

Diatoms modify the relationship between dissolved silicon and bicarbonate in impounded rivers

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ABSTRACT

In order to investigate the relationship between dissolved silicon (DSi) and bicarbonate (HCO_3^-), we analysed water samples from the main Changjiang channel and its main tributaries in August 2006 and, during 2007-2009, the contrasting tributaries Wujiang (WJ) – mainly dominated by carbonate weathering, and Ganjiang – mainly dominated by silicate weathering. The DSi: HCO_3^- ratio ranged from 0 to 0.67, which is in agreement with mixing between the weathering of carbonate or silicate. A negative correlation between DSi and HCO_3^- was observed and interpreted as the imprint of regional geology on water chemistry. This relationship and the DSi: HCO_3^- ratios illustrate the predominant role of carbonate weathering on the riverine HCO_3^- , even for catchments where silicate rocks are dominant. In contrast, a cascade of dams in WJ tributary influenced the concentrations of DSi and HCO_3^- and resulted in a positive correlation between DSi and HCO_3^- . This is because reservoirs allowed populations of diatoms to develop and diatoms stoichiometrically sequestered DSi and HCO_3^- causing a negative feedback regulation on the DSi: HCO_3^- ratio. Our study demonstrates that the relationship between DSi and HCO_3^- can reflect not only their geological background but also the influence of algal activity (diatom uptake) within rivers.

Key words: chemical weathering, dam effect, diatoms, negative regulation, Changjiang river.

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INTRODUCTION

Continental minerals comprise elements in different stoichiometric ratios. This is the basis of using riverine dissolved solutes to understand mineral chemical weathering. Dissolved silicon (DSi) only originates from silicate weathering, while bicarbonate (HCO_3^-) is mainly derived from weathering of silicate and carbonate (Berner *et al.*, 1983; Berner, 1992; Gaillardet *et al.*, 1999; Mortatti and Probst, 2003). Therefore, the critical stoichiometric ratio of DSi to HCO_3^- is predictable for a river with a certain geological background.

River is an important passage for transporting matter from land to sea. Human activities such as damming influence riverine transport function. With the increase of dam construction, primal rivers evolve into impounded rivers. Phytoplankton is often the main primary producer in impounded rivers. Algae assimilate nutrient elements from their habitats in characteristic stoichiometric ratios (Sternler and Elser, 2002) and the elemental composition of a phytoplankton assemblage depends on their species composition (Quigg *et al.*, 2003).

Diatoms are one of the most important algal groups in rivers and have an obligate requirement for silicon as their

cell wall contains this element. The average ratio of silicon to carbon is 0.79 for freshwater diatoms (Sicko-Goad *et al.*, 1984) and 0.13 for marine diatoms (Brzezinski, 1985). Diatom elemental stoichiometry links the biogeochemical cycles of Si and carbon (C) and thus the water ratio of DSi to HCO_3^- will be influenced by diatom uptake of DSi and dissolved inorganic carbon (DIC). Furthermore, the widespread construction of dams has changed riverine aquatic ecosystems (Conley *et al.*, 1993; Humborg *et al.*, 1997; Wang *et al.*, 2008) and increased the potential for phytoplankton growth, consequently increasing their potential effect on elemental geochemical cycles.

Although the Si and C cycles are linked biologically and geochemically in river-reservoir systems, so far little attention has been paid to the Si:C ratio. However, the stoichiometry in rivers of other elements such as Ca:Na, K:Na and Cl:Na are widely used as markers of chemical weathering processes (Gaillardet *et al.*, 1999; Han and Liu, 2004; Chetelat *et al.*, 2008). Therefore, in this study we present the hypothesis that DSi and HCO_3^- are linked in river-reservoir system, this biogeochemical linkage reflecting not only their external geological resource but also the influence of algal activity within the river.

The Changjiang (CJ) is particularly well suited to verifying our hypothesis because it is one of the largest rivers in the world covering a huge surface area (1.81×10^6 km², 19% of China). The water chemistry of CJ has been studied extensively. The major element chemistry is mainly controlled by rock weathering (mainly carbonates; Gaillardet *et al.*, 1999; Chen *et al.*, 2002), and four major resources (carbonates, silicates, evaporites and agriculture/urban effluents) contribute to the total dissolved solutes (Chetelat *et al.*, 2008). However, the catchments of its tributaries differ in their geology and hence the tributaries differ in chemical composition (Zhang *et al.*, 2003). Its specific HCO₃⁻ flux is strongly correlated with the carbonate mineral content of river basin (Cai *et al.*, 2008). Its DSi concentration has decreased due to a reduction in suspended sediment loading and increasing diatom consumption following dam construction (Duan *et al.*, 2007). Nowadays more than 50,000 dams have been constructed in CJ basin rivers (Yang *et al.*, 2011) and stoichiometric change of nutrients due to damming has already had a significant effect on ecosystems in the CJ estuary and the East China sea (Gong *et al.*, 2006; Chai *et al.*, 2009).

In the following, we report data on concentrations of DSi, HCO₃⁻ and related chemical and biological variables to understand the relationship between DSi and HCO₃⁻ and the mechanism that controls this relationship in CJ river and its tributaries.

METHODS

Study areas

Changjiang river springs from the Qinghai-Tibet plateau and flows 6300 km to the East China sea, draining an area of 1.8×10^6 km² and falling over 5400 m. The CJ basin is usually divided into three well described physiographic provinces (Chetelat *et al.*, 2008). Carbonate rocks are widely spread over the CJ basin and are particularly widespread in the southern part and sub-basin of the Han-shui river; evaporites are mainly present in the upper reach of CJ; and metamorphic rocks mainly underlie the sub-basin of the Ganjiang (GJ) river (Fig. 1B). Ten major tributaries joining the main CJ channel were studied. The main river receives water from Xiangjiang and Yuanjiang rivers *via* lake Dongting and from GJ river *via* lake Poyang. Wujiang (WJ) (mainly dominated by carbonate weathering) and GJ (mainly dominated by silicate weathering) rivers were selected for a specific comparative investigation.

Wujiang is a southern tributary of CJ, with a length of 1037 km and a drainage area of 8.8×10^4 km². It has an average runoff of 53.4×10^9 m³ per year with a fall of 2124 m and is the largest river in Guizhou province. The WJ basin is dominated by Permian and Triassic carbonate

rocks, which are interbedded with clastic rocks (Han and Liu, 2004). In WJ, a series of reservoirs have been constructed and the river is now a major hydropower source for China's massive *West-to-East power transmission project*.

Ganjiang river springs from Ganyuandong in the city of Ruijin and flows from south to west to the CJ. The river has a length of 766 km and a drainage area of 8.1×10^5 km² (51% of the Jiangxi province). The GJ river basin mainly overlies metamorphic and igneous rocks such as granite (Li *et al.*, 2007; Chetelat *et al.*, 2008).

Sampling and analytical methods

The CJ main channel and its largest tributaries (Fig. 1) were sampled in August 2006 using the sampling and analytical methods described in Chetelat *et al.* (2008). Wujiang river was sampled monthly from July 2007 to June 2008 for the investigation of water chemistry and sampled seasonally in July and October 2007 and in January and April 2008 for the investigation of phytoplankton. The GJ and its tributaries were sampled in May 2009 at 23 stations. Surface water (upper 0.5 m) samples were collected. Water profiles were investigated in the reservoirs Hongjiadu (W3), Dongfeng (W7), Wujiangdu (W15), Shangyoujiang (G2), and Wanan (G12 and G13). Water samples for depth profiles were taken with 5 L Niskin bottles. Sampling depths were 0, 5, 15, 30 and 60 m in W3, W7 and W15; 0, 5, 15, 25 and 35 m in G2; 0 and 5 m in G12; and 0, 5, 10 and 20 m in G13.

Temperature (T), dissolved oxygen, pH and chlorophyll *a* (Chl *a*) were measured *in situ* using a calibrated water quality probe (model YSI 6600) in WJ and GJ rivers. Samples for major cation and anion analysis were filtered through 0.45- μ m Sartorius® cellulose acetate filters (Sartorius, Göttingen, Germany). Samples for cation analysis were acidified to pH 2 with ultrapurified HNO₃. Major cations (Ca²⁺, Mg²⁺, K⁺, and Na⁺) and Si were analysed by inductively coupled plasma optical emission spectrometry (ICP-OES) with a precision better than 5%. Anions (SO₄²⁻, NO₃⁻, and Cl⁻) were determined by ionic chromatography Dionex 120 (Dionex, Sunnyvale, CA, USA) with a precision of 5%. Alkalinity was manually titrated with HCl in the field. Concentrations of HCO₃⁻ and dissolved CO₂ were calculated based on alkalinity, pH and T field data with constants corrected for temperature and ionic strength (Maberly, 1996; Barth and Veizer, 1999). The ionic strength was calculated from the major ions. A small portion of each sample was stored for measuring total nitrogen (TN) and total phosphorus (TP). Both TN and TP were determined spectrophotometrically (Unico UV-2000; Unicosci, Wanchai, Hong Kong) after alkaline potassium persulfate digestion (EPA, 1988).

One and a half L of water sample were preserved with Lugol's solution for analysis of phytoplankton in the sea-

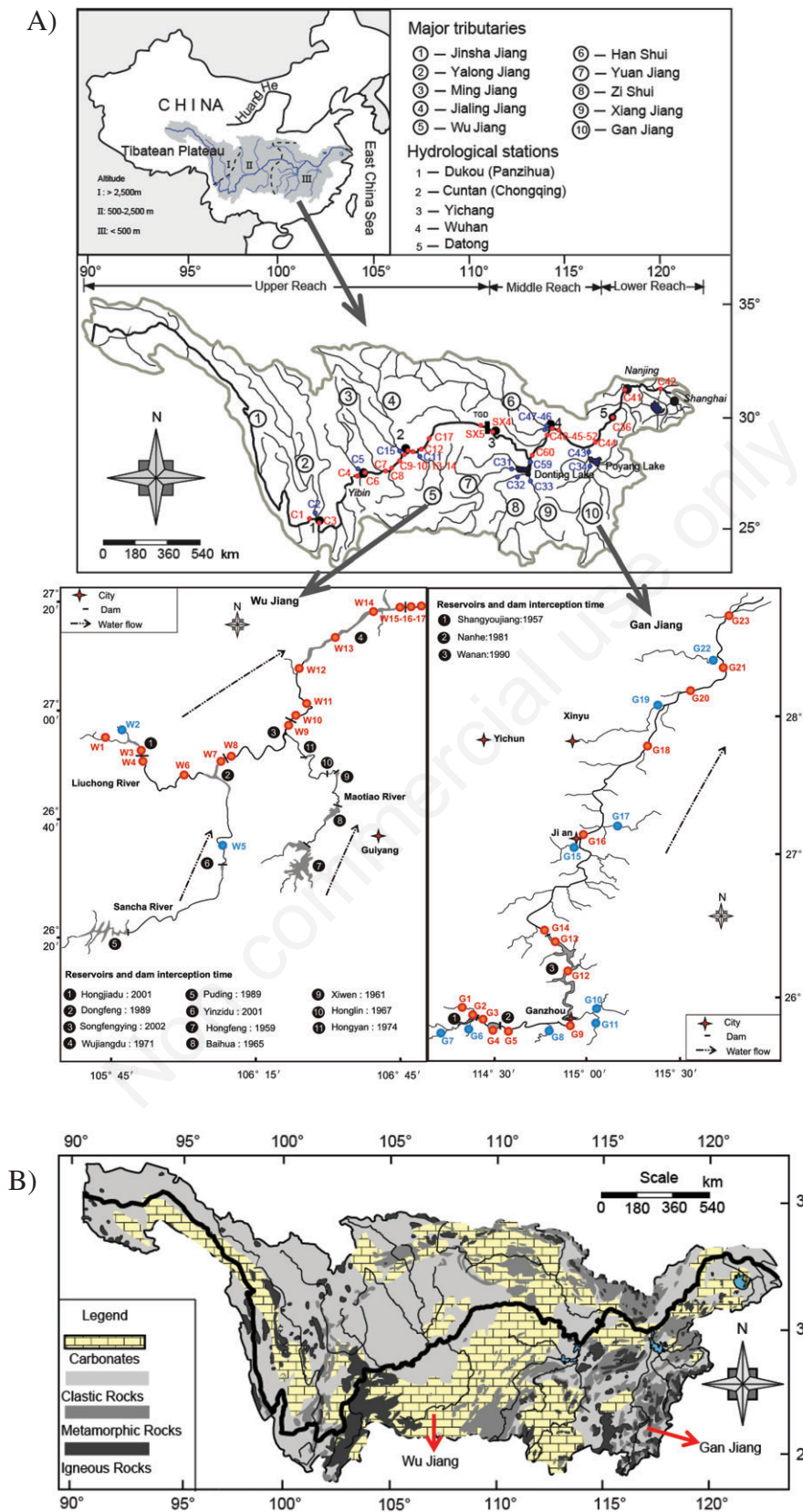


Fig. 1. Map showing sampling numbers and locations (A). Geological maps of the Changjiang basin, modified from Chen *et al.* (2002) and Li *et al.* (2007) (B).

sonal investigation of WJ. The method by Zhang and Huang (1991) was used for taxon identification, counting and cell dimensions using a standard light microscope. The wet weight (mg L^{-1}) of phytoplankton biomass was calculated according to its biovolume and cell density (Zhang and Huang, 1991).

Data collection and statistical analysis

A long term data of annual mean of DSi, HCO_3^- and dissolved CO_2 at the hydrological station of Datong near Nanjing were collected by Wang *et al.* (2007) and integrated by us to increase the dataset for analysis. Pearson's correlation coefficient analyses were carried out with the software SPSS (version 11.5; SPSS Inc., Chicago, IL, USA). Descriptive analyses were carried out with the Minitab 16 statistical software (Minitab Inc., State College, PA, USA). Contour graphs were produced using software Surfer (version 10; Golden Software Inc., Golden, CO, USA).

RESULTS

Basic physical and chemical properties

Calcium (Ca^{2+}) was the main cation in the rivers of the CJ basin (Fig. 2). Wujiang river had a higher mean concentration ($1504 \mu\text{mol L}^{-1}$) than CJ ($853 \mu\text{mol L}^{-1}$) and GJ ($586 \mu\text{mol L}^{-1}$). Calcium accounted for 70, 52 and 63% of the total mean cations in WJ, CJ and GJ, respectively. The mean Ca^{2+} concentration in CJ was similar to the mean value ($808 \mu\text{mol L}^{-1}$) reported by Chen *et al.* (2002). Magnesium (Mg^{2+}) was another main cation in CJ (mean value of $323 \mu\text{mol L}^{-1}$) and WJ (mean value of $418 \mu\text{mol L}^{-1}$), but was lower in GJ (mean value of $85 \mu\text{mol L}^{-1}$). The concentration of sodium (Na^+) varied widely in CJ, ranging from 94 (Yuanjing at Taoyuan, C31) to $1393 \mu\text{mol L}^{-1}$ (Jinshajiang at Panzhuhua, C1) (Fig. 2).

Bicarbonate was the main anion in the CJ basin rivers (Fig. 2). Wujiang river had a mean value ($2242 \mu\text{mol L}^{-1}$) similar to CJ ($1857 \mu\text{mol L}^{-1}$), but the concentration of HCO_3^- was lower in GJ ($615 \mu\text{mol L}^{-1}$). Bicarbonate accounted for 66, 74 and 64% of the total mean anions in WJ, CJ and GJ, respectively. The mean HCO_3^- concentration in CJ was similar to the mean value ($2110 \mu\text{mol L}^{-1}$) reported by Chen *et al.* (2002). Wujiang river had higher concentrations of SO_4^{2-} and NO_3^- than CJ and GJ, and this is consistent with the reports of Chen *et al.* (2002) and Duan *et al.* (2007). The Cl^- concentration in CJ ranged from 30 to $1134 \mu\text{mol L}^{-1}$ and, like for Na^+ , was more variable in CJ than in GJ and WJ. The mean Si concentration in GJ ($157 \mu\text{mol L}^{-1}$) was similar to that in CJ ($161 \mu\text{mol L}^{-1}$) and almost three times larger than that in WJ ($56 \mu\text{mol L}^{-1}$).

The median pH in CJ was similar to that in WJ and a little higher than that in GJ (Fig. 2). Wujiang river showed a larger variation of Chl *a* concentration than GJ due to the presence of reservoirs in WJ, some of which were eu-

trophic (Wang *et al.*, 2008). There was a very high concentration of dissolved CO_2 ($1389 \mu\text{mol L}^{-1}$) in the release water of Wujiangdu reservoir (W16) since release water derived from the deep water next to the dam (W15) and bottom water had a high CO_2 concentration due to respiration of organic matter (Wang *et al.*, 2011).

Spatio-temporal variation in chemical constituents in Wujiang tributary

There was a clear seasonality in concentrations of DSi, HCO_3^- and CO_2 in the surface waters of WJ (Fig. 3). In the reservoirs, the concentration of DSi was low between April and September. During this period, the extent of DSi decrease was related to the age of the reservoir: older reservoirs sequestered more DSi because they were productive and supported large diatom populations. Therefore, DSi concentration decreased downstream because of the construction of this cascade of reservoirs. Bicarbonate also decreased at the surface of the reservoirs and the decrease was also greater in older reservoirs; nonetheless, the release water (W4, W8, W10 and W16) had high HCO_3^- concentration. The dissolved CO_2 showed a pattern of variation that was similar to HCO_3^- and this was consistent with the report on the cascade of reservoirs in its tributary – Maotiao river (Wang *et al.*, 2011). Depth profiles of DSi, HCO_3^- and CO_2 showed a clear chemical stratification from April to September and their concentrations increased with depth in Dongfeng reservoir (Fig. 4). Depth profiles in the reservoirs Hongjiadu and Wujiangdu were similar to Dongfeng reservoir and therefore the profile data are not shown.

Phytoplankton in Wujiang tributary

There were three main algal groups in WJ: chlorophytes, dinophytes and diatoms. Biomass of diatoms showed a clear spatio-temporal variation and got the highest value in Wujiangdu reservoir in April (Fig. 5A). Phytoplankton species composition and biomass differed from different reservoirs (Fig. 5B) and seasons (Fig. 5C). Chlorophytes and dinophytes were dominant in July, and the contribution of diatoms to the total phytoplankton increased from October (52.3%) and January (43.9%) and was greatest in April (78.7%). Dinophytes were important only in July 2007. The mean value of phytoplankton biomass in different stations ranged from 1.7 mg L^{-1} (W4) to 14.5 mg L^{-1} (W14), and in different season from 1.9 mg L^{-1} (January) to 8.4 mg L^{-1} (July). Unsurprisingly, phytoplankton biomass decreased with depth (Fig. 5D).

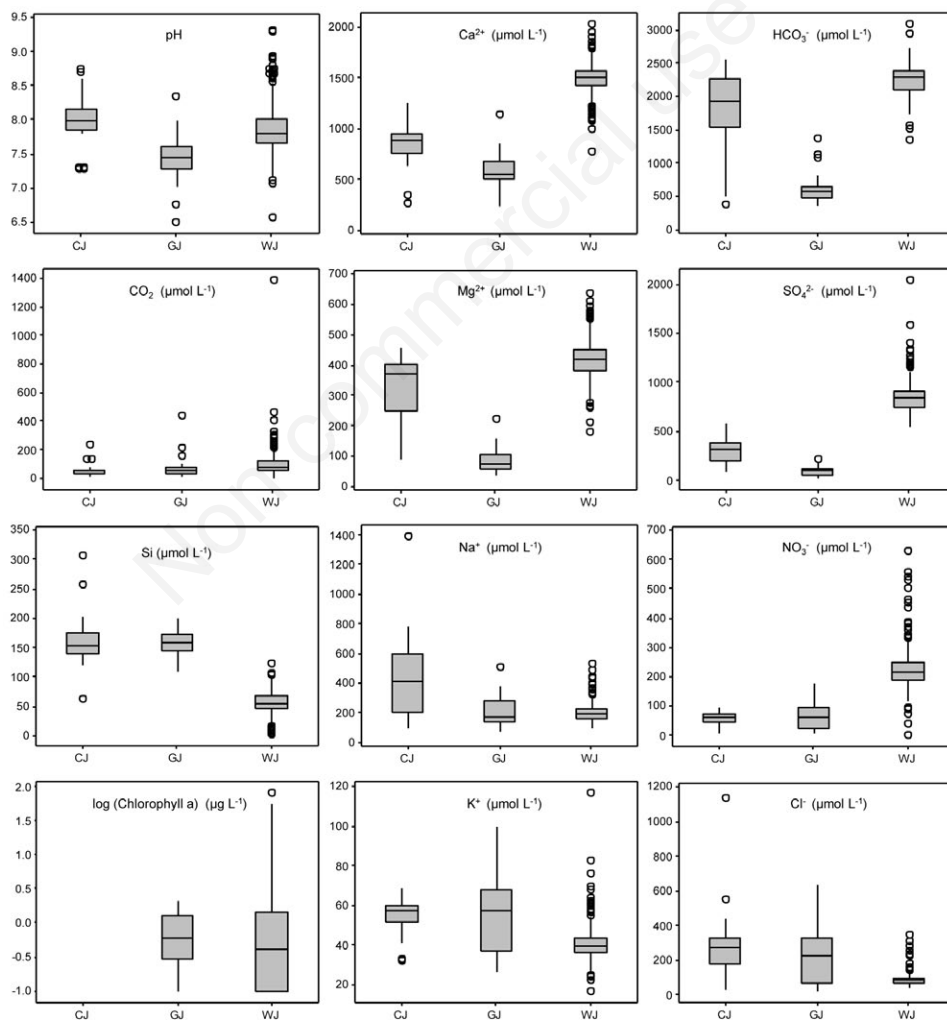
Spatial variation in Changjiang and Ganjiang rivers

The concentration of HCO_3^- in CJ decreased from upstream ($2570 \mu\text{mol L}^{-1}$ in C1) to downstream ($1563 \mu\text{mol L}^{-1}$ in C42) as a result of inflowing tributaries of lower al-

Tab. 1. Relationships between physical, chemical, and biological variables in Wujiang river in terms of the results of Pearson's correlation coefficient analysis.

	T	pH	DSi	HCO ₃ ⁻	CO ₂	TN	TP	Chl <i>a</i>	DIAT	PHY	CDTP	Si:HCO ₃ ⁻
pH	0.215***											
DSi	-0.083	-0.683***										
HCO ₃ ⁻	-0.520***	-0.187***	0.109*									
CO ₂	-0.089	-0.708***	0.468***	0.153**								
TN	-0.120*	-0.266***	0.233***	0.109*	0.203***							
TP	-0.113	0.017	-0.182**	0.129*	0.129*	-0.142*						
Chl <i>a</i>	0.388***	0.604***	-0.559***	-0.217***	-0.334***	-0.081	0.231***					
DIAT	0.137	0.369***	-0.471***	-0.227*	-0.311**	-0.050	-0.007	0.186				
PHY	0.554***	0.564***	-0.465***	-0.545***	-0.287**	0.110	0.243*	0.942***	0.327***			
CDTP	-0.413***	0.167	-0.251*	0.115	-0.280**	-0.246*	-0.017	-0.177	0.543***	-0.189		
Si:HCO ₃ ⁻	0.038	-0.640***	0.961***	-0.145**	0.416***	0.201***	-0.206***	-0.534***	-0.427***	-0.398***	-0.261**	
Si:CO ₂	0.369***	0.637***	-0.117*	-0.367***	-0.378***	-0.153**	-0.080	0.508***	-0.044	0.686***	-0.196	-0.038

T, temperature; DSi, dissolved silicon; HCO₃⁻, bicarbonate; CO₂, carbon dioxide; TN, total nitrogen; TP, total phosphorus; Chl *a*, chlorophyll *a*; DIAT, diatom biomass; PHY, phytoplankton biomass; CDTP, contribution of diatoms to the total phytoplankton; Si:HCO₃⁻, ratio between silicon and bicarbonate; Si:CO₂, ratio between silicon and carbon dioxide. *Significance at the 0.05 level (2-tailed); **significance at the 0.01 level (2-tailed); ***significance at the 0.001 level (2-tailed).

**Fig. 2.** Box plots of physical, chemical, and biological variables measured in Changjiang (CJ) river and Ganjiang (GJ) and Wujiang (WJ) tributaries. Box boundaries indicate the 25th and 75th percentiles; whiskers extend to a maximum of 1.5 times the inter-quartile range; the inner horizontal line is the median; and circles indicate outliers.

kalinity. The concentration of dissolved CO_2 had a similar pattern of variation to that of HCO_3^- , but with a smaller amplitude ($56 \mu\text{mol L}^{-1}$) (Fig. 6A). Dissolved silicon increased slightly from upstream to downstream, which was caused by input of DSi from GJ. The Three Gorges reservoir resulted in higher concentrations of CO_2 and HCO_3^- and lower concentrations of DSi.

In GJ, concentrations of dissolved CO_2 and HCO_3^- were fairly constant along the main channel except for high values in G5 and G9 (Fig. 6B). These high values may have been caused by bottom waters releasing from Nanhe reservoir and the input of its tributary (G8) with the higher bicarbonate concentrations. The concentration of DSi in GJ decreased slightly downstream.

Long term variation at Datong station

The concentrations of Si and dissolved CO_2 declined from 1964 to 1984 at the hydrological station of Datong. On the contrary, there was no significant change in the concentration of HCO_3^- (Fig. 7). The amplitude of HCO_3^- ($272 \mu\text{mol L}^{-1}$) was 10 times less than its own concentration, whereas the amplitudes of DSi ($61 \mu\text{mol L}^{-1}$) and CO_2 ($124 \mu\text{mol L}^{-1}$) were comparable, as for magnitude, to their own concentrations.

Relationship analysis

The concentrations of Si and HCO_3^- were significantly negatively correlated in CJ and GJ, but weakly positively correlated in WJ (Fig. 8). This suggests that factors influencing the concentrations of DSi and HCO_3^- in WJ were different from the two other rivers. The concentration of DSi was significantly positively correlated with CO_2 in WJ, while this was not found in CJ and GJ (Fig. 8).

The Si: HCO_3^- ratio in WJ was significantly negatively correlated with the values of diatoms, Chl *a* and TP (Tab. 1). This highlights the large effect of biological factors on the Si: HCO_3^- ratio in WJ.

DISCUSSION

Response of dissolved silicon-bicarbonate relationship to geological background

Rock weathering controls the major element chemistry of CJ. According to chemical and isotopic analysis, carbonate weathering is the main contribution to total dissolved solids (TDS), ranging from 40 to 80%. The contribution of silicate weathering to TDS ranged from 5 to 19.5% and that of evaporite weathering from 0 to 45% in CJ (Chetelat *et al.*, 2008). Evaporite weathering was most important in upper CJ but accounted for less than 10% of TDS in middle and lower CJ. The GJ was dominated by silicate weathering and WJ by carbonate weathering (Chetelat *et al.*, 2008).

The molar ratio of DSi to HCO_3^- originating from weathering of a given mineral is predictable. For calcite, one of the main carbonate minerals, this ratio is 0 (eq. 1) and for a silicate mineral plagioclase (Beaulieu *et al.*, 2010), this ratio is 0.67 (eq. 2). Therefore, the DSi: HCO_3^- ratio will vary from 0 to 0.67 for the river water chemistry controlled by carbonate and/or silicate weathering.

Our results clearly show that the Si: HCO_3^- ratio fluctuates between 0 and 0.67 (Fig. 8), increases with the dominance of silicate weathering, and decreases with the dominance of carbonate weathering.

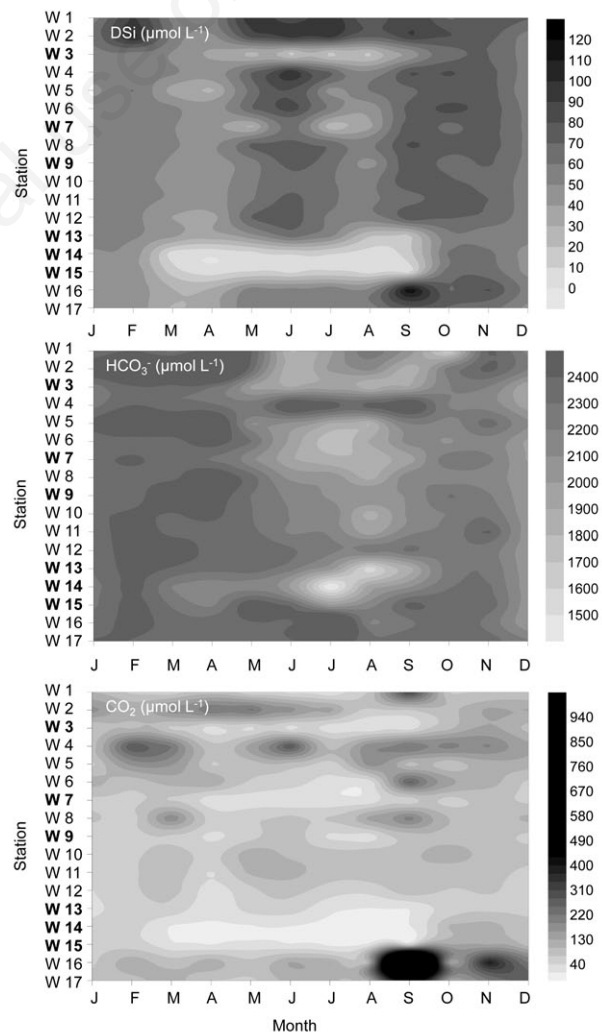
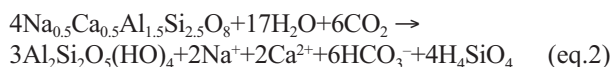
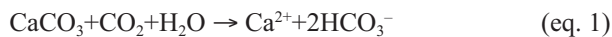


Fig. 3. Spatio-temporal variation of dissolved silicon, bicarbonate, and carbon dioxide of surface waters in Wujiang tributary. Stations are listed one after the other and stations from reservoirs are in bold.



The rate of carbonate weathering is 100 times greater than that of silicate weathering (Plummer *et al.*, 1978; Liu and Dreybrodt, 1997; Kump *et al.*, 2000). River water controlled by carbonate rock weathering will contain high HCO_3^- and low DSi concentrations, whereas river water controlled by silicate rock weathering will contain low HCO_3^- and high DSi concentrations. In addition, the inhomogeneous distributions of silicate and carbonate rocks finally resulted in a significant negative relationship between HCO_3^- and Si in CJ (Fig. 8). According to equation 2, it is expected to find a positive relationship between

HCO_3^- and DSi in a catchment dominated by silicate rock weathering. However, an inverse relationship between HCO_3^- and DSi was found and the DSi: HCO_3^- ratios deviated from 0.67 in GJ (Fig. 8), suggesting that, even in a catchment controlled by silicate weathering, HCO_3^- was still mainly derived from trace amounts of bedrock carbonate, this last being unevenly distributed in the river basin. This conclusion is consistent with other studies showing that 1% carbonate in the bedrock contributed to 82% HCO_3^- flux in the major Himalayan rivers (Blum *et al.*, 1998; Jacobson *et al.*, 2002). As for a catchment dominated by carbonate rock weathering, because DSi only originates from silicate rocks and these minor silicate rocks unevenly inlay carbonate rocks, a negative relationship between HCO_3^- and DSi is also expected. However, this expectation was not found in WJ (Fig. 8). The DSi-

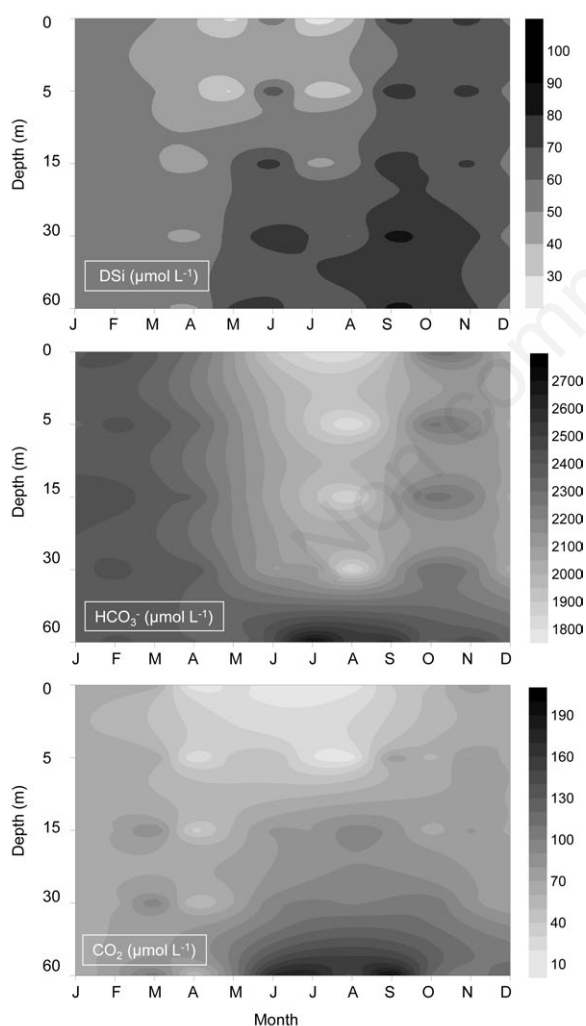


Fig. 4. Monthly profiles of dissolved silicon, bicarbonate, and carbon dioxide in Dongfeng reservoir on Wujiang tributary.

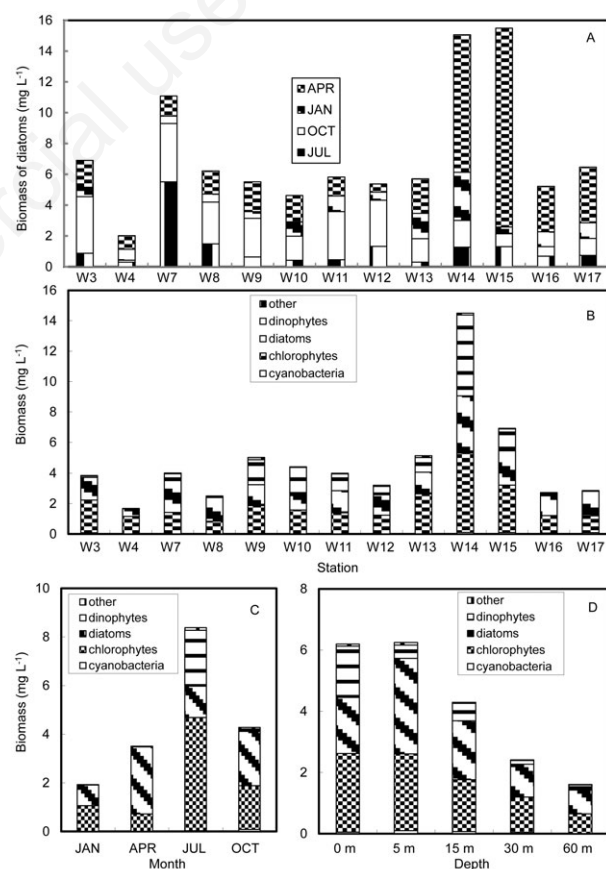


Fig. 5. Phytoplankton species composition and biomass in Wujiang tributary. Spatio-temporal variation of diatoms (A), mean phytoplankton species composition in different stations (B), seasons (C) and depths (D).

HCO_3^- relationship in the reservoir waters showed significant difference from that in the river waters (Fig. 9), and only DSi and HCO_3^- of river waters in WJ (Fig. 9B) showed similar relationship to that in CJ and GJ (Fig. 8). The strong involvement of biological activity modified the DSi- HCO_3^- relationship and therefore positive relationships between HCO_3^- and DSi and between CO_2 and DSi were found for the total dataset of WJ (Tab. 1).

Regulation of dissolved silicon-bicarbonate relationship by diatom uptake

A riverine heterotrophic ecosystem may be transformed into an autotrophic one after damming a river to produce a reservoir (Wetzel, 2001). Phosphorus was the

main factor limiting algal growth because Chl *a* and phytoplankton biomass were significantly correlated with the concentration of TP (Tab. 1). Diatoms are an important group of phytoplankton (Fig. 5). According to the correlation analysis, DSi and CO_2 were stoichiometrically assimilated by diatoms and this led to the significant positive correlation between DSi and CO_2 in WJ's reservoirs (Fig. 9A). The dissolved CO_2 in these reservoirs can only come from the conversion of HCO_3^- in place of the atmosphere because the partial pressure of CO_2 in river is much higher than that in atmosphere (Wang *et al.*, 2007, and references therein). Uptake of CO_2 by algae finally induced the decrease of HCO_3^- and this caused HCO_3^- to be strongly positively correlated to DSi in WJ's reservoirs. After regulation by diatom uptake, the relationship be-

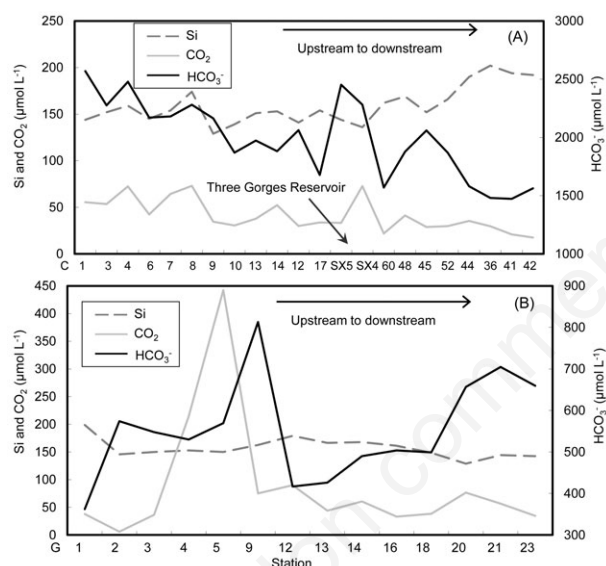


Fig. 6. Variation of silicon, bicarbonate, and carbon dioxide in the Changjiang main channel (A) and the Ganjiang main channel (B) from upstream to downstream.

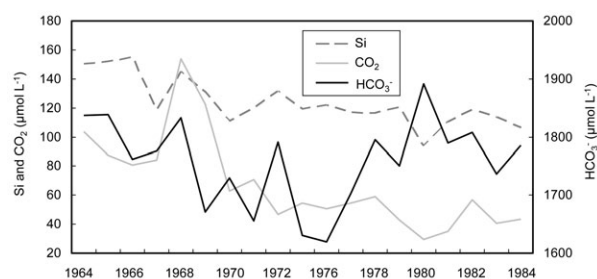


Fig. 7. Variations of annual mean of silicon, bicarbonate, and carbon dioxide at Datong station near Nanjing from 1964 to 1984. Data were collected by Wang *et al.* (2007).

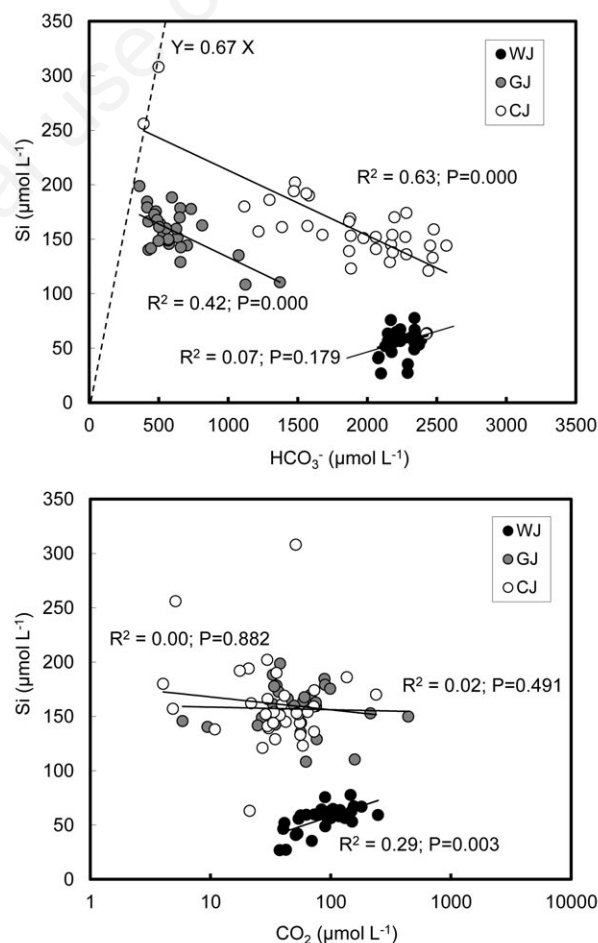


Fig. 8. Scatter plots of silicon vs bicarbonate and silicon vs carbon dioxide in Changjiang (CJ), Ganjiang (GJ) and Wujiang (WJ) rivers. Data from WJ are the monthly mean of different stations including different depths. The line $Y=0.67X$ is based on the plagioclase weathering (eq. 2).

tween DSi and HCO_3^- in the downstream river waters became different from that in the upstream ones (Fig. 9B). Therefore, the regulation by diatoms clearly changed the negative correlation between DSi and HCO_3^- that was controlled by geological background.

The concentration of DSi was an order of magnitude similar to CO_2 concentration but both were about 10 times smaller than HCO_3^- concentration. Variation in the Si: HCO_3^- ratio was more dependent on variation of DSi compared with HCO_3^- , whereas the Si: CO_2 ratio was more dependent on variation of CO_2 compared with DSi according to the correlation analysis. Therefore, the Si: HCO_3^- ratio showed a different correlation with the related variables from the Si: CO_2 ratio.

The Si: HCO_3^- ratio decreased with the abundance of diatoms, Chl *a* and TP (Fig. 10) because Si and C are needed for diatom growth, and the mean Si:C ratio of freshwater diatoms (0.79; Sicko-Goad *et al.*, 1984) was larger than the Si: HCO_3^- ratio of the river waters (<0.67). This is a negative feedback regulation. The Si: HCO_3^- ratio also showed

a significant negative correlation with the contribution of diatoms to the total phytoplankton (Tab. 1). This relationship means that the increase of non-diatom phytoplankton will tend to increase the Si: HCO_3^- ratio because non diatoms only assimilate DIC but not DSi. Therefore, the Si: HCO_3^- ratio can also respond to changes of algal species composition.

The spatial snapshot of DSi did not show a large fluctuation in CJ and GJ (Fig. 6). However, long term data showed a clear decrease in its concentration in Datong (Fig. 7) and, obviously, the Si: HCO_3^- ratio decreased with time as a response, indicating the occurrence of the regulation by diatom uptake. Although the bottom releasing water of the reservoirs can cause the increase of DSi in release waters, the net transport downstream for DSi through the reservoirs is negative because there is still a part of biological silicon staying in the sediments. This is the reason why DSi concentration decreased with time in Datong (Fig. 7). The decreasing tendency of silicon was also reported in middle and lower CJ (Duan *et al.*, 2007)

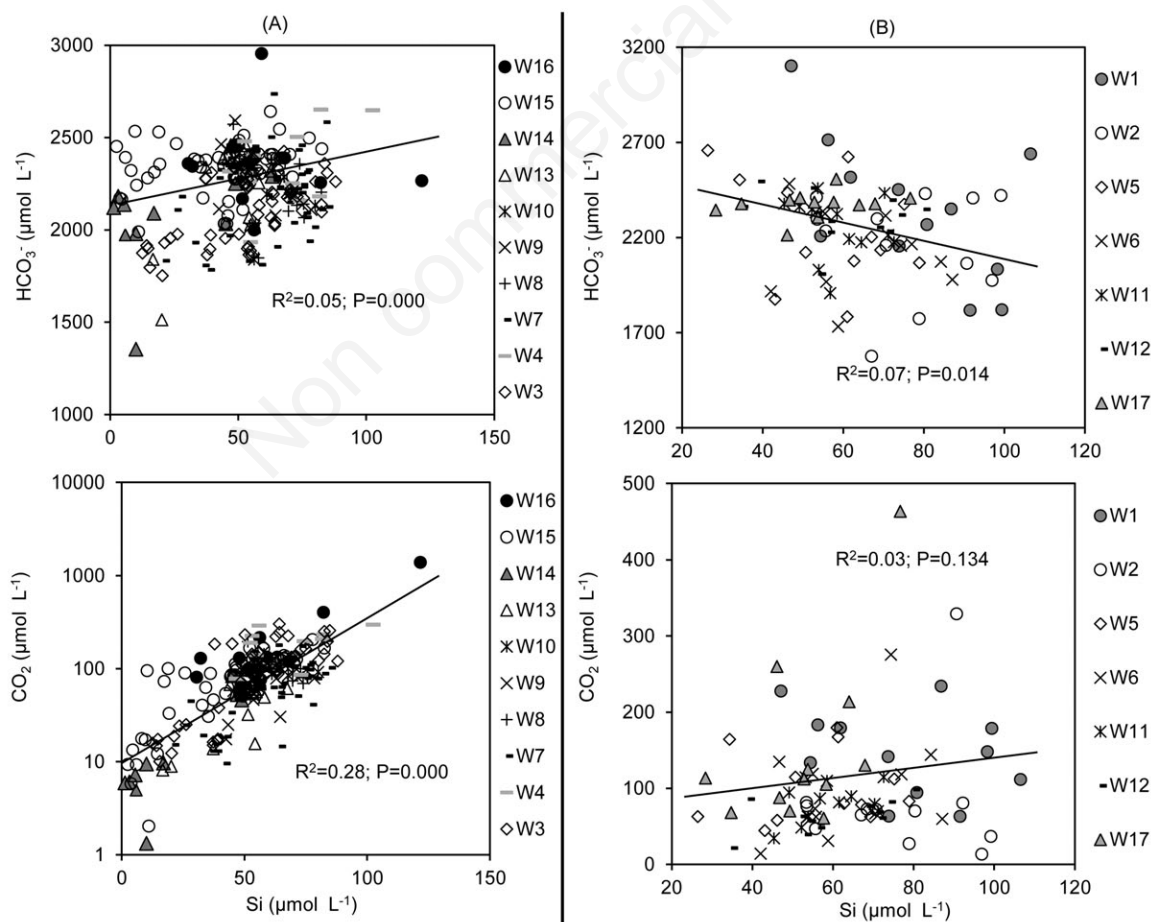


Fig. 9. Scatter plots of silicon vs bicarbonate and silicon vs carbon dioxide in Wujiang tributary. Data from reservoir and release waters are shown in (A), while data from river waters are in (B).

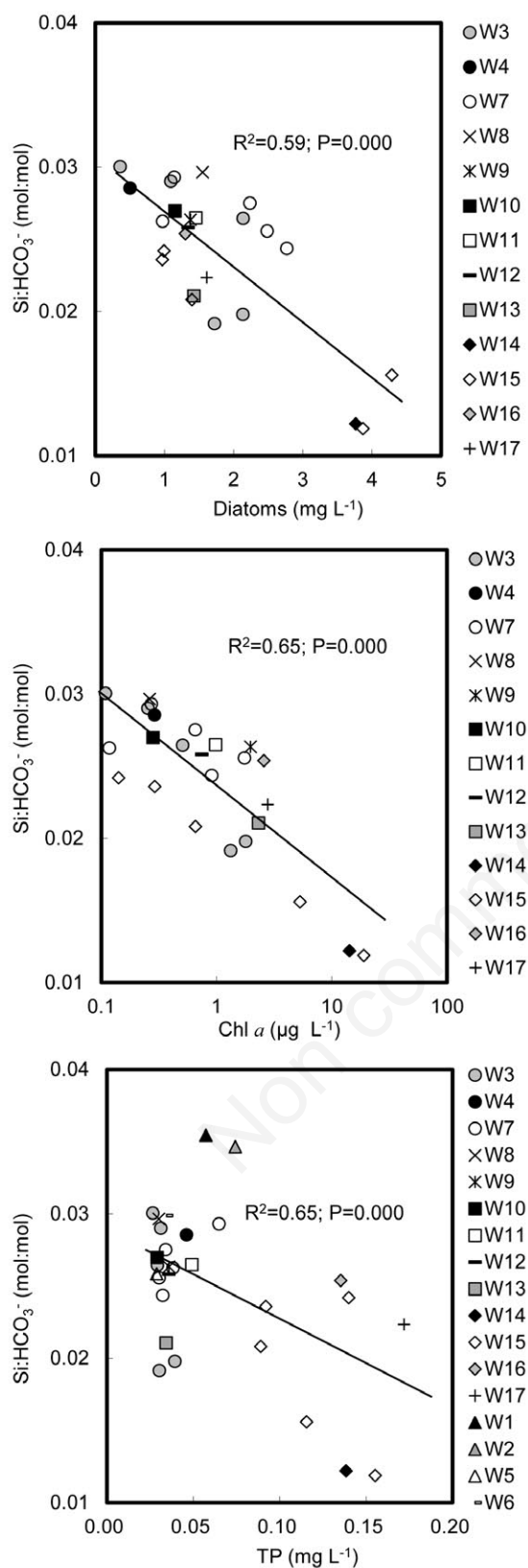


Fig. 10. Scatter plots of mean Si:HCO₃⁻ ratio vs mean diatom biomass, Chlorophyll *a* and total phosphorus, respectively, in different stations.

in a long term. This suggested that the regulation by diatom uptake can be inconspicuous in a spatial snapshot on a large scale, but the cumulative effect of this regulation with time cannot be ignored.

CONCLUSIONS

The concentrations of DSi and HCO₃⁻ in the CJ basin rivers were related to their geological origin. The DSi:HCO₃⁻ ratio ranged from 0 to 0.67, increased when silicate weathering dominated, and decreased when carbonate and/or evaporite weathering prevailed. Damming induced the increasing effect of algal activity on the elemental geochemical cycle, and diatom uptake could synchronously sequester DSi and CO₂ (from HCO₃⁻) via the mean stoichiometric ratio of 0.79, thus having a negative feedback regulation on the DSi:HCO₃⁻ ratio. Our study demonstrated that the relationship between DSi and HCO₃⁻ can reflect not only their geological background but also the influence of algal activity within rivers.

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