Electron scattering by whistler-mode ELF hiss in

plasmaspheric plumes

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- 3 Abstract. Non-adiabatic loss processes of radiation belt energetic elec-
- trons include precipitation loss to the atmosphere due to pitch-angle scat-
- 5 tering by various magnetospheric plasma wave modes. Here we consider elec-
- 6 tron precipitation loss due to pitch-angle scattering by whistler-mode ELF
- ⁷ hiss in plasmaspheric plumes. Using wave observations and inferred plasma
- densities from the Plasma Wave Experiment on the Combined Release and

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Radiation Effects Satellite (CRRES), we analyze plume intervals for which well determined hiss spectral intensities are available. We then select 14 repesentative plumes for detailed study, comprising 10 duskside plumes and 4 11 non-duskside plumes, with local hiss amplitudes ranging from maximum values of above 300 pT to minimum values of less than 1 pT. We estimate the electron loss timescale τ_{loss} due to pitch-angle scattering by hiss in each cho-14 sen plume as a function of L-shell and electron energy; τ_{loss} is calculated from quasi-linear theory as the inverse of the bounce-averaged diffusion rate evaluated at the equatorial loss cone angle. We find that pitch-angle scattering by hiss in plumes can be efficient for inducing precipitation loss of outer-18 zone electrons with energies throughout the range 100 keV - 1 MeV, though the magnitude of τ_{loss} can be highly dependent on wave power, L-shell, and electron energy. For 100 keV - 200 keV electrons, typically $\tau_{loss} \sim \! 1$ day while 21 the minimum loss timescale $(\tau_{loss})_{min} \sim \text{hours}$. For 500 keV - 1 MeV electrons, typically $(\tau_{loss})_{min} \sim \text{days}$, while $(\tau_{loss})_{min} < 1$ day in the case of 23 large wave amplitude (~ 100 's pT). Apart from inducing direct precipitation loss of MeV electrons, scattering by hiss in plumes may reduce the generation of MeV electrons by depleting the lower-energy electron seed population. Models of the dynamical variation of the outer-zone electron flux should incorporate electron precipitation loss induced by ELF hiss scattering in plasmaspheric plumes.

1. Introduction

In order to understand and quantify energetic electron flux variations in the inner magnetosphere, it is necessary to assess both the electron energization and loss processes. Loss mechanisms may be adiabatic, which are temporary, or non-adiabatic which result in 32 net particle loss. Non-adiabatic loss processes include precipitation loss to the atmosphere due to pitch-angle scattering by plasma waves, and loss due to particle drift across the magnetospheric boundary. Radiation belt electrons can undergo gyroresonant pitch-angle scattering by various wave-modes, including whistler-mode VLF chorus, plasmaspheric ELF hiss, and electromagnetic ion cyclotron (EMIC) waves, e.g., see Summers et al. [2007a, 2007b] and references therein. In the present paper we analyze a particular form of electron precipitation loss, namely that due to pitch-angle scattering by ELF hiss in plasmaspheric 'plumes'. The plasmasphere is a cold (a few eV), dense $(10-10^4 \text{ cm}^{-3})$ plasma torus surrounding 41 the Earth in the innermost magnetosphere [e.g., Carpenter, 1963; Chappell et al., 1970; Carpenter and Park, 1973; Horwitz et al., 1990; Carpenter and Lemaire, 1997; Lemaire and Grinquiz, 1998; Ganguli et al., 2000; Goldstein, 2006; Dent et al., 2006]. The multiion (H⁺, He⁺, O⁺) plasma comprising the plasmasphere derives from the ionosphere and co-rotates with the Earth. The region of cold plasma rotation and the overall shape of the plasmasphere is controlled by the interaction of the co-rotational electric field and the solar wind influenced dawn-to-dusk cross-tail electric field. During intense geomagnetic storms the plasmaspheric boundary layer, or plasmapause, can lie inside L=2 for several days [Baker et al., 2004], while during prolonged periods of quiet geomagnetic conditions the

plasmasphere can extend to beyond geosynchronous orbit ($L \sim 6.6$) and possess no distinct outer boundary [Goldstein et al., 2003]. Following geomagnetically disturbed periods, and as a result of interplay between forces driving the plasma sunward and corotational 53 forces, plasma typically drains from the body of the plasmasphere in the afternoon local time sector. The resulting large-scale plasma structures which stretch toward the outer magnetosphere are usually attached to the plasmasphere and are called plasmaspheric plumes or plasmaspheric drainage plumes. Historically, they have been called tails [Taylor] et al., 1971 or detached plasma elements (or blobs) [Chappell, 1974]. Using plasma density data inferred from the Plasma Wave Experiment on the Combined Release and Radiation Effects Satellite (CRRES), Moldwin et al. [2004] found that plumes can exist at all local times under all levels of geomagnetic activity, but that most were observed in the noonto-dusk sector following enhanced geomagnetic activity. In many of the methods hitherto used it should be noted that whether or not the observed plasma structures were attached 63 to the plasmasphere could not easily be determined. Excellent global images of evolving plasmaspheric plumes have been provided by the extreme ultraviolet (EUV) imager of the 65 Imager for Magnetopause-to-Aurora Global Exploration (IMAGE) satellite [e.g., Sandel et al., 2003; Goldstein et al., 2003, 2004, 2005; Spasojevic et al., 2003, 2004; Burch, 2006; 67 Goldstein, 2006. In situ measurements from the four CLUSTER satellites confirm that plumes rotate around the Earth, with their feet attached to the main plasmasphere fully 69 co-rotating, but with their tips often rotating more slowly and moving outward away from the Earth [Darrouzet et al., 2006]. 71

Plasmaspheric hiss is a broadband ELF electromagnetic whistler-mode emission which occurs in the frequency range from ~ 100 Hz to several kHz. Hiss is present over a

broad region of the plasmasphere even during geomagnetically quiet periods and intensifies during storms or substorms [Smith et al., 1974; Thorne et al., 1974; Meredith et al., 2004]. Broadband amplitudes of hiss range from 10 pT or below during quiet periods to ~ 100 's pT during disturbed times [Smith et al., 1974; Tsurutani et al., 1975; Meredith et al., 2004. Hiss is generally field-aligned near the magnetic equator and tends to propagate more obliquely at higher latitudes [Parrot and Lefeuvre, 1986; Santolik et al., 2001]. There are extensive observations of plasmaspheric hiss, e.g., see Hayakawa and Sazhin [1992], Meredith et al. [2004], Masson et al. [2004] and references therein. Whistler-mode hiss 81 has also been observed in plasmaspheric plumes [Chan and Holzer, 1976; Cornilleau-Wehrlin et al., 1978; Hayakawa et al., 1986; Parrot and Lefeuvre, 1986. Analyzing 83 CRRES wave and particle data, Meredith et al. [2004] found that plasmaspheric hiss peaks in particular equatorial ($|MLAT| < 15^{\circ}$) and midlatitude ($15^{\circ} < |MLAT| < 30^{\circ}$) regions, mainly on the dayside, and that generally his amplitudes depend on L-shell, MLT and magnetic latitude, as well as substorm activity. Plasmaspheric hiss, together with other 87 whistler-mode emissions [Abel and Thorne, 1998], plays an important role in controlling the structure of the Earth's radiation belts. Lyons and Thorne [1973] showed that the formation of the quiet-time 'slot' region between the inner (1.3 < L < 2.5) and outer (3 < L < 7) radiation belts can be explained as an equilibrium balance between inward radial diffusion and pitch-angle scattering loss of energetic electrons to the atmosphere 92 induced by plasmaspheric hiss [Lyons et al., 1972]. Plasmaspheric hiss can also cause 93 scattering loss of MeV electrons from the outer radiation belt over a timescale of days, or less, under appropriate conditions [Tsurutani et al., 1975; Albert, 1994, 2003; Summers et al., 2007b]. Meredith et al. [2006a] used CRRES data to measure the gradual decay

of energetic (214 keV - 1.09 MeV) electron fluxes in the outer zone following enhanced

geomagnetic activity. Meredith et al. [2006a] and Summers et al. [2007b] found that scattering by plasmaspheric hiss propagating at zero or small wave normal angles could account for the measured electron decay rates over a wide range of energies and L-shells. The generation mechanism of plasmaspheric hiss has not been fully resolved and remains 101 controversial. There are two leading theories for the source of plasmaspheric hiss, namely, 102 in situ natural instability in the magnetosphere [e.g., Etcheto et al., 1973; Thorne et 103 al., 1979; Huang et al., 1983], and lightning-generated whistlers [e.g., Sonwalkar and 104 Inan, 1989; Draganov et al., 1992; Green et al., 2005. The analysis by Green et al. 105 [2005] supporting lightning as the dominant source for plasmaspheric hiss was disputed 106 by Thorne et al. [2006]; see also the reply by Green et al. [2006]. Meredith et al. [2006b] subsequently analyzed the entire CRRES database of plasmaspheric hiss together with 108 the global distribution of lightning and concluded that while higher-frequency hiss (2 -109 5 kHz) is generated by lightning, lower-frequency hiss (100 Hz - 2 kHz) is generated by 110 natural instability in space. Evidence that lower-frequency hiss intensifies during enhanced 111 geomagnetic activity [e.g., Meredith et al., 2004] points to natural instability as the origin of lower-frequency hiss. 113 There is increasing interest in wave-particle interactions occurring in plasmaspheric plumes with respect to their role in influencing particle dynamics in the inner magneto-115 sphere. Pitch-angle scattering by EMIC waves in plumes can cause significant precipita-116

tion loss of energetic protons [Burch et al., 2002; Spasojevic et al., 2004; Burch, 2006].

Summers et al. [2007b] found that an assumed realistic spatial distribution of EMIC waves

and hiss in an empirically measured plume could induce rapid scattering loss of outer zone

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electrons. The relative contributions to electron scattering by hiss and EMIC waves in plumes depend on the electron energy and L-shell, as well as the wave properties; see Figures 21 and 22 of Summers et al. [2007b]. In the present paper we analyze hiss-electron 122 interaction in plasmaspheric plumes selected from the CRRES mission. Specifically, we determine intervals during which CRRES crossed a plume and select a subset of plume 124 intervals for which well-observed hiss data are available. We then use quasi-linear theory 125 to determine hiss-induced pitch-angle scattering rates at the loss cone for electrons of specified energy at given L-values. We can thereby estimate timescales for precipitation 127 loss of energetic electrons in the inner magnetosphere due to scattering by hiss in plumes. The present study is the first to determine electron precipitation loss timescales due to 129 scattering by measured hiss in observed plumes. We present our selection of plume intervals and associated hiss data in section 2. In section 3 we summarize the necessary 131 quasi-linear theory required for our calculations. In section 4 we present our estimates for 132 electron loss timescales due to scattering by hiss in the chosen plumes. Finally, in section 133 5 we discuss the significance of our results. 134

2. CRRES plume and wave observations

CRRES was launched on 25 July, 1990 and functioned until 12 October, 1991. The spacecraft had a geosynchronous transfer orbit, namely an elliptical orbit with a perigee of $1.05 R_E$ and an apogee of $6.26 R_E$ with respect to the Earth's center, with an inclination of 18.15° . The outermost L-shell reached by CRRES was $L \sim 8$. The orbital period was approximately 9 hours 55 minutes, and the apogee of CRRES precessed from 10.00 MLT to 14.00 MLT through midnight before the mission terminated. The satellite was able to provide excellent coverage of the radiation belts for nearly 15 months since it traversed

the inner magnetosphere on average about 5 times per day. The wave data and plasma 142 densities used in this study were obtained from the Plasma Wave Experiment (PWE) on board CRRES. This experiment measured electric fields from 5.6 Hz to 400 kHz, using 144 a 100 m tip-to-tip long wire antenna, with a dynamic range covering a factor of at least 10⁵ in amplitude [Anderson et al., 1992]. The electric field detector was thus able to detect waves from below the lower hybrid resonance frequency (f_{LHR}) to well above the 147 upper hybrid resonance frequency (f_{UHR}) for a large fraction of each orbit. The maximum plasma density that could be measured was $\sim 2000 \text{ cm}^{-3}$ because of the upper frequency 149 limit of the instrument. The CRRES PWE also included a boom-mounted search coil magnetometer that covered the frequency range from 5.6 Hz to 10 kHz and operated until 151 the March 1991 storm. While the electric field data were sampled with high-frequency 152 resolution by the PWE sweep frequency receiver at eight seconds per spectra above 6.4 153 kHz, the search coil data were sampled by a 14-channel analyzer that sampled the magnetic 154 field eight times per second every other 32 seconds. 155

We determine the presence of a plume by monitoring the behavior of the plasma density
as inferred from observations of the upper hybrid resonance frequency. If, while CRRES
is clearly outside the plasmasphere, the density suddenly increases, remains high for some
time, and then suddenly decreases, we identify the region as a potential plume. We refer
to such a region as a plume for simplicity, even though CRRES observations cannot determine if the identified high-density region is attached to the plasmasphere. We also use the
absence of electrostatic electron cyclotron harmonic (ECH) waves as a criterion for identifying high-density plasma regions, as described by *Meredith et al.* [2004]. Identification of
a plume and its boundary can be problematic observationally, and is, to a degree, subjec-

tive. Determination of a boundary of a plume is straightforward if it is sharp, and difficult 165 if it is gradual. This situation likewise applies to the determination of the boundary of the plasmapause itself. To complement the techniques applied for identifying plumes in 167 our study, we also make use of the rigorous plume selection criteria of Moldwin et al. [2004]. The comprehensive study of plumes during the CRRES mission by Moldwin et al. 169 [2004] employed the database of plasmapause locations identified by Moldwin et al. [2002] 170 and the empirical plasmaspheric and trough density models developed by Sheeley et al. 171 [2001]. These three studies used a common database of plasma density derived from the 172 CRRES Plasma Wave Experiment. Moldwin et al. [2002] identified the innermost steep density gradient in the density profile as the plasmapause, a factor of 5 drop within half 174 an L-shell being required. In order to select 'plasmaspheric' intervals located outside the plasmapause, Moldwin et al. [2004] used L=3 as a dividing line for whether to use the plasmaspheric or trough density model. If the plasmapause is located earthward of L=3, 177 plasmaspheric plume intervals are defined as those whose density exceeds the trough plus 178 one standard deviation density of the Sheeley et al. [2001] model. If the plasmapause is 179 located outside of L=3, plasmaspheric plume intervals are defined as those whose density exceeds the Sheeley et al. [2001] plasmaspheric model. These models are scaled to each 181 orbit to account for the wide variability in the plasmaspheric density from day to day. The criterion used by Moldwin et al. [2004] to select a plume is that the density throughout 183 the requisite interval must exceed the model value of the plasmaspheric density (or trough 184 plus one standard deviation) over a minimum of 8 consecutive observations (a duration of $\sim 1 \text{ min}$).

In our study we choose 14 plume intervals which we specify in Table 1, according to 187 orbit number, by giving the start and end values of universal time (UT), magnetic local time (MLT), L-shell, and magnetic latitude (MLAT). We have chosen 10 plumes with a 189 duskside MLT location, namely crossed by outbound CRRES orbits 605, 672, 673, 674, 810, 869, 871, 939, 941, and 977. The remaining 4 plumes, crossed by outbound orbits 191 302 and 446, and inbound orbits 297 and 446, are non-duskside. The 14 chosen plume 192 crossings are illustrated in Figure 1 in which we also show the approximate trajectory for CRRES orbit 446. Our chosen plumes were likewise identified as plumes by Moldwin et 194 al. [2004], with the exception of the 3 plumes associated with orbits 297 and 446. These latter plumes were not selected in the Moldwin et al. [2004] study because the density 196 did not satisfy their conservative plasmaspheric density criteria. We nevertheless regard these features as representative non-duskside "plumes" because of their distinctly elevated densities compared to the surrounding trough. 199

In some of the selected plumes common to the present study and that of *Moldwin et al.* [2004], the specified start and end of the plume interval, as for instance given by L-shell, differ slightly because of the differing plume boundary criteria used in the two studies. This issue does not lead to serious difficulties in our investigation since we base the conclusions of our analysis on electron loss timescales that are calculated 'well inside' each plume. Thus, possibly spurious 'edge effects' are readily eliminated.

In Figure 2 we show measured CRRES electron density profiles corresponding to the outbound (blue) and inbound (red) portions of orbits 446, 869, 939, and 977 in the respective panels (a), (b), (c), and (d). For comparison purposes, in each panel we also show upper and lower black curves representing respectively the saturated plasmasphere

density and trough density given by the Carpenter and Anderson [1992] model. The
plume intervals for orbits 446, 869, 939, and 977 as specified by L-shell range in Table 1
can be observed to match the corresponding intervals of elevated density in panels (a) (d) of Figure 2.

In Figure 3 the measured wave electric field spectral intensities (in $V^2m^{-2}Hz^{-1}$) are 214 plotted as a function of UT for the complete CRRES orbits 446, 869, 939, and 977 in the 215 respective panels (a), (b), (c), and (d). The magnetic local time, magnetic latitude, and L-shell are given at hourly intervals. The solid white line shows the value of the electron 217 gyrofrequency f_{ce} , determined from the measured ambient magnetic field, and the dashed white lines below f_{ce} represent $0.5f_{ce}$, $0.1f_{ce}$, and the lower hybrid resonance frequency, 219 f_{LHR} . The dotted white lines above f_{ce} correspond to the first four harmonics of f_{ce} . The 220 solid red line denotes the upper hybrid resonance frequency $f_{UHR} = (f_{pe}^2 + f_{ce}^2)^{1/2}$ (where 221 f_{pe} is the electron plasma frequency) calculated from the lower-frequency cut-off of the 222 electromagnetic continuum, and the red dashed line represents f_{UHR} calculated from wave 223 emissions at f_{UHR} inside the plasmapause. The chosen plumes in the orbits 446, 869, 939, 224 and 977 are indicated in Figure 3, together with their associated hiss emissions. Profiles 225 of the AE index are provided at one-minute time resolution. The empirical position of 226 the plasmapause as defined by Carpenter and Anderson [1992] is also marked.

We base the calculations in our study on hiss in the frequency range 104 < f < 1040Hz. The general criterion used in this paper to identify hiss in plumes is that used by

Meredith et al. [2004] to identify plasmaspheric hiss, namely, ECH wave amplitudes for

frequencies in the range $f_{ce} < f < 2f_{ce}$ must be less than 0.0005 mVm⁻¹ in order for

wave emissions below f_{ce} in the frequency band 104 < f < 1040 Hz to be identified

as hiss. Whistler-mode chorus has a frequency range 0.3 < f < 30 kHz in the region 233 3 < L < 7. Consequently, chorus can lie in our chosen hiss frequency band at higher L-shells. However, whistler-mode chorus is usually observed outside the plasmasphere 235 and high-density regions so can be excluded from consideration in our chosen plumes. At lower L-shells, magnetosonic waves can also fall into our chosen hiss frequency band. 237 These waves, which are closely confined to the equatorial region, are enhanced during 238 active conditions below the lower hybrid resonance frequency f_{LHR} , represented by the lowest dashed line in the spectrograms in Figure 3. We find no evidence of enhanced 240 magnetosonic waves within our chosen plumes, except possibly during orbit 871. For the chosen plume in this orbit, his intensities may be slightly over-estimated in the region 242 4.25 < L < 5.25 as a result of contamination by magnetosonic waves.

In order to convert observed hiss electric field spectral intensities to magnetic field intensities, we use a cold-plasma dispersion relation for parallel-propagating whistler-245 mode waves (equation (4) of section 3, with $\varepsilon = 0$), Maxwell's induction equation, and expression (1) given by Meredith et al. [2004]. Magnetic field wave intensities over the 247 frequency range 104 < f < 1040 Hz are then defined as an integral of the averaged wave spectral intensity (nT²Hz⁻¹). The corresponding wave amplitudes are obtained by taking 249 the square root of the wave intensities, as detailed in section 3. Conversion from electric to magnetic fields is relatively insensitive to wave normal angle for wave normal angles 251 less than 50° , if $f < 0.5 f_{ce}$ [Meredith et al., 2004]. We discuss our assumption of parallel 252 wave propagation further below. 253

In Figure 4 we present hiss spectral intensities in nT²Hz⁻¹ within four chosen plumes during orbits 446, 869, 939, and 977, in the respective panels (a), (b), (c), and (d). In

each panel the spectral intensity is shown at a range of specified L-shells within the plume. In Figure 5 we show the measured values of the local hiss amplitudes at the given L-shells within each of our 14 chosen plume intervals. Local wave amplitudes range from maximum values that exceed 300 pT, for the plumes in orbits 810 and 939, to minimum values of less than 1 pT, in orbits 297, 446(In), and 810. This probably represents the widest range of hiss amplitudes to be expected in plasmaspheric plumes. Further, since we have mainly chosen more-commonly occurring duskside plumes while also including a selection of non-duskside plumes, we can consider that the choice of plume intervals for our study is reasonably general. For each of the 14 chosen plumes, we present in Table 2 an average value for the hiss amplitude $\overline{\Delta B}$ (pT) calculated by averaging the measured spectral intensity along each plume crossing.

3. Theory

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From Summers [2005] (equations (10) and (17)), we can write the local pitch-angle diffusion coefficient for electron cyclotron resonance with field-aligned R-mode electromagnetic waves in the form,

$$D_{\alpha\alpha} = \frac{\pi}{2} \frac{|\Omega_e|^2}{B_0^2} \frac{1}{(E+1)^2} \sum_{j=1}^N \left(1 - \frac{\omega_j \cos \alpha}{v \, k_j} \right)^2 \frac{I(k)}{|v \cos \alpha - d\omega_j/dk_j|} \tag{1}$$

for broadband waves of intensity I(k) or $\hat{I}(f)$ (nT²/Hz), defined on the frequency range $\omega_1 < \omega < \omega_2$, where

$$\Delta B^2 = \int_{-\infty}^{\infty} I(k) \, dk = \int_{f_1}^{f_2} \hat{I}(f) \, df \,, \tag{2}$$

and ΔB is the wave amplitude; $f = \omega/2\pi$, $f_1 = \omega_1/2\pi$, and $f_2 = \omega_2/2\pi$; α is the particle pitch-angle and v is the particle speed; E is the dimensionless particle kinetic energy given by $E = E_k/(m_e c^2) = \gamma - 1$ where $\gamma = (1 - v^2/c^2)^{-1/2}$ is the Lorentz factor (c is

the speed of light), and m_e is the electron rest mass; $|\Omega_e| = eB_0/(m_ec)$ is the electron gyrofrequency, where e is the unit charge and B_0 is the magnitude of the uniform static magnetic field; the wave frequency ω_j and wavenumber k_j (where $j = 1, 2, \dots, N$) satisfy the gyroresonance condition

$$\omega_j - v \, k_j \cos \alpha = |\Omega_e|/\gamma \,, \tag{3}$$

as well as the dispersion relation,

$$\left(\frac{ck}{\omega}\right)^2 = 1 - \frac{(1+\varepsilon)/\alpha^*}{(\omega/|\Omega_e|-1)(\omega/|\Omega_e|+\varepsilon)},$$
(4)

284 where

$$\alpha^* = \Omega_e^2 / \omega_{pe}^2 \tag{5}$$

is an important cold-plasma parameter; $\varepsilon = m_e/m_p$ where m_p is the proton rest mass; and $\omega_{pe} = (4\pi N_0 \, e^2/m_e)^{1/2}$ is the plasma frequency where N_0 is the electron number density.

It is convenient to express formula (1) in terms of the practical wave intensity $\hat{I}(f)$ (nT²/Hz). Then, also introducing the variables,

$$x = \omega_j / |\Omega_e|, \quad y = c k_j / |\Omega_e|, \tag{6}$$

we thereby obtain the result,

$$D_{\alpha\alpha} = \frac{1}{4} \frac{|\Omega_e|^2}{B_0^2} \frac{1}{(E+1)^2} \sum_{j=1}^N \left(1 - \frac{x \cos \alpha}{y \beta} \right)^2 \frac{\hat{I}(f) |F(x,y)|}{|\beta \cos \alpha - F(x,y)|}, \tag{7}$$

where (from (3))

$$y = (x - 1/\gamma)/(\beta \cos \alpha). \tag{8}$$

In (7), the function F(x,y) is given by expression (C1) in Summers [2005]; $\beta = v/c = [E(E+2)]^{1/2}/(E+1)$; and x satisfies the quartic equation (A1) given also in Summers DRAFT

In order to apply (7) to the assumed dipole magnetic field of the inner magnetosphere, it remains to carry out bounce-averaging of (7) to take account of the magnetic mirror-like geometry. Using the formalism given by *Summers et al.* [2007a], we write the bounceaveraged diffusion coefficient $\langle D_{\alpha\alpha} \rangle$ as

$$\langle D_{\alpha\alpha} \rangle = \frac{1}{S(\alpha_{eq})} \int_0^{\lambda_m} D_{\alpha\alpha} (\alpha) \frac{\cos \alpha \cos^7 \lambda}{\cos^2 \alpha_{eq}} d\lambda, \qquad (9)$$

304 where

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$$S(\alpha_{eq}) = 1.3 - 0.56 \sin \alpha_{eq}.$$
 (10)

In (9), α_{eq} is the equatorial pitch-angle of a particle, and λ is the magnetic latitude of a particle with pitch-angle α at any point along a field line; α_{eq} , λ , and α satisfy the relation,

$$\sin^2 \alpha = f(\lambda) \sin^2 \alpha_{eq}, \tag{11}$$

310 where

$$f(\lambda) = (1 + 3\sin^2 \lambda)^{1/2} / \cos^6 \lambda$$
. (12)

 λ_m is the latitude of the mirror point of the particle and is given by the equation,

$$X^{6} + (3\sin^{4}\alpha_{eq})X - 4\sin^{4}\alpha_{eq} = 0, \tag{13}$$

with $X = \cos^2 \lambda_m$.

We substitute the local diffusion coefficient $D_{\alpha\alpha}(\alpha)$ given by (7) into (9), and regard α as a function of α_{eq} and λ , as given by (11). Thus, the bounce-averaged diffusion coefficient $\langle D_{\alpha\alpha} \rangle$ is a function of α_{eq} . The background magnetic field B_0 occurring in $D_{\alpha\alpha}(\alpha)$ in (9)

is replaced by the value,

$$B_0 = B_{eq} f(\lambda), \qquad (14)$$

320 where

$$B_{ea} = B_{local} / f(\lambda_{local}). \tag{15}$$

 B_{eq} is the equatorial magnetic field, and B_{local} is the locally observed magnetic field at the observed magnetic latitude λ_{local} , corresponding to the observed L-value. In the absence 323 of other data to show latitudinal variations in density, we assume that the background electron number density N_0 is constant along a field line $(N_0 = N_{eq} = N_{local})$. We likewise 325 assume that the hiss spectral intensity is constant along a field line. From a statistical survey of CRRES data Meredith et al. [2004] found that hiss peaks near the equatorial 327 $(|MLAT| < 15^{\circ})$ and midlatitude $(15^{\circ} < |MLAT| < 30^{\circ})$ regions. CRRES data are not 328 available at high latitudes ($|MLAT| > 30^{\circ}$). Our assumption that hiss is also present at high latitudes is partially justified by other studies. For example, Thorne et al. [1973], 330 using OGO5 search coil magnetometer data, found that hiss was present on almost every pass through the plasmasphere. Thorne et al. [1973] found little distinction between 332 lower latitude ($|MLAT| < 30^{\circ}$) and high latitude ($|MLAT| > 30^{\circ}$) plasmaspheric hiss emissions, and concluded that properties of hiss remain largely constant throughout the 334 plasmasphere. We make the assumption that his has constant spectral intensity along a field line on the basis of the best information available. Nevertheless, we recognize that if the wave power is confined to a lower range of latitudes then our calculations may over-337 estimate the higher-energy loss rates since the waves resonate with higher-energy electrons at higher latitudes. Dependence of electron loss timescales on the latitudinal distribution

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of hiss, for a given energy and L-value, is examined by Summers et al. [2007b] (section 3).

Evaluation of the integral in (9) can be carried out by standard numerical quadrature which requires evaluation of the integrand at a set of λ -values (quadrature points) in the range $0 < \lambda < \lambda_m$. This requires, in particular, determination of the local diffusion coefficient $D_{\alpha\alpha}(\alpha)$ at the quadrature points. Therefore, at each quadrature point the relevant resonant roots x of the above-noted quartic equation must be found.

We take as an estimate of the electron loss timescale,

$$\tau_{loss} = (1/\delta)(1/\langle D_{\alpha\alpha}^{LC} \rangle), \tag{16}$$

where $\langle D_{\alpha\alpha}^{LC} \rangle$ is the bounce-averaged diffusion coefficient (9) evaluated at $\alpha_{eq} = (\alpha_{LC})_{eq}$ where $(\alpha_{LC})_{eq}$ is the equatorial loss cone angle given by

$$\sin(\alpha_{LC})_{eq} = [L^5(4L-3)]^{-1/4}. \tag{17}$$

In order to account for the limited angular (MLT) spread of the observed hiss in a given plume, we have inserted into (16) a drift-averaging factor δ which we specify in the following section. The value of τ_{loss} depends on the kinetic energy E, L-shell, the measured hiss spectral intensity \hat{I} , the drift-averaging factor δ , and the equatorial value of the parameter α^* , namely $\alpha^*_{eq} = (\Omega_e^2/\omega_{pe}^2)_{eq}$. The local electron gyrofrequency $|\Omega_e|$ was determined from the CRRES fluxgate magnetometer instrument [Singer et al., 1992]. The local electron plasma frequency ω_{pe} was estimated from CRRES data on electrostatic waves at the upper hybrid frequency and the low-frequency cut-off of electromagnetic continuum radiation, as described by Meredith et al. [2002]. In the upper panel of Figure 6 we show the variation of α^*_{eq} with L-value, for the chosen plume intervals, deduced from local CRRES values

for $|\Omega_e|$ and ω_{pe} . In the lower panel of Figure 6 we show minimum electron energies for 362 cyclotron resonance with hiss at the frequency 1040 Hz, as a function of L-value, corresponding to the upper panel. The minimum resonant energy for electron resonance with 364 hiss is obtained by setting $\sigma = e$ (for electrons) and s = 1 (for R-mode waves) in formula (16) of Summers et al. [2007a]. 366 As described in this section, the determination of the electron loss timescale τ_{loss} in 367 our study assumes that the observed whistler-mode hiss is strictly field-aligned. CRRES data do not provide information on the wave-normal angle or angular spread of the waves. 369 While the assumption that the waves are field-aligned is likely to be an approximation, we consider that our method for calculating τ_{loss} yields reasonably reliable results based 371 on the relatively limited available data. In support of our method, we cite the recent 372 analysis by Summers et al. [2007b] who calculated electron loss timescales due to scat-373 tering by plasmaspheric hiss during low geomagnetic activity in the region 3 < L < 5. 374 Summers et al. [2007b] assumed field-aligned hiss with zero wave-normal distribution and 375 predicted electron loss timescales in good agreement with the measured values obtained 376 from CRRES Medium Electrons A data by Meredith et al. [2006a]. It should nevertheless be pointed out that inclusion of higher-order scattering could significantly alter 378

Specifically, we would expect increased loss timescales if the wave-normal angle is large,

the scattering rates near the edge of the loss cone if the hiss becomes strongly oblique.

as demonstrated by Meredith et al. [2006a].

4. Electron loss timescales

Electron loss timescales calculated in this paper must of course be considered in the context of plume lifetimes overall. Plumes have been observed over the duration of many

consecutive CRRES orbits, e.g., see Figure 8 of Moldwin et al. [2004]. Since the CRRES orbital period is about 10 hours, this indicates that plumes can last from 10 hours to more than 1 day. Global imaging by the EUV imager of the IMAGE satellite has tracked the 386 evolution of various plumes over several hours to more than 1 day [e.g., Spasojevic et al., 2003; Goldstein et al., 2004. Very few studies have measured the full global evolution of a 388 plasmaspheric plume from its creation to its complete dissipation. It is possible that some 389 plumes persist for several days. For practical purposes, we take an upper limit for the lifetime of a plume to be 5 days, in which case a value of τ_{loss} exceeding 5 days indicates 391 that electron scattering by hiss is ineffective for that particular plume at the L-shell and electron energy under consideration. Plume formations exceeding 5 days in duration are 393 likely to consist of multiple plumes formed in succession. However, at geosynchronous orbit $(L \sim 6.6)$ cold dense regions in narrow MLT channels are commonly observed over 10-day intervals or longer. In an investigation using multiple geosynchronous satellites, 396 Moldwin et al. [1994] found that plasmaspheric plasma was absent on only 13% of the 397 days in the study interval. 398 For electrons of a given energy E, we determine the loss timescale τ_{loss} due to scattering 399

by hiss at a given L-shell in a chosen plume as the inverse of the bounce-averaged diffusion coefficient $\langle D_{\alpha\alpha} \rangle$ evaluated at the equatorial loss cone angle (formulae (16) - (17)). An orbiting energetic electron traverses a plume only for a fraction of its orbit. To take account of the azimuthal (MLT) spread of a plume we have included a drift-averaging factor δ in (16). The azimuthal spread of a particular plume varies during its evolution and is typically 0.1 R_E to 1.5 R_E or more [e.g., Spasojevic et al., 2003; Darrouzet et al., 2006]. Hereinafter, we take the drift-averaging factor δ as 6% since this appears to correspond approximately to the 'typical' observed azimuthal width of a plume.

Shprits et al. [2006a] recently analyzed the controlling effect of the pitch-angle scattering 408 rates near the loss cone on energetic electron lifetimes, and found that the electron phase space density reaches an equilibrium shape within hours of the simulation when scattering 410 rates do not drop below 1/10 of the value near the edge of the loss cone for up to a 30°-wide 411 range of pitch angles. In this case, electron lifetimes were found to be primarily controlled 412 by scattering rates near the edge of the loss cone. Further, Shprits et al. [2006a] found 413 that while a drop in the diffusion coefficients by a factor of 100 to 1000 near $\alpha_{eq} = 90^{\circ}$ results in weak scattering at high pitch angles, the lower pitch-angle particle distribution 415 decays on a timescale comparable to that determined by the diffusion rate near the edge 416 of the loss cone. Herein, we utilize the findings of Shprits et al. [2006a] and estimate 417 electron loss timescales by using the scattering rate at the edge of the loss cone only in 418 those cases in which the diffusion rate is small over a high pitch-angle range narrower than 419 $75^{\circ} < \alpha_{eq} < 90^{\circ}$. By using this criterion, we expect that our reported timescales afford 420 reasonable estimates of the decay times of at least the bulk of the electron distribution. In order to carry out accurate drift-averaging of the diffusion rates we require specifi-422 cation of the complete MLT distribution of hiss spectral intensity. However, only point measurements of the hiss intensity are made by CRRES at particular MLT-values and spe-424 cific L-shells as the satellite traverses each particular plume. We assume that the average 425 wave power determined along the satellite track is a measure of the MLT wave distribution 426 at a given L-value. Specifically, for each plume we average over all the measured profiles 427 of the hiss spectral intensity, and we use this average intensity, together with the drift-

averaging factor δ , to determine the MLT-averaged scattering rate at each L-shell. In 429 Figure 7 we show examples of profiles of the bounce-averaged and drift-averaged electron diffusion coefficient $\langle D_{\alpha\alpha} \rangle$ for the plume interval in orbit 977 for electrons of energies 100 431 keV, 200 keV, 500 keV, and 1 MeV, at the given L-shells. In Figure 8 we plot the electron loss timescale τ_{loss} at the specified energies as a function of L-shell for the chosen plumes 433 in orbits 446 (Out), 869, 939, and 977 in the respective panels (a), (b), (c), and (d). Hiss 434 intensity for orbit 446 (Out) is the strongest of the 4 non-duskside plume crossings, and hiss intensities during orbits 869, 939, and 977 are among the strongest in the 10 duskside plume crossings (see Figure 5 and Table 2). From Figure 8, and the corresponding figures for the other 10 plumes not shown, we deduce that at a fixed L-shell, τ_{loss} increases as the 438 electron kinetic energy increases from 100 keV to 1 MeV. It is also evident from Figure 8 that scattering by hiss in plumes can be especially effective for electrons of energy 100 -440 200 keV. For instance, for the plume in orbit 977 (panel (d)), for which the average wave 441 amplitude $\overline{\Delta B} = 102$ pT, τ_{loss} ranges from 2.9 to 6.3 hr for 100 keV electrons, and from 442 6.6 to 13.8 hr for 200 keV electrons. For the plume in orbit 939 (panel (c)), for which the 443 wave intensity is strong ($\overline{\Delta B} = 203 \text{ pT}$), the minimum loss timescale $(\tau_{loss})_{min}$ is 0.7 hr for 100 keV electrons and 1.1 hr for 200 keV electrons. For orbit 869 (panel (b), $\overline{\Delta B} = 34$ 445 pT), minimum timescales are $(\tau_{loss})_{min} = 15.2$, 36.9 hr for 100 keV, 200 keV electrons, and for orbit 446 (Out) (panel (a), $\overline{\Delta B} = 48 \text{ pT}$) we find $(\tau_{loss})_{min} = 1.6, 3.7 \text{ days for } 100$ 447 keV, 200 keV electrons. 448 We can also see from Figure 8 that, in general, scattering by hiss in plumes is somewhat 449 less effective for 500 keV electrons, and still more ineffective for 1 MeV electrons. Minimum 450

loss timescales for the plumes in orbits 977, 869, and 446 are respectively $(\tau_{loss})_{min} = 1$,

5.6, and 13 days, for 500 keV electrons. Corresponding respective values for 1 MeV electrons are $(\tau_{loss})_{min} = 3.1$, 18, and 41 days. Particularly rapid scattering of 500 keV - 1 MeV electrons by hiss in plumes is possible, but only in the case of intense waves, e.g., for the plume in orbit 939, $(\tau_{loss})_{min} = 4.6$ hr for 500 keV electrons and $(\tau_{loss})_{min} = 15.4$ hr for 1 MeV electrons.

In Figure 9 we plot the loss timescales for electrons of energy 100 keV, 200 keV, 500 457 keV, and 1 MeV, as a function of L-shell, in the respective panels (a) - (d), for all the 14 chosen plumes in our study. Complementary to Figure 9, we list in Table 3 the number 459 of the chosen plumes for which electrons at each of the energies 100 keV, 200 keV, 500 keV, and 1 MeV, have a loss timescale less than the specified values (0.1, 0.5, 1, and 461 2 days) at some L-shells. Figure 9 and Table 3 essentially summarize the results of our calculations of τ_{loss} for our total selection of plumes. The decrease in efficiency of 463 scattering by hiss as electron energies increase from 100 keV to 1 MeV is confirmed in 464 Figure 9 by the general upward shift of the timescale profiles from panel (a) through to 465 panel (d). Likewise, the number of profiles (or portions of profiles) located above $\tau_{loss} = 5$ 466 days, the nominal timescale above which scattering is ineffective in plumes, progressively increases from panel (a) through to panel (d). The degree of effectiveness of electron 468 scattering by hiss at each of the energies 100 keV, 200 keV, 500 keV, and 1 MeV can be particularly appreciated by viewing each of the panels (a) - (d) of Figure 9 with each 470 corresponding column of Table 3. For instance, for 100 keV electrons scattering is fairly 471 rapid (τ_{loss} < 0.5 day) at some L-shells in 6 of the 14 chosen plumes. Further, in 9 472 plumes scattering of 100 keV electrons can be regarded as at least moderately effective 473 $(\tau_{loss} < 2 \text{ days})$ at some L-shells. It is also useful to examine the entries in each particular

row of Table 3 separately, e.g., the number of plumes for which $\tau_{loss} < 1$ day, at some L-shells, progressively decreases from 7 to 1 as the electron energy increases from 100 keV to 1 MeV. The scattering of MeV electrons in less than 1 day appears to require hiss amplitudes well in excess of 100 pT. Such a statement is not straightforward to qualify accurately, however, since τ_{loss} depends in a complicated way on the various parameters occurring in the formula for the diffusion coefficient (7).

Of the 14 chosen plumes, 6 contain at least reasonably intense hiss, specifically with an average wave amplitude satisfying $\overline{\Delta B} \geq 42$ pT (see Table 2). The plume in orbit 939, with $\overline{\Delta B} = 203$ pT, contains the most intense hiss. We have also selected 3 plumes with relatively weak hiss, satisfying $\overline{\Delta B} \leq 16$ pT. Electron scattering in these plumes (in orbits 297, 302, and 673) is naturally likewise weak, in general, with minimum values of τ_{loss} of at least several days.

Overall, it is clear from the numerical results reported in this section that, under appropriate conditions, hiss in plumes can induce significant precipitation losses of energetic (100 keV - 1 MeV) electrons in the outer zone, 3 < L < 7.

5. Discussion

This is the first study to quantify the contribution of pitch-angle scattering by whistlermode hiss in plumes to the total precipitation loss of outer-zone energetic electrons using
experimental wave data in observed plumes. Understanding the acceleration and loss
mechanisms of radiation belt electrons is needed to develop models for nowcasting and
forecasting of relativistic (> 1 MeV) electrons that are a potential danger to satellites
and humans in space. A primary objective of the twin-spacecraft NASA Radiation Belt
Storm Probes (RBSP) mission [Kintner et al., 2002] and the proposed Canadian Outer

Radiation Belt Injection, Transport, Acceleration and Loss Satellite (ORBITALS) mission 497 [Mann et al., 2006] is to understand the physical processes that control the dynamical variation of outer radiation belt electron fluxes. Wave-particle interactions undoubtedly 499 play a crucial role in radiation belt electron dynamics. Electron gyroresonance with VLF chorus can lead to stochastic acceleration of seed ($\sim 100~{
m keV}$) electrons to MeV 501 energies in the low density regions outside the plasmasphere and plasmaspheric plumes 502 [Summers et al., 1998, 2002, 2004, 2007b; Horne and Thorne, 1998; Roth et al., 1999; 503 Summers and Ma, 2000; Meredith et al., 2002, 2003a; Miyoshi et al., 2003; Horne et al., 504 2005a, 2005b; Varotsou et al., 2005; Omura and Summers, 2006; Shprits et al., 2006b; Li et al., 2007. Relativistic (> 1 MeV) electrons just outside the plasmapause can be 506 scattered by VLF chorus into the loss cone, on timescales of a day, and observed at low altitudes as microburst precipitation [Lorentzen et al., 2001; Thorne et al., 2005]. 508 Scattering by EMIC waves along the duskside plasmasphere can induce precipitation loss 509 of MeV electrons on timescales of several hours to a day [Lorentzen et al., 2000; Summers 510 and Thorne, 2003; Meredith et al., 2003b; Summers et al., 2007b. In the present study 511 we have shown that whistler-mode hiss in plumes can likewise induce precipitation loss of MeV electrons in a day or less, though only in the case of exceptionally strong waves 513 (typically with amplitude 100's pT). Of particular interest in our study is the finding that electrons of energy 100 - 200 keV, which are required to form a seed population from 515 which MeV electrons are generated, can suffer rapid precipitation loss (in a timescale 516 of hours) due to scattering by hiss in plumes. Thus, while scattering by hiss in plumes 517 may not usually induce rapid precipitation loss of MeV electrons, hiss scattering may 518 reduce the generation of MeV electrons by depleting the seed electron population. To

verify the latter process, quantitative modeling of the transport of seed electrons from
the plasmasheet would be required. The general conclusion to our study is that pitchangle scattering by hiss in plumes in the frequency range 104 < f < 1040 Hz can be
efficient for inducing precipitation loss of outer-zone electrons with energies throughout
the range 100 keV - 1 MeV. However, the results in section 4 show that the magnitude
of the precipitation loss timescale can be highly dependent on wave power, L-value, and
electron energy. Further, as we have pointed out above, pitch-angle scattering rates can
be sensitive to wave-normal angle and the latitudinal distributions of density and wave
power. Accordingly, the precipitation loss timescales computed in this paper could be
conservatively regarded as lower bounds.

The competition between acceleration and loss of energetic electrons is determined by wave-particle interactions taking place outside and inside the dense thermal regions 531 comprising the plasmasphere and plasmaspheric plumes. Acceleration and loss of energetic 532 electrons due to gyroresonance with whistler-mode chorus take place outside these thermal 533 regions, while precipitation loss due to pitch-angle scattering by hiss and EMIC waves 534 takes place inside the thermal regions. The generation and global distribution of energetic electrons in the outer zone is therefore greatly influenced by the distribution of thermal 536 plasma. Accurate modeling of the dynamical variation of the outer radiation belt electron flux requires knowledge of the spectral intensity and temporal variation of the appropriate 538 wave modes both inside and outside the thermal plasma regions. 539

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ORBIT	UT	UT	MLT	MLT	L	L	MLAT	MLAT
	(start)	(end)	(start)	(end)	(start)	(end)	(start)	(end)
297 (In)	01:58	04:02	03:55	05:44	6.43	4.18	12.8	7.1
302 (Out)	22:01	23:35	23:17	01:28	3.26	5.86	15.4	16.3
446 (Out)	00:31	01:15	21:22	22:18	4.41	5.63	26.7	23.7
$446 \; (In)$	05:22	07:16	01:06	02:47	6.72	4.76	13.8	6.75
605 (Out)	05:03	06:43	19:14	20:45	4.15	6.05	4.9	-2.6
672 (Out)	15:33	15:48	18:14	18:33	4.45	4.85	-8.5	-9.9
673 (Out)	00:53	01:26	17:35	18:25	3.55	4.45	6.7	3.3
674 (Out)	11:08	11:48	18:09	18:55	4.45	5.55	-15.4	-18.5
810 (Out)	17:30	19:23	15:54	17:43	4.15	6.25	-6.0	-9.0
869 (Out)	00:54	02:46	16:07	17:14	5.65	6.85	-13.9	-14.5
871 (Out)	20:33	22:44	15:05	16:54	4.25	6.35	-7.1	-7.4
939 (Out)	01:12	02:16	13:49	14:48	4.95	6.15	-20.2	-18.8
941 (Out)	21:31	22:26	13:25	14:30	4.05	5.25	-11.8	-10.3
977 (Out)	09:30	11:00	13:06	14:16	5.85	6.65	-21.8	-17.7

Table 1. Specification of the 14 CRRES plume crossings chosen in this study. The magnetic latitude (MLAT) is given in degrees; UT is universal time and MLT is magnetic local time.

ORBIT	297	302	446 (Out)	446 (In)	605	672	673	674	810	869	871	939	941	977
$\overline{\Delta B}$ (pT)	14	16	48	27	42	25	13	37	91	34	60	203	31	102

Table 2. Average hiss amplitude $\overline{\Delta B}$ (pT) calculated by averaging the measured spectral intensity along each chosen plume crossing.

	100 keV	200 keV	500 keV	$1 \mathrm{MeV}$
$\overline{\tau_{loss}} < 2 \text{ day (some } L)$	9	7	3	1
$\tau_{loss} < 1 \text{ day (some } L)$	7	5	2	1
$\tau_{loss} < 0.5 \text{ day (some } L)$	6	2	1	0
$\tau_{loss} < 0.1 \text{ day (some } L)$	1	1	0	0

Table 3. Entries in the table indicate the number of the 14 chosen plumes for which electrons at the indicated energy have a loss timescale less than the specified value at some L-shells.

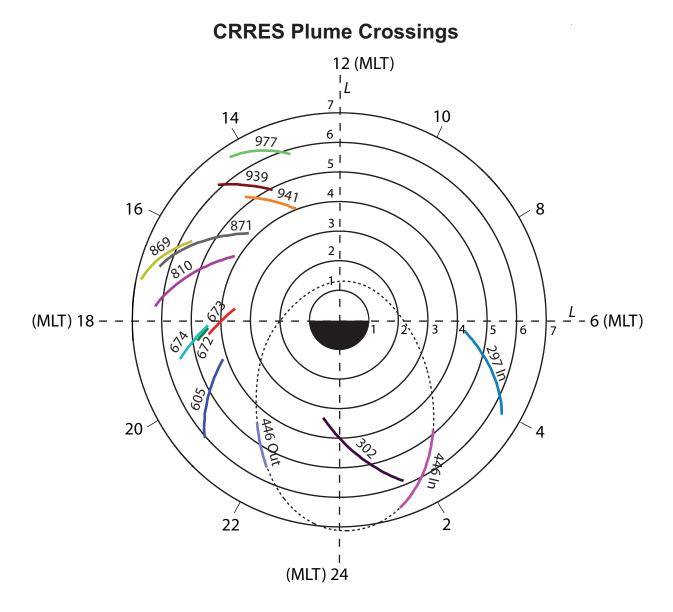


Figure 1. Diagram showing the 14 CRRES plume crossings identified by orbit number, chosen in this study. All the chosen plumes correspond to outbound portions of the specified orbits, except for the indicated 297 (In) and 446 (In) inbound crossings. Also shown is an approximate trajectory for CRRES orbit 446.

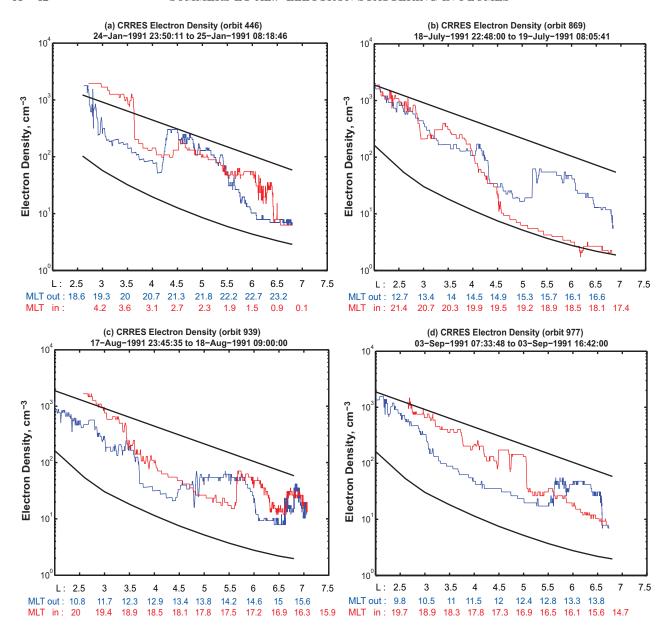
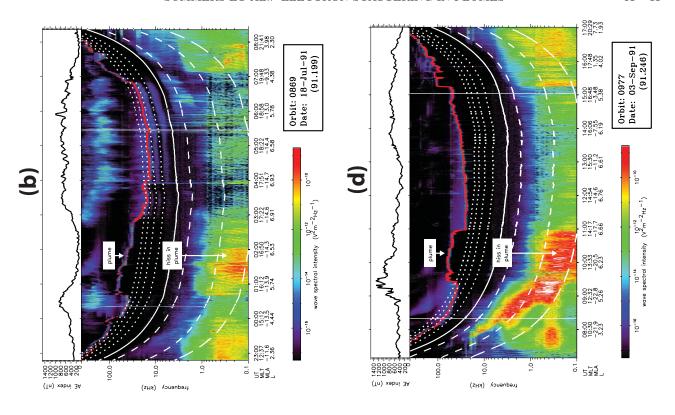


Figure 2. Measured CRRES electron density profiles for orbits 446, 869, 939, and 977. Chosen plume intervals during these orbits are specified in Table 1. The upper and lower black curves in each panel are model profiles of the saturated plasmasphere density and trough density due to *Carpenter and Anderson* [1992].



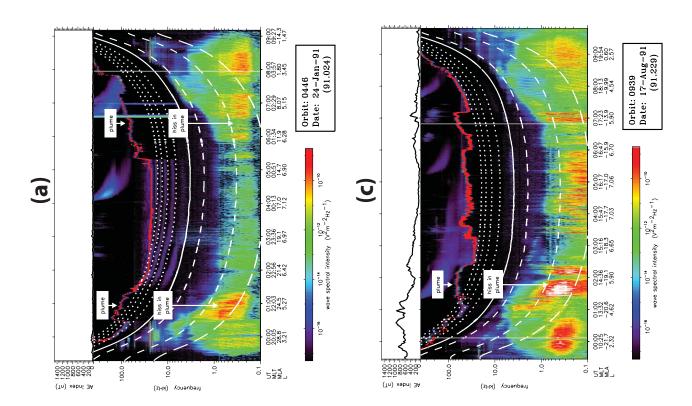


Figure 3. Survey plot of the wave spectral intensities observed on CRRES for orbits 446, 869, 939, and 977 in the respective panels (a), (b), (c), and (d).

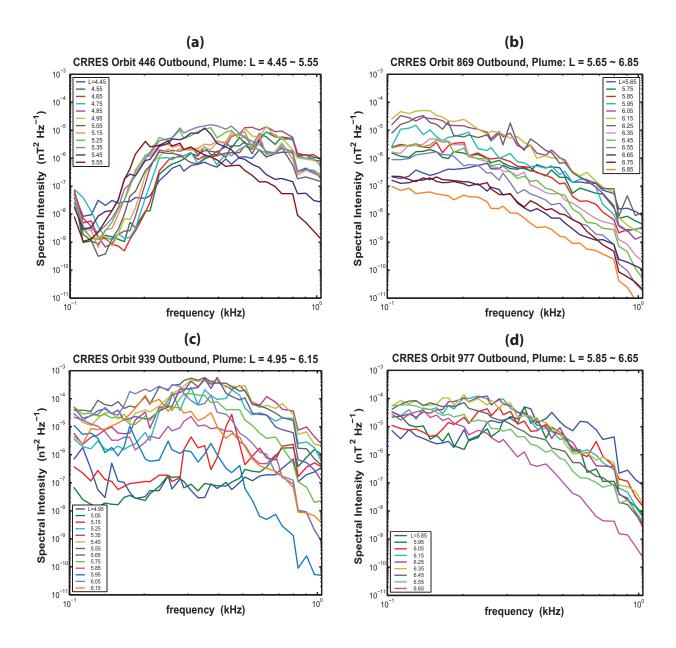


Figure 4. Corresponding to Figure 3, measured hiss spectral intensities at the indicated L-values during the chosen plume crossings from CRRES orbits 446 (Out), 869, 939, and 977.

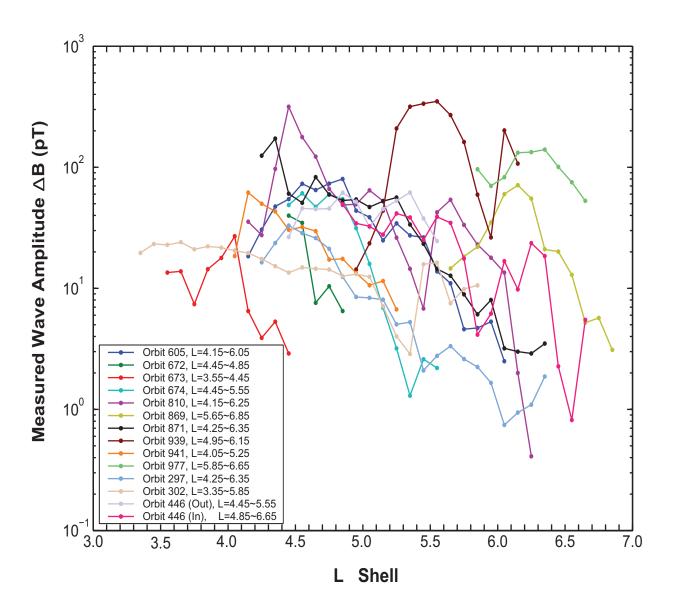


Figure 5. Local hiss amplitude in the frequency range 104 - 1040 Hz measured by CRRES along each chosen plume crossing.

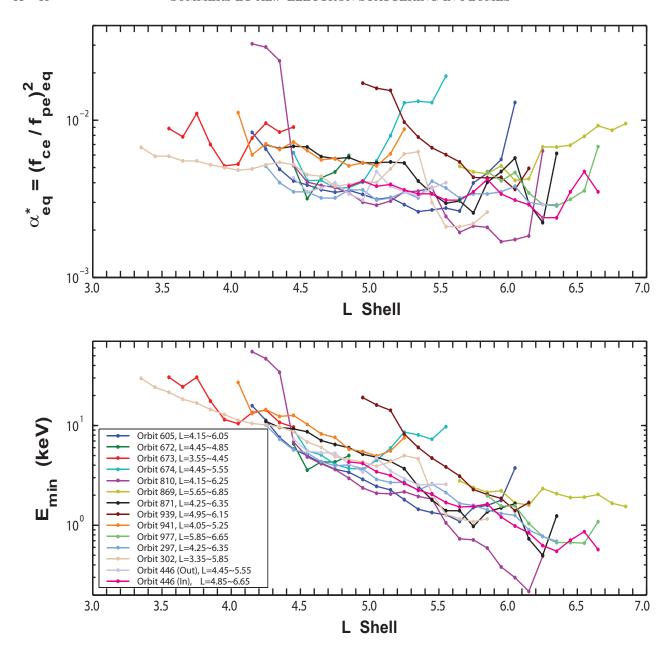


Figure 6. Variation of the equatorial value of the parameter $\alpha^* = (f_{ce}/f_{pe})^2$ with L-value, for the specified plume intervals, inferred from local observed values of f_{ce} and f_{pe} (top panel). In the bottom panel, we show minimum energies for electron resonance with hiss at the frequency 1040 Hz, as a function of L-value, calculated using the values of α^* given in the top panel.

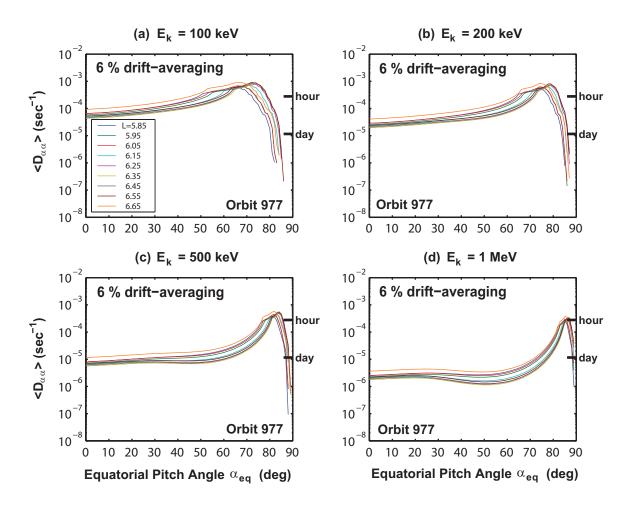


Figure 7. Bounce-averaged pitch-angle diffusion rates for electrons interacting with hiss during the chosen plume crossing for orbit 977, at the indicated L-values and electron energies. Hiss in the plume is assumed to be distributed along the whole field line.

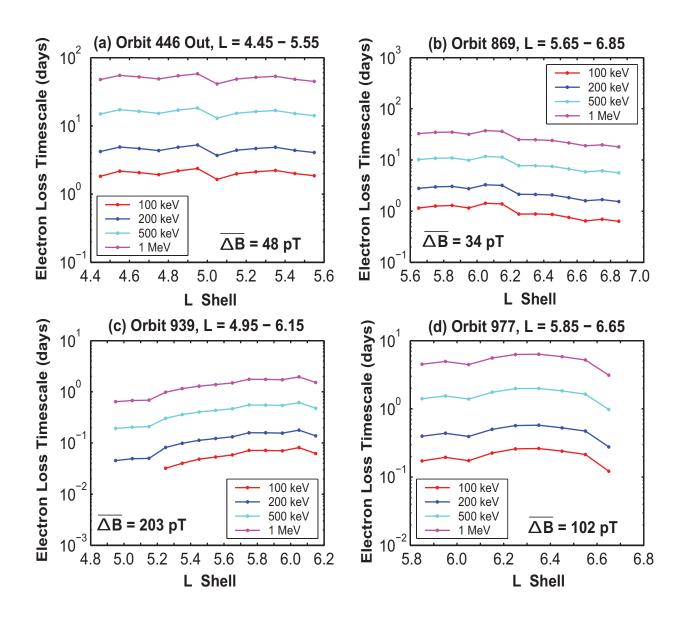


Figure 8. Electron loss timescales due to scattering by hiss at the specified energies, for the indicated CRRES plume crossings, as a function of L-value; 6 % drift-averaging has been applied.

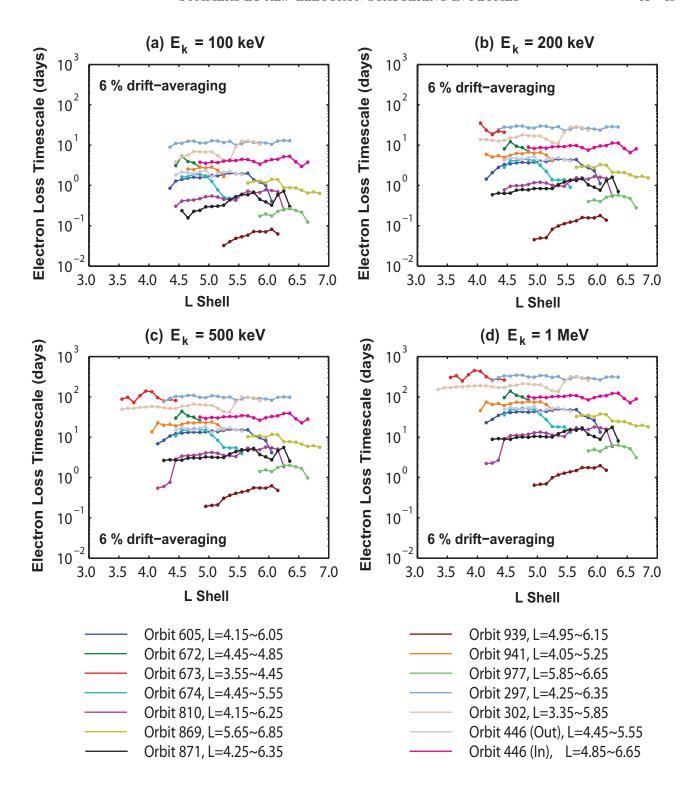


Figure 9. Summary plot of electron loss timescales due to scattering by hiss at the specified energies, for each of the 14 chosen CRRES plume crossings, as a function of L-value.