

1 **Improving the characterisation of Quaternary deposits for groundwater**
2 **vulnerability assessments using maps of recharge and attenuation potential**

3 **K J Griffiths¹, A M MacDonald², N S Robins¹, J Merritt², S J Booth³, D Johnson⁴ and P J**
4 **McConvey⁵**

5 ¹British Geological Survey, Maclean Building, Crowmarsh Gifford, Wallingford, Oxon OX10 8BB kjgr@bgs.ac.uk

6 ²British Geological Survey, Murchison House, West Mains Road, Edinburgh EH9 3LA

7 ³ British Geological Survey, Kingsley Dunham Centre, Keyworth, Nottingham NG12 5GG

8 ⁴Environment Agency, Sapphire East, 550 Streetsbrook Road, Solihull, West Midlands, B91 1QT

9 ⁵Geological Survey of Northern Ireland, Colby House, Stranmillis Court, Belfast BT9 5BF

10 **Abstract**

11 Assessing aquifer vulnerability is difficult for bedrock aquifers concealed by highly variable
12 superficial deposits such as glacial till. Many current groundwater vulnerability maps, and the
13 geological maps on which they are based, do not adequately account for regional and vertical
14 variations in the characteristics of superficial deposits. A new methodology for characterising
15 recharge potential and contaminant retardation potential of superficial deposits is presented here
16 which captures primary geological and hydrogeological expert knowledge in a systematic
17 manner. The method modifies existing superficial geology maps using Quaternary domains and
18 their descriptions, bedrock lithology, thickness of superficial deposits and applies additional
19 information on superficial geology and bedrock lithology. Central to the methodology is a matrix
20 which enables local geological and hydrogeological knowledge to be incorporated in a
21 systematic and traceable manner. The scale-independent methodology has been piloted at
22 1:625 000 scale to produce maps of recharge and attenuation potential for Great Britain.
23 Preliminary verification against several indicators (HOST data, the Scottish vulnerability
24 screening tool, and nitrate data) has been encouraging. The method is being used by the

25 Environment Agency as part of its vulnerability assessments for the characterization of
26 groundwater bodies as required by the Water Framework Directive (Council of European
27 Communities 2000).

28

29 Quaternary superficial deposits cover approximately 90% of the British landmass. Superficial
30 deposits, of which glacial till is a major component, moderate potential recharge to underlying
31 bedrock aquifers and are an important influence on vulnerability to pollution through their ability
32 to attenuate pollutants. Superficial deposits also influence groundwater quality (Shand et al.
33 2007). Characterisation of these deposits and the processes occurring within them is critical to
34 the assessment of groundwater vulnerability and the characterisation of groundwater bodies
35 required by the Water Framework Directive (Council of European Communities 2000).

36 Groundwater vulnerability maps for England and Wales were first developed for the National
37 Rivers Authority in the early 1990s and are still in use today, mainly as a planning tool. The
38 maps are at 1: 100 000 scale and cover all of England and Wales (Palmer et al. 1995). These
39 maps classify the dominant soil strata above the saturated aquifer as the vulnerability indicator,
40 recognising a matrix of three aquifer classes (weakly, moderately and highly permeable,
41 although on the England and Wales maps this was transposed to non, minor and major aquifer)
42 and three soil leaching potential classes (high, medium and low). The groundwater vulnerability
43 maps for England and Wales do not consider travel time to the water table because depth to
44 water was not then generally available on a regional basis (Robins et al. 1994). Furthermore,
45 they were designed for a conservative pollutant that would not degrade in the unsaturated zone
46 (Palmer et al. 1995).

47 Similar maps were produced for Scotland and Northern Ireland, but have recently been replaced
48 as part of the Water Framework Directive characterisation work. New groundwater vulnerability
49 screening layers were produced for Scotland at an effective scale of 1: 100 000 (Ó Dochartaigh
50 et al. 2005), and for Northern Ireland at 1:250 000 scale (Ball et al. 2005), in which the
51 underlying bedrock aquifer class does not directly influence the 'vulnerability' rating.

52 In Ireland, a slightly different approach to vulnerability mapping has been adopted (Daly and
53 Warren 1998; Misstear and Daly 2000). The topsoil was not taken into account but superficial
54 cover (known as subsoil in Ireland) is the main feature of the maps, whereas in the original

55 England and Wales vulnerability maps it was generally only considered as a secondary feature.
56 The Irish approach also considers the nature of the recharge, i.e. whether it is point or diffuse.
57 Superficial cover material posed a significant problem in developing the original vulnerability
58 maps for England and Wales. Where a superficial cover occurs as a granular deposit and where
59 it is saturated it was depicted as an intergranular aquifer. Wherever till is greater than 5 m thick,
60 it was assumed to be of sufficiently low permeability to protect the underlying aquifer and was
61 depicted as a stippled area denoting likely protection subject to field checking of the cover
62 material. Glacial till may vary in composition from silty clay to sandy silt over limited horizontal
63 and vertical distances, with corresponding variations in permeability. It may also be fractured or
64 incised.

65 Improved information on the lithology, variability and thickness of superficial material is the
66 desired next step for improved aquifer vulnerability assessments (e.g. McMillan et al. 2000; Ó
67 Dochartaigh et al. 2005; Fitzsimons and Misstear 2006). Current geological maps of superficial
68 deposits alone are not suitable for determining groundwater vulnerability, since superficial
69 deposits within a particular mapped unit may vary significantly in composition, not only on a
70 local or regional scale but also on a national scale (Anon. 2004).

71 This paper describes a new methodology for supporting assessment of groundwater vulnerability
72 on a nationwide basis using improved characterisation of the superficial deposits. Figure 1
73 illustrates how this method seeks to improve groundwater vulnerability assessments through the
74 incorporation of Quaternary domain conceptual models, local geological and hydrogeological
75 knowledge, national datasets including estimates of drift thickness (part of BGS's GeoSure
76 national dataset on natural ground instability across Great Britain) and site specific information
77 (e.g. boreholes and site investigation reports). Amalgamation of all existing information and
78 knowledge regarding the properties of the superficial deposits enables national maps of recharge
79 potential and attenuation potential to be produced. In this initial phase these maps provide
80 national coverage at 1:625 000 scale but more detailed maps i.e. 1:50 000 scale are being

81 produced by the Environment Agency which has recently adopted the method described in this
82 paper. Additionally, the maps should evolve to incorporate new information, allowing the
83 vulnerability classifications to be refined and improved.

84 Figure 1. The available resources for improving the hydrogeological characterisation of the
85 superficial deposits.

86

87 **The influence of superficial deposits on bedrock hydrogeology**

88 Superficial deposits influence the vulnerability of bedrock groundwater in two main ways (see
89 Figure 2). Firstly, they affect the recharge potential (Rushton 2005): the volume of water or
90 pollutant that can travel through the superficial deposits to the bedrock aquifer below. The
91 relative recharge potential relates to the superficial deposits and does not reflect whether the
92 underlying aquifer (in particular bedrock) could physically accept a high, moderate or even low
93 amount of recharge (poorly productive fractured bedrock aquifers underlie a significant
94 proportion of Great Britain and their ability to accept recharge due to low storage and
95 transmissivity can be significantly restricted). Secondly, superficial deposits have a large impact
96 on the potential attenuation of pollutants. Superficial deposits may often reduce the vulnerability
97 of an underlying aquifer to pollution by increasing both the travel time for pollutants to reach the
98 water table, allowing time for degradation of some contaminants, and also retarding
99 contaminants through sorption on clay surfaces or organic material.

100 Figure 2 Factors influencing recharge and attenuation in the sub-surface. The methodology
101 described in this paper is concerned only with the processes in the superficial deposits.

102 *Recharge*

103 Recharge through superficial deposits is largely determined by the thickness of the deposits,
104 lithology, vertical fracturing and the architecture of the sediments. Investigations of the
105 characteristics of low permeability superficial deposits in North America concluded that recharge

106 can occur through all Quaternary deposits, even where matrix permeability is very low (Keller et
107 al. 1988; van der Kamp 2001). In Great Britain it is consistently reported that recharge occurs
108 through glacial till (Klink et al. 1997; Foster 1998; Marks et al. 2004; Cuthbert 2005). Both
109 Rushton (2005) and Fitzsimons and Missteart (2006) identify an important factor for recharge as
110 the prevailing hydraulic gradient between till and underlying aquifer. Soley and Heathcote
111 (1998) observed recharge of 36 mm a^{-1} through till on interfluvies in East Anglia, decreasing to
112 10 mm a^{-1} in areas of thicker cover where the underlying Cretaceous Chalk bedrock aquifer is
113 locally confined.

114 The role of superficial deposit thickness and permeability in controlling recharge was
115 highlighted by Daly and Warren (1998) in the context of vulnerability mapping guidelines for
116 Ireland and further developed by Missteart et al. (2009). As unsaturated permeability can only be
117 measured experimentally, grain size is often used as an indicator. Primary permeability may also
118 be enhanced by weathering and fracture development (Gerber et al. 2001).

119 Where fractures occur in superficial deposits, e.g. in a fractured till, gravity by-pass flow may
120 occur and may be the dominant flow mechanism (Cuthbert 2005), although mainly restricted to
121 the uppermost 3 m in Ireland (Missteart et al. 2008). Studies of glacial till in situ for waste
122 disposal often indicate the importance of fracturing which can increase permeability by 1 – 3
123 orders of magnitude. These fractures are most common at shallow depths (<10 m) in oxidised
124 and weathered deposits, but can also be found at greater depths (10 – 30 m) (Keller et al. 1988;
125 Frederica 1990; van der Kamp 2001; Cuthbert 2005). Presence of fractures will, therefore, have
126 an important bearing on recharge potential and the methodology considers this on a subjective
127 basis.

128 On rare occasions, superficial deposits may increase the vulnerability of groundwater to
129 pollution through focussed recharge draining from the periphery of the deposits and leading to
130 the development of localised karst-like features in carbonate aquifers where runoff from
131 superficial deposits may be acidic (MacDonald et al. 1998).

132

133 *Attenuation*

134 The processes promoting contaminant attenuation depend on the type of soil and rock, the types
135 of contaminant and the associated activity. Attenuation is generally most active in the soil,
136 where bacterial activity is greatest. The unsaturated zone, is nevertheless, of importance as it
137 represents the most widespread line of defence against pollution of groundwater (e.g. Rodvang
138 and Simpkins 2001). The key indicators in the unsaturated zone relate to the process of sorption,
139 ion exchange, filtration and precipitation. Of these, ion exchange is the main overall process, the
140 others being dependent on the nature of the pollutant as much as the nature of the unsaturated
141 zone medium. In this way a single value describing the potential for the medium to attract
142 cations, the cation exchange capacity (CEC), is the most useful parameter in assigning its
143 attenuation potential. CEC describes the process of attracting cations to a negatively charged
144 surface – usually clay minerals (Appelo and Postma 1993). However, CEC values are
145 commonly not available for the vadose zone and not universally available for the soil zone. It is
146 useful, therefore, to use clay mineral content of the vadose zone as a surrogate for CEC.

147 A second part of the attenuation process is controlled by the availability of carbon as a catalyst
148 for adsorption and precipitation in the medium (Smith and Lerner 2007). Thus the two key
149 indicators used to derive the attenuation potential of the superficial material making up the
150 vadose zone are clay mineral and organic contents.

151 Increased travel time can help the progression of attenuation processes in the Quaternary
152 deposits and allow time for the contaminants to degrade. Therefore, vertical permeability and
153 thickness are also factors to be considered (Foster 1998).

154

155 **Quaternary domains mapping**

156 Subdivision of the Quaternary has always been problematic. Although an approach to
157 standardisation of mapping is currently being promoted by the BGS (Anon. 2004; McMillan
158 2005), to date there has not been a standard approach for Great Britain. The traditional
159 stratigraphic subdivision of the Quaternary paid little regard to either the processes under which
160 deposition took place or the lithologies that were deposited (see Figure 3 for 1:625 000
161 Quaternary map).

162 Figure 3 Superficial deposits of Great Britain at 1:625 000 scale (copyright BGS).

163 An alternative approach to classification is to define characteristic regional provinces which can
164 be subdivided into smaller domains (Anon. 2004). The landmass may be divided into two
165 provinces, Glaciated and Non-glaciated based on landscape evolution, geomorphology and the
166 nature and distribution of superficial deposits. The Glaciated province includes upland and
167 lowland Britain and the whole of Northern Ireland. In this province much of the Quaternary
168 record has been removed or modified by subaerial erosion and successive Quaternary
169 glaciations. In the Non-glaciated Province, which lies to the south of the southern limit of the
170 Anglian glaciation, there remains a more complete record of the processes of weathering, erosion
171 and deposition which were driven by climate changes throughout the Quaternary (Anon. 2004).

172 Within the two main Provinces, eleven domains have been differentiated (Figure 4), based partly
173 on their geomorphology and assemblages of superficial deposits and partly on genetic linkages to
174 the surface processes which formed them. A number of domains have been subdivided into sub-
175 domains (40 in total) to allow for local variations which largely relate to the characteristics of the
176 underlying bedrock, especially in southern England.

177 Detailed domain assessment has yet to be made for Northern Ireland but a preliminary review
178 identified six domains, the most extensive of which is the Till dominant domain which covers 80
179 % of the land surface of Northern Ireland. The Till dominant domain (TD) is subdivided into two
180 sub-domains, TD1 and TD2 on the basis of the proportion of till to rock near surface.

181 Figure 4 Map of the eleven main Quaternary Domains as identified by Anon. (2004). The
182 boundaries of the sub domains are also outlined.

183 **Methodology**

184 Systematic use of additional information and expert knowledge has been applied to reinterpret
185 the existing superficial deposits geological map (Figure 5). Central to the method is a data
186 matrix. Within the matrix, the expert geological knowledge is captured and then re-interpreted by
187 hydrogeologists in terms of recharge potential and attenuation.

188 Figure 5 An outline of the methodology described in this paper to develop recharge and
189 attenuation potential maps by using expert knowledge to interpret existing superficial geology
190 maps.

191 The first step is to overlay the eleven Quaternary Domains (see Figure 4) on top of the digital
192 Superficial Deposits Geological Map (the method is scale-independent but for the purposes of a
193 trial it was applied at 1: 625 000 scale). This allows the origin and thus possible nature of, for
194 example, sandy till in the Scottish Highlands and clayey till mapped in East Anglia to be
195 identified as separate entities. For each domain there follows a description of the associated
196 landforms, lithological deposits and depositional processes. This information is developed into a
197 description of the likely hydrogeological processes pertaining to each domain, and uses case
198 study information where available to support this analysis. These descriptions serve two
199 purposes: they are there to inform users of the methodology about the characteristics of the
200 domains, and they provide background information to assist in classifying lithological units
201 within each domain.

202 The matrix describes the recharge and attenuation potential for each geological unit in each
203 domain. This has been done by asking geological experts a series of questions for the deposits in
204 each domain (Table 1). The matrix provides a framework for meaningful and systematic
205 discussion between hydrogeologists and Quaternary geologists. The lithology of glacial till is
206 governed across much of Great Britain by the nature of the bedrock. Glaciers scoured material

207 from bedrock and in general did not move it far before depositing it to form till. In consequence,
208 glacial tills developed on sandstones or coarse crystalline rocks tend to be sandy and permeable,
209 while those developed on mudstones or strata such as the Coal Measures are clay rich and have
210 low permeability. The matrix describes whether the lithology of the till within a particular
211 domain largely reflects the bedrock lithology. If it does, the soil parent map (a simplified
212 geological map based on bedrock lithology) was used to determine recharge potential and
213 attenuation.

214 Table 1. Framework for discussion between Quaternary geological experts and hydrogeologists

215 Assessment of recharge potential has been made largely on the basis of the composition and
216 grain size of the deposits, and the likelihood of encountering continuous low permeability layers.
217 For most of the lithotypes the composition is likely to vary and the assessments have been based
218 on the relative proportions of the different grades of material. The recharge potential is based on
219 the perceived permeability of the deposits. Recharge potential assessments were made as High
220 (H), Medium (M) or Low (L) and a rationale for the assessments provided. This assessment is
221 carried out twice, both for the majority (primary) situation and for a subordinate (secondary)
222 situation; for example, the primary recharge potential through clayey till is Low but the
223 secondary potential where the till is locally more granular may be Medium. At this stage the
224 fracturing has not been considered – but a general likelihood of fracturing can be indicated by
225 the thickness of the deposits (see below). There is no quantification of the amount of recharge
226 that the deposits could transmit, merely an assessment of their relative ability to transmit
227 infiltrating water (recharge) in general. The method does not take into account other known
228 elements of recharge processes such as the effects of runoff from less permeable deposits with
229 focussed recharge around the periphery and does not consider the thickness of the deposit.

230 Assessment of the attenuation capacity has been made by estimating the presence of clay
231 minerals (as a surrogate for cation exchange capacity) and organic material such as peat. The soil
232 cover is not assessed – just the likely clay and organic material within the superficial deposit.

233 The same primary and secondary classification is also made for attenuation potential although
234 many primary and secondary classifications may be the same. Peat, for example, has High
235 primary and High secondary attenuation potential. In this way the primary classification relates
236 to the predominant lithology within each lithotype. For example, an alluvial deposit may be
237 predominantly sand and gravel, but there may also be minor clay layers within that deposit. The
238 predominant (primary) attenuation potential is low, however, where clays are present the
239 (secondary) attenuation potential is high. It is, therefore, the interpreted understanding of the
240 nature and variability of the superficial geological deposit which generates the assessment of
241 recharge and attenuation potential.

242 Thickness can be important to recharge and attenuation potential in several ways: very thin
243 deposits may be discontinuous, thinner deposits may be fractured throughout whereas increased
244 travel times through a thicker deposit allow more time for degradation of contaminants. Rather
245 than including thickness implicitly within the assessment (as part of the rating) it was left to be
246 included explicitly as a separate layer as required. Five categories were chosen:

- 247 • Absent, or not sufficiently thick to be mapped.
- 248 • 1 – 3 m – superficial deposits can be variable or discontinuous, recharge may occur in
249 holes within the cover.
- 250 • 3 - 10 m – tills can be fractured within the upper 10 m, therefore, where the till is less
251 than 10 m thick there is a higher potential for bypass flow with higher recharge potential
252 and lower attenuation potential.
- 253 • 10 - 30 m – fracture and bypass flow can occur but are less likely.
- 254 • more than 30 m thick. There is a greater chance of heterogeneity and saturated deposits,
255 including tills, are likely to behave as aquifers in their own right unless the entire
256 sequence is impermeable.

257 Central to the methodology is the assignment of glacial till lithologies derived from bedrock
258 lithology for various Quaternary Domains. Table 2 shows the attenuation and recharge potential
259 predicted for superficial material derived from each of 15 generic bedrock types. A variety of
260 bedrock lithologies fall under the Sedimentary Mixed classification and the attenuation and
261 recharge potential properties are variable within this category. The values are amended wherever
262 the properties of the Sedimentary Mixed bedrock are known. The general rule of underlying
263 bedrock directly influencing overlying till properties is not universal, and exceptions are
264 captured in the matrix, for example, the till cover along parts of the north-east coast of England
265 is smeared with low permeability marine clays which reduce the recharge potential from
266 Medium to Low. Other exceptions occur where two ice sheets converge, e.g. the Irish Sea Ice
267 and the Welsh Ice meet and interdigitate over central Wales, and a similar situation is present
268 over much of East Anglia where one till may overlies another. In these instances, especially
269 where the till is thick, the younger deposits in the sequence may bear little relationship to the
270 underlying bedrock.

271 Table 2 Generic classification of superficial cover based on underlying bedrock lithology (L –
272 Low, M – Medium, H – High)

273 The rationale for each classification is given in the matrix so that the assessment is transparent.
274 Revisions can be made as local users add their knowledge to the matrix. The classifications and
275 maps can mature and evolve as more data are incorporated.

276

277 **GIS application of the methodology**

278 The different geological datasets (superficial deposits, Quaternary Domains and bedrock
279 geology) and the data matrix were integrated to create 1: 625 000 recharge and attenuation
280 potential maps using Access™ and ArcGIS™ (version 9.1) (Figure 6). Analyses for Northern
281 Ireland have not been presented because of the preliminary nature of the domains mapping. The
282 shapefile of superficial deposits at 1: 625 000 scale was used as the main layer. Each polygon

283 within this shapefile was attributed with the domain/sub domain into which it falls (Figure 4).
284 Where the superficial polygon intersected more than one domain it was divided accordingly. The
285 resulting polygons were then attributed according to the bedrock classifications, again splitting
286 them if more than one bedrock type was intercepted (Table 2).

287 Figure 6 Process diagram of GIS application of the methodology.

288 Once the polygons had been created, recharge and attenuation potential were assigned to each
289 polygon based on the classifications given in the matrix table for all the lithotypes within each
290 domain. Modifications were made where local variations had been noted by the geologist in their
291 assessment, e.g. to distinguish between low permeability clayey silt and clay 'carse' type beach
292 deposits and more permeable raised beach deposits, both identified only as coastal deposits on
293 the superficial deposits map. The classification for the underlying bedrock type was used (with
294 exceptions, where known) where the lithotype was till.

295 The output contains nine categories for both recharge and attenuation potential which show the
296 natural variability displayed by combining the primary and secondary classifications. The nine
297 classes (as primary classification/secondary classification) for both recharge and attenuation
298 potential are: HH, HM, HL, MH, MM, ML, LH, LM and LL.

299 Each shape file was converted into a grid or raster file of 1 km² grid size to provide recharge
300 potential (Figure 7) and attenuation (Figure 8). Where two or more classifications occurred
301 within the 1 km², a majority rule was applied (i.e. the dominant classification was assigned to the
302 whole grid square). Two grids are available, one for recharge potential and the other for
303 attenuation potential.

304 Figure 7 Map of recharge potential using 1:625 000 data. For high recharge potential, more
305 water would be expected to pass through the superficial deposits. Information on the thickness
306 of the deposits is not incorporated.

307 Figure 8 Map of attenuation potential using 1:625 000 scale data. For high attenuation,
308 negligible contamination is expected through the superficial deposits. Information on drift
309 thickness is not incorporated.

310 **Outputs and Validation**

311 The national 1: 625 000 scale dataset of superficial deposits was used to provide national
312 coverage, albeit at a coarse scale, enough for proof of method. The output is at 1 km grid size.
313 Now that the methodology has been adopted by the Environment Agency larger scale data are
314 being used and a more meaningful output derived, however, the results so far are a guide for
315 comparison with existing perceptions of regional recharge and vulnerability of groundwater.
316 The current output should not, therefore, be applied to support specific environmental
317 judgements. The map outputs have been compared with a number of other datasets including
318 HOST data (Hydrology of Soil Types), the new groundwater vulnerability screening maps for
319 Scotland and the results from a number of regional studies on recharge through superficial
320 deposits. One of the best methods for testing the maps is also to compare the assessment of a
321 particular area with the experience and knowledge of local hydrogeologists.

322 HOST is a dataset which is based on the 1 x 1 km National Soil Map data together with data
323 from the Centre for Ecology & Hydrology on the behaviour of river catchments (Boorman et al.
324 1995). There are 29 HOST classes, which describe the dominant pathways of water movement
325 through the soil and substrate. Standard Percentage Runoff (SPR) is given for each HOST soil
326 class. These were used to calculate an estimate of the Standard Percentage Infiltration (SPI) for
327 each HOST class ($100\% - \text{SPR} = \text{SPI}$).

328 The calculation of SPR is based on the analysis of flood event data, i.e. collated river flow and
329 rainfall data for storm events. SPR is the percentage of rainfall that causes a short term increase
330 in river flow at the catchment outlet and, therefore, the SPI is not directly comparable with
331 infiltration. Nonetheless the SPI map provides an indication of the general response of different
332 catchments to rainfall. In general there is a broad agreement between the datasets, but areas of

333 till which have Sedimentary Mixed bedrock parent material offer a poor match suggesting that
334 additional local knowledge input is desirable.

335 Figure 9 shows the results of the statistical comparison with the HOST data for SPI. There is
336 broad agreement i.e. areas with lower percentage SPI have a greater proportion of superficial
337 cover assigned as low recharge potential while those with higher percentage SPI have a greater
338 proportion of high recharge potential superficial cover.

339 Figure 9 Plot of recharge potential against the HOST Standard Infiltration potential.

340 The recharge and attenuation potential maps were also compared with the vulnerability screening
341 maps of superficial deposits determined for Scotland (Ó Dochartaigh et al. 2005). The map of
342 attenuation potential shows a similar pattern to the Scottish vulnerability screening map for the
343 superficial deposits – low and medium values are prevalent over much of Scotland except where
344 peat or other clay/organic rich materials occur. There is particularly good agreement with the
345 lower vulnerability rating. There is also good agreement between the recharge potential map and
346 the superficial aquifer productivity maps generated for Scotland using pumping test data, HOST
347 data and expert knowledge (MacDonald et al. 2005).

348 The maps have also been examined against more detailed recharge studies reported in the
349 literature. For example, the East Anglian Till is known to permit little recharge to the underlying
350 Chalk, particularly in interfluvial locations (Marks et al. 2004). Till has a significant impact on
351 recharge quantity and distribution to the underlying Chalk aquifer. Beneath the interfluvial
352 recharge appears to be lower than previous estimates of 20 – 40 mm a⁻¹ (Soley and Heathcote,
353 1998), maybe as low as 5 mm a⁻¹ (Marks et al. 2004). The recharge potential map (Figure 7)
354 shows that the Low Medium recharge potential classification is consistent with these studies.
355 Similarly in the Cheshire Basin, recharge to the unconfined Sherwood Sandstone aquifer has
356 been estimated at approximately 350 mm a⁻¹, but where till cover is present, recharge to the
357 underlying aquifer is only about 52 mm a⁻¹ (Vines 1984). These variations are reflected by

358 recharge potential classifications for the till covered areas of LM and LL and for the areas where
359 till is absent recharge potential is classed as HM.

360 Initial testing of the output data layers against national nitrate concentrations in groundwater for
361 England and Wales (Johnson 2006) shows that the GIS output reflects the expected influence of
362 poorly permeable superficial deposits on nitrate fate and transport. However, the correlation is
363 weak in some places largely due to insufficient local detail.

364 The nitrate comparison indicated that key properties of the poorly permeable superficial deposits
365 that are required to predict nitrate concentration in groundwater are the overall coverage of the
366 cover material (expressed as percentage cover per km²) and its thickness, e.g. >10 m superficial
367 cover, both of which are available as digital datasets held by BGS. The importance of these
368 parameters varies according to the domain under consideration, e.g. depth of superficial deposits
369 is more important for the Till Dominant than for the Dissected Till domain. It is likely that both
370 these parameters need to be incorporated into the matrix if the characterisation of superficial
371 deposits is to be improved significantly.

372

373 **Discussion**

374 The methodology enables rapid assessment of the hydrogeological properties of superficial
375 deposits on a national/regional scale. Use of the 1: 625 000 scale datasets provides a broad
376 overview that is sufficient for proof of method. Now that the methodology is established it is
377 intended that it should be applied to more detailed data. It is also intended that input of local
378 knowledge will be invaluable for assessing the success of the methodology and improving the
379 resultant maps where anomalies or shortfalls exist.

380 Inevitably there are a number of constraints imposed by the methodology. These include scale
381 issues and use of classifications at 1: 625 000 scale which inevitably simplifies otherwise
382 heterogeneous systems. However, the heterogeneity is captured by the primary and secondary
383 estimates of recharge and attenuation potential, i.e. a highly variable 3D deposit is classed as HL,

384 while homogenous deposits may be HH. In addition it has only been possible to capture some of
385 the available three dimensional data in the matrix, for example, for the areas of thick superficial
386 deposits in the Cheshire Basin. Application at a larger scale will enable the characterisation of
387 the 3D data in more detail.

388 There are also mapping issues, for example, some areas of raised beach/estuarine deposits,
389 particularly around the Wash, have been mapped as alluvium on the 1: 625 000 superficial
390 geology map. For many areas moraine and till have not been subdivided and variations in the
391 morphology and composition of these deposits have implications for permeability and
392 attenuation potential. The matrix methodology is helpful here, since explanations can be made in
393 the matrix by local experts so that these mapping issues can be overcome by adjusting
394 classifications manually.

395 There is also the question of whether there are enough assessment categories. For some deposits
396 a highly variable category may be more appropriate than a fixed descriptor. There is also an
397 argument that thickness should be integrated into the classifications rather than examined as a
398 separate dataset. These issues can be addressed readily in subsequent usage, given that the
399 database can be modified easily.

400 The coarse scale used in the trial does not allow recharge and attenuation properties to be
401 assessed confidently at local, or even small catchment scale. However, the methodology is
402 essentially scale independent, and can be readily adapted to apply the most commonly used
403 geological dataset at 1: 50 000 scale, although it should be considered that the quality and precise
404 type of mapping varies from one area to another at that scale.

405 The methodology represents a means of capturing existing and developing geological knowledge
406 and extracting relevant elements to improve conceptual hydrogeological understanding. The
407 importance of this interaction between hydrogeologists and mapping geologists who have an
408 extensive understanding of the nature and variability of superficial deposits in a particular area
409 cannot be over-estimated. Further development of the associated GIS and incorporation of more

410 detailed data will allow application at local catchment/groundwater body scale. In addition, the
411 GIS could be used as an ongoing archive of local understanding and knowledge, the layers and
412 datasets developed complementing and enhancing existing vulnerability maps and water
413 resource models.

414 Although there are a number of additional datasets available, outputs need to be evaluated before
415 they are formally incorporated into the assessment procedure. The initial testing provides a
416 baseline from which to judge the relative improvement in outputs due to the inclusion of any
417 new dataset into the method. New data sets could include: soils data, land drainage data, rock
418 head elevation, nitrate monitoring and loading data, and seasonal variability of loading.

419

420 **Conclusions**

421 Superficial deposits have been evaluated using Quaternary domain classifications to provide
422 improved geological and hydrogeological characterisation. Geological and hydrogeological data
423 have been integrated with expert local geological and hydrogeological knowledge using a data
424 matrix to provide improved assessment of the impact of superficial deposits on groundwater
425 vulnerability. New groundwater vulnerability maps of recharge potential and attenuation
426 potential within superficial cover material over bedrock aquifers have been developed at
427 1:625 000 scale.

428 A qualitative assessment of the results compares favourably with existing recharge studies and
429 regional hydrogeological knowledge. Favourable comparison with data from HOST provides a
430 quantitative assessment of the method. The methodology excludes direct consideration of the
431 soil zone and other intrinsic processes which are encapsulated in recharge estimation and in the
432 assessment of attenuation potential of the vadose zone. These exclusions do not impact on the
433 assessment for the superficial strata and the output is consistently robust.

434 As the methodology is scale independent it is recommended that it is applied at 1: 50 000 scale
435 subject to data licensing issues. This will enable more detailed characterisation and assessment

436 of groundwater vulnerability at a local level. Rescaling will also provide an opportunity to add
437 additional components to the methodology as required.

438 Although the method is constrained by the availability of data and local knowledge, one of its
439 strengths is that it allows new information to be incorporated at a later date so that assessments
440 are improved and validated. This will enable a flexible, holistic and dynamic approach to
441 groundwater vulnerability assessments and ensure that they are based on the best available
442 information. The methodology is currently being trialled by the Environment Agency.

443

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567

Figure 1 The available resources for improving the hydrogeological characterisation of the superficial deposits.

Figure 2 Factors influencing recharge and attenuation in the sub-surface. The methodology described in this paper is concerned only with the processes in the superficial deposits.

Figure 3 Superficial deposits of Great Britain at 1:625 000 scale (copyright BGS).

Figure 4 Map of the eleven main Quaternary Domains as identified by QMT (2004). The boundaries of the sub domains are also outlined.

Figure 5 An outline of the methodology described in this paper to develop recharge and attenuation potential maps by using expert knowledge to interpret existing superficial geology maps.

Figure 6 Process diagram of GIS application of the methodology.

Figure 7 Map of recharge potential using 1:625 000 data. For high recharge potential, more water would be expected to pass through the superficial deposits. Information on the thickness of the deposits is not incorporated.

Figure 8 Map of attenuation potential using 1:625 000 scale data. For high attenuation, then negligible contamination is expected though the superficial deposits. Information on drift thickness is not incorporated.

Figure 9 Plot of recharge potential against the HOST Standard Infiltration potential.

Table 1 Framework for discussion between Quaternary geological experts and hydrogeologists

Table 2 Generic bedrock classification (L – Low, M – Medium, H – High)

Matrix fields to be completed on each mapped superficial deposit within each domain type	Assessor
General Description	
Map issues	
Sub-domains	
General Thickness	
Architecture	
Lithology	
Clay content	
Organic content	
Horizontal permeability	
Vertical permeability	
Primary adsorption potential	
Secondary adsorption potential	
Rationale for adsorption potential	
Primary recharge potential	
Secondary recharge potential	
Rationale for recharge rating	

Geological expert

Hydrogeologist

Bedrock classification	Attenuation Potential		Recharge Potential	
	Primary	Secondary	Primary	Secondary
Chalk	H	M	L	L
Crystalline	L	M	H	M
Crystalline Coarse	L	M	H	M
Crystalline Fine	H	M	L	M
Metamorphic	M	H	M	L
Meta Argillaceous	H	L	L	H
Meta Limestone	M	L	M	M
Meta Rudaceous	M	L	M	H
Sedimentary Arenaceous	L	M	H	M
Sedimentary Argillaceous	H	M	L	L
Sedimentary Coal	H	M	L	M
Sedimentary Limestone	M	L	L	M
Sedimentary Mixed	M	M	M	M
Sedimentary Rudaceous	M	M	H	M
Volcanoclastic	L	M	M	M

Current situation

National geological map of the superficial deposits
Isolated case studies of groundwater movement through superficial deposits

Available Resources

Conceptual Models (Quaternary Domains)

Hydrogeologists

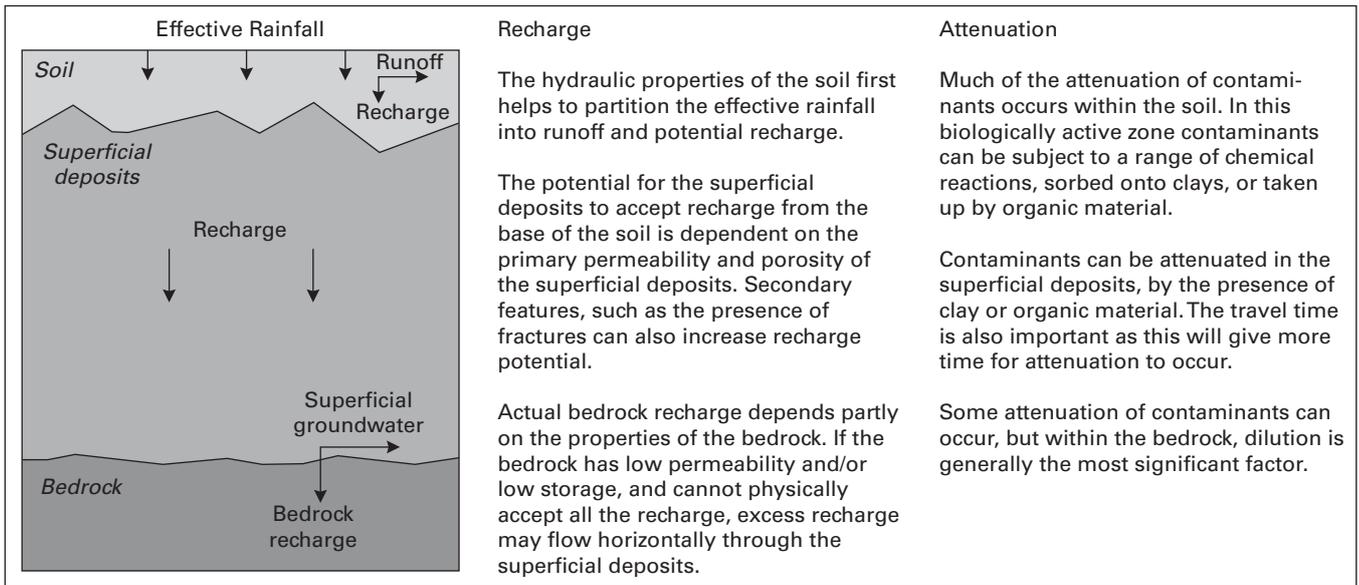
Regional Geologists

Information on thickness

Borehole Information National datasets (e.g. Soil, Groundwater
Vulnerability maps, geochemical data

Desired Destination

National map of the attenuation potential of the superficial deposits
National map of the recharge potential of the superficial deposits



Recharge

The hydraulic properties of the soil first helps to partition the effective rainfall into runoff and potential recharge.

The potential for the superficial deposits to accept recharge from the base of the soil is dependent on the primary permeability and porosity of the superficial deposits. Secondary features, such as the presence of fractures can also increase recharge potential.

Actual bedrock recharge depends partly on the properties of the bedrock. If the bedrock has low permeability and/or low storage, and cannot physically accept all the recharge, excess recharge may flow horizontally through the superficial deposits.

Attenuation

Much of the attenuation of contaminants occurs within the soil. In this biologically active zone contaminants can be subject to a range of chemical reactions, sorbed onto clays, or taken up by organic material.

Contaminants can be attenuated in the superficial deposits, by the presence of clay or organic material. The travel time is also important as this will give more time for attenuation to occur.

Some attenuation of contaminants can occur, but within the bedrock, dilution is generally the most significant factor.

