3 m depth. Bouldery or gravelly  
292 glacial and periglacial deposits were encountered in 10 of the pits, gravelly relict alluvial fan  
293 deposits, with buried peat lenses containing wood fragments, were encountered in the other;  
294 all provided samples that enabled calibration of the geophysical signatures of the deposits by  
295 their geotechnical and sedimentary characteristics.

296 A shallow cored borehole was sunk on the floodplain of the Talla Water and 4 others were  
297 drilled within the areas of thickest peat interpreted from the ARP and ERT data. Drilling was  
298 undertaken, using a light weight Dando Terrier drilling rig. The drilling equipment was  
299 unable to reach rockhead, but the boreholes provided data on the nature and thickness of the  
300 postglacial sediments where pitting was impossible. The boreholes are now instrumented to  
301 provide real-time measurements of shallow groundwater level fluctuations within the lower  
302 portions of the Talla site.

303 ***Light-weight penetrometer***

304 Geotechnical techniques are well known for enhancing characterisation of glacial sequences  
305 (Gerrard, 1981) and understanding the formation of peat bogs (Moore and Bellamy, 1974;  
306 Hobbs, 1986). Characterisation of peat types and peat stratification is aided by using  
307 properties, such as undrained shear strength values, as a proxy for amount of compaction and  
308 consistency of the sediments. Although it does not provide measurements of undrained shear

309 strength, the Panda ® Dynamic Cone Penetrometer is a tool commonly used for compaction  
310 control testing and for investigating natural soils in geotechnical site investigations (Langton,  
311 1999). It provides vertical profiles of cone resistance, which relate to the variations in strength  
312 and density of superficial materials.

313 The penetrometer apparatus weighs around 20 kg and can be operated by one person making  
314 it a useful rapid, low-cost probing of soils on sites where terrain is restrictive. The apparatus  
315 measures the frictional resistance of a soil, measured in 'qd' ( $1\text{qd} = 1\text{MPa}$ ). It works by  
316 driving a 2 cm diameter cone into the ground to depths of up to 6m. Penetrometer  
317 measurements have been previously used to characterise cohesive glacial till deposits (Gunn,  
318 Pearson, Chambers, Nelder, Lee, Beamish, Busby, Tinsley R., Tinsley W., 2006) but, at the  
319 Talla site, they were used to characterise the stratigraphy of the peat deposits, alluvium and  
320 glacial deposits across the low lying portions of the area. The technique is restricted to 'very  
321 soft to very-stiff' fine grained soils (as defined by British Standard 5930:1999) and it was  
322 physically unable to penetrate very densely packed sand and gravels or competent cobbles.

323 At Talla penetrometer profiles were taken at 20 m intervals along the geophysical lines and  
324 also at borehole locations. The profiles all intercept a highly resistive boundary between 0.2  
325 and 2.5 m characterised by a sudden increase in the cone resistance. Comparison with the  
326 borehole logs indicates this response is commonly due to a change in resistance at the base of  
327 the topsoil. Similar minor variations in the cone resistance at greater depths equate with  
328 lithological changes with in the peat and alluvium as illustrated in Figure 11c, d and e. The  
329 interpolation of peaks above 1MPa, indicating materials with the engineering strength of rock,  
330 provides a continuous cross section of the base of peat deposits.

331 *Comparison of Panda, GPR and ERT data:* The GPR results (Figure 11a and b) were  
332 validated by comparison with Panda penetrometer profiles and borehole descriptions (Figure

333 11c and e). The strength profiles and borehole logs were overlaid onto the geophysical  
334 pseudo-sections to provide a validation check for the estimated velocity parameters used in  
335 the geophysical inversion. The peaks in the penetrometer data correspond well to reflectors in  
336 the GPR image indicating that the GPR can be used to map the base of the peat in detail  
337 across the section. The resulting geological sequence (Figure 11d) shows lateral and vertical  
338 variability of these deposits.

339 <Figure 11>

340 Calibrated radargram (Figure 11a) also proved the existence of three distinct peat-filled  
341 palaeochannels (c. 50 m wide and 2-3.5 m deep), cut into the underlying moraine ridge  
342 (labelled as major peat accumulations). The most northerly palaeochannel (Figure 10b) shows  
343 a 2 m thick upper reflective layer related to fibrous peat overlying a basal, more amorphous  
344 layer (Figure 11c). This suggests two peat layers of different water content, which is due to  
345 the variation in the density and its degree of humification (Leopold and Volkel, 2003). The  
346 aerobic to anaerobic contact between growing sphagnum moss and underlying peat, appears  
347 to be marked by the large amplitude reflector at a depth of 0.6m (between 490 to 520 m)  
348 caused by the peat's high level of water saturation. A comparable strong reflection, caused by  
349 a decrease in water content indicates the contact of the peat with the underlying unsaturated  
350 morainic deposits.

## 351 **3D MODELLING AND VISUALISATION**

### 352 **Methodology**

353 The modelling package used in this study is Geological Surveying and Investigation in 3D  
354 (GSI3D) software, which has been co-developed by BGS and INSIGHT GmbH since 2001.  
355 The GSI3D methodology is easy to use, simple and intuitive and its success is based on the

356 fact that it utilises exactly the same data and methods, albeit in digital forms, that geologists  
357 have traditionally used to make geological maps and cross-sections (Kessler and Mathers,  
358 2004).

359 Basic data formats used in GSI3D are geo-registered raster images such as topographic base  
360 maps and air photos, DTM loaded as standard ASCII grid file, digital borehole data,  
361 geological map data loaded as ESRI shape files and geo-registered planar vertical sections and  
362 horizontal slices. The backbone of the software is the GVS text file, which contains all  
363 geological units in their correct and unique stratigraphical order, which defines the “stack”  
364 that is calculated to make the 3D geological model. A more detailed description of the  
365 software methodology is given in Kessler, Mathers, Sobisch (2008). In short, the software  
366 normally combines a DTM, geological surface linework and downhole (borehole, trial pit,  
367 auger hole) geological data to enable construction of regularly spaced intersecting cross  
368 sections, by correlating boreholes and the outcrops-subcrops of units, to produce geological  
369 fence diagrams. The Talla study combines the results of a suite of geophysical and  
370 geotechnical surveys to produce additional interpreted cross-sections to constrain the model  
371 beyond the depth reached by the more usual methods of ground investigation. Mathematically  
372 interpolating between nodes along the geological, geophysical and geotechnical cross-  
373 sections, together with the mapped surface extent of the units, produces a 3D model. This is  
374 built from a series of stacked Triangulated Irregular Networks (TINs), corresponding to the  
375 top and basal surface of each geological unit.

### 376 **Talla 3D Model**

377 Constructing the model initially involved collating the main geological, geophysical and  
378 geotechnical datasets described above, and ensuring all of the data were in a compatible geo-  
379 referenced digital format for import into the GSI3D software. This was a relatively simple

380 task for many of the data sets collected during the first and second phases of the study. Most  
381 of the geological log data (from auger holes, trial pits and boreholes) were recorded digitally  
382 in the field on tablet pc's with integral GPS positioning systems. The data was loaded into  
383 BGS customised Microsoft Access 2003 geodatabases and linked to location and cartographic  
384 data generated in a customised ArcMap 9.1 © ESRI based GIS. The methodology is described  
385 in more detail in Jordan, Bee, Smith, Lawley, Ford, Howard, Laxton (2005). During survey of  
386 the third phase, all of the geophysical field data and most of the geotechnical were recorded  
387 digitally. Accurate locations, which were required for the import of these traverses and data  
388 points into GSI3D, were provided by dGPS.

389 The GSI3D model of Talla incorporated: a merged high-resolution Lidar/NextMap DTM,  
390 black and white aerial photography, British Ordnance Survey 1: 10 000 scale topographic  
391 data, 1:10,000 scale geological mapping of bedrock and superficial deposits, logs of auger  
392 holes, boreholes and trial pits, geophysical profiles and maps, and Panda penetrometer  
393 profiles. Screen shot in Figure 12 shows the model construction in GSI3D with its three  
394 interactive map and section windows in which the user draws the intersecting cross-sections  
395 and envelope units. By combing the information from the cross-sections and the envelopes,  
396 the software calculates the volumes of each geological unit and stacks them in the correct  
397 stratigraphical order. The methodology on incorporating geophysical data in GSI3D is  
398 described in more detail in Williams and Scheib (2009).

399 <Figure 12>

## 400 **Results**

401 The geological units modelled include (in stratigraphical order) alluvium, peat, talus, head,  
402 alluvial terraces, alluvial fans, glacio-fluvial deposits, moraine and glacial till. The model is  
403 particularly successful for characterising and visualising the extent and thickness of these near

404 surface deposits. The best results were obtained when the data from geological, geophysical  
405 and geotechnical transects were combined to provide the interpretation along given sections  
406 (such as TAL-D, Figure 10a). Figure 12b shows geophysical sections along with correlated  
407 geological sections in a 3D fence diagram. They were also used in the section window (Figure  
408 12c), where the modeller used the resistivity profile as guidance to correlate a more accurate  
409 and refined extent of peats (low resistivity) and gravelly morainic deposits (high resistivity).  
410 This integrated approach produced the highest degree of confidence in the modelled accuracy  
411 of the depth and extent of the Quaternary successions and enabled better interpretation of the  
412 hidden geology.

413 The GSI3D package also produced realistic areal extents and calculated volumes for several  
414 of the shallow deposits. For example, the areal extent of peat within the model area is 406,000  
415 m<sup>2</sup>, the calculated volume is 569,444 m<sup>3</sup>, giving the deposit an average thickness of 1.4 m;  
416 comparable figures for the morainic deposits were 1,088,000 m<sup>2</sup>, 6,770,776 m<sup>3</sup> and 6.2 m.  
417 These appear to be geologically realistic values, based on empirical observations of the  
418 features and deposits in the field.

419 The modelling was less successful constraining rockhead at depth. It was limited by a lack of  
420 direct geological evidence (from drilling and pitting) in the central parts of the observatory  
421 site, to calibrate geophysical interpretation of rockhead beneath the thickest parts of the  
422 Quaternary sequence. This is a consequence of a combination of factors, particularly the  
423 ground conditions (thick waterlogged blanket mire), which precluded the use of heavy weight  
424 drilling equipment and limited mechanical excavations to the flanks of the mire. Nevertheless,  
425 the geophysical results (particularly from the ERT and GPR surveys) combined with  
426 inferences of bedrock morphology at depth, from evidence at the more peripheral localities,  
427 provide a much improved estimation of rockhead across the site area than is possible from  
428 'traditional' walkover surveys alone.

## 429 **DISCUSSION AND OUTLOOK**

430 The Talla Observatory is remote and environmentally sensitive, with limited access possible  
431 for heavy equipment. It is however typical of many upland catchment sites in northern  
432 Britain, in both the diversity of landforms and sediments that it contains. Consequently, to  
433 characterise the area so that its environmental baselines could be established, an innovative  
434 combination of conventional geological, geomorphological, mapping and invasive point-  
435 measurement techniques, were employed, in conjunction with shallow geophysical methods.  
436 Invasive techniques, such as drilling and pitting, had to be used sparingly, as both are  
437 destructive and the ground conditions meant that they provide an incomplete characterisation  
438 of the subsurface (Lapen, Moorman, Price, 1996). The integration of these 'traditional'  
439 methods with shallow geophysical techniques (GPR, ARP, ERT) and geotechnical  
440 measurements enabled a more complete evaluation of the extent and thickness of the deposits  
441 than would otherwise be possible. The use of terrestrial LiDAR scanning with remotely  
442 sensed NEXTMap data also provide a high accuracy DTM that enabled detailed and accurate  
443 modelling of the site as a whole as well as characterisation of individual landform elements.

444 When the survey data were combined in a 3D geological modelling and visualisation package,  
445 a unique picture emerged of the relationship between the landscape and the glacial and  
446 postglacial sediments at the Talla Observatory site. This provides a rigorous framework for  
447 recording its past environmental evolution and monitoring its future environmental changes.

448 Some of this work has already begun.  $^{14}\text{C}$  dating of plant macro fossils from peat cores taken  
449 from the deepest parts of Talla Moss has already been undertaken and an extensive  
450 programme of analyses pollen and insect fossil assemblages are underway. Cosmogenic  $^{10}\text{Be}$   
451 dating of erratic boulders from the moraines being used to establish the date of the last  
452 glaciation of the site and monitoring of shallow groundwater levels linked to precipitation

453 records has commenced. Further investigations of the permeability of the glacial sediments by  
454 means of Guelph permeameter measurements are planned, with the eventual aim of  
455 constructing a quantified model of the shallow groundwater component of baseflow to the  
456 hydrology of the Talla Water Burn.

#### 457 **ACKNOWLEDGEMENTS**

458 The authors would like to thank the landowner Dr. R. Glatt from Hearthstanes Estates for  
459 granting access to the Talla site. This paper is published with the permission of the Director of  
460 the British Geological Survey (NERC).

461

462 **REFERENCES**

463 Ballantyne, C. K., 1984: The Late Devensian periglaciation of upland Scotland. Quaternary  
464 Science Reviews, 3, 311-343.

465

466 Ballantyne, C.K., 1987: The present-day periglaciation of upland Britain. [In:] Boardman, J.,  
467 (ed), Periglacial processes and landforms in Britain and Ireland. Cambridge University Press,  
468 Cambridge, 113-126.

469

470 Benn, D.I., 1992: The genesis and significance of 'hummocky moraine': evidence from the  
471 Isle of Skye, Scotland. Quaternary Science Reviews, 11, 781-799.

472

473 Besson, A., Cousin, I., Dabas, M., Boizard, H., Richard, G., 2005: Electrical resistivity  
474 feasibility to characterize structural heterogeneity of cultivated soils: from 1D to 3D.  
475 Geophysical Research Abstracts, 7, 02826.

476

477 Buckley, S.J., Howell, J.A., Enge, H.D., Kurz, T.H., 2008: Terrestrial laser scanning in  
478 geology: data acquisition, processing and accuracy considerations. Journal of the Geological  
479 Society, London, 165, 625-638.

480

481 Chambers, F.M., Barber, K.E., Maddy, D., Brew, J., 1997: A 5500-year proxy climate and  
482 vegetation record from blanket mire at Talla Moss, Borders, Scotland. *The Holocene*, 7, 391-  
483 399.

484

485 Chambers, J.E., Kuras, O., Meldrum, P.I., Ogilvy, R.D., Hollands, J., 2006: Electrical  
486 resistivity tomography applied to geologic, hydrogeologic, and engineering investigations at a  
487 former waste-disposal site. *Geophysics*, 71, B231–B239.

488

489 Chiverrell, R.C., Harvey, A.A., Foster, G.C., 2007: Hillslope gullying in the Solway Firth –  
490 Morecambe Bay region, Great Britain: Responses to human impact and/or climatic  
491 deterioration? *Geomorphology*, 84, 317-343.

492

493 Comas, X., Slater, L., Reeve, A., 2004: Geophysical evidence for peat basin morphology and  
494 stratigraphic controls on vegetation observed in a Northern Peatland. *Journal of Hydrology*,  
495 295, 173-184.

496

497 Comas, X., Slater, L., Reeve, A., 2005: Stratigraphic controls on pool formation in a domed  
498 bog inferred from ground penetrating radar. *Journal of Hydrology*, 315, 40-51.

499

500 Dahlin, T., 2001: The development of DC resistivity imaging techniques. *Computers &*  
501 *Geosciences*, 27, 1019-1029.

502

503 Davis, J.L., Annan, A.P., 1989: Ground-penetrating radar for high-resolution mapping of soil  
504 and rock stratigraphy. *Geophysical Prospecting*, 37, 531–551.

505

506 Floyd, J.D., 2001: The southern Uplands Terrane: a stratigraphical review. *Transactions of the*  
507 *Royal Society of Edinburgh: Earth Sciences*, 91, 349-362.

508

509 Foster, G.C., Chiverrell, R.C., Harvey, A.A., Dearing, J.A., Dunsford, H., 2008: Catchment  
510 hydrogeomorphological responses to environmental change in the Southern Uplands of  
511 Scotland. *The Holocene*, 18, 935-950.

512

513 Gerrard, A.J., 1981: *Soils and landforms*. George Allen & Unwin, London. 219pp.

514

515 Gunn, D.A., Pearson, S.G., Chambers, J.E., Nelder, L.M., Lee, J.R., Beamish, D., Busby, J.P.,  
516 Tinsley, R.D. and Tinsley, W.H. 2006: An evaluation of combined geophysical and  
517 geotechnical methods to characterize beach thickness. *Quarterly Journal of Engineering*  
518 *Geology and Hydrogeology*, 39, 339-355.

519

520 Hobbs, N.B., 1986: Mire morphology and the properties and behaviour of some British and  
521 foreign peats. *Quarterly Journal of Engineering Geology*, London, 19, 7-80.

522

523 Jordan, C.J., Bee, E.J., Smith, N.A.S., Lawley, R.S., Ford, J., Howard, A.S., Laxton, J.L.,  
524 2005: The development of Digital Field Data Collection systems to fulfil the British  
525 Geological Survey mapping requirements. Proceedings of the International Association of  
526 Mathematical Geology 2005: GIS and Spatial Analysis, Toronto, 2, 886-891.

527

528 Kessler, H., Mathers, S.J., 2004: Maps to Models. *Geoscientist*, 14/10, 4-6.

529

530 Kessler, H., Mathers, S., Sobish, H.-G., 2008: The capture and dissemination of integrated  
531 geospatial knowledge at the British Geological survey using GSI3D software and  
532 methodology. *Computers & Geosciences*, doi:10.1016/j.cageo.2008.04.005.

533

534 Langton, D.D., 1999: The Panda lightweight penetrometer for soil investigation and  
535 monitoring material compaction. *Ground Engineering*, 32, 33-37.

536

537 Lapen, D.R., Moorman, B.J., Price, J.S., 1996: Using ground-penetrating radar to delineate  
538 subsurface features along a wetland catena. *Soil Science Society of America Journal*, 60, 923-  
539 931.

540

541 Leopold, M., Volkel, J., 2003: GPR images of periglacial slope deposits beneath peat bogs in  
542 the Central European Highlands, Germany. [In:] Bristow, C.S., Jol, H.M., (eds), *Ground  
543 Penetrating Radar in Sediments*. Geological Society, London, Special Publications, 211, 181-  
544 189.

545

546 Moore, P.D., Bellamy, D.J., 1974: Peatlands. Paul Elek, London.

547

548 Sauer, D., Felix-Henningsen, P., 2004: Application of ground penetrating radar to determine  
549 the thickness of Pleistocene periglacial slope deposits. *Journal of Plant Nutrition and Soil*  
550 *Science*, 167, 752-760.

551

552 Slater, L.D., Reeve, A., 2002: Investigating peatland stratigraphy and hydrogeology using  
553 integrated electrical geophysics. *Geophysics*, 67, 365-378.

554

555 Theimer, B.D., Nobes, D.C., Warner, B.G., 1994: A study of the geoelectrical properties of  
556 peatlands and their influence on ground penetrating radar surveying. *Geophysical*  
557 *Prospecting*, 42, 179-209.

558

559 The Survey Association, 2008: Best Practice Guidance notes for Network RTK surveying in  
560 Great Britain. Download at <http://www.tsa-uk.org.uk/>.

561

562 Williams J.D.O., Scheib A.J., 2009: Application of near-surface geophysical data in GSI3D:  
563 case studies from Shelford and Talla Linnfoots. British Geological Survey, 29pp.  
564 (OR/08/068). [http://nora.nerc.ac.uk/5347/1/OR\\_08\\_068.pdf](http://nora.nerc.ac.uk/5347/1/OR_08_068.pdf)