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3 **A HISTORY OF PROTEROZOIC TERRANES IN SOUTHERN SOUTH**  
4 **AMERICA: FROM RODINIA TO GONDWANA**  
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26 **Keywords**

27 *Paleoproterozoic; cratons; Grenvillian; Neoproterozoic rifting; SW Gondwana assembly*  
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31 **Abstract**

32 The role played by Paleoproterozoic cratons in southern South America from the  
33 Mesoproterozoic to the Early Cambrian is reconsidered here. This period involved protracted  
34 continental amalgamation that led to formation of the supercontinent Rodinia, followed by  
35 Neoproterozoic continental break-up, with the consequent opening of Clymene and Iapetus oceans,  
36 and finally continental re-assembly as Gondwana through complex oblique collisions in the late  
37 Neoproterozoic to Early Cambrian. The evidence for this is based mainly on a combination of  
38 precise U-Pb SHRMP dating and radiogenic isotope data for igneous and metamorphic rocks from a  
39 large area extending from the Rio de la Plata craton in the east to the Argentine Precordillera in the  
40 west and as far north as Arequipa in Peru. Our interpretation of the paleogeographical and  
41 geodynamic evolution invokes a hypothetical Paleoproterozoic block (MARA) embracing basement  
42 ultimately older than 1.7 Ga in the Western Sierras Pampeanas (Argentina), the Arequipa block  
43 (Peru) the Rio Apa block (Brazil), and probably also the Paraguaia block (Bolivia).  
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56 **Introduction**  
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3 The role of southern South American cratons in Rodinia reconstructions, particularly those of  
4 Amazonia and Rio de la Plata, is a long-debated issue (Hoffman, 1991; Dalziel, 1997; Weil et al.,  
5 1998; Omarini et al., 1999; Loewy et al., 2003; Li et al., 2008; Trindade et al., 2006; Cordani et al.,  
6 2010; Santosh et al., 2009, among others). The debate was stimulated by the ideas that 1) Eastern  
7 Laurentia was juxtaposed to Amazonia and the Rio de la Plata craton in Rodinia at ca. 1 Ga as a  
8 result of Grenvillian orogeny, and that 2) Laurentia rifted away to its present position in the  
9 northern hemisphere (present coordinates) in the late Neoproterozoic, accompanied by opening of  
10 the Iapetus Ocean and the final amalgamation of West Gondwana (Hoffman, 1991; Dalziel, 1997  
11 and references therein). However, the models derived from these studies took only minor account of  
12 the relatively small outcrops with Paleoproterozoic basement south of Amazonia and west of the  
13 Rio de la Plata craton. The wealth of data now available from detrital zircon ages and crystallization  
14 ages of many igneous rocks has transformed this situation. These outcrops are scattered over a very  
15 large region (Fig. 1), with an extensive cover of Mesozoic to Cenozoic sedimentary rocks, which  
16 hinders correlation between them. The Paleoproterozoic rocks crop out as inliers within the Andean  
17 belt (e.g., the Arequipa block in Peru; Loewy et al., 2004, and references therein; Casquet et al.,  
18 2010), in the Andean foreland (Sierra de Maz in the Western Sierras Pampeanas of central  
19 Argentina; Casquet et al., 2006, 2008a) and in the stable mainland far from the Andean active  
20 margin (e.g., Rio Apa and Paraguá in southern Brazil; Cordani et al., 2010) (Fig. 1). Other  
21 occurrences may be hidden farther south in Argentine Patagonia. In consequence, the role of these  
22 outcrops in Rodinia reconstructions has been largely underestimated, hindering understanding of the  
23 role played by cratons in the Neoproterozoic-to-Early Paleozoic evolution of southern South  
24 America after the break-up of Rodinia.

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26 The presence of a pre-Grenvillian continental mass called Pampia has been suggested  
27 (Ramos, 1988; Ramos & Vujovich, 1993), initially as a block embracing most of the present Sierras  
28 Pampeanas realm, with a late Neoproterozoic turbidite basin (the Puncoviscana Formation) along  
29 the western passive margin that eventually collided with the Arequipa-Antofalla block. This view  
30 was largely retained by Ramos et al. (2010) in a recent review of Pampia.

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32 Based largely on our own work since the 1990s in the pre-Andean basement, we proposed  
33 (Casquet et al., 2009) that by the end of the Paleoproterozoic the basement outcrops referred to  
34 above, i.e., Arequipa, Rio Apa and Maz, constituted a single larger continental mass (the MARA  
35 craton, Fig. 1). Part of this craton was involved in Mesoproterozoic orogenies along its northern and  
36 western margins that led first to its accretion to Amazonia at ca. 1.3 Ga or earlier, and then  
37 amalgamation with Laurentia between ca 1.3 and 1.0 Ga. The latter event involved accretion of  
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2 juvenile arcs and continental collision with reworking of older continental crust. After Rodinia  
3 break-up the larger continental mass embracing MARA + Laurentia + Amazonia (and probably  
4 other still unconstrained cratons) underwent protracted Neoproterozoic rifting, as exemplified by A-  
5 type granite and carbonatite-syenite intrusions. Opening of an oceanic basin was eventually  
6 followed by oblique collision with some West Gondwana cratons (including Rio de la Plata and  
7 Kalahari) to produce the Pampean-Paraguay-Araguaia orogeny. This process was coeval with  
8 rifting-drifting of Laurentia and the opening of the Iapetus Ocean along the western margin of the  
9 large continental mass and represents the final stage in the formation of SW Gondwana. We provide  
10 here a detailed account of this evolution.  
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## 21 **1. The Paleoproterozoic MARA craton**

22 In the Western Sierras Pampeanas of Argentina, the Maz terrane (comprising the sierras of  
23 Maz and Espinal) (Fig. 1) consists of a metamorphic Andean-type magmatic arc (1.33 – 1.26 Ga)  
24 and older metasedimentary rocks. (Casquet et al., 2006, Rapela et al., 2010). The latter contain  
25 detrital zircons older than 1.7 Ga and have Nd model ages of between 1.7 and 2.6 Ga and very  
26 radiogenic Pb, from which we infer that the protoliths were probably cover to a Paleoproterozoic  
27 basement older than 1.7 Ga (Casquet et al., 2008a). The Maz terrane was further reworked in the  
28 Ordovician and Silurian by the Famatinian orogeny (Casquet et al., 2005).  
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35 The Arequipa Massif in southern Peru (Cobbing and Pitcher, 1972) (Fig. 1) consists for the  
36 most part of Paleoproterozoic metasedimentary rocks and orthogneisses that record orogenic events  
37 (magmatism, sedimentation and metamorphism) between ca. 1.79 and 2.1 Ga (Loewy et al., 2004;  
38 Casquet et al., 2010). These rocks underwent intense Grenville-age (*sensu* Rivers, 2008)  
39 metamorphism between 1.04 and 0.84 Ga (Martignole and Martelat, 2003; Loewy et al., 2004;  
40 Casquet et al., 2010 and references therein). UHT metamorphism first recorded in the Arequipa  
41 Massif by Martignole and Martelat (2003) remains of disputed age, either Paleoproterozoic or  
42 Grenvillian (Martignole and Martelat, 2003; Casquet et al., 2010). Mixing of Paleoproterozoic and  
43 juvenile Grenvillian sources was recognized farther south in northern Chile and Argentina (Loewy  
44 et al., 2004), lending support to the idea of a continuous basement of Paleoproterozoic age under  
45 this part of the Central Andes, corresponding to the northern part of the Arequipa-Antofalla craton  
46 of Ramos (1988). A link between the Maz terrane and the northern part of the Arequipa-Antofalla  
47 craton was first established by Casquet et al. (2008a) on the basis of detrital zircon evidence and Pb  
48 and Nd isotope compositions.  
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3 Finally the Rio Apa block, south of present-day Amazonia (Fig. 1), consists of a suite of  
4 Paleoproterozoic orthogneisses recording igneous events at 1.95 Ga, 1.84 Ga, and between 1.72 and  
5 1.77 Ga, as well as medium- to high-grade metamorphism of ca. 1.7 Ga (Cordani et al., 2010). The  
6 Rio Apa block was overprinted by a thermal event at ca. 1.3 Ga (Cordani et al., 2010) coincident  
7 with the San Ignacio orogeny along the southern margin of Amazonia (Böger et al., 2005; Cordani  
8 and Teixeira, 2007).

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10 The three cratonic outcrop areas referred to above, although separated by hundreds of  
11 kilometres, show evidence of common geological processes evidenced by U-Pb geochronology and  
12 similar Nd model ages (Fig. 2). This leads to the idea that all three formed part of a common  
13 continental mass before the onset of the Mesoproterozoic orogenies. We call this craton MARA  
14 (after **M**az - **A**requipa - **R**io **A**pa), consisting of rocks formed between 1.7 and 2.1 Ga and with Nd  
15 residence ages ( $T_{DM}$ ) between 1.7 and 2.6 Ga (Casquet et al., 2009).  
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## 26 **2. The Mesoproterozoic Evolution of the Mara Craton**

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28 No evidence has so far been recognized in any of the three Paleoproterozoic outcrops for  
29 significant igneous or metamorphic activity between ca. 1.3 and 1.6 Ga, although they experienced  
30 contrasting igneous and metamorphic events in the second half of the Mesoproterozoic (Fig. 3).  
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33 The Maz terrane records an Andean-type magmatic arc (1.33 – 1.26 Ga) and intermediate P/T  
34 amphibolite to granulite facies metamorphism between 1.23 and 1.17 Ga followed by emplacement  
35 of AMCG complexes at 1.07-1.09 Ga (Casquet et al., 2005, 2006; Rapela et al., 2010). Moreover in  
36 the nearby Sierra de Pie de Palo, and in minor outcrops south of it (Fig. 1), a late Mesoproterozoic  
37 juvenile arc/back-arc oceanic complex has been identified (the Pie de Palo Complex) that records  
38 protracted subduction between ca. 1.24 and 1.03 Ga (Kay et al., 1996; Vujovich et al., 2004; Rapela  
39 et al., 2010). The Pie de Palo Complex is the basement of the enigmatic, supposedly Laurentia-  
40 derived, Precordillera terrane (Thomas and Astini, 1996; for a review see Ramos, 2004):  
41 alternatively the terrane might have been para-autochthonous (Aceñolaza and Toselli, 2000; Galindo  
42 et al., 2004; Finney, 2007). In any hypothesis, docking of this terrane to the margin of Gondwana  
43 occurred in the mid-Ordovician during the Famatinian orogeny (Ramos et al., 1998; Ramos, 2004;  
44 Casquet et al., 2001; Galindo et al., 2004). Overlying the Pie de Palo Complex is an imbricate thrust  
45 system that reworked basement consisting of late Mesoproterozoic orthogneisses and  
46 metasedimentary rocks overlain by a Neoproterozoic sedimentary cover (Casquet et al., 2001;  
47 Mulcahy et al., 2011). This basement underwent pre-Famatinian metamorphism under conditions  
48 close to those of the Maz terrane, to which it is probably equivalent (Casquet et al., 2001). Orogenic  
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2 activity between ca 1.3 and 1.0 Ga in the Maz terrane and the Pie de Palo Complex can be  
3 interpreted as resulting from the approach and eventual collision of the MARA craton (and the  
4 juxtaposed Amazonia) with Laurentia to produce the middle to late Mesoproterozoic orogenic belt  
5 that fringes Amazonia on the west, with outcrops as far north as Colombia (Cardona et al., 2010,  
6 and references therein) (Fig. 1). Paleomagnetic data for ca. 1.2 Ga are compatible with this  
7 interpretation (Tohver et al., 2004). The relative positions of the oceanic Pie de Palo Complex and  
8 the continental Maz terrane in the Mesoproterozoic orogen are difficult to retrieve because of  
9 Famatinian oblique thrusting in Sierra de Pie de Palo and protracted post-Paleozoic activity along  
10 the Bermejo-Desaguadero fault that separates the block containing the Sierra de Pie de Palo and the  
11 Argentine Precordillera from that containing the sierras of Maz and Espinal (Fig. 1).  
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15 The Rio Apa block underwent a strong thermal episode at ca. 1.3 Ga with temperatures above  
16 300°C, which affected the entire region (Cordani et al., 2010). This corresponds to the San Ignacio  
17 orogeny (1.34–1.32 Ga; Böger et al., 2005), the main belt of which developed farther north, along  
18 southern Amazonia (Fig. 1), and is one of several 1.56–1.3 Ga orogenic belts constituting the  
19 Rondonia-San Ignacio province of southern Amazonia (Bettencourt et al., 2010, and references  
20 therein). The lack of evidence in the Rio Apa block for the late Mesoproterozoic Sunsás orogeny s.l.  
21 (1.20–1.07 Ga; Cordani and Teixeira, 2007; Böger et al., 2005) suggests that it - and consequently  
22 the MARA craton - was accreted to the southern Amazonia margin during the San Ignacio orogeny,  
23 and was a mainly stable region in late Mesoproterozoic times. Deformation associated with the  
24 Sunsás orogeny, long considered the main representative of the Grenville orogeny in southern South  
25 America, occurred farther north along branched transcurrent belts and pull-apart basins (Fig. 1)  
26 involving local metamorphism and granitic magmatism (for a review see Teixeira et al., 2010 ). The  
27 Paraguá block of Eastern Bolivia (Fig. 1) also has Paleoproterozoic basement older than ca. 1.7 Ga  
28 (Böger et al., 2005) that was accreted to SW Amazonia during the San Ignacio orogeny  
29 (Bettencourt et al., 2010) and could thus also have been part of MARA.  
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33 The history of the Arequipa massif differs in that it shows evidence for true Grenville-age  
34 (*sensu* Rivers, 2008) low-P high-T regional metamorphism between ca 1.04 and 0.85 Ga, younger  
35 than in the Maz terrane and Rio Apa block (Loewy et al., 2004; Casquet et al., 2010). The massif  
36 was probably an inlier in the middle to late Mesoproterozoic orogenic belt that only underwent late-  
37 orogenic metamorphism. However its pre-orogenic location and the geodynamic setting of  
38 metamorphism remain uncertain. With respect to location Dalziel (1992, 1994) and Sadowsky and  
39 Bettencourt (1996, and references therein) proposed that the Arequipa massif was the tip of a  
40 promontory of NE Laurentia. Subsequently Loewy et al. (2004) suggested that the Arequipa massif  
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2 and its southward extension may have been part of a larger craton in collision with Amazonia. With  
3 respect to metamorphism an extensional setting in the Grenvillian hinterland has been hypothesized  
4 on the grounds that extension was widespread over southern Amazonia at this time (equivalent to  
5 the Rigolet event of the Grenville orogeny; Casquet et al., 2010). Significantly, accretion of MARA  
6 to Amazonia during the San Ignacio orogeny would explain the input of detrital zircons with ages  
7 between 1.2 and 1.6 Ga to the late Mesoproterozoic Atico basin in Arequipa, for which no local  
8 sources have been recognized (Casquet et al., 2010). Amalgamation of Laurentia and the MARA  
9 craton (with Amazonia) in the Mesoproterozoic at ca. 1.2 Ga (Tohver et al., 2002) was an  
10 important contribution to the formation of Rodinia.  
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### 21 **3. Neoproterozoic to Early Paleozoic evolution**

22 The Neoproterozoic to Early Paleozoic history is summarized focusing on evidence from the  
23 Sierras Pampeanas of Argentina.  
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#### 27 *3.1 Rifting events and the Clymene Ocean*

28 Protracted rifting of Rodinia took place throughout the Neoproterozoic. Two early aborted  
29 rifting events, at ca. 840 and 760 Ma, are represented by A-type granitoids in Sierra de Maz and  
30 Sierra de Pie de Palo, bearing zircons with juvenile Hf and O isotopic signatures (Baldo et al., 2006;  
31 Colombo et al., 2009; Rapela et al., 2011). Further rifting occurred at ca. 570 Ma (Ediacaran),  
32 represented by a carbonatite-nepheline syenite complex in the Sierra de Maz (Casquet et al.,  
33 2008b). We suggest that this latter event probably initiated opening of the Clymene Ocean (Fig. 4).  
34 This ocean was named by Trindade et al. (2006), who argued on the basis of paleomagnetism for  
35 such a late Neoproterozoic ocean between Amazonia + Laurentia on the one hand and West  
36 Gondwana cratons, such as Rio de la Plata and Kalahari, on the other. As an alternative to the  
37 Pampia model of Ramos (1988) and Ramos et al. (2010), we envisage that MARA was attached to  
38 the former continental mass and that the closure of this ocean led to the Pampean orogeny (Casquet  
39 et al., 2009). 570 Ma is just within error of the Sm-Nd age of  $647 \pm 77$  Ma for alleged Pampean  
40 ophiolite relics (whole-rock errorchron with MSWD = 7.6), obtained by Escayola et al. (2007).  
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#### 55 *3.2 The Difunta Correa Sedimentary Sequence*

56 Further evidence for the Clymene Ocean comes from the Sr-isotope composition of platform  
57 carbonates of the Difunta Correa Sequence, which was deposited on the Paleoproterozoic and  
58 Mesoproterozoic basement of the Western Sierras Pampeanas in the late Neoproterozoic (Varela et  
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2 al., 2001; Galindo et al., 2004; Rapela et al., 2005; Murra et al., 2011). From comparison with the  
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4 Sr-isotope composition of seawater through time, Galindo et al. (2004) deduced a maximum age of  
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6 580 Ma (Ediacaran) for the sequence in Sierra de Pie de Palo. This accords with the similar findings  
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8 of Varela et al. (2001) for equivalent carbonate cover in Sierra de Umango (Western Sierras  
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10 Pampeanas) and of Murra et al. (2011) for marbles from Sierra de Ancasti (Eastern Sierras  
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12 Pampeanas). Similar Ediacaran shallow-marine carbonates that were post-glacial with respect to the  
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14 Marinoan (ca. 635 Ma) and Gaskiers (ca. 580 Ma) events are recorded elsewhere in southern South  
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16 America (Misi et al., 2007). This evidence for extensive carbonate platforms at ca. 580 Ma is  
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18 compatible with the existence of the Clymene Ocean during Ediacaran time (Fig. 4).  
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### 20 21 *3.3 The Puncoviscana Formation*

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23 The Puncoviscana Formation (Turner, 1960) is a thick, mainly siliciclastic partly turbiditic,  
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25 succession (Ježek, 1990; Omarini et al., 1999, Zimmermann, 2005 and references therein) that crops  
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27 out in northern Argentina and allegedly throughout most of the eastern Sierras Pampeanas (e.g.,  
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29 Schwartz and Gromet, 2004; Rapela et al., 2007). It has been the subject of much controversy in  
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31 terms of its age and tectonic setting of sedimentation. The formation is important in that it shows  
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33 the main evidence for the Early Cambrian Pampean orogeny, in the form of penetrative deformation  
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35 and metamorphism, the grade of the latter increasing from the Puna and Sierras Orientales in the  
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37 north (very low- to low-grade) to the Sierras de Córdoba in the south (high-grade). The formation  
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39 was also host to the Pampean plutonic arc in the south, formed between ca. 550 and 530 Ma (Rapela  
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41 et al., 1998; Schwartz et al., 2008; Iannizzotto et al., 2011). The term Puncoviscana Formation in  
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43 the literature embraces sedimentary rocks probably older, coeval and younger than the Pampean  
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45 magmatic arc with the only constraint that they are older than the unconformably overlying Middle  
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47 to late Cambrian Meson Group (e.g., Omarini et al., 1999; Adams et al., 2008, 2011; Escayola et al.,  
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49 2011, and references therein). We restrict our treatment here to that part of the siliciclastic  
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51 succession that hosts the magmatic arc in the south, which is of relevance to the early history of the  
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53 Puncoviscana sedimentary basin. This southern tract is mainly pelitic and contains characteristic  
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55 detrital zircons populations with major peaks at 1100–960 Ma and 680–570Ma and lacks grains  
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57 derived from the nearby Rio de la Plata craton (2.02–2.26 Ga) (Schwartz and Gromet, 2004;  
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59 Escayola et al., 2007; Rapela et al., 2007). Sedimentation here took place on the eastern margin of  
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61 the Clymene Ocean between ca. 570 Ma (the approximate age of the youngest detrital zircons) and  
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63 ca. 530 Ma (Fig. 4); the older age being coincident with that of the anorogenic carbonatite-syenite  
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65 event referred to above. The paleogeographical position was probably distant from the Rio de la

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2 Plata craton, to which the Puncoviscana Formation basement became juxtaposed through right-  
3 lateral displacement during Pampean subduction and collision (Schwartz and Gromet, 2004; Rapela  
4 et al., 2007; Verdecchia et al., 2011). However, the sedimentary setting of this tract of the  
5 Puncoviscana Formation remains uncertain; a fore-arc basin was suggested Rapela et al. (2007) but  
6 a passive margin setting for the older part of the tract can not be discounted.  
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#### 12 13 **4. The Pampean orogeny and the amalgamation of SW Gondwana** 14

15 Subduction started in the Late Neoproterozoic or Early Cambrian along the eastern margin of  
16 the Clymene Ocean, giving rise to the Pampean orogeny. An I-type Andean-type magmatic arc  
17 developed between ca. 550 and 530 Ma (Rapela et al., 1998; Schwartz et al., 2008; Iannizzotto et  
18 al., 2011) (Fig. 4). At the same time Laurentia rifted away from MARA + Amazonia in the west  
19 (present coordinates), resulting in opening of the Iapetus Ocean (Dalziel, 1997) and with  
20 development of passive margin sedimentary sequences well preserved along the Appalachian  
21 margin of Laurentia and in the Precordillera terrane of western Argentina (e.g., Astini et al., 1995;  
22 Thomas and Astini, 1996). Final closure of the Clymene Ocean occurred between 530 and 520 Ma  
23 as implied by the ages of intermediate P/T Barrovian-type collisional metamorphism and coeval S-  
24 type plutonism (Rapela et al., 1998; Rapela et al., 2002; Otamendi et al., 2009) (Fig. 4).  
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33 At the start of the Pampean orogeny (ca. 550 Ma) the Paleoproterozoic and Mesoproterozoic  
34 basement of the Western Sierras Pampeanas was part of a large but ephemeral continental mass  
35 rifted from Laurentia (Rapela et al., 2007; Casquet et al., 2009), consisting of MARA together with  
36 Amazonia (Fig. 4). The Western Sierras Pampeanas probably formed the southern tail of this  
37 landmass (Rapela et al., 2007). Participation of the Pie de Palo complex in this new continental  
38 assemblage cannot be ruled out in the hypothesis of a para-autochthonous Precordillera terrane  
39 (Galindo et al., 2004; Finney, 2007).  
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46 The Pampean orogeny involved oblique closure of the Clymene Ocean between MARA +  
47 Amazonia and other West Gondwana cratons (Rio de la Plata, Kalahari,...), ultimate collision of  
48 these continental masses bringing to an end the formation of Gondwana (Trindade et al., 2006;  
49 Rapela et al., 2007). We suggest that these collisions were responsible for the formation of a  
50 continuous mobile belt embracing the Pampean orogen in the south and the Paraguay and Araguaia  
51 belts further north (Fig. 1), all of which have igneous, metamorphic and structural features in  
52 common (Rapela et al., 2007; Moura et al., 2008; McGee et al., 2011, Bandeira et al., 2011).  
53 Moreover, no evidence of Cambrian orogeny has yet been convincingly demonstrated for the  
54 western margin of Amazonia (e.g., Chew et al., 2007). Consequently, this orogenic belt was  
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2 probably part of the Terra Australis orogen of Cawood (2005), although the South American tract  
3 that we describe here follows a different trend along eastern Amazonia (Fig. 3). The Trans-  
4 Brasiliano lineament of Cordani et al. (2003) (Fig. 1) is interpreted as a late Pampean mega-fault  
5 equivalent to the Córdoba Fault (e.g., Rapela et al., 2007) (Fig. 4) responsible for the final assembly  
6 of eastern South America continental masses before Pangea. On the other hand the western margin  
7 of MARA + Amazonia (now part of Gondwana) facing the Iapetus Ocean remained passive until  
8 the Early Ordovician when it evolved into an Andean-type orogeny that persisted throughout the  
9 Paleozoic and the Mesozoic, evolving into the present Andean margin of South America.  
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## 19 **5. Conclusions**

20 Southern South America contains an outstanding record of Rodinia formation, further  
21 supercontinent break-up in the Neoproterozoic and final re-assembly of continental blocks in SW  
22 Gondwana in the Early Cambrian. We suggest here that several minor Paleoproterozoic blocks,  
23 such as the Maz terrane in the Western Sierras Pampeanas, the Arequipa block including its  
24 southern extension, and the Rio Apa block, at least, formed a major continental mass, i.e., the  
25 MARA craton, which collided with Amazonia at ca. 1.3 Ga. The resulting continent further  
26 amalgamated to Laurentia during middle and late Mesoproterozoic orogenies as part of Rodinia  
27 formation. Protracted break-up of Rodinia took place in the Neoproterozoic as recorded by episodic  
28 anorogenic magmatism and eventual opening of the Clymene Ocean in Ediacaran times. Post-  
29 glacial platform carbonates formed in this ocean followed by deposition of the largely turbiditic  
30 Puncoviscana Formation along the eastern margin in late Ediacaran to Early Cambrian times.  
31 Eastward right-lateral subduction led to closure of the Clymene Ocean coeval with Laurentia  
32 drifting away to the west to open the Iapetus Ocean. The proto-Andean margin formed at this time  
33 and remained passive till the start of the Andean-type Famatinian orogeny in the Early Ordovician.  
34 Final closure of the Clymene Ocean led to oblique collision of the large continental mass formed by  
35 MARA+Amazonia with other West Gondwana cratons (Kalahari, Rio de la Plata..) to produce the  
36 transpressional Pampean-Paraguay-Araguai orogenic belt in the Early Cambrian, and brought to an  
37 end the assembly of SW Gondwana.  
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6

## 7 **References**

- 9  
10 Aceñolaza, F.G., Toselli, A.J. 2000. Argentine Precordillera. Allocthonous or autocthonous  
11 Gondwanic? *Zentralblatt für Geologie Paläontologie*, 1999, 743-756.  
12  
13 Adams, C.J., Miller, H., Toselli, A.J., Griffin, W. 2008. The Puncoviscana formation of northwest  
14 Argentina: U-Pb geochronology of detrital zircons and Rb-Sr metamorphic ages and their  
15 bearing on its stratigraphic age, sediment provenance and tectonic setting. *Neues Jahrbuch für*  
16 *Geologie und Paläontologie Abhandlungen*, 247, 341-352.  
17  
18 Adams, C.J., Miller, H., Aceñolaza, F.G., Toselli, A.J., Griffin, W.L. 2011. The Pacific Gondwana  
19 margin in the late Neoproterozoic-early Paleozoic: Detrital zircon U-Pb ages from  
20 metasediments in northwest Argentina reveal their maximum age, provenance and tectonic  
21 setting. *Gondwana Research*, 19, 71-83.  
22  
23 Astini, R.A., Benedetto, J.L., Vaccari, N.E. 1995. The early Paleozoic evolution of the Argentine  
24 Precordillera as a Laurentian rifted, drifted and collided terrane: a geodynamic model.  
25 *Geological Society of America Bulletin*, 107, 253-273.  
26  
27 Bandeira, J., McGee, B., Nogueira, A.C.R., Collins, A.S., Trindade, R. 2011. Sedimentological and  
28 provenance response to Cambrian closure of the Clymene ocean: The upper Alto Paraguai  
29 Group, Paraguay belt, Brazil, *Gondwana Research*, doi: 10.1016/j.gr.2011.0  
30  
31 Baldo, E., Casquet, C., Pankhurst, R.J., Galindo, C., Rapela, C.W., Fanning, C.M., Dahlquist, J.A.,  
32 Murra, J. 2006. Neoproterozoic A-type granitic magmatism in the Western Sierras Pampeanas  
33 (Argentina): evidence for Rodinia break-up along a proto-Iapetus rift?. *Terra Nova*, 18, 388-  
34 394.  
35  
36 Bettencourt, J.S., Leite Jr., W.B., Ruiz, A.S., Matos, R., Payolla, B.L., Tosdal, R.M. 2010. The  
37 Rondonian-San Ignacio Province in the SW Amazonian Craton: An overview. *Journal of*  
38 *South American Earth Sciences*, 29, 28-46.  
39  
40 Böger, S.D., Raetz, M., Giles, D., Etchart, E., Fanning, C.M. 2005. U-Pb data from the Sunsás  
41 region of Eastern Bolivia, evidence for an allocthonous origin of the Paragua block.  
42 *Precambrian Research*, 139, 121-146.  
43  
44 Cardona, A., Chew, D., Valencia, V.A., Bayona, G., Mišković, A. Ibañez-Mejía, M. 2010.  
45 Grenvillian remnants in the Northern Andes: Rodinian and Phanerozoic paleogeographic  
46 perspectives. *Journal of South American Earth Sciences*, 29, 92-104.  
47  
48  
49  
50  
51  
52  
53  
54  
55  
56  
57  
58  
59  
60  
61  
62  
63  
64  
65

- 1  
2  
3 Casquet, C., Baldo, E., Pankhurst, R.J., Rapela, C.W., Galindo, C., Fanning, C.M., Saavedra, J.  
4 2001. Involvement of the Argentine Precordillera Terrane in the Famatinian mobile belt:  
5 Geochronological (U–Pb SHRIMP) and metamorphic evidence from the Sierra de Pie de  
6 Palo. *Geology*, 29, 703–706.  
7  
8  
9  
10 Casquet, C., Rapela, C.W., Pankhurst, R.J., Galindo, C., Dahlquist, J., Baldo, E.G., Saavedra, J.,  
11 González Casado, J.M., Fanning, C.M. 2005. Grenvillian massif-type anorthosites in the  
12 Sierras Pampeanas. *Journal of the Geological Society, London*, 162, 9–12.  
13  
14  
15 Casquet, C., Pankhurst, R.J., Fanning, C.M., Baldo, E., Galindo, C., Rapela, C., González-Casado,  
16 J.M., Dahlquist, J.A. 2006. U–Pb SHRIMP zircon dating of Grenvillian metamorphism in  
17 Western Sierras Pampeanas (Argentina): correlation with the Arequipa Antofalla craton and  
18 constraints on the extent of the Precordillera Terrane. *Gondwana Research*, 9, 524–529.  
19  
20  
21  
22 Casquet, C., Pankhurst, R.J., Rapela, C., Galindo, C., Fanning, C.M., Chiaradia, M., Baldo, E.,  
23 González-Casado, J.M., Dahlquist, J.A. 2008a. The Maz terrane: a Mesoproterozoic domain  
24 in the western Sierras Pampeanas (Argentina) equivalent to the Arequipa-Antofalla block of  
25 southern Peru? Implications for Western Gondwana margin evolution. *Gondwana Research*,  
26 13, 163-175.  
27  
28  
29  
30  
31  
32 Casquet, C., Pankhurst, R.J., Galindo, C., Rapela, C.W., Fanning, C.M., Baldo, E., Dahlquist, J.,  
33 González Casado, J.M., Colombo, F. 2008b. A deformed alkaline igneous rock – carbonatite  
34 complex from the Western Sierras Pampeanas, Argentina: evidence for late Neoproterozoic  
35 opening of the Clymene Ocean?. *Precambrian Research*, 165, 205–220.  
36  
37  
38  
39 Casquet, C., Rapela, C.W., Pankhurst, R.J., Baldo, E., Galindo, C., Fanning, M., Saavedra, J. 2009.  
40 Proterozoic terranes in southern South America: Accretion to Amazonia, involvement in  
41 Rodinia formation and further west Gondwana accretion. In: *Rodinia: Supercontinents,*  
42 *Superplumes & Scotland.* Geological Society of London, Fermor Meeting, Edinburgh,  
43 Abstract.  
44  
45  
46  
47  
48 Casquet, C., Fanning, C.M., Galindo, C., Pankhurst, R.J., Rapela, C.W., Torres, P. 2010. The  
49 Arequipa Massif of Peru: New SHRIMP and isotope constraints on a Paleoproterozoic inlier  
50 in the Grenvillian orogen. *Journal of South American Earth Sciences*, 29, 128–142.  
51  
52  
53  
54 Cawood, P. 2005. Terra Australis Orogen: Rodinia breakup and development of the Pacific and  
55 Iapetus margin of Gondwana during the Neoproterozoic and Paleozoic. *Earth-Science*  
56 *Reviews*, 249-279.  
57  
58  
59  
60  
61  
62  
63  
64  
65

- 1  
2 Chew, D.M., Schaltegger, U., Kosler, J., Whitehouse, M.J., Gutjahr, M., Spikings, R.A., Miskovic,  
3 A. 2007. U-Pb geochronologic evidence for the evolution of the Gondwanan margin of the  
4 north-central Andes. *Geological Society of America Bulletin*, 119, 697-711.  
5  
6  
7 Cobbing, E.J., Pitcher, W.S. 1972. Plate tectonics and the Peruvian Andes. *Nature*, 246, 51-53.  
8  
9 Colombo, F., Baldo, E. G., Casquet, C., Pankhurst, R.J., Galindo, C., Rapela, C.W., Dahlquist, J.A.,  
10 Fanning, C.J. 2009. A-type magmatism in the sierras of Maz and Espinal: A new record of  
11 Rodinia break-up in the Western Sierras Pampeanas of Argentina. *Precambrian Research*,  
12 175, 77–86.  
13  
14  
15  
16 Cordani, U.G., Teixeira, W. 2007. Proterozoic accretionary belts in the Amazonian Craton. In:  
17 Hatcher, R.D., Jr., Carlson, M.P., McBride, J.H., & Martinez Catalán, J.R. (eds.) 4-D  
18 Framework of Continental Crust. *Geological Society of America Memoir*, 200, 297–320.  
19  
20  
21 Cordani, U.G., D'Agrella-Filho, M.S., Brito-Neves, B.B., Trindade, R.I.F. 2003. Tearing up  
22 Rodinia: the Neoproterozoic palaeogeography of South American cratonic fragments. *Terra*  
23 *Nova*, 15, 350–359.  
24  
25  
26  
27 Cordani, U.G., Teixeira, W., D'Agrella-Filho, M.S., Trindade, R.I. ( 2009) The position of  
28 Amazonian Craton in supercontinents. *Gondwana Research*, 15, 396-407.  
29  
30  
31 Cordani, U.G., Teixeira, W., Tassinari, C.G., Coutinho, J.M.V., Ruiz, A.S. 2010. The Rio Apa  
32 craton in Mato Grosso do Sul (Brazil) and northern Paraguay: Geochronological evolution,  
33 correlations and tectonic implications for Rodinia and Gondwana. *American Journal of*  
34 *Science*, 310, 981-1023.  
35  
36  
37  
38 Dalziel, I.W.D. 1992. On the organization of American plates in the Neoproterozoic and the  
39 breakout of Laurentia. *GSA Today*, 2, 1-2.  
40  
41  
42 Dalziel, I.W.D. 1994. Precambrian Scotland as a Laurentia-gondwana link: Origin and significance  
43 of cratonic promontories. *Geology*, 22, 589-592.  
44  
45  
46 Dalziel, I. W. D. 1997. Overview. Neoproterozoic-Paleozoic geography and tectonics: Review,  
47 hypothesis, environmental speculation. *Geological Society of America Bulletin*, 109, 16-42.  
48  
49 Escayola, M.P., Pimentel, M., Armstrong, R. 2007. Neoproterozoic backarc basin: Sensitive high-  
50 resolution ion microprobe U–Pb and Sm–Nd isotopic evidence from the Eastern Pampean  
51 Ranges, Argentina. *Geology*, **35**, 495–498.  
52  
53  
54 Escayola, M.P., van Staal, C.R., Davis, W.J. 2011. The age and tectonic setting of the Puncoviscana  
55 Formation in northwestern Argentina: An accretionary complex related to Early Cambrian  
56 closure of the Puncoviscana Ocean and accretion of the Arequipa-Antofalla block. *Journal of*  
57 *South American Earth Sciences*, doi: 10.1016/j.jsames.2011.04.013.  
58  
59  
60  
61  
62  
63  
64  
65

- 1  
2  
3 Finney, S.C. 2007. The parautochthonous Gondwanan origin of the Cuyania (greater Precordillera)  
4 terrane of Argentina: A re-evaluation of evidence used to support an allochthonous Laurentian  
5 origin. *Geologica Acta*, 5, 127-158.  
6  
7 Galindo, C., Casquet, C., Rapela, C., Pankhurst, R., Baldo, E., Saavedra, J. 2004. Sr, C and O  
8 isotope geochemistry and stratigraphy of Precambrian and lower Paleozoic carbonate  
9 sequences from the Western Sierras Pampeanas of Argentina: tectonic  
10 implications. *Precambrian Research*, 131, 55-71.  
11  
12 Hoffman, P.F. 1991. Did the breakout of Laurentia turn Gondwanaland inside out?. *Science*, 252,  
13 1409-1412.  
14  
15 Iannizzotto, N.F., Rapela, C.W., Baldo, E. 2011. Nuevos datos geocronológicos, geoquímicos e  
16 isotópicos del batolito de Sierra Norte - Ambargasta en su sector más austral, provincia de  
17 Córdoba. *Actas del XVIII Congreso Geológico Argentino*, Neuquén, 190-191.  
18  
19 Jezek, P. 1990. Análisis sedimentológico de la Formación Puncoviscana entre Tucumán y Salta. In:  
20 Aceñolaza, F.G., Millar, H., Toselli, A.J. (eds.), *El Ciclo Pampeano en el noroeste argentino*.  
21 Serie Correlación Geológica nº 4. Universidad Nacional de Tucumán, pp. 9-35.  
22  
23 Kay, S.M., Orrell, S., Abbruzzi, J.M. 1996. Zircon and whole rock Nd-Pb isotopic evidence for a  
24 Grenville age and a Laurentian origin for the basement of the Precordillera in Argentina. *The*  
25 *Journal of Geology*, 104, 637-648.  
26  
27 Li, Z.X., Bogdanova, S.V., Collins, A.S., Davidson, A.S., De Waele, B., Ernst, R.E., Fitzsimons,  
28 I.C.W., Fuck, R.A., Gladkochub, D.P., Jacobs, J., Karlstrom, K.E., Lu, S., Natapov, L.M.,  
29 Pease, V., Pisarevsky, S.A., Thrane, K., Vernikovsky, V. 2008. Assembly, configuration,  
30 and break-up history of Rodinia: A synthesis. *Precambrian Research*, 160, 179-210  
31  
32 Loewy, S.L., Connelly, J.N., Dalziel, I.W.D., Gower, C.F. 2003. Eastern Laurentia in Rodinia:  
33 constraints from whole-rock Pb and U/Pb geochronology. *Tectonophysics*, 375, 169-197.  
34  
35 Loewy, S., Connelly, J.N., Dalziel, I.W.D. 2004. An orphaned block: The Arequipa-Antofalla  
36 Basement of central Andean margin of South America. *Geological Society of America*  
37 *Bulletin*, 116, 171-187.  
38  
39 Martignole, J., Martelat, J.E. 2003. Regional-scale Grenvillian-age UHT metamorphism in the  
40 Mollendo-Camana block (basement of the Peruvian Andes). *Journal of Metamorphic*  
41 *Geology*, 21, 99-120.  
42  
43 McGee, B., Collins, A.S., Trindade, R. I.F. 2011. G'day Gondwana—the final accretion of a  
44 supercontinent: U-Pb ages from the post-orogenic São Vicente Granite, northern Paraguay  
45 Belt, Brazil, *Gondwana Research*, doi:10.1016/j.gr.2011.04.011  
46  
47  
48  
49  
50  
51  
52  
53  
54  
55  
56  
57  
58  
59  
60  
61  
62  
63  
64  
65

- 1  
2 Misi, A., Kaufman, A.J., Veizer, J., Powis, K., Azmy, K., Boggiani, P.C., Gaucher, C., Teixeira,  
3 J.B.G., Sanches, A.L., Iyer, S.S.S. 2007. Chemostratigraphic correlation of Neoproterozoic  
4 successions in South America. *Chemical Geology*, 237, 143-167.  
5  
6  
7 Moura, C.A.V., Pinheiro, B.L.S., Nogueira, A.C.E., Gorayeb, P.S.S., Galarza, M.A. 2008.  
8 Sedimentary provenance and palaeoenvironment of the Baixo Araguaia Supergroup:  
9 constraints on the palaeogeographical evolution of the Araguaia Belt and assembly of West  
10 Gondwana. In: Pankhurst, R.L., Trow, R.A.J., Brito Neves, B.B. & de Wit, M.J. (eds.), *West*  
11 *Gondwana: Pre-Cenozoic Correlations Across the South Atlantic Region*. Geological Society,  
12 London, Special Publications, 294, 173-196.  
13  
14  
15  
16  
17  
18  
19 Mulcahy, S.R., Roeske, S.M., McClelland, W.C., Jourdan, F., Iriondo, A., Renne, P.R., Vervoort,  
20 J.D., Vujovich, G.I. 2011. Structural evolution of a composite middle to lower crustal section:  
21 The Sierra de Pie de Palo, northwest Argentina. *Tectonics*, 30, TC1005,  
22 doi:10.1029/2009TC002656  
23  
24  
25  
26  
27  
28  
29  
30  
31  
32  
33  
34  
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36  
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52  
53  
54  
55  
56  
57  
58  
59  
60  
61  
62  
63  
64  
65
- Murra, J., Baldo, E., Galindo, C., Casquet, C., Pankhurst, R., Rapela, C., Dahlquist, J. 2011. Sr, C and O isotope composition of marbles from the Sierra de Ancasti, Eastern Sierras Pampeanas, Argentina: age and constraints for the Neoproterozoic–Lower Paleozoic evolution of the proto-Gondwana margin. *Geologica Acta*, 9, 1, 1–23.
- Omarini, R.H., Sureda, R.J., Götze, H.J., Seilacher, A., Pflüger, F. 1999. Puncoviscana folded belt in northwestern Argentina: testimony of Late Proterozoic Rodinia fragmentation and pre-Gondwana collisional episodes. *International Journal of Earth Sciences*, 88, 76-97.
- Otamendi, J.E., Vujovich, G.I., de la Rosa, J.D., Tibaldi, A.M., Castro, A., Martino, R.D., Pinotti, L.P. 2009. Geology and petrology of a deep crustal zone from the Famatinian paleo-arc, Sierras Valle Fértil-La Huerta, San Juan, Argentina. *Journal of South American Earth Sciences*, 27, 258–279.
- Ramos, V.A. 1988. Late Proterozoic-Early Paleozoic of South America – a collisional history. *Episodes*, 11, 168-174.
- Ramos, V.A. 2004. Cuyania, an exotic block to Gondwana: review of an historical success and the present problems. *Gondwana Research*, 7, 1009-1026.
- Ramos, V.A., Vujovich, G.I. 1993. The Pampia Craton within Western Gondwanaland. In: Ortega-Gutiérrez, F., Coney, P., Centeno-García, E. & Gómez-Caballero, A. (eds) *Proceedings of The First Circum-Pacific and Circum-Atlantic Terrane Conference*, Mexico, 113–116.
- Ramos, V., Dallmeyer, R. D., Vujovich, G. 1998. Time constraints on the Early Palaeozoic docking of the Precordillera, central Argentina, in Pankhurst, R. J., and Rapela, C.W. (eds.), *The*

- 1  
2 Proto-Andean margin of Gondwana. Geological Society, London, Special Publications, 142,  
3 143-158.  
4  
5  
6 Ramos V.A., Vujovich G., Martino R. & Otamendi J. 2010. Pampia: a large cratonic block missing  
7 in the Rodinia supercontinent. *Journal of Geodynamics*, 50, 243-255.  
8  
9  
10 Rapela, C.W., Pankhurst, R.J., Casquet, C., Baldo, E., Saavedra, J., Galindo, C., Fanning, C.M.  
11 1998. The Pampean Orogeny of the southern proto-Andes: evidence for Cambrian continental  
12 collision in the Sierras de Córdoba. In: Pankhurst, R.J. & Rapela, C.W. (eds.), *The Proto-*  
13 *Andean Margin of Gondwana*. Geological Society, London, Special Publications, 142, 181–  
14 217.  
15  
16  
17  
18 Rapela, C.W., Baldo, E., Pankhurst, R.J., Saavedra, J. 2002. Cordierite and leucogranite formation  
19 during emplacement of highly peraluminous magma: the El Pílon Granite Complex (Sierras  
20 Pampeanas, Argentina). *Journal of Petrology*, 43, 1003-1028.  
21  
22  
23  
24 Rapela, C.W., Pankhurst, R.J., Casquet, C., Fanning, C.M., Galindo, C., Baldo, E. 2005. Datación  
25 U-Pb SHRIMP de circones detríticos en parafibrolitas neoproterozoicas de la secuencia  
26 Difunta Correa (Sierras Pampeanas Occidentales, Argentina). *Geogaceta*, 38, 227-230.  
27  
28  
29 Rapela, C.W., Pankhurst, R.J., Casquet, C., Fanning, C.M., Baldo, E.G., González-Casado, J.M.,  
30 Galindo, C., Dahlquist, J. 2007. The Río de la Plata craton and the assembly of SW  
31 Gondwana. *Earth Science Reviews*, 83, 49-82.  
32  
33  
34  
35 Rapela, C.W., Pankhurst, R.J., Casquet, A., Baldo, E., Galindo, C., Fanning, C.M., Dahlquist,  
36 J.M. 2010. The Western Sierras Pampeanas: Protracted Grenville-age history (1330–1030  
37 Ma) of intra-oceanic arcs, subduction–accretion at continental-edge and AMCG intraplate  
38 magmatism. *Journal of South American Earth Sciences*, 29, 105–127.  
39  
40  
41  
42 Rapela, C.W., Pankhurst, R.J., Fanning, C.M., Galindo, C., Baldo, E.G., Casquet, C., Colombo, F.,  
43 Dahlquist, J. 2011. Hf and O isotopic signature of 840-730 Ma A-type granitic and felsic  
44 igneous rocks in southern South America. VII Hutton Symposium on Granites and Related  
45 Rocks, Avila, Spain. Abstract 121.  
46  
47  
48  
49 Rivers, T. 2008. Assembly and preservation of lower, mid and upper orogenic crust in the Grenville  
50 Province – Implications for the evolution of large hot long-duration orogens. *Precambrian*  
51 *Research*, 167, 237-259  
52  
53  
54  
55 Sadowski, G.R., Bettencourt, J.S. 1996. Mesoproterozoic tectonic correlations between eastern  
56 Laurentia and the western border of the Amazon Craton. *Precambrian Research*, 76, 213-227.  
57  
58  
59  
60  
61  
62  
63  
64  
65

- 1  
2 Santosh, M., Maruyama, S., Yamamoto, S. 2009. The making and breaking of supercontinents:  
3 some speculations based on superplumes, super downwelling and the role of tectosphere.  
4 Gondwana Research, 15, 324-341.  
5  
6  
7 Schwartz, J.J., Gromet, L.P. 2004. Provenance of a late Proterozoic – early Cambrian basin, Sierras  
8 de Córdoba, Argentina. Precambrian Research, 129, 1–21.  
9  
10 Schwartz, J.J., Gromet, L.P., Miró, R. 2008. Timing and duration of the calcalkaline arc of the  
11 Pampean Orogeny: implications for the Late Neoproterozoic to Cambrian evolution of  
12 Western Gondwana. Journal of Geology, 116, 39–61.  
13  
14  
15  
16  
17 Teixeira, W., Geraldes, M.C., Matos, R., Ruiz, A.S., Saes, G., Vargas-Mattos, G. 2010. A review of  
18 the tectonic evolution of the Sunsás belt, SW Amazonian Craton. Journal of South American  
19 Earth Sciences, 29, 47-60.  
20  
21  
22 Thomas, W.A., Astini, R.A. 1996. The Argentine Precordillera: a traveler from the Ouachita  
23 embayment of North America Laurentia. Science, 273, 752–757.  
24  
25  
26 Tohver, E., Van der Pluijm, B.A., Van der Voo, R., Rizzotto, G., Scandolara, J.E. 2002.  
27 Paleogeography of the Amazon craton at 1–2 Ga: early Grenvillian collision with the Llano  
28 segment of Laurentia. Earth Planetary Science Letters, 199, 185-200.  
29  
30  
31 Tohver, E., Bettencourt, J.S., Tosdal, R., Mezger, K., Leite, W.B., Payolla, B.L. 2004. Terrane  
32 transfer during Grenville orogeny: tracing the Amazonian ancestry of southern Appalachian  
33 basement through Pb and Nd isotopes. Earth and Planetary Science Letters, 228, 161-176.  
34  
35  
36  
37 Turner, J.C. 1960. Estratigrafía de la Sierra de Santa Victoria (Argentina). Boletín de la Academia  
38 Nacional de Ciencias en Córdoba, 41, 163-196.  
39  
40  
41 Trindade, R.I.F., D’Agrella-Filho, M.S., Epof, I., Brito Neves, B.B. 2006. Paleomagnetism of the  
42 early Cambrian Itabaiana mafic dikes, NE Brazil, and implications for the final assembly of  
43 Gondwana and its proximity to Laurentia. Earth Planetary Science Letters, 244, 361-377.  
44  
45  
46 Varela, R., Valencio, S., Ramos, A., Sato, K., González, P., Panarello, H., Roverano, D. 2001.  
47 Isotopic strontium, carbon and oxygen study on Neoproterozoic marbles from Sierra de  
48 Umango, and their foreland, Argentina. III South American Symposium on Isotope Geology,  
49 450–453.  
50  
51  
52  
53 Verdecchia, S.O., Casquet, C., Baldo, E.G., Pankhurst, R.J., Rapela, C.W., Fanning, C.M., Galindo,  
54 C. 2011. Docking of the Río de la Plata craton to southwestern Gondwana: age constraints  
55 from U–Pb SHRIMP detrital zircon ages from Sierras de Ambato and Velasco (Sierras  
56 Pampeanas, Argentina). Journal of the Geological Society, London, 168, 1061-1071.  
57  
58  
59  
60  
61  
62  
63  
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65



- 1  
2  
3 Vujovich, G., van Staal, C., Davis, W. 2004. Age constraints on the tectonic evolution and  
4 provenance of the Pie de Palo Complex, Cuyania Composite Terrane, and the Famatinian  
5 Orogeny in the Sierra de Pie de Palo, San Juan, Argentina. *Gondwana Research*, 7, 1041-  
6 1056.  
7  
8  
9  
10 Weil, A.B., Vand der Voo, R., Mac Niocaill, C., Meert, J.G. 1998. The Proterozoic supercontinent  
11 Rodinia: paleomagnetically derived reconstructions from 1100 to 800 Ma. *Earth and*  
12 *Planetary Science Letters*, 154, 13-24.  
13  
14  
15 Zimmermann, U. 2005. Provenance studies of very low- to low-grade metasedimentary rocks of the  
16 Puncoviscana complex, northwest Argentina. In: Vaughan, A.P.M., Leat, P.T. & Pankhurst,  
17 R.J. (eds), *Terrane Processes at the Margins of Gondwana*. Geological Society, London,  
18 Special Publications, 246, 381-416.  
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## 24 **Figure Captions**

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26 Figure. 1: Sketch map of South America showing Paleoproterozoic to Archean cratons and the  
27 middle-to-late Mesoproterozoic, and Neoproterozoic to Early Cambrian orogenic belts. The  
28 MARA craton reached its present position after right-lateral oblique accretion to the Rio de la  
29 Plata craton during the Pampean orogeny and further displacement along the Córdoba fault.  
30 Outcrops in red are Paleoproterozoic and Mesoproterozoic outcrops referred to in the text.  
31 DBF: Desaguadero-Bermejo Fault; CF: Córdoba Fault.  
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37 Figure 2: The Paleoproterozoic record in the three outcrops forming the hypothetical MARA craton.  
38 Data from Casquet et al. (2008a, 2010) and Cordani et al. (2010).  
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40 Figure 3: Middle-to-late Mesoproterozoic evolution of the MARA craton. From Rapela et al.  
41 (2010), Casquet et al. (2010) and Cordani et al. (2010)  
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44 Figure. 4: 3-D diagrams showing geotectonic evolution during the Neoproterozoic and the Early  
45 Cambrian that led to the Pampean orogeny and final amalgamation of MARA to Gondwana.  
46 See text for explanation. The figure highlights the role played by the opening of the Clymene  
47 Ocean in the late Neoproterozoic and its subsequent closure in the Early Cambrian coeval  
48 with drifting of Laurentia in the west *en route* to the northern hemisphere  
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Figure 1

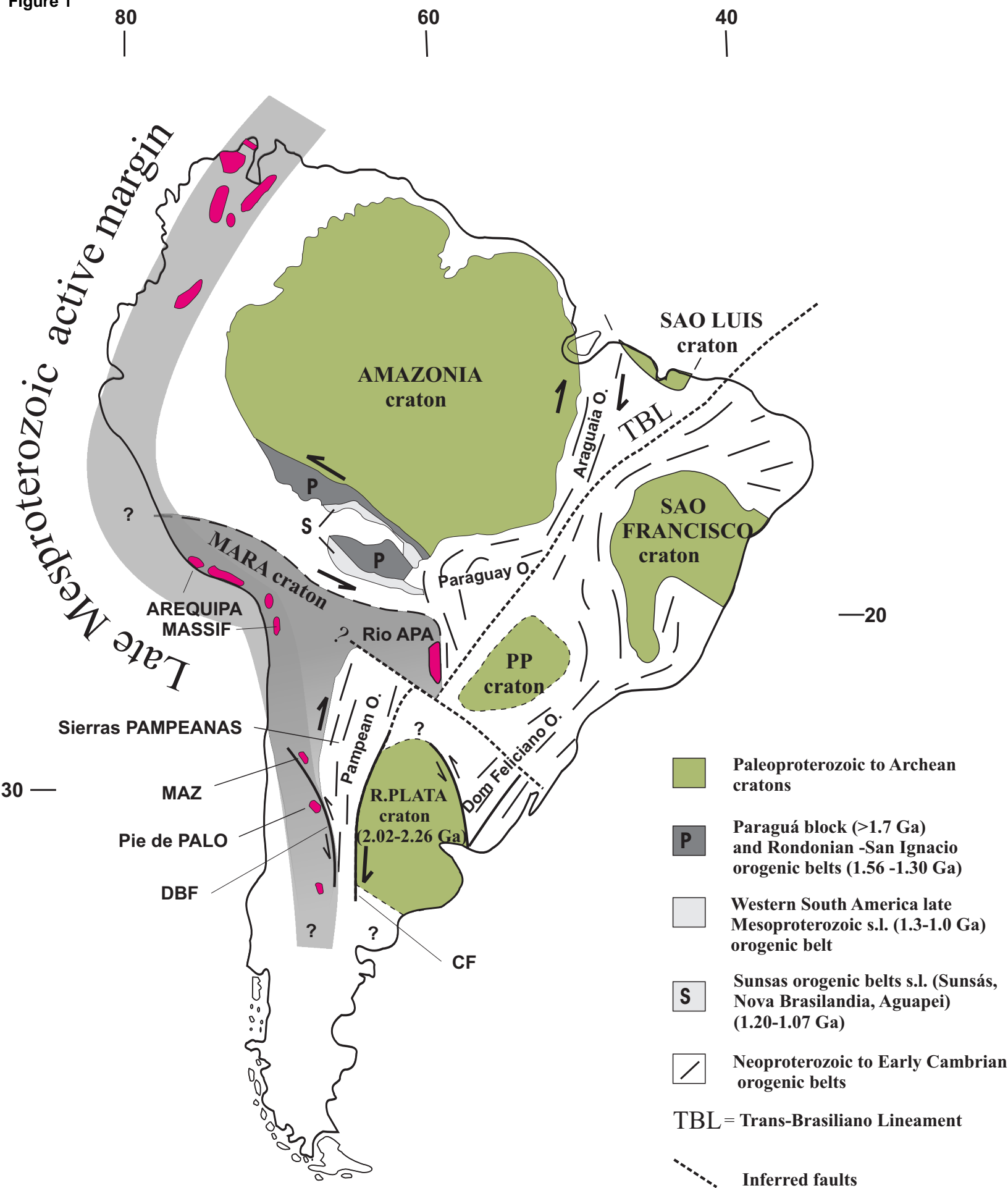


Figure 1

Figure 2

**COMPARATIVE EVOLUTION OF PALEOPROTEROZOIC BLOCKS**

	AREQUIPA		RIO APA		MAZ TERRANE	
<b>U-Pb Geochronology</b>	<b>1.7 - 1.79 Ga</b>	<i>Late felsic magmatism</i>	<b>1.72 - 1.76 Ga</b>	<i>Granite magmatism &amp; regional metamorphism</i>	<b>1.3 - 1.7 Ga</b>	<i>Sedimentation on Paleoproterozoic basement</i>
	<b>~1.87 Ga</b>	<i>UHT metamorphism</i> <i>Sedimentation</i>			<b>1.7 Ga</b>	<i>Magmatism (inferred from detrital zircons in metasediments)</i>
	<b>1.9 - 2.1 Ga</b>	<i>Magmatism (inferred from detrital zircons in metasediments)</i>	<b>1.94 - 1.83 Ga</b>	<i>Granite magmatism</i>	<b>1.9 Ga</b>	
<b>Nd model ages (TDM)</b>	<b>1.9 - 2.5 Ga</b>		<b>1.9 - 2.5 Ga</b>		<b>1.7 - 2.6 Ga</b>	

Figure 2

Figure 3

### GRENVILLIAN EVOLUTION IN THE MARA CRATON

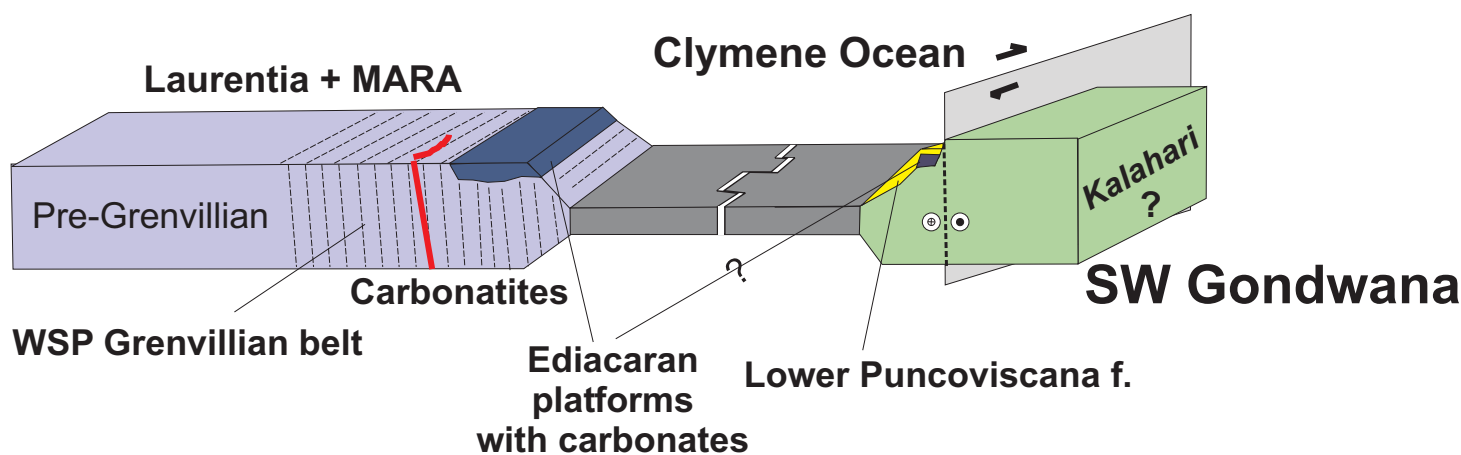
Age (Ma)	Maz terrane (continental)	Pie de Palo terrane (oceanic)	Arequipa block (continental)	Rio Apa block (continental)
900 -			~ 1.04 - 0.85 Ga <i>Diachronous low-P high-T regional metamorphism</i>	
1000 -		~1.03 Ga <i>Emplacement of TTG suites</i>		
1100 -	~1.09-1.07 Ga <i>Climax of lithospheric extension, emplacement of AMCG complexes. Exhumation of high-grade terranes</i>	~1.11 Ga <i>New subduction-related acid and basic magmatism</i>	~ 1.2 - 1.0 Ga <i>Sedimentation (Atico Basin)</i>	
1200 -	~1.23-1.17 Ga <i>Arc-continent collision, lithospheric thickening and high-grade metamorphism at the continental edge</i>	~1.17 Ga <i>Arc-related magmatism</i>		
1300 -	~1.33-1.26 Ga <i>Andino-type magmatism Emplacement of cordilleran granites in Paleoproterozoic sediments</i>	~1.20 Ga <i>Arc/back-arc oceanic complex N-MORB and oceanic arc lavas</i> ~1.24 Ga <i>Intraoceanic arc subduction (Las Matras block)</i>		
	<b>GRENVILLIAN ACTIVE MARGIN FACING LAURENTIA?</b>		<b>GRENVILLIAN HINTERLAND EXTENSION?</b>	<b>STABLE GRENVILLIAN</b>

Figure 3

Figure 4

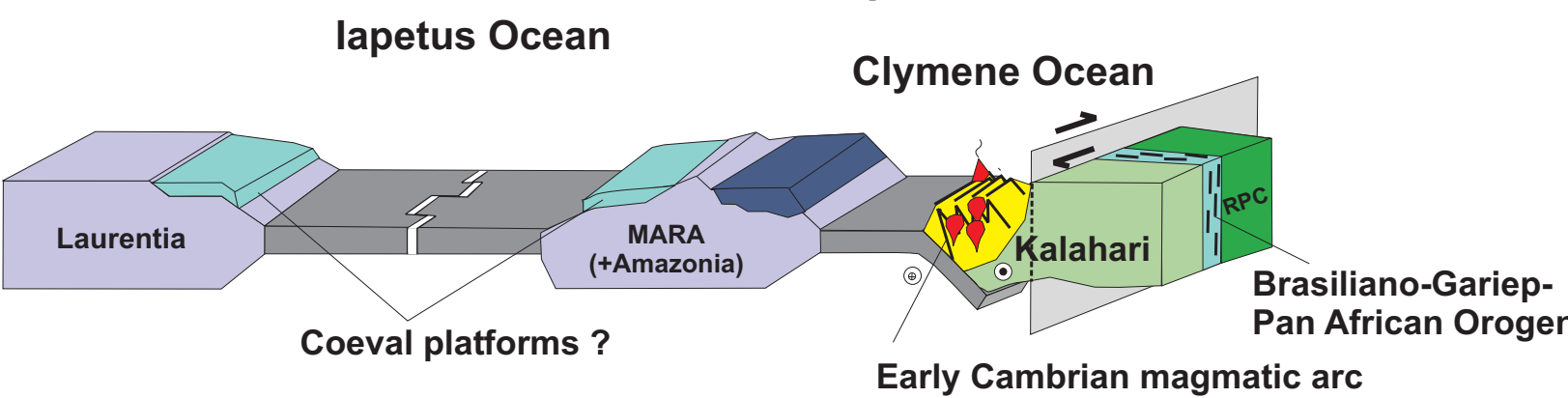
a) 570-540 Ma

### Opening of the Clymene Ocean



b) 540-530 Ma

### Opening of the Iapetus Ocean and Pampean subduction



c) 530-520 Ma

### Pampean oblique collision

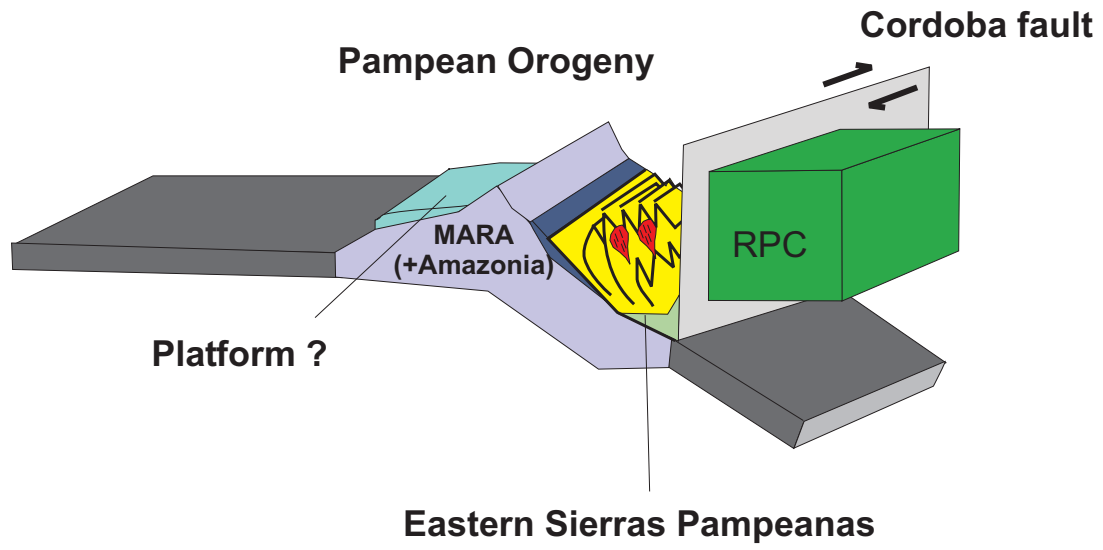


Figure 4