Predicting Extremes in European Geomagnetic Activity

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Abstract
Rapid variations in the geomagnetic field constitute a natural hazard, e.g. for navigation and to power grids and pipeline networks. In order to better allocate resources towards mitigating these risks, we must be able to model the recurrence of extremes of geomagnetic activity over many years. However, the data we have from which to develop such a model, in the form of continuous series of 1-minute samples of the geomagnetic field, typically stretch back less than 40 years. Without a longer record, it is difficult to construct a clear picture of the magnitude and frequency of extremes in geomagnetic activity. We therefore apply the statistical technique of ‘extreme value analysis’ on a number of decades of geomagnetic data recorded at observatories across Europe, and in doing so arrive at an estimate of the maximum field strength and time-variation that might be observed once in every 100 and 200 years.

Extreme Value Theory
We use a Generalised Pareto Distribution (GPD) to describe the tail of the distribution of geomagnetic activity (see e.g. Coles, 2004). The GPD is a unification of the Gumbel, Frechet and Weibull distributions, widely used in the scientific literature when modelling extremes in variables. By fitting a GPD curve to the extreme values in our data set, we can make estimates about the probability of geomagnetic conditions more extreme than those in our data set, and thus provide estimates of the largest extremes likely to be observed over a given period.

To accurately characterise the probability of extreme values, we must exclude all non-extreme samples from our data. This can be achieved using a ‘point over threshold’ approach, where some value is chosen as a threshold for extreme geomagnetic activity.

There are some assumptions implicit in the theory behind the GPD which must be examined: namely that the data are stationary (show no time-dependancy) and the probability of a particular sample exceeding the threshold is not dependant on the value of previous samples; that is, the data are independent. Clearly, geomagnetic data are not independent; storms lasting several hours are likely to produce a number of extreme values, which cannot be considered to be independent of one-another. This may be dealt with through de-clustering the data; this technique is described later.

Neither are the data stationary; the geomagnetic field exhibits cyclical behaviour with periods from days to years. However, we assume there to be no significant time-dependancy on the scale of centuries and that any non-stationarity does not affect the results at these time scales.

Data Preparation
One-minute geomagnetic time series of $H$ (horizontal field) and $D$ (declination) were downloaded from the World Data Centre for Geomagnetism in Edinburgh (wdc.bgs.ac.uk) for 29 European observatories. The observatories were chosen to provide a representative spread of locations across the continent, covering a range of magnetic latitudes and for which there exists continuous data spanning a number of years.

From $D$ and $H$, we computed time series of the variations from quiet levels due to the consecutive sample in the original and time series for each of the 29 observatories. A further two time series, $D_H$ and $D_H$, were constructed by computing the difference between successive sample in the original $D$ and $H$ time series.

Threshold Selection
Before fitting a GPD curve to the data, we first define the threshold which determines which values are to be considered extreme. All data points below this threshold are discarded. An appropriate threshold for each of the variables was determined by plotting the scale and shape parameters of the resulting GPD for a range of thresholds. The ideal threshold should be low enough to allow for a meaningful number of samples, but high enough that the modified scale parameter is constant and the shape parameter linear (within error-margins), above the chosen threshold.

We found that setting the threshold at the 98.6th percentile was reasonable for each variable at most observatories.

De-clustering
Clusters of extreme values occur during geomagnetic storms. This results in statistical dependency in the data, which must be eliminated to meet the assumptions of the model. We identified clusters by looking for extreme values that were not separated by at least one day. Only the peak value from each cluster was retained.

Determine GPD parameters
Applying the threshold and de-clustering reduces the data set by more than 99.9%. The remaining data are fitted to a GPD using the ‘R’ statistics software, along with the ‘eXtremes’ package. The model can be plotted as a ‘return-level plot’ (Fig 3).

Results
For each observatory we have computed the peak variation and rate-of-change predicted by the GPD to occur over periods of 100 and 200 years. The results for two of the four time series are summarised in Figure 4. The results from the UK observatories at Lerwick, Eskdalemuir and Harland are shown.

Future Work
This statistical analysis could be improved by, for example, treating variables $D$ (H) and $D/dt$ (dH/dt) as components of the same multi-variate statistic. It should also be possible to extend the application of this technique to look at extremes in global activity levels, dependent on data availability.

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References
http://www.assessment.ucar.edu/toolkit/.

Figure 1. (H, D) (degrees) and 1-minute rate-of-change of the sidereal mid-latitude observatory (one minute data, 1993-2005). Daily maxima of the absolute residuals are shown, as an integer number, to identify any solar cycle dependence.

Figure 2. Return periods for observed Harland (H) and (D) (lower) residuals (circles) and the fitted and simulated (GPD) (red) to each of $H$ and $D$. Vertical scales are degrees and nT respectively; horizontal scale is time in years. The blue lines are the approximate symmetric ±95% confidence limits from the fit of model to data, via an 'R' extreme function.

Figure 3. Response of modified scale and shape parameters of GPD function to threshold. The red arrows indicate a reasonable threshold to use for these data (residuals $dH/dt$ at Eskdalemuir observatory).

Figure 4. The measured maximum, 100 year return-level and 200 year return-level for (H) and (D) (residuals). The ±95% confidence interval is represented by the translucent segments of each column.