

Gas emissions and crustal deformation from the Krýsuvík high temperature geothermal system, Iceland.

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Abstract

The Krýsuvík volcanic system is located on the oblique spreading Reykjanes Peninsula, SW Iceland. Since early 2009 the region has been undergoing episodes of localized ground uplift and subsidence. From April–November 2013, we operated near-real time monitoring of gas emissions in Krýsuvík, using a Multi-component Gas Analyzer System (Multi-GAS), collecting data on gas composition from a fumarole (H_2O , CO_2 , SO_2 , H_2S). The dataset in this study, comprises a near-continuous gas composition time series, the quantification of diffuse CO_2 gas flux, analytical results for direct samples of dry gas, seismic records, and GPS data. Gas emissions from the Krýsuvík geothermal system were examined and compared with crustal deformation and seismicity.

The gas emissions from the Krýsuvík system are H_2O -dominated, with CO_2 as the most abundant dry gas species, followed by smaller concentrations of H_2S . The average subsurface equilibrium temperature was calculated as 278 °C. This is consistent with previous observations made through sporadic spot sampling campaigns. In addition, the semi-continuous Multi-GAS dataset reveals higher variations in gas composition than previously reported by spot sampling.

The diffuse soil CO_2 flux is found to be variable between the three studied degassing areas in Krýsuvík, ranging from 10.9–70.9 T/day, with the highest flux in Hveradalir where the Multi-GAS station is located. The total flux is estimated as 101 T/day.

Comparison between Multi-GAS and geophysical data shows that peaks of H_2O -rich emissions appears to follow crustal movements. Coinciding with the H_2O -rich peaks, SO_2 is detected in minor amounts (~ 0.6 ppmv), allowing for calculations of $\text{H}_2\text{O}/\text{SO}_2$, CO_2/SO_2 and $\text{H}_2\text{S}/\text{SO}_2$ ratios. This is the first time SO_2 has been detected in the Krýsuvík area.

The large variations in $\text{H}_2\text{O}/\text{CO}_2$ and $\text{H}_2\text{O}/\text{H}_2\text{S}$ ratios are considered to reflect variable degassing activity in the fumarole. The activity of the fumarole appears less intense during intervals of low or no recorded seismic events. The $\text{H}_2\text{O}/\text{CO}_2$ and $\text{H}_2\text{O}/\text{H}_2\text{S}$ ratios are lower, presumably due to H_2O condensation affecting the steam jet before reaching the Multi-GAS inlet tube.

Keywords: Krýsuvík, volcanic gas, volcano monitoring, geothermal gas, crustal deformation, volcanic CO₂ flux

1 Introduction

Monitoring volcanic and geothermal gases, along with seismicity and ground deformation, can lead to better understanding of volcano behavior, and provide an early warning of volcanic activity. Several studies have focused on quiescent degassing from active volcanic and geothermal systems, detecting peaks of increased gas emissions prior to eruptions (e.g., Young et al., 1998, Aiuppa et al., 2010). It has been shown that, under certain conditions, seismicity and ground deformation may help in releasing gases into these systems and increase fumarolic emissions (e.g., Watson et al., 2000, Italiano et al., 1998, Toutain and Baubron, 1999, Chiodini et al., 2012, 2015, and references therein). Monitoring of soil CO₂ diffuse degassing in geothermal areas has also proven to give reliable information on the mass/energy budget of utilized and non-utilized geothermal systems (Brombach et al., 2001; Chiodini et al., 2001, Fridrikson et al., 2006, Óladóttir, 2012). In Iceland, there has been limited research on the relationship between degassing from high temperature geothermal areas (where temperature at 1 km depth is greater than 200°C) and seismic energy release and ground deformation. The Krýsuvík volcanic system, located on the oblique spreading Reykjanes Peninsula (RP), is characterized by abundant degassing through soil and fumaroles within a high-temperature geothermal area. It has high seismic activity, characterized by swarms of micro-earthquakes as well as main-shock/aftershock sequences (Klein et al., 1977 and references therein, Ward and Björnsson, 1971, Kristjánsdóttir, 2013). In the last decade the region has undergone episodes of uplift and subsidence, with high seismic activity occurring during periods of uplift (Michalczevska et al., 2012).

Here we present semi-continuous, near-real time gas measurements in the Krýsuvík geothermal system (April-November 2013). Such measurements are relatively new in Iceland with the first station installed on top of Mt. Hekla volcano in 2012 (Ilyinskaya et al., 2015; Di Napoli et al., 2016). We evaluate the composition of the gas emitted from the Krýsuvík geothermal system and interpret its origin. This is done through analysis of a semi-continuous time series of gas composition (Multi-GAS sensor system, e.g., Aiuppa et al., 2009), direct steam sampling of fumaroles, and quantification of the diffuse CO₂ flux from the geothermally active areas (accumulation chamber method, e.g., Fridriksson et al., 2006). The gas time series are compared to observations of ground deformation and seismicity in Krýsuvík.

1.1 Regional settings

Krýsuvík is one of five active, NE-SW trending, volcanic systems located on the oblique spreading Reykjanes peninsula, SW Iceland (Saemundsson et al., 2010, Hreinsdóttir et al., 2001, Einarsson, 2008) (Figure 1). The Krýsuvík system is thought to be in an early stage of evolution, dominated by rift volcanism with no major magma chamber (Arnórsson et al., 1975). Volcanic activity is periodic with roughly 1000 years intervals between eruptive episodes, each eruptive episode lasting for 400-500 years (Saemundsson and Jóhannesson, 2006, Jónsson, 1978). The last volcanic eruption in Krýsuvík took place in the 12th century

(Saemundsson and Sigurgeirsson, 2013). Krýsuvík currently hosts a high-temperature geothermal system, the heat source of which is considered to be dyke intrusions (Arnórsson et al., 1975, Arnórsson, 1987). Recent resistivity measurements within the Krýsuvík system indicate a conductive body at approximately 2 to 5 km depth (Didana, 2010, Hersir et al., 2013). This body is located near the central part of the Krýsuvík geothermal area, with an approximate size of 10 km², and coincides with the source of the inflation and deflation observed with GPS and InSAR measurements (Michalczyewska et al., 2012, Hersir et al., 2013). The lack of S-wave attenuation in the region has been used as an argument against the presence of large volumes of molten materials, with Adelinet et al., (2011) suggesting the presence of a gaseous or supercritical fluid at 6 km depth from the analysis of tomography data.

The Krýsuvík geothermal system is usually divided into six subareas: Sandfell, Trölladyngja, Köldunámur, Seltún, Hveradalir, and Austurengjar (Figure 2). This study is focused on the last three, which have the highest and most continuous geothermal surface activity including hot and altered ground with mud pools, fumaroles, and solfataras. The main surface activity is confined to the Vesturháls and Sveifluháls hyaloclastite ridges, including Seltún and Hveradalir, and Austurengjahver, east of the Sveifluháls ridge (Markússon and Stefánsson, 2011). The results of resistivity measurements indicate that the geothermal subareas within Krýsuvík, originate from one and the same system of an approximate size of 40-60 km² (Gudmundsson et al., 1975, Saemundsson and Sigurgeirsson, 2013) (Figure 2).

2 Methodology and data processing

2.1 Gas composition analysis using MultiGAS and direct fumarole sampling

A Multi-component Gas Analyzer System station (Multi-GAS, INGV-type, see e.g., Aiuppa et al., 2009 and references therein, Ilyinskaya et al., 2015) was installed on 26 April 2013 at Hveradalir in Krýsuvík (Figure 3), next to a fumarole in an area of high and persistent surface geothermal activity. Data collection was discontinued over a short period from 27 June – 5 July 2013 when the station was needed for another project. The sampling inlet was ~20 cm above ground level (Figure 3), which was necessary to avoid saturation of the CO₂ sensor.

The inlet was tested at several heights above the ground, and this height gave the most reliable data. The setup of the station can vary between sampling sites. The site in Krýsuvík has much more steam coming from the ground, than the site at e.g., Mt. Hekla, where saturation problems with the sensors were not an issue, and the station was located inside a small hut (Ilyinskaya et al. 2015).

The station was powered by a permanent wind turbine and was configured to acquire data in cycles of 200 samples, each being the median of 9 measurements at 1 Hz (30 minutes per sampling cycle) with a time interval of 6 hours. A 3G radio modem was used for telemetry, and data was retrieved remotely using custom-made software (Ratiocalc 2.0, Tamburello, 2015) which allows for the automatic creation of a gas species scatter plot from the data acquired. The gas molar ratios were calculated from the gradient of best fit

regression lines (Aiuppa et al., 2009, 2010) and calculations were restricted to intervals when measured concentrations showed a substantial excess relative to ambient air (see e.g., Ilyinskaya et al., 2015). Overall uncertainty in the derived ratios is $\leq 20\%$ (Aiuppa et al., 2009).

A total of eight samples of dry gas were collected in two campaigns with an eight week interval (Figure 2) from fumaroles with a focussed steam flow (four samples in Hveradalir, two in Seltún, and two in Austurengjar). The samples were collected into evacuated double port bottles with 25 ml of a 10M NaOH solution. The ground temperature next to the fumarole was recorded and was in all cases $>97^{\circ}\text{C}$.

The dry gas samples were analysed for gas composition using standard procedures at the Iceland GeoSurvey (ÍSOR), Reykjavík, Iceland. Headspace gases (N_2 , CH_4 , Ar, H_2 , and O_2) were analysed for using gas chromatography in a Perkin-Elmer Arnel 4019 gas chromatograph. CO_2 and H_2S were analysed for by titration of the caustic solution and the gas to condensate ratio determined by weighing.

2.2 Soil temperature and diffuse CO_2 flux through soil

The soil CO_2 flux was measured using a West Systems fluxmeter in the three sub-areas of Krýsuvík studied using the accumulation chamber method (e.g., Fridriksson et al., 2006). The method has proven to be an accurate way to measure soil CO_2 fluxes in volcanic and geothermal areas, since it does not require assumptions or corrections that depend on soil characteristics (Chiodini et al., 1998). The measurements were carried out on an approximately 25×25 meter grid, where possible. The total number of measurement points was 435; 217 in Hveradalir, 136 were in Seltún, and the remaining 82 were in Austurengjar. Most of the measurements were taken in late summer or autumn, during dry and calm weather conditions to avoid the influence of weather factors and water saturation of the soil. The average time of each measurement was 1-2 minutes, depending on the time the rate of CO_2 concentration increase stabilized. To evaluate the total CO_2 emission from the areas measured the Kriging algorithm (Cardellini et al., 2003, and references therein) was used for interpolation. The soil temperature was measured with a handheld digital thermometer with a 15 cm long probe.

2.3 Geophysical data

The gas measurements was compared to seismic and GPS data from Krýsuvík. The seismic data were provided by the Icelandic Meteorological Office (IMO) from the SIL seismic network and included 217 events in the Krýsuvík region from late April through November 2013. The seismic moment was used to estimate the moment magnitude, M_w (Hanks and Kanamori, 1979):

$$M_w = \frac{2}{3}(\log_{10}(M_0) - 9.1) \quad (1)$$

The largest recorded seismic event for the study period was M_w 2.2 on 11 July 2013. The frequency-magnitude distribution (Gutenberg and Richter, 1944) for the Krýsuvík catalogue gives a magnitude of completeness, M_w 0.75 and the slope or b-value of 1.6. All events with $M_w < 0.75$ were discarded from further analysis, leaving a total of 172 events.

The global average b-value is around 1 but ranges locally from 0.5-2 depending on factors like the type of tectonic environment and stress. Keiding et al., (2009) used data spanning from 1997 to 2006 to evaluate the b-value for the Krýsuvík region, giving the considerably lower value of 0.9. The dataset used here (172 events) is considerably smaller which could bias the estimate. However, high b-values are often observed in volcanic and geothermal regions and are assumed to be related to heterogeneous crust as well as local stress perturbations and fluids (Wyss, 1973, Schorlemmer et al., 2005).

The daily cumulative seismic moment was estimated showing several peaks of increased seismic activity over the study period, with the largest one occurring in mid-July 2013.

The GPS station MOHA started continuous operation in 2010 to monitor crustal deformation in the Krýsuvík region. The station is located just north of the center of uplift observed from 2010 to 2011 (Figure 1) (Michalczyewska et al., 2012).

GPS data was analysed using GAMIT/GLOBK version 10.4 using over 100 continuous global GPS stations to evaluate average daily site positions in the ITRF08 reference frame. In the processing we solve for station coordinates, satellite orbit and earth rotation parameters, atmospheric zenith delay every two hours, and three atmospheric gradients per day. The IGS08 azimuth and elevation dependent absolute phase center offsets were applied to all antennas and ocean loading was corrected for using the FES2004 model. To minimize the common mode signal in the time series the de-trended time series from the GPS station NYLA, outside the deforming region, were subtracted from the data. Figure 4 shows the local inflation/deflation periods from the last decade of vertical GPS data (2007-2016) at station MOHA. During the year of 2013 the estimated total subsidence was 21 mm. We used a running weighted average of seven days for the dataset when comparing to gas data.

3 Results

3.1 Assessment of the influence of meteorological conditions on the MultiGAS data

Before interpreting the MultiGAS data, we assessed the influence of meteorological conditions on the measurements. The gas ratios ($\text{H}_2\text{O}/\text{CO}_2$, $\text{H}_2\text{O}/\text{H}_2\text{S}$, $\text{H}_2\text{O}/\text{SO}_2$, $\text{CO}_2/\text{H}_2\text{S}$, CO_2/SO_2 , $\text{H}_2\text{S}/\text{SO}_2$) were compared to wind speed (m/s) and precipitation (mm/day) data. Wind data from the three weather stations closest to the location of the MultiGAS station was investigated: Festarfjall, Selvogur and Straumsvík (IMO monitoring data). Based on the topography of the area it was concluded that the atmospheric conditions at the MultiGAS station were most similar to those at the location of the Festarfjall weather station. (Figure 1). The wind dataset from the Festarfjall station consists of hourly measurements. This resolution allows for an accurate comparison between the wind speed and each MultiGAS acquisition. MultiGAS data was acquired for 30 min starting at 00:00, 06:00, 12:00 and 18:00 each day. The precipitation data (IMO monitoring data) was obtained from the Keflavík airport station (Figure 1), since the Festarfjall station does not collect data on precipitation. The precipitation data in Keflavík is collected twice per day

(09:00 and 18:00). The representative precipitation is the total accumulated precipitation per day (where one day is defined from 18:00-18:00 UTC), which is used to compare with all MultiGAS acquisitions made that same day.

All gas molar ratios showed greatest variability during dry periods (< 2 mm/day). However, high values for $\text{CO}_2/\text{H}_2\text{S}$, CO_2/SO_2 and $\text{H}_2\text{S}/\text{SO}_2$ ratios are not confined to the dry periods since these gas species are less affected by condensation during rainy days than H_2O (Figure 5).

During periods of low wind speed (< 5 m/s) the $\text{CO}_2/\text{H}_2\text{S}$, CO_2/SO_2 and $\text{H}_2\text{S}/\text{SO}_2$ ratios show the greatest fluctuation. Progressive decrease in ratios obtained and decrease in fluctuation of the ratios is observed with higher wind speeds (Figure 5). No visual correlation is apparent between the $\text{H}_2\text{O}/\text{X}$ ratios (where X stands for CO_2 , H_2S and SO_2) with wind speed below 10 m/s (Figure 5).

To account for the influence of meteorological conditions we eliminate all gas molar ratios obtained during days of more than 2 mm/day precipitation in the analysis. Additionally, $\text{CO}_2/\text{H}_2\text{S}$, CO_2/SO_2 and $\text{H}_2\text{S}/\text{SO}_2$ ratios collected in periods of wind speed above 5 m/s were eliminated.

3.2 Gas composition

The fumarole gas samples from Krýsuvík were dominated by H_2O (96.6-99.6 vol%) with CO_2 as the dominant dry gas component (on average 83.8 vol%), followed by much smaller amounts of H_2S (on average 9.17 vol%), hydrogen (H_2), nitrogen (N_2), methane (CH_4) and argon (Ar) (Table 1). One sample (Seltún 1) had a small component of O_2 , a result of atmospheric contamination during sampling.

The MultiGAS ratios detected (where effects of the meteorological conditions have been eliminated) are shown to be highly variable (Table 2). Minor amounts of SO_2 were measured for the first time in Krýsuvík by the MultiGAS sensor. The measured SO_2 concentrations are very low but the bulk (63%) is above the detection limit (0.05 ppmv) of the MultiGAS sensor, allowing for calculations of the X/SO_2 ratios ($\text{X} = \text{H}_2\text{O}$, CO_2 and H_2S). Concentrations below 0.05 ppmv are considered to be instrumental noise. The SO_2 sensor is not quantitative below 1 ppmv so the measurements (0.05-1 ppmv) should only be viewed as qualitative assessments of SO_2 presence with a great uncertainty.

The calculated ratios of $\text{H}_2\text{O}/\text{CO}_2$, $\text{H}_2\text{O}/\text{H}_2\text{S}$ and $\text{CO}_2/\text{H}_2\text{S}$ from the fumarole samples are compared to the MultiGAS ratios in Table 3. The average value for $\text{H}_2\text{O}/\text{CO}_2$ and $\text{H}_2\text{O}/\text{H}_2\text{S}$ from the fumarole samples (220 and 2812, respectively, excluding the Seltún 1 sample), fall within the range of the highest $\text{H}_2\text{O}/\text{CO}_2$ and $\text{H}_2\text{O}/\text{H}_2\text{S}$ ratios obtained by the MultiGAS sensor (217 and 10,300, respectively). For $\text{CO}_2/\text{H}_2\text{S}$, the average and median values for the fumarole samples (13 and 10, respectively) are close to the average and median values of the MultiGAS data (17 and 12, respectively).

3.3 Soil CO₂ flux and temperature

The highest soil CO₂ diffuse degassing values were found in areas with intense surface activity, where steam rises through fissures and cracks towards the surface, (total 0.31 km²). The CO₂ fluxes ranged between 0 and 29,200 g m⁻² day⁻¹ with an average value of 385 g m⁻² day⁻¹ and a median value of 6 g m⁻² day⁻¹. In Austurengjar (0.09 km²) 80% of the observation points showed low CO₂ fluxes (<10 g m⁻² day⁻¹) 68% in Hveradalir (0.140 km²) and 51% in Seltún (0.08 km²) (Figure 6).

To process the CO₂ measurement data and evaluate the total CO₂ emission from the area studied, sequential Gaussian simulations (sGs) were used. In this method, the sample set was used to generate a great number of equiprobable representations or realizations of the spatial distribution of the CO₂ flux. For each area, 100 realizations were performed using the sGs algorithm of the *sgsim* code by Deutsch and Journel (1998). The results of the 100 simulations are depicted in Figure 6, that shows the mean CO₂ flux of individual cells in the model which was generated from the original dataset.

The results from the three observation areas shows that the Hveradalir area had by far the greatest total flux of 70.9 T/day. Next was Seltún with 19.6 T/day and last Austurengjar with 10.9 T/day. The total CO₂ soil flux from the three observation areas was estimated as 101 T/day. The soil temperature ranged between 4.2-99.0 °C with the average value 19°C and median value 12°C.

4 Discussion

4.1 Application of gas geothermometers

Gas geothermometers (here: Arnórsson et al., 1998 and Arnórsson and Gunnlaugsson, 1985, Table 4) were used to determine the subsurface temperature of the Krýsuvík geothermal system based on the fumarole steam composition (Table 5). The gas geothermometers assume that the concentrations of CO₂, H₂S and H₂ in geothermal reservoir waters are controlled by temperature dependent equilibria with minerals (Arnórsson and D'Amore, 2000).

The lowest individual variations were observed for both the CO₂ geothermometers, indicating a sub-surface temperature of 293°C (Table 5). The agreement of other geothermometers was not very good as predicted temperatures varied for individual samples. A substantial discrepancy was observed between the various geothermometers that are known from previous studies (e.g., Arnórsson and Gunnlaugsson, 1985).

The average sub-surface temperature for all samples was estimated around 278°C which is comparable to previous studies, suggesting little or no change in sub surface temperature in the past 33 years (e.g., Arnórsson and Gunnlaugsson, 1985, Arnórsson, 1987, Yohannes, 2004).

4.2 Correlation of geophysical observations and the MultiGAS data

The waterlevel of lake Kleifarvatn (Figure 2), which has no outlet river, suggest only relatively minor changes in ground water level during the measurement period of the MultiGAS station (April-November 2013), with the highest and lowest lake level of 138.22 m and 137.85 m on 26 September and 11 November 2013, respectively. For comparison, the difference between the highest and lowest waterlevel from 2003-2017 was 3.8 m (Icelandic Met Office, 2018). Groundwater level measurements in a borehole (TD-4) located over 6 km NNW of the MultiGAS station, west of Mt Trölladyngja (Figure 2), shows similar trend but with larger fluctuations with just over 3 m water level difference during the study period compared to over 15 m difference from 2009 to 2017 (Elefsen, S, Ó, personal communications 13 February 2017). Based on this, changes in groundwater in the area over the study period, were not considered to be a major factor in the gas composition, nor the activities of the fumaroles in the area. The MultiGAS measurements were conducted during a deflation period with relatively low seismic activity. Here we attempt to compare the gas ratios to the observed deformation, to observed variations in seismic and deformation activity.

4.2.1 Comparison with $\text{H}_2\text{O}/\text{CO}_2$, $\text{H}_2\text{O}/\text{H}_2\text{S}$ and X/SO_2 MultiGAS ratios

Distinct periods with peaks in $\text{H}_2\text{O}/\text{CO}_2$ and $\text{H}_2\text{O}/\text{H}_2\text{S}$ gas ratios are observed from the MultiGAS time series. SO_2 was detected almost exclusively at the same time as these detected peaks, allowing for calculations of X/SO_2 ratios (Figure 7). The $\text{H}_2\text{O}/\text{CO}_2$ and $\text{H}_2\text{O}/\text{H}_2\text{S}$ MultiGAS peaks appear to occur during crustal movements (micro-seismicity and ground deformation detected with continuous GPS measurements, Figure 7). The largest seismic events recorded for the period of this study were M_w 2.2 (11 July 2013) and M_w 1.5 (26 April 2013). Following these events and throughout the aftershock period (11 July to 7 August and 26 April to 25 May 2013, respectively) the most extensive increase in the $\text{H}_2\text{O}/\text{CO}_2$ and $\text{H}_2\text{O}/\text{H}_2\text{S}$ MultiGAS ratios were observed. Similarly, moderate increases for the same gas ratios were observed during periods of moderately sized peaks of an accumulative seismic moment. Peaks of high $\text{H}_2\text{O}/\text{CO}_2$ and $\text{H}_2\text{O}/\text{H}_2\text{S}$ MultiGAS ratios also seemed to follow periods of ground uplift when associated with the seismic events recorded. No gas peaks were observed during period of subsidence associated with low seismic activity (Figure 7).

The $\text{H}_2\text{O}/\text{CO}_2$ and $\text{H}_2\text{O}/\text{H}_2\text{S}$ MultiGAS ratios showed greater variations (1-217 and 9-10,300, respectively) than the ratios in our fumarole samples ($\text{H}_2\text{O}/\text{CO}_2$ range 184-269 and $\text{H}_2\text{O}/\text{H}_2\text{S}$ range 1080-4527). During periods of recorded crustal movements, the $\text{H}_2\text{O}/\text{CO}_2$ and $\text{H}_2\text{O}/\text{H}_2\text{S}$ MultiGAS ratios increase, and are closer, or equal to the gas ratios in the fumarole samples (Table 3).

It is proposed that the great variations in $\text{H}_2\text{O}/\text{CO}_2$ and $\text{H}_2\text{O}/\text{H}_2\text{S}$ MultiGAS ratios are proposed to be related to the intensity of fumarole activity. The fumarole activity is low during periods of low or no recorded seismic events and land subsidence. During these periods low $\text{H}_2\text{O}/\text{CO}_2$ and $\text{H}_2\text{O}/\text{H}_2\text{S}$ MultiGAS ratios are obtained due to a significant H_2O condensation (i.e., removal of H_2O from the gas phase) before the steam reaches the inlet tube. Due to the MultiGAS setup in Krýsuvík (inlet 20 cm above ground) the low $\text{H}_2\text{O}/\text{CO}_2$

and H₂O/H₂S MultiGAS ratios (<180 and <1000, respectively) are not representative of the emitted fumarole gas composition.

The fumaroles are interpreted to be more active during periods of recorded crustal movements. During these periods the steam rises faster, because of opening of new pathways and increased boiling by pressure release caused by seismicity. As a result, H₂O is less affected by pre-sampling condensation, which results in higher H₂O/CO₂ and H₂O/H₂S gas ratios measured by the MultiGAS.

4.2.2 Comparison with CO₂/H₂S MultiGAS ratio

No visual correlation is observed between CO₂/H₂S ratios obtained by the MultiGAS and periods of recorded crustal movements (Figure 8). These gas species are significantly less affected by condensation processes than H₂O and their detection was less dependent on the variations in the activity of the fumarole. The CO₂/H₂S ratios obtained from the MultiGAS data generally fall within the range of ratios determined in fumarole steam from Krýsuvík. However, the MultiGAS data shows more variability than the fumarole data (this study, Arnórsson, 1987 and data from the Iceland GeoSurvey database). This is considered to be the result of the MultiGAS station running near-continuously for over 7 months, thereby picking out short-timescale variations that may be missed by point sampling. The variations observed in CO₂/H₂S ratios are therefore believed to be linked to variations in the degassing behavior of the system.

4.3 Origin of the Krýsuvík gas emissions

The degassing regime in Krýsuvík is shown to be highly variable over short timescales (hours and days) with changes in fumarole activity and fluctuations in gas composition which is linked to variations in seismicity along with ground deformation. The results from the MultiGAS and fumarole samples resemble a typical composition of low temperature (<100°C) fumarole steam (Lee et al., 2005) from a liquid dominated system (Goff and Janik, 2000). The gas composition from this study is in good agreement with results of previous gas measurements in Krýsuvík (e.g., Arnórsson, 1987) and indicates no increased magmatic gas contribution. Gas geothermometers from this study also show no obvious change in sub surface temperature compared to previous studies (Arnórsson and Gunnlaugsson, 1985, Arnórsson, 1987, Yohannes, 2004), indicating a steady heat flow from the heat source over the time period. Therefore, we cannot conclude that the SO₂ observed is connected with new magma intruding into the systems' roots.

SO₂ concentrations are not expected to be high in geothermal systems like Krýsuvík, due to abundant water within the system where hydrolysis reactions change SO₂ into H₂S, sulfuric acid (H₂SO₄) and elemental sulfur (S) (e.g., Ármannsson et al., 1981, 1989, Ármannsson and Hauksson, 1980, Gíslason et al., 1978, Óskarsson, 1978, 1984). However, the SO₂ detected here (0.05-0.5 ppmv) is not considered to be a false signal or interference from other gas species. This is based on two main reasons: (1) SO₂ is observed independently from high concentrations of H₂S and therefore the SO₂ measurements are unlikely to be interference between the H₂S and SO₂ sensors. (2) The MultiGAS station is equipped with highly sensitive sensors that are capable of detecting very low concentrations. It is considered possible that small amounts of SO₂ may be present in emissions from Krýsuvík. Magmatic SO₂ might be able to ascend rapidly to the surface during periods of elevated seismicity. Another potential, and probably more likely, source

of the SO₂, is near surface oxidation of H₂S to SO₂ (Arnórsson, 1987), with the formation of elemental sulfur (S) at the fumarolic vent (Lee et al., 2005). The presence and source of SO₂ in Krýsuvík, needs further studying, where different methods are applied.

The results of several studies (e.g., Noguchi and Kamiya 1963, Casadevall et al., 1983, Fischer and Arehart, 1996) have shown changes in the fumarole gas composition prior to and during eruptive events. Studies from the rifting episodes in Krafla, NE-Iceland during 1975-1984 (a volcanic system which bears many similarities to Krýsuvík), revealed changes in local fumarolic gas composition (e.g., Óskarsson, 1984, 1978, Ármannsson et al., 1989, Gíslason et al., 1984). The gas composition was CO₂-rich during the first weeks of rifting and remained unchanged until 1983. The outgassing CO₂ was released from the deep aquifers beneath the area by the interaction of magmatic gas with the hydrothermal system.

A comprehensive record and monitoring in the Krýsuvík region would be a great asset in understanding the gas source and the degassing regime of Krýsuvík. When new magma intrudes into the roots of the Krýsuvík system, the fumarolic gas composition might change in similar way to what was seen at Krafla 1975-1984 (e.g., Óskarsson, 1984, 1978, Ármannsson et al., 1989, Gíslason et al., 1984). The gas composition in Krýsuvík would be expected to become CO₂ richer, resulting in lowering of H₂O/CO₂, and increase in CO₂/H₂S MultiGAS- and spot sampling ratios, resulting in greater CO₂ gas fluxes through soil.

Several smaller localities with apparent surface activity (mostly related to the Trölladyngja subarea) were not included in this study. We therefore conclude that the total soil CO₂ flux from the Krýsuvík geothermal system during this study is greater than 101 T/day. Furthermore, the total CO₂ soil flux estimated here should be considered as the minimum value given that the amount of CO₂ dissolved in groundwater is unknown.

The total measured soil CO₂ flux from the neighboring Reykjanes (~0.4 km²) and Hengill volcanic systems (168 km²) (using the accumulation chamber method) indicate a total soil CO₂ flux of 78.5±13.9 T/day (Óladóttir, 2014) and 1,526 ± 160 T/day of which 453 T/day of volcanic/hydrothermal origin (Hernández et al., 2012), respectively.

Studies of the volcanic systems along the RP during the last two decades have shown little evidence of magmatic contributions with the possible exception of the Hengill volcano (Hreinsdóttir et al., 2001, Einarsson, 2008, Keiding et al., 2008), where minor magma injection into the roots of the volcano has been suggested to have taken place from 1995 to 1998, triggering an intense seismic swarm (Sigmundsson et al., 1997).

The last eruptive activity took place in Krýsuvík during the 1151-1188 Krýsuvík fires. Based on the eruption history of RP, with several hundred years of volcanic quiescence between volcanically active periods, Krýsuvík may enter the next episode within the next decades. The recent episodes of uplift or subsidence could indicate that the system is heating up. It is important to continue the monitoring of gas emissions in Krýsuvík, at least over the next episode of ground uplift and elevated seismicity to compare with the dataset from this study. It would be of particular interest to observe if changes in the gas composition will occur during an inflation episode, as that will give important information on the source of the inflation.

5 Conclusions

The gas composition from Krýsuvík geothermal system was studied (April-November 2013) using the MultiGAS method and compared with geophysical observations. The gas composition is H₂O dominated with CO₂ as the dominant dry gas species, with smaller concentrations of H₂S and trace concentrations of SO₂.

The gas emissions (in the form of diffuse soil degassing) were measured in three areas of high geothermal surface activity within the Krýsuvík system (Seltún, Hveradalir and Austurengjar). The total emission is estimated as 101 T/day but this is a minimum value for the total emission, as the amount of CO₂ dissolved in groundwater is not known and several smaller localities with apparent surface activity were not studied.

The time series of gas composition (MultiGAS data) identified short-lived episodes of elevated and highly variable H₂O/CO₂ and H₂O/H₂S ratios, the highest H₂O/CO₂ and H₂O/H₂S ratios followed the periods of highest accumulative seismic moment per day 26 April 2013 and 11 July 2013. SO₂ was detected at the same time. Comparison with seismicity and ground deformation shows that the ratio peaks follow increased seismic activity, with the highest H₂O/CO₂ and H₂O/H₂S ratios following elevated seismicity and crustal movements. We propose that these peaks represent periods of elevated fumarolic activity that is responding to the seismicity and land uplift due to opening of new pathways in the crust and increased boiling within the system. During these periods the steam reaches the inlet tube faster and is less affected by pre-sampling H₂O condensation. The CO₂ and H₂S gases are significantly less affected by condensation and the CO₂/H₂S ratio does not change according to the same pattern. Most of the CO₂/H₂S ratios determined fall within the known range for Krýsuvík fumaroles. However, several markedly higher values are detected, demonstrating more variability in the degassing system than previously known.

It is considered crucial to continue the MultiGAS measurements in the Krýsuvík geothermal due to recent episodes of ground uplift and subsidence as well as future volcanic activity. During the span of the MultiGAS measurements in 2013 the total subsidence in the area was ~21 mm.

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Figure 1 Volcanic systems on the RP (purple) and seismic zone across the Peninsula that marks the axis of the plate boundary (brown) (Einarsson, 2008). High-temperature geothermal areas within the volcanic systems (striped). Modified from (Sæmundsson and Sigurgeirsson, 2013). The five volcanic systems are arranged en echelon along the peninsula, spaced approximately 5 km apart (Clifton and Katterhorn, 2006). Continuous GPS stations in operation in 2013 on the RP including the region of the earthquake data (dashed box), and the center of uplift (orange circle). Blue and dark red dots show the location of Festafjall and Keflavík airport weather stations, respectively.

Figure 2 Outlines of the Krýsuvík high-temperature geothermal system identified by resistivity surveys (orange line) (Gudmundson et al., 1975). Krýsuvík sub-areas and the two hyaloclastite ridges, Sveifluháls and Vesturháls, with which the geothermal activity in Krýsuvík is mostly associated. The green dot shows the location of the MultiGAS station, and the purple dots show the location of the fumarole samples collected.

Figure 3 MultiGAS station in Krýsuvík. The station is located next to a fumarole in an area of continuous geothermal activity in Hveradalir, Krýsuvík, and is powered by a wind turbine. The sampling inlet of the MultiGAS station is located ~20 cm above the ground to avoid saturation of the CO₂ sensor.

Figure 4 The local inflation/deflation periods from GPS time series (2007-2016) at station MOHA, along with daily cumulative seismic moment. The red band corresponds to the time period of this study.

Figure 5 (Upper) CO₂/H₂S and H₂O/CO₂ molar ratios as a function of precipitation (mm/day). The CO₂/H₂S ratios are obtained predominantly during periods with <2 mm/day rainfall. However, high ratio values are not confined to the dry periods. All H₂O/CO₂ ratios >19 are obtained during dry periods (< 2mm/day). The most frequently obtained values for H₂O/CO₂ (<19) are visually evenly distributed between ~ 0-5 mm/day. (Lower) CO₂/H₂S and H₂O/CO₂ molar ratio as a function of wind speed (m/s). The highest ratios of CO₂/H₂S (>75) and largest variations are obtained during relatively low wind speed (approximately <5 m/s). Lower values of CO₂/H₂S (<20) are detected more frequently than the higher values and there is no visible relation to wind speed <10 m/s where marked decrease is seen in the frequency of ratios obtained. Wind speed <10 m/s does not appear to affect detection of H₂O/X molar ratios (X = CO₂, H₂S and SO₂) in >10 m/s markedly fewer ratios are detected.

Figure 6 The results of soil CO₂ flux measurements from the three observation areas, Hveradalir, Seltún, and Austurengjar. The highest flux was observed in Hveradalir, 70.9 T/day. In Seltún the flux was found to be 19.6 T/day and the lowest flux was calculated in Austurengjar, 10.9 T/day. The total flux from the three measured areas was calculated as 101 T/day.

Figure 7 Normalized variations in gas composition as measured by the MultiGAS station (measurements affected by metrological conditions have been eliminated) correlated with geophysical observations. There are distinct intervals with peaks of increased H₂O/CO₂ and H₂O/H₂S ratios (upper figure). SO₂ is detected during the same intervals allowing calculation of X/SO₂ ratios (lower figure). Red line: cumulative seismic

moment (Nm). Black line: seismic moment per day (Nm/day). Blue curve: Vertical crustal movements measured with GPS (mm). Light grey intervals: rainy days (>2 mm/day). Dark grey interval: station not operating. Peaks of increased $\text{H}_2\text{O}/\text{CO}_2$ and $\text{H}_2\text{O}/\text{H}_2\text{S}$ ratios appear to follow episodes of recorded seismic events and crustal deformation.

Figure 8 Time series (April-November 2013) of $\text{CO}_2/\text{H}_2\text{S}$ MultiGAS molar ratios (where measurements effected by metrological conditions have been eliminated) compared with crustal movements. The $\text{CO}_2/\text{H}_2\text{S}$ ratio does not show any visible variations related to seismicity or crustal deformation. The yellow band corresponds to $\text{CO}_2/\text{H}_2\text{S}$ ratios (3-41) from fumaroles in the Krýsuvík area (this study, Arnórsson, 1987, data from the Iceland GeoSurvey database).

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Table 1 Vol% concentration of selected gases in fumarolic steam from the 8 samples. Values within () for CO₂ and H₂S refer to % of total gas volume.

Sample	Date	H ₂ O%	CO ₂ %	H ₂ S%	H ₂ %	Ar%	O ₂ %	N ₂ %	CH ₄ %
Seltún 1	18.02.2014	96.9	75.5 (2.3)	13.9 (0.430)	9.32	0.02	0.15	1.03	0.08
Seltún 2	23.04.2014	99.4	79.6 (0.48)	15.3 (0.092)	4.80	-	-	0.32	0.02
Hverahvammur 1	18.02.2014	99.5	88.9 (0.44)	8.69 (0.043)	2.02	-	-	0.40	-
Hverahvammur 2	23.04.2014	99.4	89.3 (0.54)	9.26 (0.056)	0.97	-	-	0.49	-
Austurengjar 1	18.02.2014	99.4	89.3 (0.54)	3.80 (0.023)	6.13	-	-	0.67	0.10
Austurengjar 2	23.04.2014	99.5	81.5 (0.41)	12.9 (0.065)	5.36	-	-	0.19	0.06
Hverahöfði 1	18.02.2014	99.4	73.4 (0.44)	4.00 (0.024)	19.7	0.04	-	2.65	0.22
Hverahöfði 2	23.04.2014	99.6	93.2 (0.37)	5.54 (0.022)	0.76	-	-	0.49	0.02

Table 2 Variations of molar gas ratios measured by the MultiGAS station in Hveradalir, Krýsuvík (excluding data affected by meteorological factors).

	Max	Min	Average	Median
H₂O/CO₂	217	1	27	9
H₂O/H₂S	10300	9	640	218
H₂O/SO₂	107000	2380	26900	25100
CO₂/H₂S	107	4	17	12
CO₂/SO₂	3010	58	767	662
H₂S/SO₂	148	11	42	35

Table 3 The calculated molar ratios of $\text{H}_2\text{O}/\text{CO}_2$, $\text{H}_2\text{O}/\text{H}_2\text{S}$ and $\text{CO}_2/\text{H}_2\text{S}$ for the fumarole samples and from the MultiGAS data. Sample Seltún 1 is excluded from the average calculations due to condensation in the sampling train. The maximum $\text{H}_2\text{O}/\text{H}_2\text{S}$ values in the MultiGAS data (10300) is by far the highest value, the second highest value (3850) is, much closer to the average value for the fumarole samples.

Ratio	MultiGAS				Fumaroles								
	Max	Min	Average	Median	Seltún 2	Hverahvammur 1	Hverahvammur 2	Austur-engjar 1	Austur-engjar 2	Hverahöfi 1	Hverahöfi 2	Average	Median
$\text{H}_2\text{O}/\text{CO}_2$	217	1	27	9.0	207	226	184	184	243	226	269	220	226
$\text{H}_2\text{O}/\text{H}_2\text{S}$	10300 (3850)	9	640	218	1080	2313	1775	4321	1530	4141	4527	2812	2313
$\text{CO}_2/\text{H}_2\text{S}$	107	4	17	12	5	10	10	24	6	18	17	13	10

Table 4 Gas geothermometers applied to Krýsuvík fumarole samples.

Geothermometer	Temperature function ¹	Reference
CO ₂	$4.724Q^3 - 11.068Q^2 + 72.012Q + 121.8$	(Arnórsson et.al 1998)
H ₂ S	$4.811Q^2 + 66.152Q + 177.6$	(Arnórsson et.al 1998)
H ₂	$6.630Q^3 + 5.836Q^2 + 56.168Q + 227.1$	(Arnórsson et.al 1998)
CO ₂	$-44.1 + 269.25Q - 76.88Q^2 + 9.52Q^3$	(Arnórsson and Gunnlaugsson 1985)
H ₂ S	$173.2 + 65.04Q$	(Arnórsson and Gunnlaugsson 1985)
H ₂	$212.2 + 65.04Q$	(Arnórsson and Gunnlaugsson 1985)

¹Q refers to the logarithm of the respective gas concentration or ratio in moles per kg of steam.

Table 5 The results from the gas geothermometers applied. The temperature value are in °C.

Sample	nr.	¹CO₂	¹H₂S	¹H₂	²CO₂	²H₂S	²H₂	Average
Hverahvammur 1	1	292	276	274	288	261	259	275
Hverahvammur 2	2	303	268	256	295	270	243	272
Austurengi 1	3	301	255	323	294	244	295	285
Austurengi 2	4	289	290	309	286	273	286	289
Hverahöfði 1	5	292	257	281	288	245	266	271
Hverahöfði 2	6	292	259	244	288	247	231	260
Seltún 2	7	297	304	313	291	284	289	296

¹(Arnórsson et al, 1998) ²(Arnórsson and Gunnlaugsson, 1985).

Highlights

- The gas emissions from the Krýsuvík system are H₂O-dominated, with CO₂ as the most abundant dry gas species, followed by smaller concentrations of H₂S. Gasgeothermometers indicate subsurface temperature of 278°C.
- The degassing regime in Krýsuvík is highly variable over short timescales (hours and days) with changes in fumarole activity and fluctuations in gas composition which is linked to variations in seismicity and ground deformation.
- The diffuse soil CO₂ flux is variable between the three studied degassing areas in Krýsuvík, ranging from 10.9-70.9 T/day.

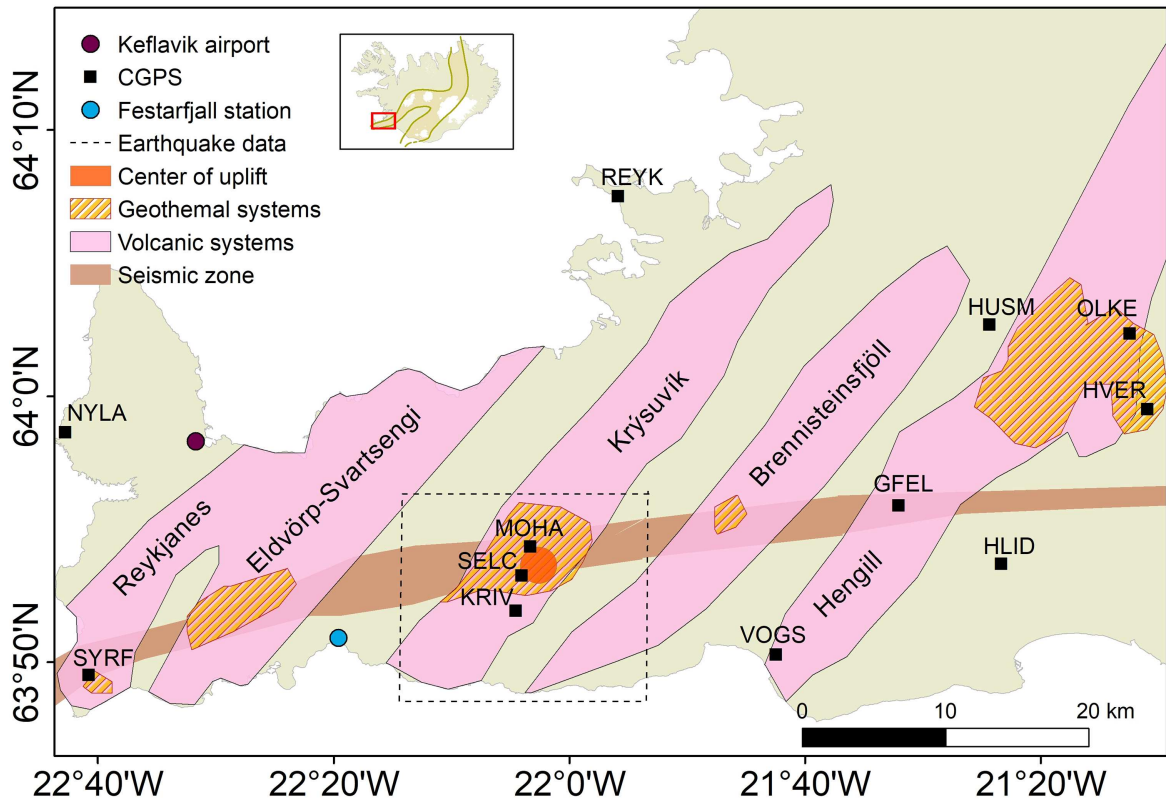


Figure 1

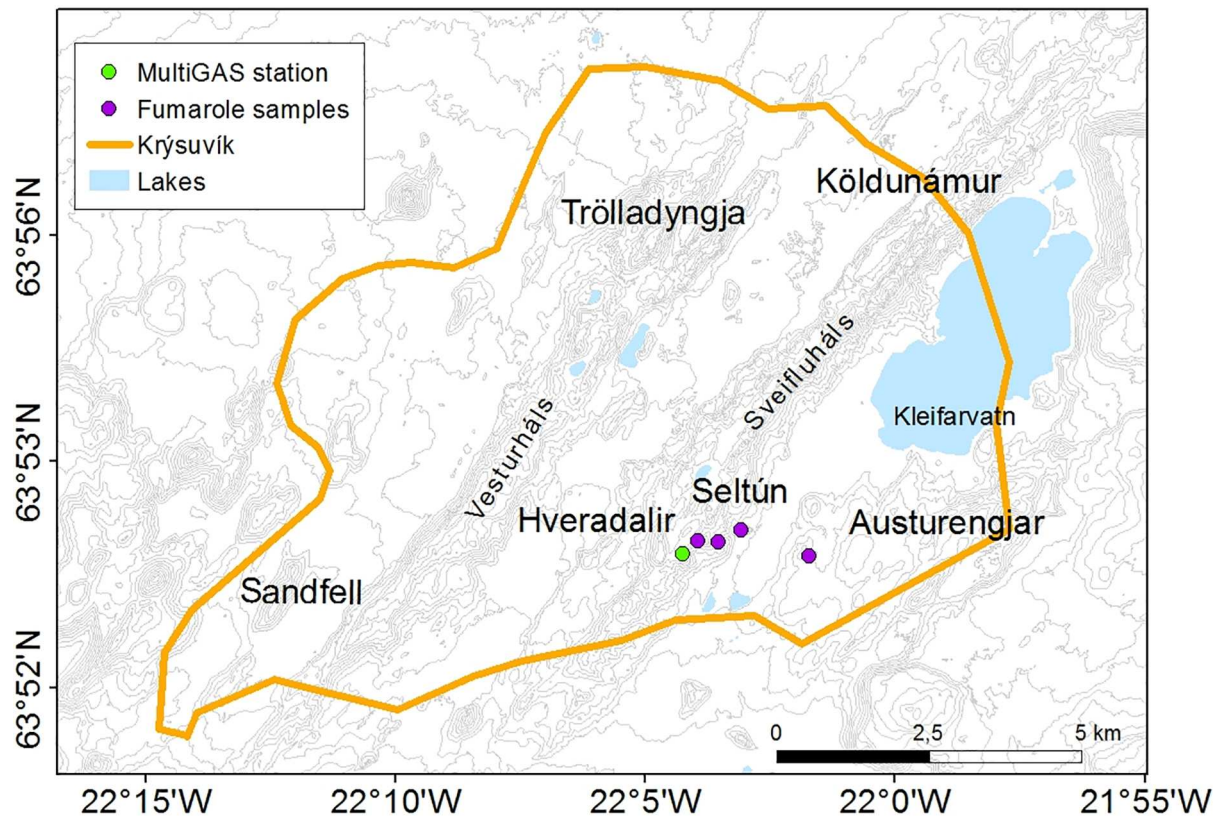


Figure 2

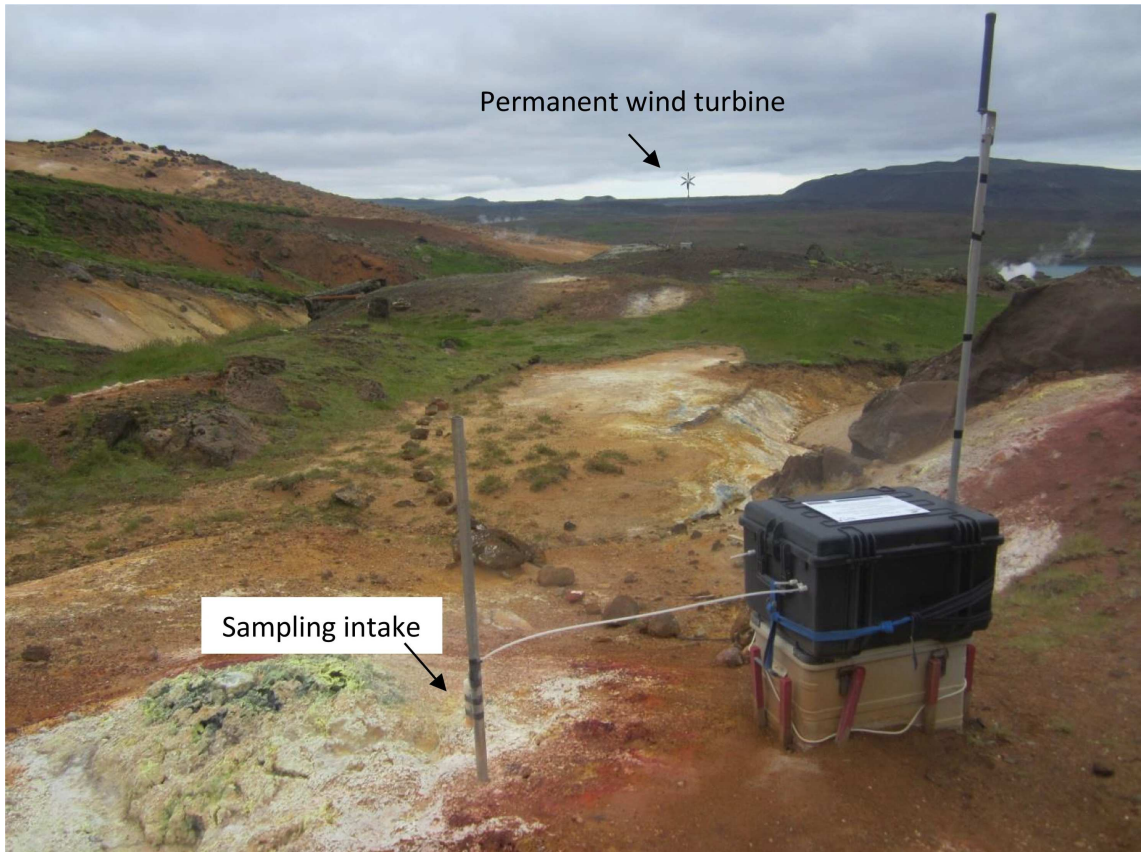


Figure 3

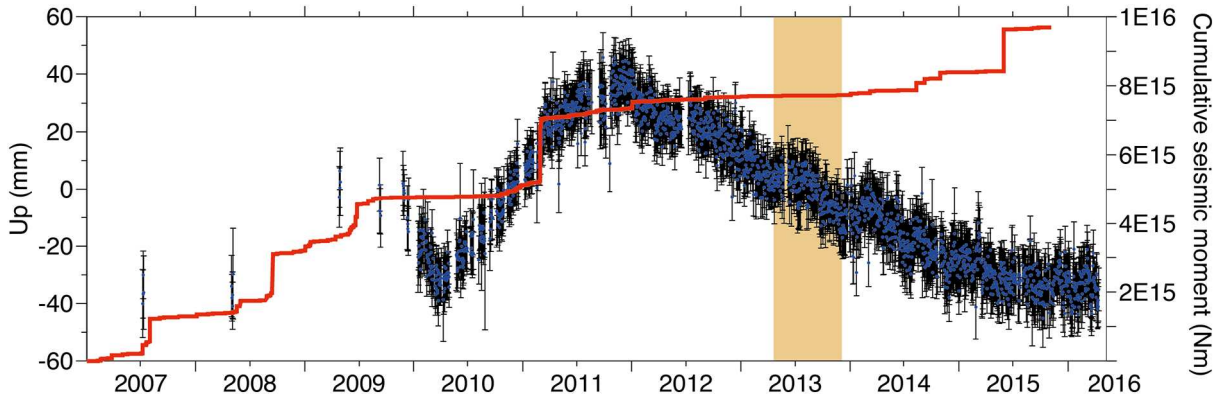


Figure 4

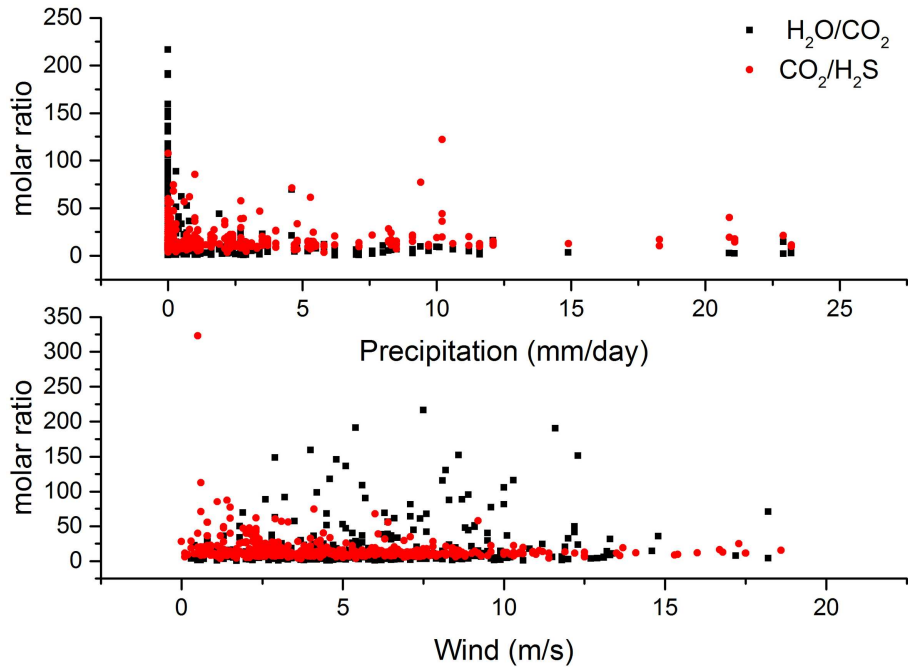


Figure 5

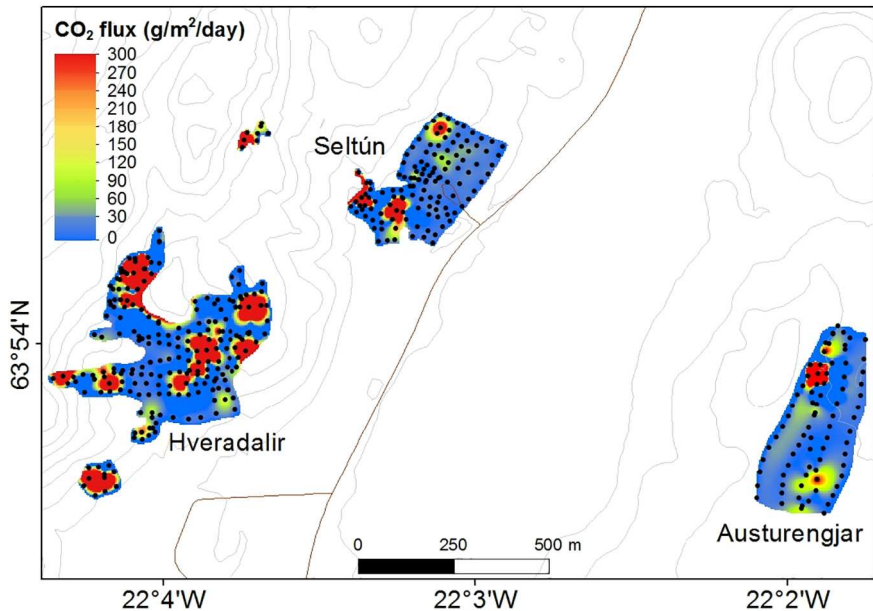


Figure 6

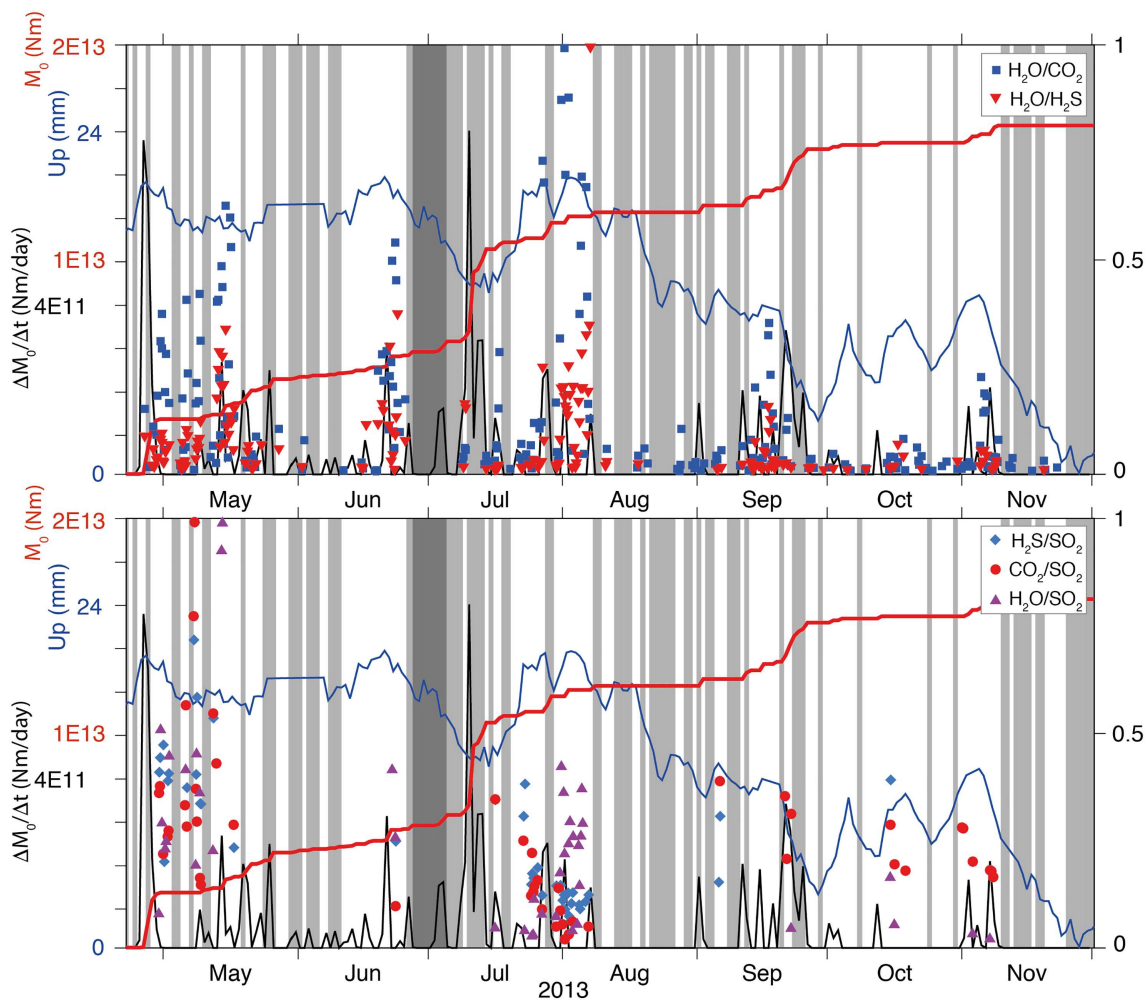


Figure 7

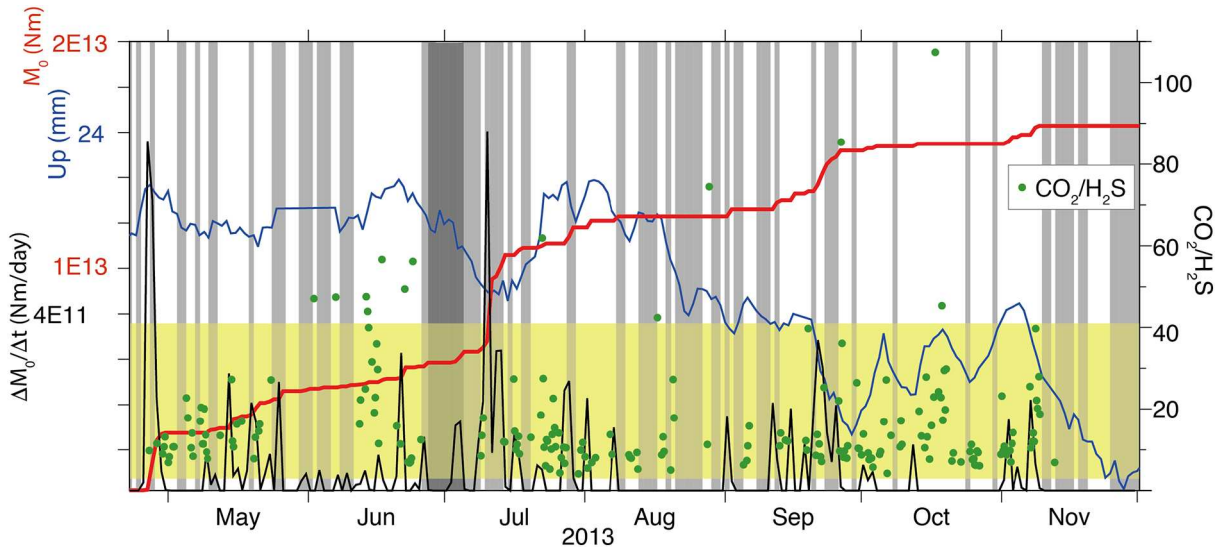


Figure 8