Magmatism, serpentinization and life: Insights through drilling the Atlantis Massif (IODP Expedition 357)

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1. Introduction

It is now well recognized that slow spreading ridges are formed by interlinked processes of magmatism, asymmetric extension, and detachment faulting that result in the exposure and alteration of lower crustal and mantle-derived rocks in oceanic core complexes (OCCs) (Andreani et al., 2007; Boschi et al., 2006; Cannat, 1993; Früh-Green et al., 2004; Karson et al., 2006; Kelemen et al., 2007; Rouméjon et al, 2015). OCCs contain olivine-rich rocks that interact with seawater to produce serpentinite over a range of temperatures (Andreani et al., 2007; Boschi et al., 2006a; Boschi et al., 2006b; Cannat, 1993; Früh-Green et al., 2004; Karson et al., 2006; Kelemen et al., 2007; Rouméjon et al, 2015). Serpentinization is a fundamental process that controls rheologic and geophysical properties (Escartin et al, 2008; Schroeder et al, 2002) and is associated with the uptake or release of many major and minor components (Alt and Shanks, 2003; Boschi et al., 2008; Delacour et al., 2008; Früh-Green et al., 2004; Schwarzenbach et al., 2012). Serpentinization reactions also lead to highly reduced, alkaline (pH 9-12) fluids with high concentrations of hydrogen, methane and formate, and which have important consequences for long-term global geochemical fluxes and for biogeochemical cycles (Holm and Charlou, 2001; Konn et al., 2009; Lang et al., 2018; Proskurowski et al., 2006, 2008).

The Atlantis Massif (30°N, Mid-Atlantic Ridge) is one of the best-studied OCCs and hosts the off-axis Lost City hydrothermal field (LCHF) on its southern wall (Fig. 1). Serpentinization reactions in the underlying mantle rocks produce high pH (9-11), <91°C fluids that form large carbonate-brucite structures upon venting on the seafloor (Kelley et al., 2001, 2005; Ludwig, et al., 2006). The fluids have high concentrations of H_2, CH_4, C_2+ alkanes and formate ($HCOO^-$) that support novel microbial communities dominated by CH_4-cycling archaea in the hydrothermal carbonate deposits (Brazelton et al., 2006; Brazelton and Baross, 2009; Lang et al., 2010; Méhay et al., 2013; Proskurowski et al., 2006, 2008; Schrenk et al., 2004). Formate and low molecular weight hydrocarbons in the Lost City hydrothermal vents are believed to be formed by abiogenic processes during serpentinization at depth (Lang et al., 2012, 2018; Proskurowski et al., 2008). Thus, the Atlantis Massif provides a natural laboratory to study the links between serpentinization processes and microbial activity in the shallow subsurface of ultramafic and mafic rock sequences that have been uplifted to the seafloor along a major detachment fault zone (Blackman et al., 2002; Cann et al., 1997; Boschi et al., 2006; Karson et al., 2006; Kelley et al., 2001, 2005; Schroeder and John, 2004). The processes controlling fluid flow and a deep biosphere are intimately linked; however, the spatial scale of lithologic variability, the implications for fluid flow paths and geochemical exchange, and the consequences for subsurface ecosystems supported by these systems remain poorly constrained.

Here we present an overview of Expedition 357 of the International Ocean Discovery Program (IODP), which cored seventeen shallow holes at nine sites (Figs. 1, 2) across the Atlantis Massif (Früh-Green et al, 2016). Expedition 357 was implemented by the ECORD Science Operator (ESO) as a Mission Specific Platform (MSP) expedition and consisted of an offshore phase on board the RRS James Cook in fall 2015 and a two-week onshore phase at the IODP Bremen Core Repository in January-February 2016 (Früh-Green et al., 2017a). A major aim of drilling was to investigate seawater infiltration and alteration processes, and their influence on the nature and distribution of microbial communities in lithologically heterogeneous domains of an oceanic core complex. Drilling along a spreading-parallel, east-west profile with seven sites targeted the serpentinite basement at varying distances away from the ridge axis and the Lost City vent field

(Fig. 1, Table 1; see also Früh-Green et al., 2015). Two sites were drilled on the eastern part of the southern wall (Sites M0068 and M0075), three sites in the central section north of Lost City (Sites M0069, M0072, and M0076), and two sites on the western end (Sites M0071 and M0073, with no recovery at M0073). This 8.5 km long profile allows us to explore the extent and activity of the subsurface biosphere in an actively serpentinizing environment and assess how abiotic and biotic processes change with aging of the lithosphere, variations in rock type, and with time of exposure on the seafloor. Two further shallow sites towards the central dome of the massif (Sites M0070 and M0074) targeted the mafic, plutonic domain drilled at IODP Site U1309. Penetration and core recovery were limited at these northern sites, and the recovered sequences were dominated by carbonate sediments and sedimentary breccias. The cores obtained during IODP Expedition 357 are the first continuous sequences of fault rocks recovered along a major detachment fault that has an inferred thickness of ~100 m (e.g., Karson et al., 2006; Schroeder and John, 2004). These cores provide a unique opportunity to study the interaction of magmatism, deformation and fluid-rock interaction during the evolution of the Atlantis Massif and the impact these processes on habitability for microorganisms.

2. Expedition strategy and methods

To obtain a comprehensive view of active serpentinization, fluid circulation and microbial activity, a strategy was developed based on the use of seabed drills that combined coring with water sampling and in situ geochemical measurements during drilling (Früh-Green et al., 2015). To enable continuous operations, two seabed drills were used: the British Geological Survey (BGS) RockDrill2 (RD2) and the Meeresboden-Bohrgerät 70 (MeBo) from the Center for Marine Environmental Sciences (MARUM; University of Bremen, Germany). This was the first time that seabed drill technology was used in the ocean drilling program. Both drills are remotely operated systems that are lowered onto the seabed, with power and control maintained from the ship via an umbilical and using multiple rods and core barrels to progressively penetrate into the seabed (Früh-Green et al., 2017b). They are both based on an HQ-size, diamond coring system, producing between 61 and 62 mm diameter cores, similar in size to the standard IODP core diameter, while cutting a smaller diameter hole. By sitting on the seabed, they do not require heave compensation and consequently have good control on bit weight, analogous to land-based coring, and bottom seawater is used as the drilling fluid.

The expedition included engineering developments that allowed continuous measurement of geochemical parameters during drilling, sampling of bottom water after drilling, and the injection of synthetic contamination tracers during drilling. To evaluate the composition of fluids emanating from the flushed boreholes in real-time, a suite of in situ sensors mounted on the drills measured dissolved oxygen, hydrogen and methane, temperature, pH, and oxidation-reduction potential (ORP) during coring operations. Bottom water was collected prior to drilling using the ship's CTD Niskin bottle rosette and after drilling using Niskin bottles mounted on the drills. Each rock drill was also equipped with a pump system to deliver perfluoromethylcyclohexane (PFC) tracer during drilling to assess seawater contamination of the cores (Orcutt et al., 2017). Shipboard sampling also evaluated contamination potential of the drilling equipment itself, including greases and other lubricants. When recovered to deck, water samples were immediately collected for dissolved H_2 and CH_4 concentration analyses, cell counts and PFC tracer, which were measured onboard, and subsamples were taken for shore-based geochemical and microbiological analyses (see Früh-Green et al., 2017b). Borehole plug systems were also designed to enable future sampling of borehole fluids; these were installed at Holes M0072B and M0075B (Früh-Green et al., 2017b). These will be visited on a US-led research expedition in September 2018 with the ROV Jason (funded by the National Science Foundation) to further investigate the serpentinization and microbiological processes operating in this system.

To accomplish the microbiology related objectives of the expedition and to enable preservation of ephemeral microbiological properties, whole round core (WRC) samples were collected shipboard immediately after core retrieval, curation, and scanning with the multi-sensor core logger. In total, 42 WRC samples were taken from the 17 holes drilled during the offshore phase of the expedition, yielding nearly 8 m in total length and representing ~14% of the entire core recovered. For part of these WRCs, potentially contaminated exterior surfaces were flame-sterilized on the ship in a KOACH open clean system with care to avoid potential contaminants (e.g. dust). Interior pieces of rock were collected after crushing using a flame-sterilized chisel and fixed for microbial cell detection (Früh-Green et al., 2017b). Subsamples of WRCs were used to establish 29 different enrichment experiments on the ship, with initial indications of positive activity in some of the treatments based on elevated cell counts. Remaining portions of the WRCs were immediately frozen at -80°C and shipped to the Kochi Core Center, Japan, at the end of the offshore phase. There, exteriors of the WRCs were removed under sterile conditions with a band saw system equipped in a clean booth (Orcutt et al., 2017) and the WRC interiors and exteriors were subsampled for multiple shore-based analyses.

Taking advantage of weather and operational downtime, IODP Expedition 357 generated a high-resolution multibeam bathymetry map across the Atlantis Massif. The new bathymetry, after processing, provides a grid with a resolution of 20–50 m, which is two to five times higher resolution than previously available bathymetry for this area (100 m) (Blackman et al., 2002). The survey covered the entire striated detachment fault surface of the Atlantis Massif and surrounding terrain, which included the ridge axis to the east, the Atlantis Fracture Zone to the south, the tectonized terrain off-axis and west of the detachment, and its gradual narrowing transition to the adjacent seafloor to the north (Fig. 1).

As with all IODP mission-specific platform expeditions, no cores were split during the offshore phase. A comprehensive onshore phase at the IODP Bremen Core Repository complemented the offshore phase, during which the cores were described in detail and the IODP minimum and some standard measurements were made (see Früh-Green et al., 2017b for details). Bulk rock preparation and geochemical analyses deviated from standard IODP procedure and were conducted in the ultraclean laboratories of the Pacific Centre for Isotopic and Geochemical Research at the University of British Columbia (Canada). Major elements were measured using an inductively coupled plasma-optical emission spectrometer (ICP-OES) and trace elements (including Cr and Ni) were determined with a high-resolution inductively coupled plasma-mass spectrometer (HR-ICP-MS), using international standards and an in-house standard (see Geochemisty section in Früh-Green et al., 2017b).

3. Expedition highlights

In brief, comparison of the recovered rock types (Fig. 2), cross-cutting relationships and the newly produced bathymetric map (Fig. 1) indicate that the central sites recovered in situ sequences that provide a record of (from oldest to youngest): early magmatism emplaced in the shallow mantle, multiphase progressive seawater penetration, serpentinization and metasomatism, dolerite intrusions, and recent basaltic volcanism. As the boreholes were located across the detachment fault zone, the cores likely sampled different structural levels which were tectonically transposed over the thickness of the detachment fault zone (~100 m). The two eastern Sites M0075 and M0068 and the western Site M0071 recovered fault scarp deposits. The scarp deposits and rubble intervals provide no information as to the orientation of structures or veins; however, the rock types, deformation and alteration characteristics, as well as contact relationships preserved in coherent portions of the cores, are similar to those recovered in the in situ sequences at the central sites and provide information about the magmatic, metamorphic and deformation history at shallow levels of the detachment fault zone. Sedimentary breccias with variably rounded clasts (predominantly basalt with some serpentinite, gabbro, and dolerite) in a foraminiferous carbonate matrix were recovered at Site M0070 to the north of the east-west transect, and only one short highly disturbed sediment core was recovered from Site M0074.

3.1 Multibeam bathymetry

New bathymetric mapping revealed a striated detachment surface with variations in geometry along-axis from south to north (Fig. 1a). Near the transform wall, the detachment fault surface summits near the Lost City hydrothermal field at <800 meters below sea level (mbsl); from there it dips ~8° east toward the ridge axis, ~6° to the west, and ~8°–10° to the north toward the central dome. The detachment deepens to the north to >1500 mbsl. This deepening is associated with a change in shape; the detachment shows curvature along the spreading direction with slopes of up to 15° at its termination toward the ridge axis, becomes subhorizontal, and dips up to 10° away from the axis.

The hanging wall cutoff (termination) is well preserved toward the northern, deeper part of Atlantis Massif, corresponding to a sharp contact between the ridgeward-dipping striated fault plane and the adjacent seafloor volcanic terrain. This volcanic terrain has a <10° tilt westward, away from the ridge axis, and hosts volcanic cones and a bathymetric texture typical of volcanic terrain along the rift valley floor (Fig. 1b). This portion of volcanic seafloor is bound further east by a fault scarp and corresponds to the top of a back-tilted tectonic block, as previously identified (Blackman et al., 2002; Cann et al., 1997).

The termination is not preserved to the south, ridgeward of the elevated portion of the detachment. Instead, the striated surface is dissected by a major scarp, with a relief of up to 3000 m from the rift valley floor and the top of the scarp. This scarp shows a lobate structure indicating mass-wasting processes (slumping). Laterally, it links north with the fault scarp bounding the tectonically uplifted volcanic seafloor, which corresponds to the present-day rift valley wall fault. The striated surface is also affected by extensive mass wasting along its southern boundary toward the transform valley (Fig. 1b). The mass wasting produces scarps that are concave at their subvertical headwall. The transform wall south of Atlantis Massif has an average slope of ~20° and numerous channels bound by high ridges channelling debris at the base of the transform valley. Widespread mass wasting at smaller scales is also observed on steeper slopes of the detachment fault surface along the flanks of larger-scale striations. Late small-slip faults (centimeters to tens of meters) cut the striated detachment surface and are subparallel to oblique to the ridge axis. Most have scarps facing away from the ridge axis, with scarps up to ~10–20 m in vertical relief and which have irregular traces.

3.2 Lithology, alteration, and structure

3.2.1. Lithology

More than 57 m of core were recovered, with borehole penetration depths ranging from 1.3 to 16.44 meters below seafloor (mbsf). Core recoveries ranged from 23 to 75% of total penetration, with 100% recovery in some intervals (Table 1; Figure 2). This significant recovery of exhumed mantle peridotite at the surface of a major detachment fault zone is unprecedented in the history of ocean drilling and provides a new window into understanding interlinked processes of crustal accretion, deformation and alteration that to date could not be deduced with conventional rotary drilling, dredging or submersible sampling. Many lithologic and intrusive contacts, deformation features and alteration characteristics are preserved in the cores, even in highly fractured and/or sheared intervals. The cores highlight a highly heterogeneous lateral and vertical distribution of ultramafic and mafic rocks that host a range of alteration styles and extent of deformation (Figs. 2 and 3).

Of the core recovered from the six sites across the southern wall (from west to east: M0071, M0072, M0069, M0076, M0068, M0075; Figs. 1a and 2), serpentinized harzburgite and dunite make up 44% of the core by length. Other major rock types include basaltic rocks and metadolerites (combined 24%) and schistose metasomatic rocks with varying proportions of talc, amphibole and chlorite (11%). Minor lithologies include calcareous sedimentary units (8%), and gabbroic rocks (4%). In comparison, previous dredging and Alvin dive campaigns at the southern Atlantis Massif recovered a similar percentage of ultramafic rocks (45% of total samples collected), but a higher percentage of gabbroic rocks (24%), metasomatic rocks (22%) and sediments (15%), and less basaltic and doleritic rocks (5%) (Schroeder and John, 2004; Boschi et al., 2006a; Karson et al, 2006). The proportion of gabbro to peridotite was less than at IODP Site U1309, where 1408m of gabbroic rocks were recovered; however, the proportion of dolerite was comparable (Blackman et al., 2006; McCaig and Harris, 2012). A synthesis of the lithological contacts, mineralogies and off shore analyses of the key sites are given in Appendixes A through J.

The ultramafic rocks are dominated by harzburgites punctuated by intervals of dunite and minor pyroxenite veins. Gabbroic rocks occur locally as zones of melt impregnation (tens of centimeters thick) and veins at Sites M0068, M0071, M0072 and M0076 (Fig. 2). The harzburgites and dunites exhibit protogranular textures and are extensively serpentinized. Intervals of weakly porphyroclastic serpentinized peridotites were rare and localized. Serpentinized dunites are found exclusively as discrete intervals alternating with harzburgite and likely represent relict mantle melt channels or domains of melt-rock interaction (Kelemen et al., 1992; Nicolas, 1986). Mantle melt-rock reaction textures including vermicular to subhedral spinels (e.g., Nicolas and Prinzhofer, 1983) and pyroxene veins were also observed. The distribution of gabbro rocks is highly heterogeneous downhole and there was an absence of continuous, coherent sections of gabbroic core. Rare magmatic fabrics characterized by diffuse but planar centimeter-scale banding/layering of igneous minerals in gabbroic rocks were recovered in Hole M0068B. Dolerite intrusions, basaltic rocks and local domains of hyaloclastites represent the latest stage of magmatic activity. Metamorphosed dolerite intrusions ranging from a few cm to several meters in thickness were recovered at the central and eastern sites, and some dolerite intervals in Hole M0075B show chilled margins against fault rocks (see Fig. F5 in Früh-Green et al., 2017c). Dolerites and metadolerites as well as poorly vesicular aphanitic to microcrystalline basalts with glassy margins also occur as mm-dm sized components within the sedimentary breccias. The dolerites were variably altered, while basalts were often fresh, with no sign of metamorphism.

All types of variably altered and deformed ultramafic and mafic rocks occur as clasts in sedimentary breccias overlying the basement sequences and as fault scarp deposits. The sedimentary cap rocks include basaltic breccias with foraminiferous carbonate sand and/or lithified foraminiferous carbonate matrix. Fresh and partially palagonitized glass in basaltic components and hyaloclastites were observed in some of the breccias and in some cores containing carbonate sediment.

3.2.2 Serpentinization and metasomatism

The IODP Expedition 357 cores reveal a high degree of alteration and intervals with variable metasomatic overprinting in the ultramafic rocks. Different types of alteration are distinguished depending on protolith: serpentinization, talc-amphibole-chlorite metasomatism and oxidation in the ultramafic rocks; and hydration, oxidation, and local Ca-metasomatism/chloritization along contacts between doleritic and gabbroic domains and serpentinites. The sequence of alteration textures and the associated mineralogical assemblages vary between sites and downhole in some cases (Früh-Green et al., 2017c, d,e and f).

Serpentinization is common in the ultramafic rocks at all sites and occurs as pervasive alteration with extensive to complete replacement of the primary mineralogy, forming mesh textures after olivine, bastites (lizardite ± chrysotile) after orthopyroxene, and different generations of serpentine veins (Fig. 4). A detailed study of the textures and mineralogies of the serpentinized serpentinites combined with in situ major and trace element analyses in primary phases and serpentine minerals is given in Rouméjon et al. (2018). These are used to make a model for the development of alteration heterogeneities at the Atlantis Massif and are summarized briefly here. Hydration of olivine led to a typical serpentine mesh texture, which is characterized by mesh cells, tens to hundreds of microns in size, delimited by microfractures that initially crosscut the olivine. Lizardite mesh rims form the outer part of cells, whereas the mesh cores are made up of poorly crystallized lizardite and/or chrysotile. Magnetite tends to crystallize in the mesh rims and concentrates along microfractures. Progressive fluid infiltration and serpentinization is evident by recrystallization of mesh textures to chrysotile-polygonal serpentine or antigorite, and by multiple sets of veins with variable infillings (Fig. 4; see also Rouméjon et al., 2018). The orthopyroxenes also show overprinting relationships with replacement by serpentine, talc, tremolite and/or chlorite.

Although the depth of penetration was limited, the abundance of gabbroic intrusions appears to increase from west to east and is associated with talc-amphibole-chlorite metasomatism and in some cases chlorite blackwall formation. Multiple generations of amphibole are observed in the gabbroic domains, reflecting progressive alteration from amphibolite to greenschist-facies conditions during exhumation, as described in previous studies (Schroeder and John, 2004; Boschi et al., 2006a). Metasomatism is characterized by varying proportions of talc, chlorite and tremolitic amphibole and is a common feature at the central and eastern sites, evident as pale greenish-white domains or vein networks (Figs. 3b and d). Talc-metasomatism is rare at the western Site M0071 and is most prevalent in Hole M0072B near Lost City (Fig. 2). It develops both as pervasive, irregular patches in the serpentinites or as localized alteration associated with mafic intervals, enclosing serpentinized harzburgite on centimeter to decimeter scales (Figs. 5 and 6). The metasomatic domains are locally deformed and the talc-rich zones are commonly sheared, forming intervals of talc-amphibole-chlorite schists. Talc generally replaces mesh textures or forms in veins in the serpentinites, whereas tremolitic amphibole and chlorite assemblages are found in mafic domains and at contacts between serpentinite and gabbro or metagabbro, or in domains that have been infiltrated by mafic melts (Figs. 5 and 6). In some sections, amphibole and chlorite appear to have formed prior to talc. Relict olivine is also found in a number of cores in the central and eastern sites (Holes M0068, M0072, M0076) where talc-amphibole-chlorite metasomatism and/or chlorite-rich alteration is most prevalent.

Metasomatism is particularly pronounced in Hole M0072B, where complex zoned intervals (approximately 5 cm thick) with mafic lenses adjacent to chlorite-rich blackwalls are repeatedly present over a few meters. Exceptional relationships between mafic intrusions (possibly doleritic or microgabbro), talc ± amphibole ± chlorite zones, and serpentinized dunite and harzburgite are observed in Cores M0072B-6R1 (Fig. 5d) and M0072B-7R1 (Fig. 6). The mafic intrusions in these cores have pale brown to pinkish-brown central domains that are surrounded by external dark green domains made up mostly of chlorite (chlorite blackwall), which in turn grade into talc-amphibole-rich domains at the contact to the serpentinites. The pinkish brown domains were originally described as rodingites (Früh-Green et al., 2017d), which have been found during previous sampling campaigns along the southern wall of the Atlantis Massif (Boschi et al., 2006a). However, subsequent analyses have yet to identify typical Ca-Al silicates, such as epidote (clinozoisite), diopside, prehnite, (hydro)garnet or vesuvianite, which are commonly found in rodingites in similar associations with serpentinites. Instead, preliminary X-ray diffraction (XRD), micro-Raman spectroscopy and microprobe analyses (unpublished data) indicate that these zones are indeed Ca-rich but are made up of chlorite and fine-grained aggregates of anorthite ± tremolitic and/or pargasitic amphibole. The association of chlorite and anorthite in these domains could result from higher temperatures of alteration than are typically associated with rodingite (> ~350°C). Anorthite may also form from fluids with higher CO_2 concentrations (Rice, 1983).

Although brucite occurs in the actively venting carbonate structures at Lost City (Kelley et al., 2001; Früh-Green et al., 2003; Ludwig et al., 2006) and is a common product of serpentinization reactions, it has not been found in previous studies of the basement rocks of the Atlantis Massif (e.g., Boschi et al., 2006a; 2008). In the IODP Expedition 357 cores, brucite could not be detected visually, microscopically or with XRD on bulk rock samples (Früh-Green et al., 2017c,d,e). In addition, a brucite signature is absent in micro-Raman spectra, which together with nearly stoichiometric serpentine compositions of the mesh texture serpentine minerals (Rouméjon et al., 2018; Rouméjon et al., this issue) strongly suggest that brucite is absent in the serpentinized peridotites that make up the southern wall of the Atlantis Massif. The absence (or dissolution) of brucite and abundance of talc in the metasomatic assemblages may be a consequence of high Si activities in the fluids during progressive hydrothermal alteration along the detachment fault zone and/or high alteration temperatures (above 350°C) during denudation of the mantle.

Finally, later-stage oxidation of the serpentinized harzburgites and dunites is characterized by reddish to brown alteration, occurring as both pervasive and localized features, and is commonly associated with calcium carbonate veins (Fig. 3a). Overprinting relationships in the ultramafic rocks demonstrate an overall progression from local amphibole-chlorite alteration to serpentinization to talc ±amphibole ±chlorite metasomatism and later oxidation.

Hydration of the dolerites and basalts manifests as pervasive background alteration with moderate to high intensity accompanied by alteration halos that flank veins. Secondary minerals vary depending on the temperature of alteration, with dolerites dominated by greenschist-facies minerals (chlorite, amphibole, and epidote), and basalts by low-temperature oxidation to iron oxyhydroxides and clays. Epidote occurs as a dominant vein mineral in metadolerites in Hole M0069A often with vein halos dominated by chlorite. Chilled margins in dolerite dikes that have intruded into talc-amphibole-chlorite schists are observed at the most eastern Site M0075. Hydration of gabbros is generally associated with chlorite-amphibole assemblages.

Hydrothermal veins are present in all rock types. Vein minerals include serpentine, talc, chlorite, amphibole, epidote, quartz, and calcium carbonate. The veins are often complex, with multiple infillings and internal textures, highlighting a protracted formation history. Crosscutting relationships are also complex, with the same veins observed both crosscutting and being crosscut by a second vein type. The occurrence of calcium carbonate veins was surprisingly limited in the recovered cores. Carbonate veins are more prevalent in the sites around the Lost City hydrothermal field, where they occur mostly within entirely serpentinized dunites and harzburgites. At the western Site M0071, calcium carbonate veins in the serpentinites predate fractures that are infilled with foraminiferous carbonate sand (Fig 3a), suggesting open fractures at the top of the detachment fault zone, as described by Schroeder et al. (2002) based on Alvin dive samples.

3.2.3 Structures and deformation history

The drilled sites are located along a roughly spreading-parallel, 8.5 km transect (west–east) in various positions (trough or wall/flank) relative to individual corrugations of the detachment fault over the southern wall of Atlantis Massif (Fig. 1). Despite the fact that a number of the holes recovered rocks that are considered not to be in situ, generalizations can be made about the structural history recorded (Früh-Green et al., 2017c,d,e,f). As in IODP Hole U1309D at the central dome of the Atlantis Massif (Blackman et al., 2006), strongly deformed microstructures formed at high temperatures are rare in the IODP Expedition 357 cores. The majority of the recovered cores show amphibolite- to greenschist-facies, semibrittle and brittle deformation (Figs. 3 and 6), which contrasts with previous studies of samples recovered by submersible and by dredging that document higher temperature, high strain conditions in parts of the southern wall of the massif (Boschi et al., 2006a; Karson et al., 2006; Schroeder and John, 2004). Fault rocks in shear zones preserved in the cores are dominated by anastomosing intervals of variable intensity, with schistose amphibole ± talc ± chlorite zones up to tens of centimeters thick. The schistose shear zones contain undeformed dolerite intrusions with preserved chilled margins; elsewhere, dolerite sheets record brittle and semibrittle deformation textures indicating repeated magmatism and faulting. Extensive intervals of flattened breccia are associated with dolerites but often contain clasts of fault rocks derived from other lithologies. Some breccia clasts show relicts of higher temperature amphibolite facies deformation, as do serpentinized intervals in the margins of talc-tremolite-chlorite schist zones. Intense cataclastic intervals and possible fault gouge occur within some breccias and also as thin intervals within the schistose shear zones. Discrete fault planes occur in most cores with a range of orientations, but lineations are generally shallow on both steep and shallow fault planes. An important observation is that the serpentinites are almost invariably statically altered, with no schistose serpentine developed and only occasional cataclastic seams. Strain within serpentinite intervals seems to be almost entirely localized within metasomatic talc-tremolite-chlorite horizons.

3.3 Bulk rock geochemistry

A wide range of major and trace element bulk rock compositions reflect the differences in rock type as well as the type and extent of alteration (Table 2, Figs. 7, 8 and 9). Independent of site location, the talc-amphibole-chlorite schists typically have high SiO_2 contents, ranging from 50-60 wt%, and low MgO/SiO_2 ratios (0.45 – 0.51) as well as lower loss on ignition (LOI: 4.3 – 5.3 wt%) than the serpentinites (LOI: 11.95 – 13.8 wt%). The serpentinized ultramafic rocks have the highest MgO/SiO_2 ratios (0.96 – 1.19) and variable but high Cr (up to 29,698 ppm in Hole M0069) and Ni (up to 14,590 ppm in Hole M0071A) contents. Overall, the talc-amphibole-chlorite schists (and in some cases the impregnated/metasomatized ultramafic rocks) are richer in Al_2O_3, Na_2O, CaO, TiO_2, and depleted in Fe_2O_3 (Fig. 7). The talc schists are also enriched in Cr and Ni relative to the gabbroic rocks and dolerites but have lower concentrations than the ultramafic lithologies (Fig. 8). Samples from Hole M0068B exhibit the highest SiO_2, CaO, and Na_2O contents, but the lowest Al_2O_3 and Fe_2O_3 contents. The most altered dolerites and gabbros have characteristically low SiO_2 concentrations (26.2 – 31.6 wt%), high Fe_2O_3 (18.8 –32.1 wt%) and low Ni and Cr (Fig. 8), which reflects the high modal abundance of chlorite in these rocks and suggests Si mobility and loss during alteration (Fig. 9). The Mg and Ni concentrations of the IODP Expedition 357 serpentinites and impregnated serpentinites are higher than those recovered during IODP Expeditions 304-305 and likely reflects the more primitive nature of the mantle peridotites recovered along the southern wall. The gabbroic compositions are similar to the IODP Hole U1309D gabbros, but the dolerites and metadolerites have higher Ni concentrations and may be the result of a higher primary modal abundance of olivine (Fig. 9).

The Rare Earth Element (REE) patterns group by lithology and show a weakly defined enrichment from west to east (Fig. 10) and some variations downhole. The serpentinized ultramafic rocks have relatively flat to slightly light REE (LREE) depleted chondrite-normalized patterns (i.e., typically centered around 1 or below). The impregnated/metasomatized serpentinites from Hole M0072B exhibit values slightly higher than 1. Dolerites and gabbros exhibit moderate LREE depletions with values ranging between 1 and 10. Two of the talc-amphibole-chlorite schists have REE patterns resembling the impregnated/metasomatized samples. Positive and negative europium anomalies were observed but do not correlate with a particular lithology or site. Along with correlated Mg# and Ni abundances (Fig. 8), geochemical trends in the serpentinized ultramafic rocks include a common uranium positive anomaly (the intensity of which decreases in impregnated / metasomatized samples) (Früh-Green et al., 2017c,d,e) and enriched lithium, cerium, and strontium anomalies in the central sites (Table 2). Such anomalies are commonly related to alteration processes, either from hydrothermal alteration or from late interaction with seawater on the seafloor. Rouméjon et al. (2018) document regional trends in trace and REE element compositions in serpentine minerals compared to primary olivine and attribute the regional and downhole variations to mobilization of elements during the successive stages of exhumation as a result of early melt emplacement, serpentinization-related fluid-rock interaction, and later fluid-rock interaction. LREE enrichments due to the proximity with metagabbros or metadolerites are particularly observed in samples from Holes M0068B and M0072B (see also Boschi et al, 2006a) and contribute to the downhole variations.

3.4 Volatile concentrations

Elevated bottom water gas concentrations recorded by the sensor package and water sampling confirmed that serpentinization is on-going at the Atlantis Massif (Figs. 11 and 12). Water samples before and after drilling indicated "hot spots" of dissolved hydrogen over Sites M0068, M0072, M0069, M0070 and M0071, with the highest concentrations of 323 nM measured in Hole M0072B. Elevated concentrations of methane were found over Sites M0072, M0070, and M0071 (Fig. 11, Table 1; see also Table T12 in Früh-Green et al., 2017c). A CTD cast directly over the Lost City hydrothermal vents (Site M0072) just south of the central drill sites had significantly elevated methane and hydrogen (35–48 nM and 196–267 nM, respectively). On a regional scale, hydrogen concentrations tended to be highest in the central sites and at the eastern Site M0068, which may reflect active serpentinization in the vicinity of the Lost City hydrothermal field (Fig. 11; Table 1). However, the interpretation of the regional-scale influence on methane and hydrogen fluxes out of the basement is ambiguous since the depth of penetration into the basement was limited to <20 mbsf.

In addition to elevated dissolved gas concentrations measured in the fluids, gas bubbles were observed issuing from the hole and around the drill base during operations at Site M0070, even when coring had stopped (Fig. 13). The bubbles could not be sampled directly with the seabed drills and thus their composition remains unknown. Bathymetry indicates that Site M0070 lies west of the western limit of the preserved striated detachment surface of Atlantis Massif (Fig. 1) at the foot of a ~30 m high irregular mound (Figure F2A in Früh-Green et al., 2017f). The three holes penetrated the same structural unit composed of either loose or cemented basalt clasts with vesicles and glass within a carbonate matrix. The mound is likely a volcanic cone that has undergone faulting and/or mass wasting and, thus, we cannot exclude volcanic gases as a source of the bubbles observed at this site.

In addition to the water sampling observations, the drill-mounted sensors recorded peaks in methane and pH that correlated with sharp decreases in oxidation-reduction potential (ORP) at many sites (Figure 12, Früh-Green et al., 2017c,d,e,f). Low ORP (or Eh) reflects reducing conditions and can be interpreted as elevated hydrogen concentrations and/or other reduced components (such as reduced iron and hydrogen sulphide) in the fluid. The ORP sensor does not respond to methane. In some cases, excursions in the sensor signals were observed while drilling, which suggests that horizons that were penetrated released reduced basement fluids and volatiles into the drilling fluid. In other cases, we observed variations in the methane, pH and ORP signals even when no drilling operations were underway or when the drills touched down on the seabed, suggesting that diffuse reduced fluids may be present at the top of the massif. In many cases we observed strong negative spikes in the ORP signals without a corresponding methane signal, which points to hydrogen and/ or other reduced phases being released into the drilling fluids. Due to limited core recovery, we were not able to clearly correlate the excursions in sensor data with specific horizons or rock types. On a regional scale, negative spikes in ORP were observed in most of the holes in the central sites, which is consistent with the higher dissolved H_2 and CH_4 concentrations at these sites and may reflect hydrothermal circulation related to the Lost City hydrothermal field.

It is worth noting that the dissolved methane concentrations were monitored with a Franatech METS sensor. Post-cruise evaluation of this sensor revealed that it responds to both CH_4 and H_2 with a response factor of 1 to 0.02, respectively. This complicates interpretations of the output of this sensor because H_2 concentrations typically exceed those of CH_4 in this environment. For example, in Lost City hydrothermal fluids, the H_2/CH_4 ratio varies from 0.5 to 9.2 (Proskurowski et al, 2008). Where we measured bottom water concentrations from CTD casts, CH_4 was often below our detection limit (0.7 nM); however, at some sites both H_2 and CH_4 were present and the H_2/CH_4 ratio ranged from 5.5 to 20.9. In samples taken in the Lost City plume, the average ratio was 5.3. Samples from the drill-mounted Niskin bottles yielded H_2/CH_4 ratios ranging from 1.2 at Site M0070A to 167 at Site M0068B (see Table T12 in Früh-Green et al., 2017c). Although we were unable to make quantitative estimates of volatile concentrations from the sensor data, the METS sensor likely recorded both H_2 and CH_4, and it is possible that the output values we observe represent H_2 concentrations that are a factor of 50 times higher than the actual recorded values given as CH_4 concentrations. Horizons with high H_2 concentrations are also indicated by the fact that elevated CH_4 signals often correlated with strong decreases in ORP.

3.5 Microbiology sampling

To accomplish the microbiology-related objectives of the expedition, an extensive program was carried out on board the ship to collect whole-round core samples immediately after core retrieval, curation, and scanning with the multisensor core logger to enable preservation of ephemeral microbiological properties. This program included (1) frozen preservation of core material for DNA- and lipid-based analyses in shore-based laboratories, (2) establishment of enrichment incubations on the ship (at ambient or in situ pressure) to assess the potential for various microbial metabolisms, (3) collection of samples to evaluate the performance of the contaminant tracer delivery, (4) preservation of samples for biomass determination via cell counting, and (5) collection of parallel samples for spatial and isotopic geochemical determination, particularly focused on carbon and minerals.

A major technical development for this expedition to enable microbiological analysis was establishing the delivery system for adding a synthetic tracer (PFC) into the drilling fluids to monitor the possibility of drilling-induced contamination (Orcutt et al., 2017). Samples of core barrel liner fluids, sensor package Niskin bottles, and exterior and interior pieces of whole-round core were collected to quantify the concentration of PFC tracer added during drilling operations and track its potential distribution into samples. After overcoming some technical difficulties with the metering pump in the delivery system, we established that PFC was delivered at saturating (>1 mg/L) concentrations into the drilling fluids (Orcutt et al., 2017). Moreover, appropriate handling conditions combined with coherent core samples resulted in the absence of tracer from the interior of core samples (whereas less coherent materials suffered potential contamination from intrusion of tracer). Overall, implementation of the tracer injection system for seabed drill systems proved to work, and PFC concentrations on the exterior and interior of core samples could be used as a measure to assess the quality of the sample material for detailed microbiological and geochemical analyses (Orcutt et al., 2017).

To obtain an initial assessment of microbial biomass in the core samples, cell abundance was determined on the ship and onshore at the Kochi Core Center (Japan) in an ultraclean laboratory. Direct counting was made with an epifluorescence microscope following cell separation from flame sterilized interior portions of subsamples. To enable low levels of cell detection, great care was taken onshore and offshore to minimize contamination of samples (Früh-Green et al., 2017b; Morono et al., 2017), resulting in a limit of detection of 9.8 cells cm^{-3}. Cell abundance in the core samples was variable and relatively low, ranging from tens to thousands of cells/cm^3, with many of the basement samples often below the minimum quantification limit of 9.8 cells cm^{-3} (Fig. 14). Cell counts in the interior portions of the basement rocks ranged from <10 to 6.5 x 10^2 cells cm^{-3}, with one sample from Hole M0071A yielding 4.1 x 10^3 cells cm^{-3}. Excluding the short core obtained at Site M0074 (because of contamination issues with core handling), the highest cell counts were found in the sediments in Hole M0069A near the contact to the basement, reaching up to 1.6 x 10^4 cells cm^{-3} at 5.46 mbsf, and decreased rapidly to <10^2 cells cm^{-3} in the underlying basement rocks. The deepest samples were from this hole (at 14.6 mbsf), where 10-24 cells cm^{-3} were measured in the serpentinites. A similar trend was observed at Hole

M0072B, with up to 5 x10^2 cells cm^{-3} within the top meter of the hole and decreasing to <20 cells cm^{-3} below 6.5 mbsf (Fig. 14).

The cell densities in the IODP Expedition 357 drill cores are distinctly lower than in the actively venting Lost City carbonate towers (10^7 to 10^8 per gram of wet weight; Kelley et al., 2005). They are also low in comparison to cell densities in fluids sampled in actively serpentinizing environments on land, which are typically less than 10^5 cells ml^{-1}, and as low as 10^2 cells mL^{-1}, although continental sites of serpentinization represent different niches within the subsurface ecosystem (e.g., Schrenk et al., 2013; Brazelton et al., 2017). These cell densities are also lower than in mafic subseafloor cores, which have been estimated at ~10^4 cells per gram of rock (Jørgensen and Zhao, 2016). Overall, the strict sampling handling protocols allowed for very low limits of microbial cell detection, and our results show that the Atlantis Massif subsurface contains a relatively low density of microbial life compared to other subseafloor crustal and serpentinizing systems. This low density suggests that something may be limiting life in this subsurface habitat compared to the other habitats, such as energy availability, high pH, or low carbon dioxide availability, but further analyses are required to determine this.

4. Implications for understanding oceanic core complex processes

Expedition 357 was the first IODP expedition to successfully use seabed drills to acquire intact shallow mantle sequences at the top of the footwall of an oceanic detachment fault zone and to monitor borehole fluids while drilling. This expedition provides insights into magmatic, tectonic and alteration processes of an oceanic core complex that is actively undergoing serpentinization and has the potential to sustain a unique subsurface biosphere. The cores have exceptionally well-preserved contacts and show strong lateral and vertical variations (from cm to m scale) in rock type and alteration assemblages that are a consequence of multiple phases of magmatism, fluid-rock interaction and mass transfer along the detachment fault zone. The results of this expedition are expected to address fundamental questions that were part of the motivation for the expedition (Früh-Green et al., 2015), such as: How are seafloor spreading and mantle melting linked to ocean crustal architecture? How do oceanic detachment faults develop and facilitate hydrothermal circulation? How do they affect the development of alteration patterns and the evolution of the deep biosphere in these environments?

IODP Expedition 357 sampled only the very shallowest level of the detachment fault zone and overlying talus blocks at the top of the massif. However, this is the first time that clear relationships of gabbro and dolerite hosted by mantle peridotite along the southern wall of Atlantis Massif have been documented. These relationships imply that melts are generated beneath volcanic-poor ridge segments at ridge-transform intersections, but much of the melt may be trapped in the mantle as it turns into lithosphere beneath the ridge axis, rather than migrating upward to form a continuous magmatic crust. Based on high-resolution ion microprobe (i.e., SHRIMP) U-Pb zircon ages from IODP Hole 1309D and broadly spaced samples collected along the southern ridge of Atlantis Massif, Grimes et al. (2008) document a protracted history of accretion in the footwall. They calculate a detachment fault slip rate of 28.7 ± 6.7 mm/a, which implies significant asymmetric plate spreading (up to 100% on the North American plate) for at least 200 ka during core complex formation. Our results are consistent with previous studies that indicate that ongoing magmatic activity associated with asymmetric plate spreading results in a heterogeneous mafic and ultramafic lithosphere with late dolerite intrusions exposed in the denuded footwall, whereas accretion of volcanic seafloor persists in the hanging wall (Cannat et al., 2006; Grimes et al., 2008; Ildefonse et al., 2007; John and Cheadle, 2010; Karson et al., 2006; McCaig and Harris, 2012; Smith et al., 2006).

The volume of gabbros in the southern wall of the Atlantis Massif and their mode of intrusion as thin lenses are distinct from the thick gabbroic sequence recovered at IODP Site U1309 (IODP Expeditions 304 and 305) at the central dome (Blackman et al., 2006; Ildefonse et al., 2007; McCaig et al., 2010; McCaig and Harris, 2012). Although a direct comparison of the two drilling campaigns is difficult to make because of depth of penetration, and the possible tectonic control on emplacement of rock sections, both campaigns yield important information about accretion and alteration processes as well as regional heterogeneities associated with the architecture and evolution of OCCs. The surface of the central dome was cored at IODP Hole U1309B, where dike rocks and basalts were recovered, and a few pebbles of talc schist together with highly altered basalt and dolerite were recovered in IODP Hole U1309H (Blackman et al., 2006; John et al 2009). In addition, Alvin sampling during cruise AT3-60 in 2000 (MARVEL expedition; Blackman et al., 2002) recovered one talc schist sample (sample 3642-1309; see Boschi et al., 2008) along dive tracks in the vicinity of IODP Site U1309. Metasomatic talc-amphibole-chlorite rocks are considered key components of detachment fault zones (e.g., Escartin et al., 2003; Boschi et al., 2006a,b; McCaig et al., 2010) and pre-date dolerite diking events and basaltic eruptions (Karson et al. 2006; McCaig and Harris, 2012). Although not abundant, the occurrence of talc schists in the central dome of the Atlantis Massif hints at the presence of a thin detachment fault zone in this area. However, on a regional scale, the newly acquired multibeam data (Fig. 1) clearly allow the corrugated surface related to the detachment fault zone to be distinguished.

The mineralogical assemblages, alteration textures, and bulk rock chemistries recorded in the IODP Expedition 357 drill cores indicate progressive seawater infiltration along the detachment fault and into the footwall, pointing to an important role of the mafic intrusions in controlling fluid chemistry and metasomatism. Early high temperature, amphibolite-facies alteration and ductile deformation features have been reported from studies of dredged and submersible sampling of the southern wall (Boschi et al., 2006a; Karson et al., 2006; Schroeder and John, 2004), but such features are less common in the IODP Expedition 357 drill cores. In contrast, alteration in the shallow IODP Expedition drill cores is dominated by serpentinization processes, brittle deformation and mass transfer between mafic and ultramafic lithologies under greenschist-facies conditions.

The occurrence of gabbroic intrusions is associated with talc-amphibole-chlorite metasomatism and local blackwall formation and appears to increase from west to east. Metasomatism and talc precipitation are most prevalent at contacts between mafic and ultramafic domains (Figs. 5 and 6.). A systematic overprinting of serpentinite by talc- and chlorite-rich assemblages is associated with the occurrence of variably thick (micro)gabbroic lenses and points to silica mobility and channelled fluid flow at varying depths within the detachment fault zone (see also Boschi et al., 2006, 2008). The geochemical influence of the gabbroic intrusions and progressive fluid-rock interaction is also evident from REE enrichments measured in serpentine minerals and tends to increase from west to east (Rouméjon et al., 2018). The general trend to slightly larger volumes of gabbroic intrusions from west to east (assuming the position of the drill holes roughly reflect differing original depths in the lithosphere) suggests that magmatic activity may have been greater at depth within the detachment fault zone before emplacement to their current locations.

The textural sequences and mineralogical assemblages in the ultramafic rocks reveal a transition between an initial pervasive phase of hydration along grain boundaries to produce mesh-textures in the serpentinites, with subsequent serpentinization and metasomatism focused along localized fluid pathways (Rouméjon et al, 2018). Alteration commences as the peridotites and gabbros are subjected to active hydrothermal circulation, but alteration of the dominant phase, olivine, to produce serpentine minerals will be limited to temperatures below approximately 500°C (Chernosky, 1973). Serpentinization of olivine becomes more effective below 350-400°C (Evans, 2004) and reaches maximum rates between 250°-300°C (Andreani et al., 2007; Martin and Fyfe, 1970; Malvoisin et al., 2012; McCollom 2016). Hydration is intense directly along the detachment fault zone, where permeability is expected to be highest (McCaig et al., 2007; McCaig et al., 2010), and progresses inside the footwall. When the fluids reach temperatures below ~350°C, efficient serpentinization commences and is recorded by the development of mesh texture at all sites. Based on zircon analyses and multicomponent magnetic remanence data in the central dome, Schoolmeesters et al. (2012) proposed a model for the thermal structure of the Atlantis Massif in which the 350°C isotherm corresponds to a depth of approximately 5 km below the surface. Thus, initiation of serpentinization would have occurred at significant depths and early in the exhumation history of the massif. The infiltration of seawater-derived hydrothermal fluids is facilitated by the closely-spaced microfracture networks that crosscut the olivine and result from combined thermal and tectonic stresses, enhanced by reaction-induced permeability at the onset of serpentinization (Rouméjon and Cannat, 2014; Rouméjon et al., 2018). As the footwall reaches shallower crustal levels, fluid flow will likely be dominated by more continuous fracture planes that can channel hydrothermal fluids through the peridotite and form veins (Andreani et al., 2007; Rouméjon et al., 2018). The transition from more pervasive grain-boundary flow to localized or channeled flow is indicated by recrystallization of the mesh texture to chrysotile-dominated serpentine and by banded veins (Rouméjon et al., 2018; Rouméjon et al., this issue).

Talc formation postdates an early phase of serpentinization, and in some cases amphibole formation (see also Boschi et al., 2006a), but predates late-stage intrusions and alteration of some dolerite dikes and the extrusion of basalt, indicating that basaltic magmatism continued as the variably altered basement sequences where emplaced on the seafloor. Alternating metasomatic and serpentinized domains as well as irregular cross-cutting vein relationships in the IODP Expedition 357 cores from the central (M0072 and M0076) and eastern sites (M0068) emphasize the dynamic nature of the system with similar composition of veins forming at multiple times. Textural relationships and the lateral and vertical distribution of metasomatic assemblages indicate that Si ± Ca ± Al mass transfer occurred locally at peridotite/gabbro or peridotite/dolorite contacts as well as through infiltration and interaction with Si-rich fluids along fractures to form talc-rich assemblages (see also Boschi et al., 2006a; McCaig et al., 2010; Rouméjon et al., 2018). In addition, the volume of carbonate veins was surprisingly low in the recovered cores, even in the sites directly above the Lost City hydrothermal field. This suggests that present-day fluid flow and hydrothermal activity at Lost City is localized by late normal faults that cut the southern wall (Denny et al., 2015).

The presence of the mafic lenses within the serpentinites – and their alteration products to mechanically weak minerals, such as talc, serpentine and chlorite – may also be critical to the development of the detachment fault zone and may enhance unroofing of upper mantle peridotites and lower crustal gabbroic rocks during seafloor spreading (Escartin et al., 2003; Schroeder and John, 2004; Boschi et al., 2006b). Talc in particular may be influential in lubricating and softening mylonitic shear zones and can lead to strain localization and focused hydrothermal circulation along such faults (see also McCaig et al, 2010). In fact, low-T detachment strain (< ~300°C) may actually be concentrated with time in the weak, talc-serpentine-rich rocks, creating a runaway system and allowing movement on the detachment fault zone to remain active while leaving a large portion of the exposed lithosphere undeformed. In addition, based on detailed studies of greenschist- to amphibolite-facies assemblages in metadolerites in the upper 130m of the IODP Site U1309D drill cores, McCaig and Harris (2012) argue that the detachment fault zone itself acts as a conductive boundary layer between gabbroic intrusions in the footwall and active hydrothermal circulation within the fault zone. They conclude that widespread occurrences of gabbro at high levels in the crust below detachment faults may be an expression of the same fundamental balance between magmatism and hydrothermal circulation that produces a layered structure at fast-spreading ridges.

Although alteration in the IODP drill cores is dominated by earlier phases of serpentinization and metasomatism associated with detachment faulting and denudation of mantle peridotites, wide-scale, active serpentinization at Atlantis Massif is indicated by elevated concentrations of H_2 and CH_4 in bottom water sampled before and after drilling. Even at the transform fault, H_2

concentrations in CTD casts were elevated (6.2 nM) relative to background seawater (<0.3 nM). Monitoring of the borehole fluids during drilling operations recorded numerous excursions in methane, temperature and ORP that often correlated with each other. The fact that the excursions occurred both while drilling as well as when no coring operations were taking place implies that horizons of reduced, and likely hydrogen-rich, fluids must exist in the basement rocks and that volatiles are being continuously expelled during active serpentinization at the Atlantis Massif. Active volatile expulsion was also indicated as bubbles emitting from Site M0070. The diffuse fluid flow indicated by the sensor package data and water sampling during IODP Expedition 357 contrasts strongly with the focused flow associated with the actively venting Lost City hydrothermal field. The detachment fault zone seems to play a passive role in channelling the basement fluids. Instead, present-day hydrothermal fluid flow is likely controlled by late-stage normal faults cutting the southern wall (Fig. 1; see also Denny et al., 2015). In addition, the present-day hydrothermal fluids, characterized by high pH, low Si, and low metal concentrations are controlled by serpentinization reactions and are chemically distinct from the higher temperature fluids that were involved with mass transfer and metasomatism at deeper levels of the detachment fault zone and at earlier stages in the evolution of the Atlantis Massif.

A major achievement of IODP Expedition 357 was to obtain microbiological samples along the west-east lithospheric age profile, which will provide a better understanding of how microbial communities evolve as ultramafic rocks are emplaced on the seafloor. Our results indicate that the subsurface of the serpentinite basement of Atlantis Massif has relatively low biomass. We anticipate that on-going post-cruise microbiological studies will provide important constraints to address basic questions, such as what is the nature of microbial communities hosted by serpentinizing rocks, and to what depth is microbial activity sustained? How do these vary with aging of the lithosphere? How do they differ from or interact with communities in sediments and mafic substrates in the same age crust? Because of the significant difference in volatile compositions and limited CO_2 stability at high pH, one can expect that biotopes hosted in serpentinizing environments will differ significantly from axial, basaltic-hosted vent systems in which CO_2 is a dominant volatile species. In addition, the mixing of oxidized seawater with highly reduced fluids leads to complex gradients in fluid chemistry and possibly temperature that may influence microbial distribution and activity. Substantially different habitats harboring various types of aerobic and anaerobic metabolisms may thus occur over a narrow spatial scale in these types of environments.

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Supplementary Material

Appendix A. Synthesis of the lithological contacts, mineralogies and off shore analyses of the western Hole M0071A (IODP Expedition 357).

Appendix B. Synthesis of the lithological contacts, mineralogies and off shore analyses of the western Hole M0071B (IODP Expedition 357).

Appendix C. Synthesis of the lithological contacts, mineralogies and off shore analyses of the western Hole M0071C (IODP Expedition 357).

Appendix D. Synthesis of the lithological contacts, mineralogies and off shore analyses of the central Hole M0069A (IODP Expedition 357).

Appendix E. Synthesis of the lithological contacts, mineralogies and off shore analyses of the central Hole M0072B, IODP Expedition 357 (Part 1).

Appendix F. Synthesis of the lithological contacts, mineralogies and off shore analyses of the central Hole M0072B, IODP Expedition 357 (Part 2).

Appendix G. Synthesis of the lithological contacts, mineralogies and off shore analyses of the central Hole M0076B, IODP Expedition 357 (Part 1).

Appendix H. Synthesis of the lithological contacts, mineralogies and off shore analyses of the central Hole M0076B, IODP Expedition 357 (Part 2).

Appendix I. Synthesis of the lithological contacts, mineralogies and off shore analyses of eastern Hole M0068B (IODP Expedition 357).

Appendix J. Synthesis of the lithological contacts, mineralogies and off shore analyses of eastern Hole M0075B (IODP Expedition 357).

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Table 1. IODP Expedition 357 site locations, core recovery, and maximum volatile concentrations.

Hole	Latitude	Longitude	Water depth (m)	Drill	Number of cores	Interval cored (m)	Core recovered (m)	Core recovery (%)	Interval open-holed (m)	Penetration depth (mbsf)	Maximum hydrogen (nM)*	Maximum methane (nM)*
Eastern Sites												
M0068A	30°7.49'N	42°5.74'W	1102.7	RD2	1	1.97	0.47	23.9	0	1.97	34	BDL
M0068B	30°7.51'N	42°5.75'W	1102	RD2	9	9.6	6.34	66.04	0	9.6	137	BDL
M0075A	30°7.67'N	42°3.98'W	1568	RD2	1	1.72	0.65	37.79	0	1.72	3	BDL
M0075B	30°7.65'N	42°3.97'W	1568	RD2	3	5.7	2.73	47.88	0	5.7		BDL
Central Sites												
M0069A	30°7.94'N	42°7.20'W	850.9	RD2	10	16.44	12.29	75	0	16.44	58	4
M0072A	30°7.79'N	42°7.32'W	820.3	RD2	2	2.23	0.87	39.1	0	2.23	12	2
M0072B	30°7.79'N	42°7.32'W	820.3	RD2	8	11.61	6.49	52.3	0.825	12.43	323	2
M0076A	30°7.62'N	42°7.08'W	768	RD2	1	1.72	0.4	23.26	0	1.72	–	–
M0076B	30°7.62'N	42°7.07'W	768	RD2	10	16.31	11.71	71.8	0	16.31	12	3
Western Sites												
M0071A	30°7.71'N	42°9.20'W	1390.8	MeBo	2	5.22	2.85	54.6	0	5.22	61	BDL
M0071B	30°7.72'N	42°9.19'W	1380	RD2	3	4.3	2.31	53.62	0	4.3	8	BDL
M0071C	30°7.70'N	42°9.21'W	1390	MeBo	9	12.15	4.44	30.29	0	12.15	6	BDL
M0073A	30°7.90'N	42°10.97'W	1430.2	MeBo	1	2.2	0	0	0	2.2	40	BDL
Northern Sites												
M0070A	30°8.55'N	42°8.19'W	1140.5	MeBo	3	4	2.09	52.25	0	4	73	2
M0070B	30°8.54'N	42°8.16'W	1140.5	RD2	1	1.3	0.38	29.23	0	1.3	5	5
M0070C	30°8.54'N	42°8.19'W	1140.5	MeBo	3	5.21	2.21	42.42	0	5.21	–	–
M0074A	30°9.87'N	42°7.32'W	1550	MeBo	1	2.68	0.86	32.09	0	2.68	BDL	BDL

Notes: * Maximum dissolved concentrations in waters sampled after drilling. Full data set of hydrogen and methane concentrations in Früh-Green et al., 2017c.
http://publications.iodp.org/proceedings/357/EXP_REPT/TABLES/357_103/357_103_T12.CSV
BDL = Below detection limited

Table 2. Bulk rock chemical compositions of representative lithologies from IODP Exp. 357

Hole	M0071A	M0071A	M0071C	M0071C	M0069A	M0069A	M0069A	M0069A	M0072B	M0072B	M0072B	M0072B	M0076B	M0068A	M0068B	M0068B	M0068B	M0068B	M0075B
Core	1R-2	2R-1	1R-1	2R-1	5R-1	5R-1	10R-1	10R-3	5R -1	7R-1	8R-1	8R-2	7R-1	1R-1	1R-1	1R-1	2R-1	2-1R	2R-1
Cm	120-121	64-67	11-13	74-76	29.5-32	110-113	80-87	0-2.5	37-38	72.5-75	34-38	76-77	81-83	34-35	37.5-40	134-139	31-36	52-55	66-
68																			
Top depth	1.780	3.360	0.110	3.420	7.175	7.980	15.520	16.290	6.355	9.713	11.048	12.268	10.534	0.340	0.375	1.340	2.030	2.240	2.940
Bottom depth	1.790	3.390	0.130	3.440	7.200	8.010	15.590	16.315	6.365	9.738	11.088	12.278	10.554	0.350	0.400	1.390	2.080	2.270	2.960
Rock type	Serp	Metagb	Serp.	Serp	Metadol	Metadol	Serp	Serp	Talc schist	Metasom	Metasom	Metasom	Serp	Talc schist	Serp	Gabbro	Gabbro	Talc schist	
	Metadol																		
	Dunite		Harzburgite	Harzburgite		Harzburgite	Dunite		Harzburgite	Harzburgite	Harzburgite	Harzburgite		Harzburgite					
Major elements (wt.%)																			
SiO_2	36.62	26.21	37.75	39.95	26.31	31.61	36.19	33.08	50.64	42.41	38.60	43.53	39.55	50.01	40.03	52.14	50.77	59.63	47.09
TiO_2	0.02	0.24	0.02	0.05	0.92	1.40	0.03	0.03	0.14	0.03	0.03	0.05	0.03	0.10	0.03	0.43	0.30	0.02	1.08
Al_2O_3	1.11	21.04	1.17	1.27	18.48	20.36	1.02	1.58	3.82	1.05	1.22	1.86	0.86	4.27	1.16	16.40	17.53	1.05	18.23
Fe_2O_3	9.06	32.07	9.19	9.34	27.46	18.75	8.85	10.98	3.35	9.79	9.91	6.92	8.24	8.49	9.98	8.30	6.75	6.68	12.09
MnO	0.08	0.68	0.11	0.10	0.17	0.16	0.08	0.07	0.06	0.09	0.09	0.13	0.20	0.14	0.11	0.15	0.13	0.14	0.12
MgO	40.05	14.19	39.58	39.49	16.68	17.41	37.71	39.22	26.06	37.03	37.61	33.18	37.86	22.87	38.78	9.25	8.98	28.14	11.23
CaO	0.09	0.40	0.10	0.27	1.06	3.58	1.75	0.18	7.66	0.72	0.45	2.26	1.03	2.64	0.25	11.62	11.27	1.66	7.90
Na_2O	0.00	0.28	0.00	0.10	0.14	0.66	0.09	0.00	0.21	0.15	0.09	0.35	0.14	0.26	0.15	2.82	2.75	0.22	1.88
K_2O	0.00	0.03	0.00	0.05	0.00	0.02	0.00	0.00	0.00	0.00	0.00	0.00	0.01	0.00	0.05	0.00	0.00	0.02	0.00
P_2O_5	0.01	0.02	0.02	0.02	0.06	0.12	0.00	0.01	0.02	0.01	0.00	0.01	0.02	0.04	0.00	0.04	0.01	0.01	0.07
LOI	12.10	8.49	11.82	12.06	8.31	7.86	13.58	13.81	4.73	10.82	12.38	10.27	12.97	5.34	11.95	0.41	0.30	4.30	3.33
TOTAL	99.14	103.63	99.76	102.70	99.59	101.92	99.29	98.95	96.70	102.10	100.38	98.55	100.91	94.16	102.47	101.57	98.80	101.86	
	103.02																		
Mg#	90	47	90	89	55	65	89	88	94	88	88	90	90	84	89	69	73	89	65
MgO/SiO_2	1.09	0.54	1.05	0.99	0.63	0.55	1.04	1.19	0.51	0.87	0.97	0.76	0.96	0.46	0.97	0.18	0.18	0.47	0.24
Al_2O_3/SiO_2	0.03	0.80	0.03	0.03	0.70	0.64	0.03	0.05	0.08	0.02	0.03	0.04	0.02	0.09	0.03	0.31	0.35	0.02	0.39
Trace elements (ppm)																			
Cr	21572	748	11705	3796	680	523	8363	29698	1986	2746	3614	2663	2420	1349	3079	100	142	1746	414
Ni	14590	190	6715	2987	341	436	6531	7931	1253	3687	2193	1984	2810	1640	1651	97	94	1741	226
Li	22.74	16.60	12.95	3.93	6.76	12.76	2.59	0.22	0.15	1.47	0.49	1.94	8.73	13.12	6.29	5.65	4.39	21.67	12.48
Sc	82.79	41.08	30.04	16.53	37.40	41.16	29.24	14.31	17.68	9.71	6.60	8.35	7.95	7.91	10.37	41.62	35.61	7.56	35.90
V	340	139	190	74	198	292	85	116	60	28	26	32	34	29	42	190	151	19	179
Co	622	152	315	137	191	149	282	408	33	155	101	77	116	78	95	38	35	56	53
Cu	120	0	231	16	0	1	5	14	3	1	0	1	4	18	24	49	44	18	2
Zn	318	331	395	64	44	38	88	255	7	21	21	16	24	73	57	48	39	87	24
Ga	9.62	10.49	5.55	1.58	18.54	23.42	2.96	7.10	5.56	2.13	1.74	3.17	1.23	7.34	1.21	14.86	14.46	2.03	17.29
Rb	0.20	0.16	0.10	0.03	0.07	0.13	0.00	0.00	0.00	0.10	0.05	0.31	0.07	0.11	0.13	0.36	0.16	0.11	0.27
Sr	17.24	3.01	10.09	3.06	1.73	57.97	757.58	3.61	2.72	2.13	1.02	3.01	13.60	3.96	3.29	81.98	84.97	2.36	72.47
Y	2.33	6.88	0.99	0.38	22.71	31.68	2.18	0.09	7.15	3.54	1.84	7.12	2.12	41.57	1.03	14.84	8.08	0.43	27.57
Zr	0.00	3.13	0.00	0.09	23.45	20.82	1.23	0.00	4.11	1.54	1.17	5.11	0.38	5.43	0.33	17.45	4.76	0.41	39.78
Nb	0.11	0.02	0.02	0.02	0.35	1.56	0.05	0.05	0.15	0.49	0.47	0.51	0.03	1.89	0.06	0.30	0.05	0.05	1.24
Mo	2.18	0.22	1.04	0.27	0.21	0.14	0.62	0.37	0.01	0.21	0.18	0.21	1.30	0.46	0.40	0.16	0.20	0.15	0.18
Cd	95.23	40.86	33.98	26.71	8.63	16.64	17.48	15.49	4.10	6.83	6.43	10.01	49.11	29.42	27.48	62.06	35.36	9.30	20.17
Sn	0.05	1.71	0.20	0.08	0.55	0.81	0.20	0.14	0.22	0.23	0.28	0.53	0.10	1.28	0.82	1.16	0.26	0.37	0.83
Sb	1.09	0.02	1.34	0.15	0.01	0.00	0.01	0.01	0.00	0.00	0.00	0.02	0.13	0.05	0.29	0.03	0.02	0.03	0.01
Ba	1.19	0.17	0.97	0.30	0.20	1.33	0.72	0.00	0.00	0.00	0.15	0.16	4.58	0.87	0.88	4.41	2.15	0.23	2.32
La	0.16	0.30	0.18	0.00	0.85	3.05	0.23	0.00	0.21	0.47	0.24	0.57	0.13	8.56	0.30	0.92	0.26	0.09	1.55
Ce	0.33	0.72	0.05	0.00	2.69	10.48	0.74	0.02	1.06	1.72	0.90	2.33	0.29	30.19	1.32	3.29	0.84	0.28	5.82
Nd	0.40	1.05	0.22	0.01	3.81	9.26	0.60	0.03	1.56	1.23	0.75	2.20	0.49	16.89	0.61	3.21	1.15	0.15	6.44
Sm	0.17	0.54	0.05	0.00	1.92	3.33	0.17	0.01	0.68	0.34	0.22	0.72	0.15	4.43	0.16	1.28	0.58	0.04	2.53
Eu	0.10	0.41	0.05	0.02	0.24	0.96	0.11	0.01	0.26	0.10	0.06	0.19	0.07	0.52	0.25	0.63	0.45	0.04	0.95
Gd	0.33	0.80	0.08	0.00	2.57	4.09	0.23	0.00	0.89	0.40	0.24	0.85	0.20	4.86	0.17	1.72	0.86	0.05	3.23
U	5.05	0.01	3.27	1.52	0.02	0.06	2.22	4.50	0.02	0.16	0.03	1.19	1.98	0.09	0.47	0.04	0.01	0.06	0.02

Notes:

Depths are given as meters below surface (mbsf).

Abbreviations: LOI = Loss on Ignition; Serpentinized: Serp; Metagb: metagabbro; Metasom: metasomatic overprinting

Mg# = Mg/(Mg + Fe)*100. Note: Mg#s published in Früh-Green et al., 2017a are incorrect. These have been recalculated here from the oxide weight %, assuming the atomic ratio of 100 x Mg/(Mg + Fe^{2+}), using molecular weights and a factor of 0.8998 to calculate Fe^{2+} from Fe_2O_3.

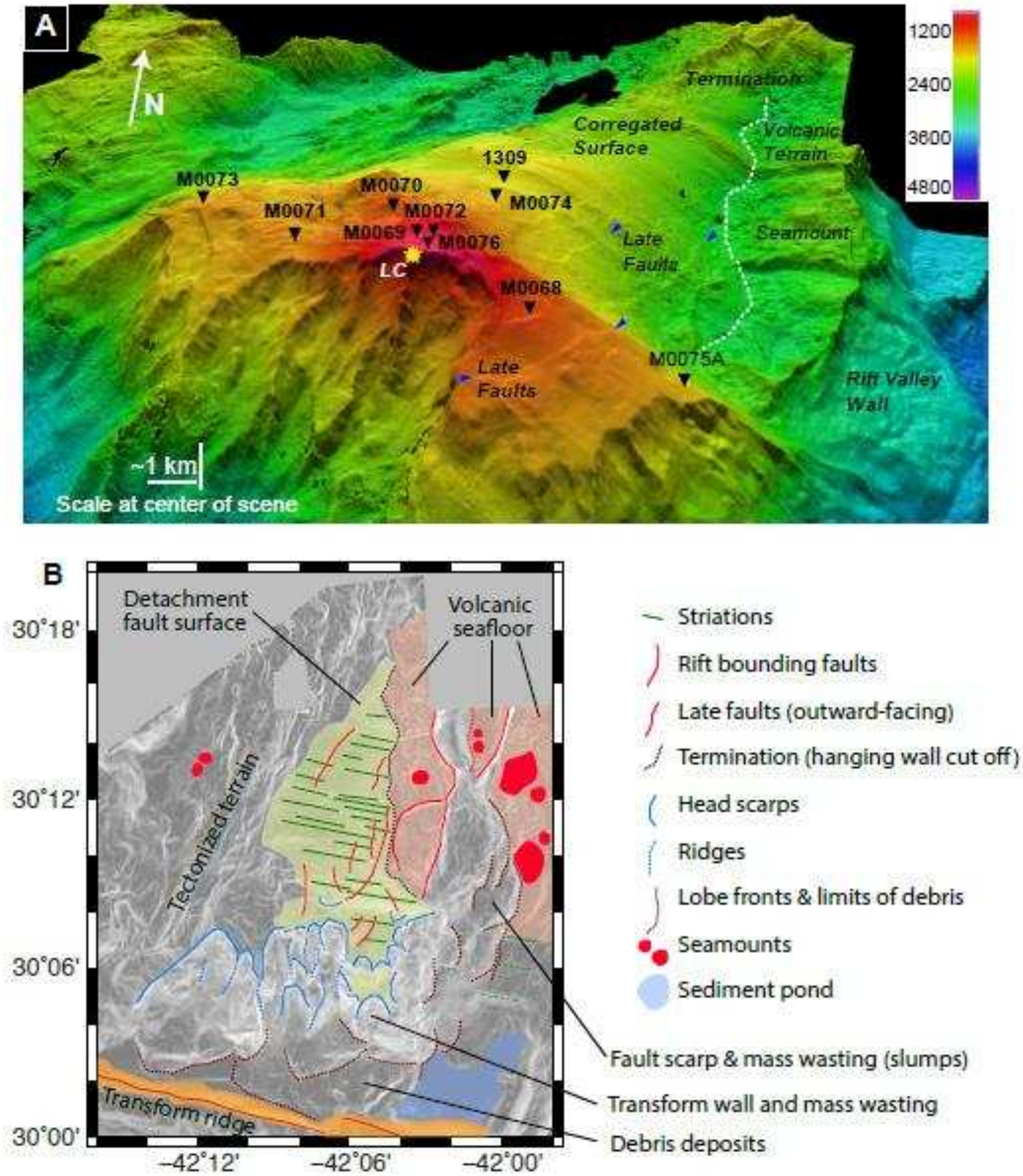

Figure 1. Bathymetry together with structural and morphological characteristics of the Atlantis Massif. (a) 3-D terrain model with a northward view of the detachment fault surface showing striations associated with detachment faulting, cross-cutting tectonic structures, with locations of the IODP Expedition 357 drill sites, the Lost City hydrothermal field (LC, yellow star) and IODP Site U1309. Based on new multibeam bathymetry acquired at 50 m resolution. (b) Interpretation of structural and morphological characteristics from new bathymetry data acquired during the expedition (reproduced from Früh-Green et al., 2017a).

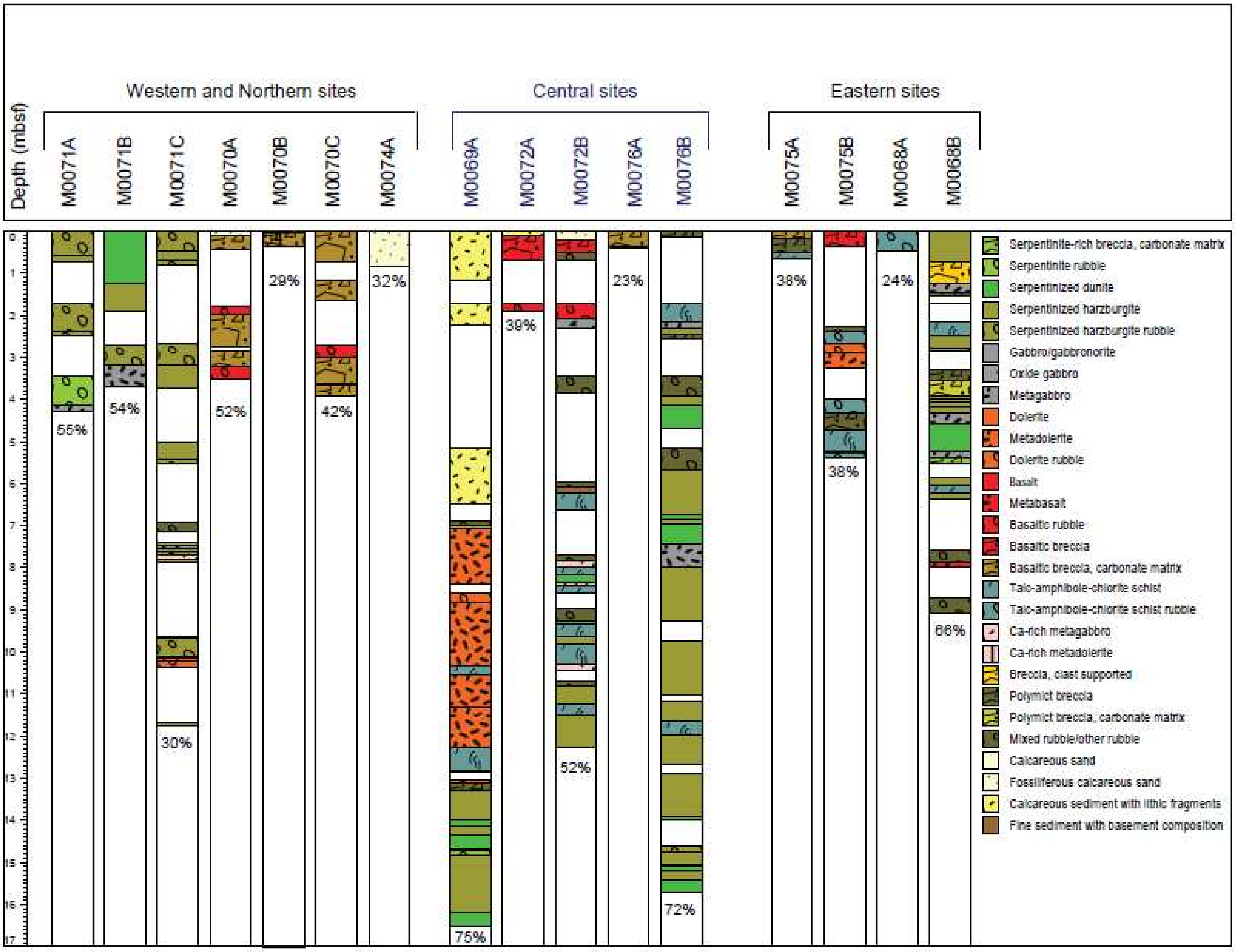

Figure 2. Lithologic variations on a regional scale and with depth in cores recovered during IODP Expedition 357. Percentages indicate overall percent core recovery for each hole. The central sites, highlighted in blue, recovered in situ sequences, whereas talus debris was recovered at the western and eastern sites along the southern ridge.

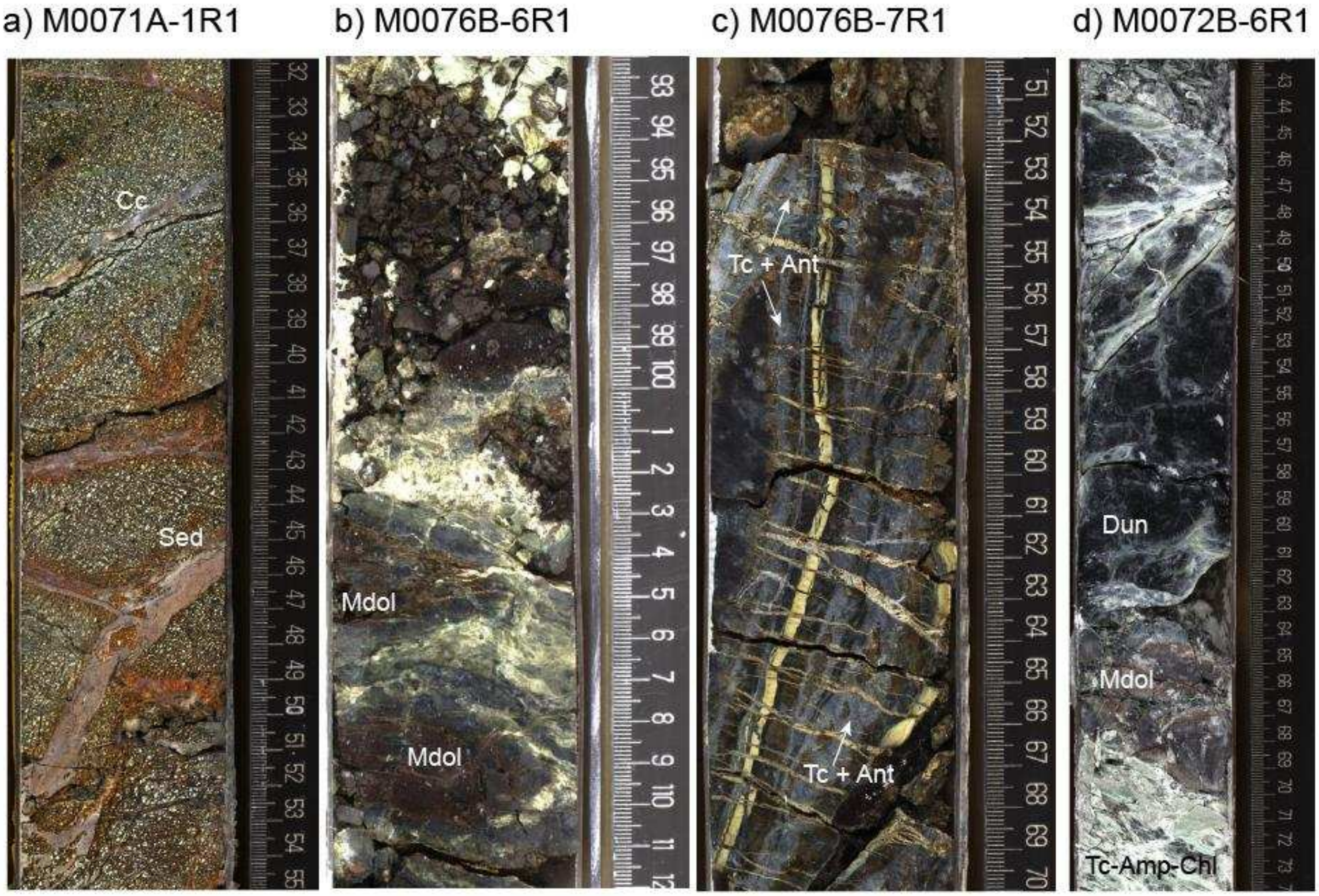

Figure 3. Examples of variations in rock type and structures in IODP Exp. 357 drill cores. (a) Serpentinized and oxidized dunite cut by moderately dipping calcite veins (Cc) and fractures filled with foraminiferous carbonate sediment (Sed). (b) Relationships between schistose zone talc-amphibole-chlorite schists (greenish-white domains) at the contact to cataclastically deformed metadolerite (Mdol). (c) Steeply dipping banded serpentine ± talc veins cutting serpentinized harzburgite. Light grey domains are previous fluid pathways resulting in metasomatic replacement of antigorite (Ant) by talc (Tc) (Rouméjon et al., 2018). (d) Metasomatic zones of talc-amphibole-chlorite schist (Tc-Amp-Chl) at contact to serpentinized dunite (Dun) intruded by dolerite and transitioning again to talc-amphibole-chlorite schist. Photos: IODP ESO.

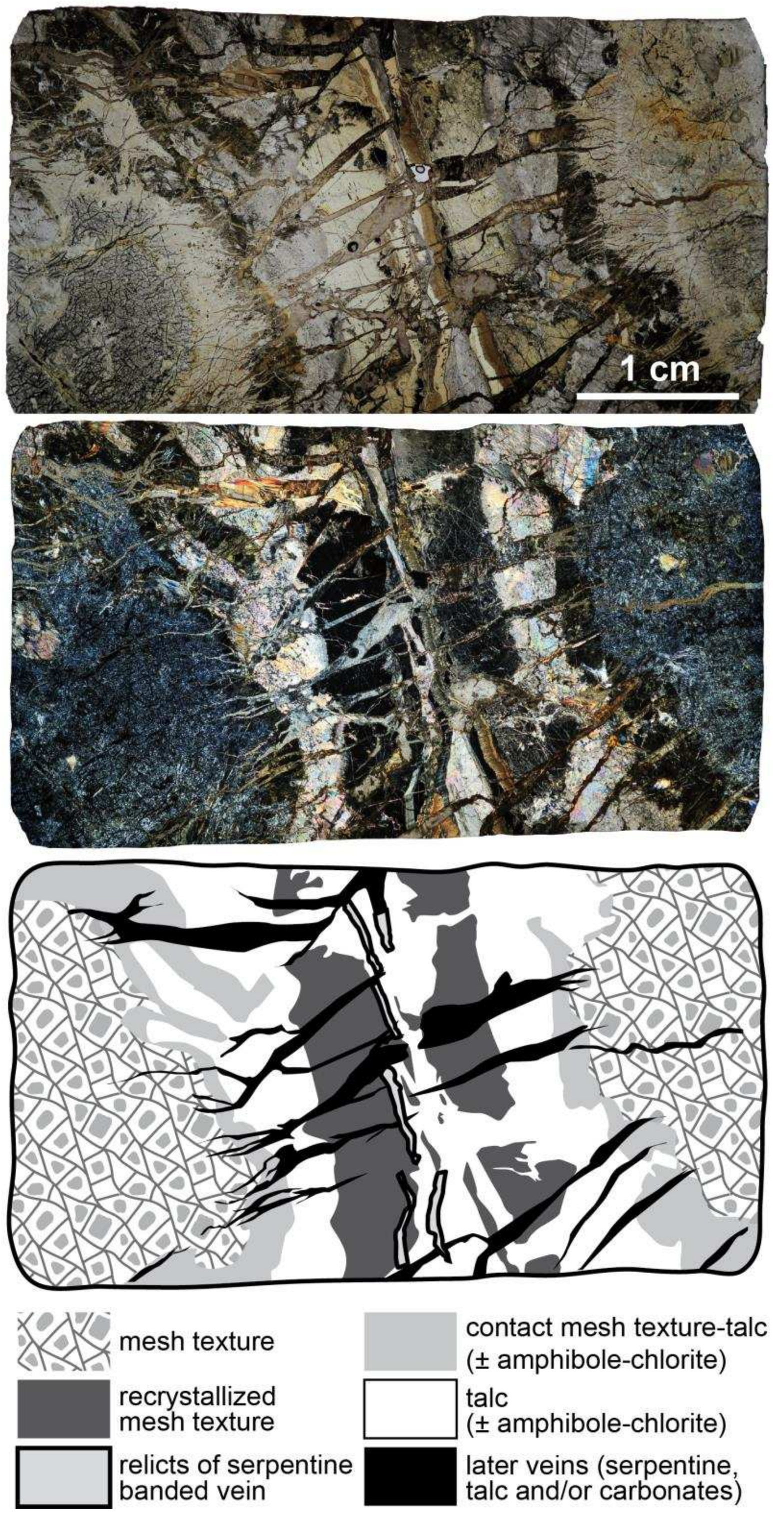

Figure 4. Characteristic serpentine textures and cross-cutting relationships associated with progressive alteration and veining in serpentinized harburzgite (example from M0076B-7R1, 43-45cm); (a) plane polarized light, (b) crossed polarized light, and (c) schematic representation of overprinting relationships. Modified from Früh-Green et al., 2017d, Fig. 7.

Figure 5

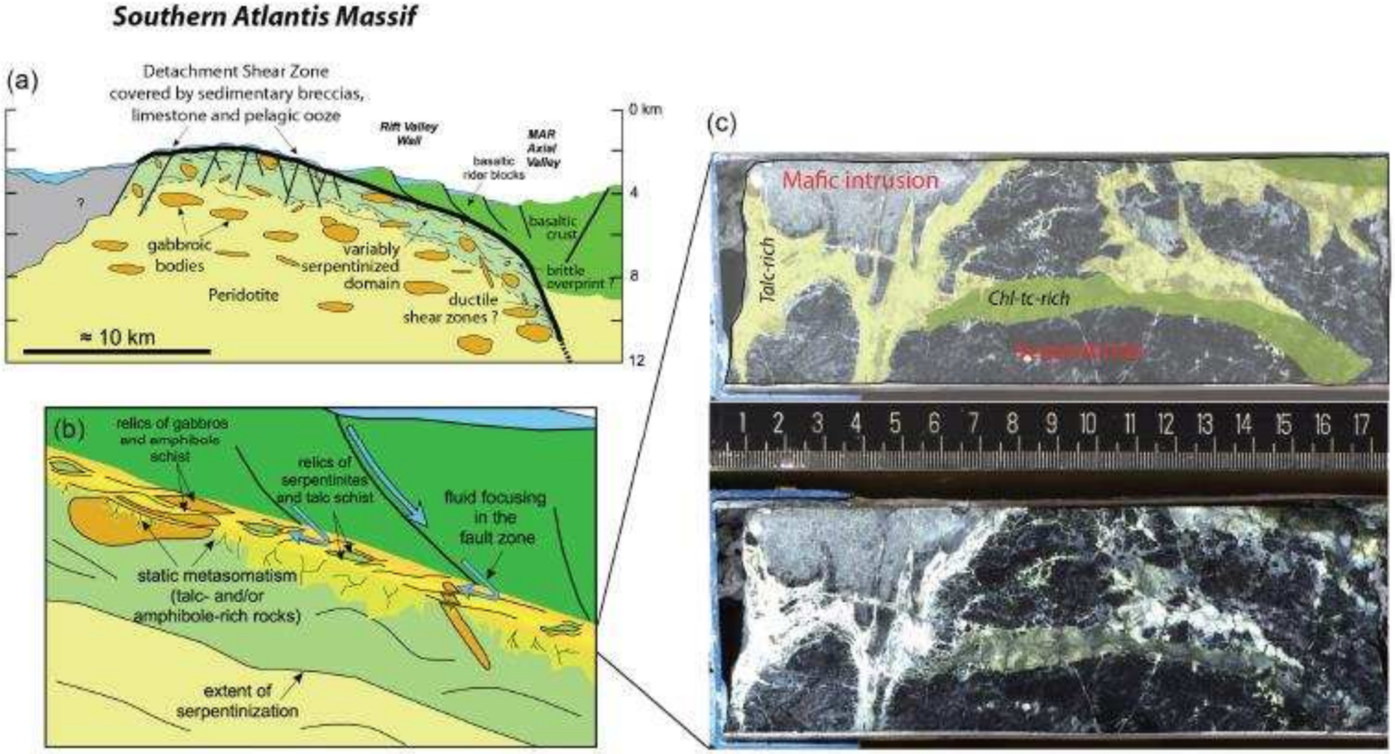

Figure 5. Model of the tectono-magmatic evolution and alteration of heterogeneous lithosphere at Atlantis Massif. (a) Interpretative cross section showing fluid pathways, metasomatic zones and extent of serpentinization (light green shaded region) related to detachment faulting and steep normal faults (modified after Boschi et al., 2006a). (b) Detail of <100 m detachment shear zone (in red-yellow) characterized by heterogeneous, variably altered and deformed gabbroic and peridotite lithologies and with extensive synkinematic metasomatism. The resulting talc-amphibole schists enclose lenses of relic, locally less deformed, serpentinite and gabbroic rocks (modified from Boschi et al., 2006a). (c) Example of magmatic intrusion in a fully serpentinized harzburgite from Core M0072-8R2, 0–18 cm. Late metasomatic alteration at the contact between the mafic/ultramafic rocks produced white and green talc-amphibole-chlorite assemblages that crosscut the previous texture. Chl = chlorite, tc = talc (reproduced from Früh-Green et al., 2017d, Fig. 13).

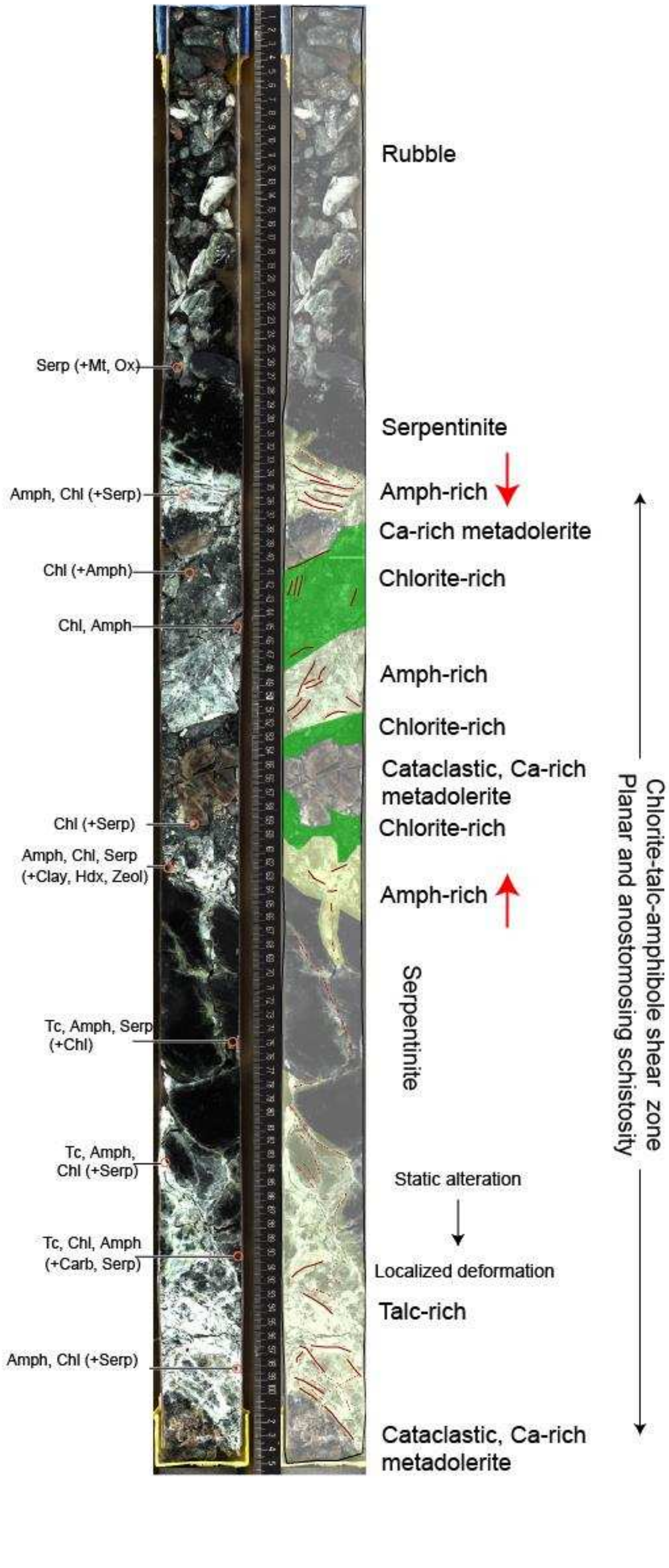

Figure 6. Example of complex lithological and deformation relationships between mafic intrusions in peridotite and metasomatic domains in the IODP Expedition 357 cores, showing a transition from static alteration to strain localization in alternating talc-, amphibole-, and chlorite-rich shear zones (from Core M0072B-7R-1, 0–105 cm). Red circles = samples taken for XRD analyses and corresponding mineral assemblages. Serp = serpentinite, Mt = magnetite, Ox = oxide, Amph = amphibole, Chl = chlorite, Hdx = hydroxide, Zeol = zeolite, Carb = carbonate. Modified from Früh-Green et al., 2017d, Fig. 12.

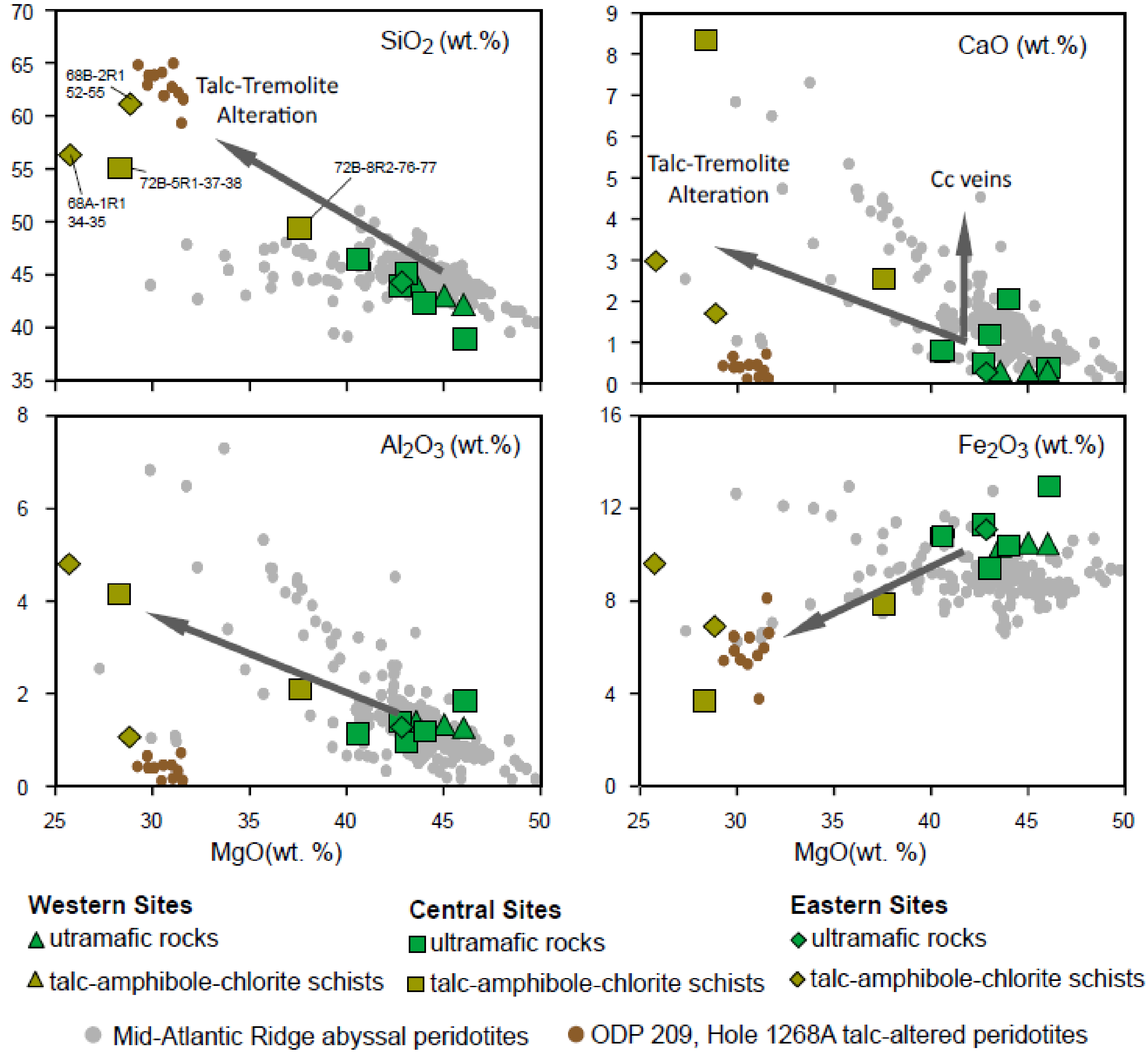

Figure 7. Selected whole-rock major elements (normalized, volatile-free compositions, and in weight % oxides, wt %) vs. MgO for serpentinized ultramafic rocks (including impregnated / metasomatized samples) and talc-amphibole-chlorite schists from Atlantis Massif, IODP Exp. 357. Data from Mid-Atlantic-Ridge abyssal serpentinized peridotites and talc-altered peridotites are shown for comparison. Talc-amphibole alteration is associated with a general trend to higher Si, Ca and Al compositions and a decrease in Mg and Fe. Global abyssal peridotite field defined by data from PetDB (http://www.earthchem.org/petdb, May 2016). Data for talc-altered peridotite field from ODP Leg 209, Hole 1268A (Paulick et al., 2006; also from PetDB). Modified from Früh-Green et al., 2017a, Fig. 11.

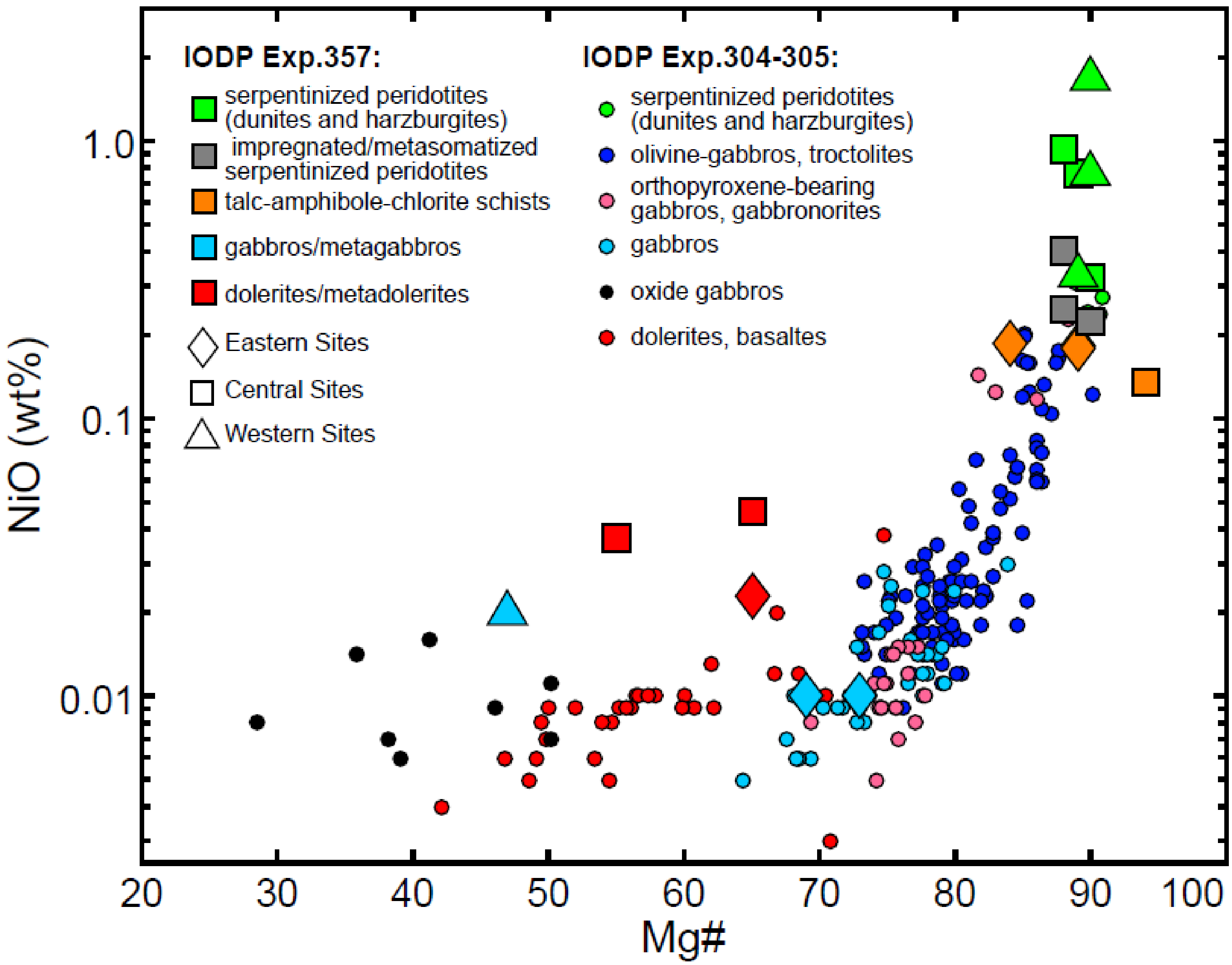

Figure 8. Ni concentrations (calculated as weight % (wt %) oxides and normalized to volatile-free concentrations, plotted on a log scale) vs. Mg# of Atlantis Massif mafic and ultramafic rocks from Expedition 357 compared with those from cores recovered at Site U1309 during Integrated Ocean Drilling Program Expedition 304/305 (Godard et al., 2009).

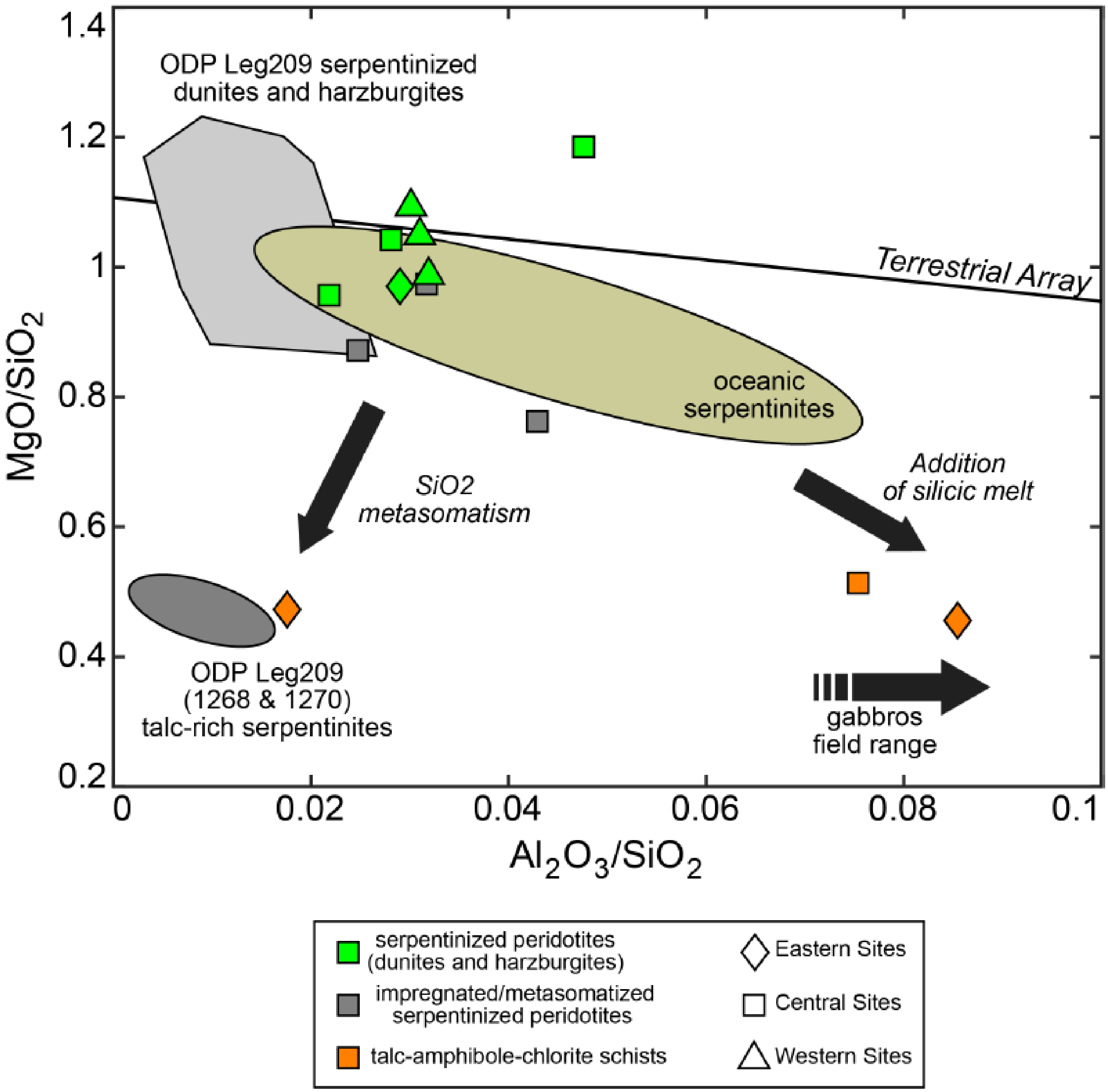

Figure 9. MgO/SiO$_2$ vs Al$_2$O$_3$/SiO$_2$ diagram showing variations in bulk rock chemistry and changes with Si-metasomatism. Atlantis Massif compositions are also compared with compositions of serpentinites and talc schists from IODP Site U1309 and from 15°20′N recovered during ODP Leg 209 (Paulick et al., 2006) as well as the global data set of abyssal peridotites reported in Niu (2004), which define a trend parallel to the terrestrial array (Jagoutz et al., 1979). The geochemistry of the Atlantis Massif samples reflects a variety of processes including modal mineralogical composition, melt impregation and multiple phases of hydrothermal alteration.

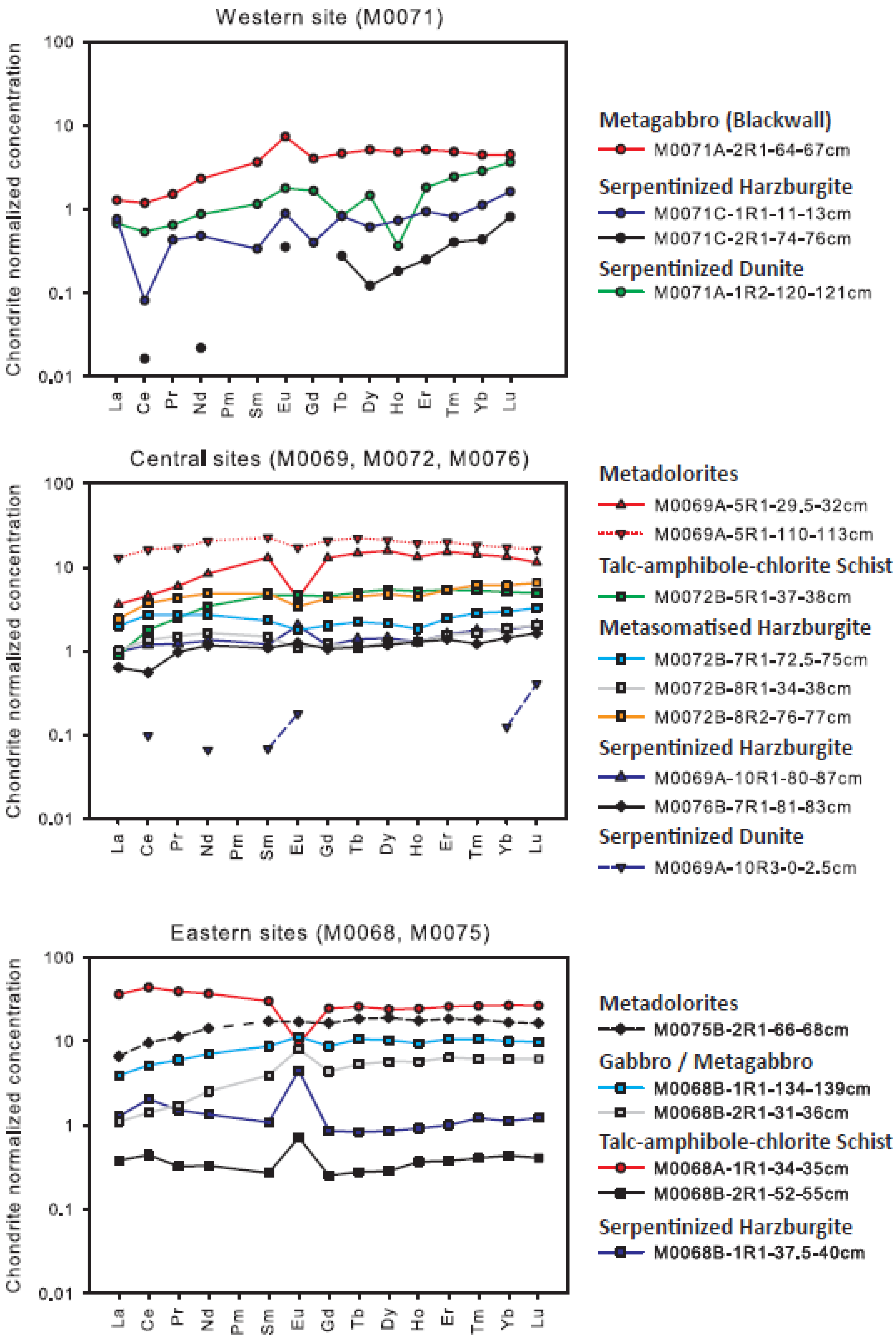

Figure 10. Compilation of chondrite-normalized REE concentrations of Atlantis Massif mafic and ultramafic rocks from samples of the IODP Expedtion 357 drill cores (see Table 2). Values for CI chondrite from McDonough and Sun, 1995.

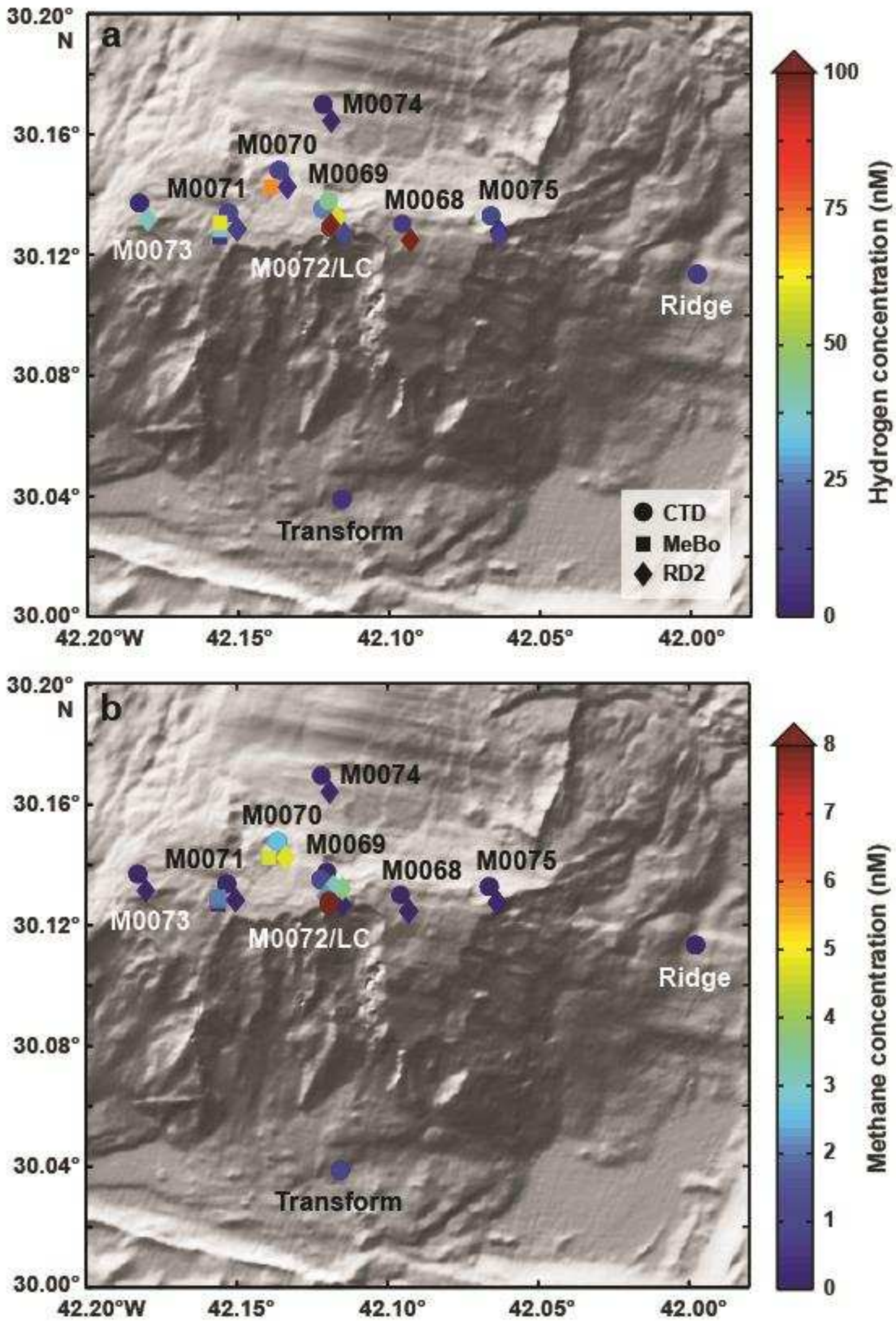

Figure 11. Highest measured hydrogen and methane concentrations in samples from CTD rosette bottom waters acquired before drilling and sensor package Niskin bottles taken by RD2 and MeBo after drilling at the Atlantis Massif drill sites during IODP Expedition 357. Dark red circles indicate samples from the Lost City (LC) plume.

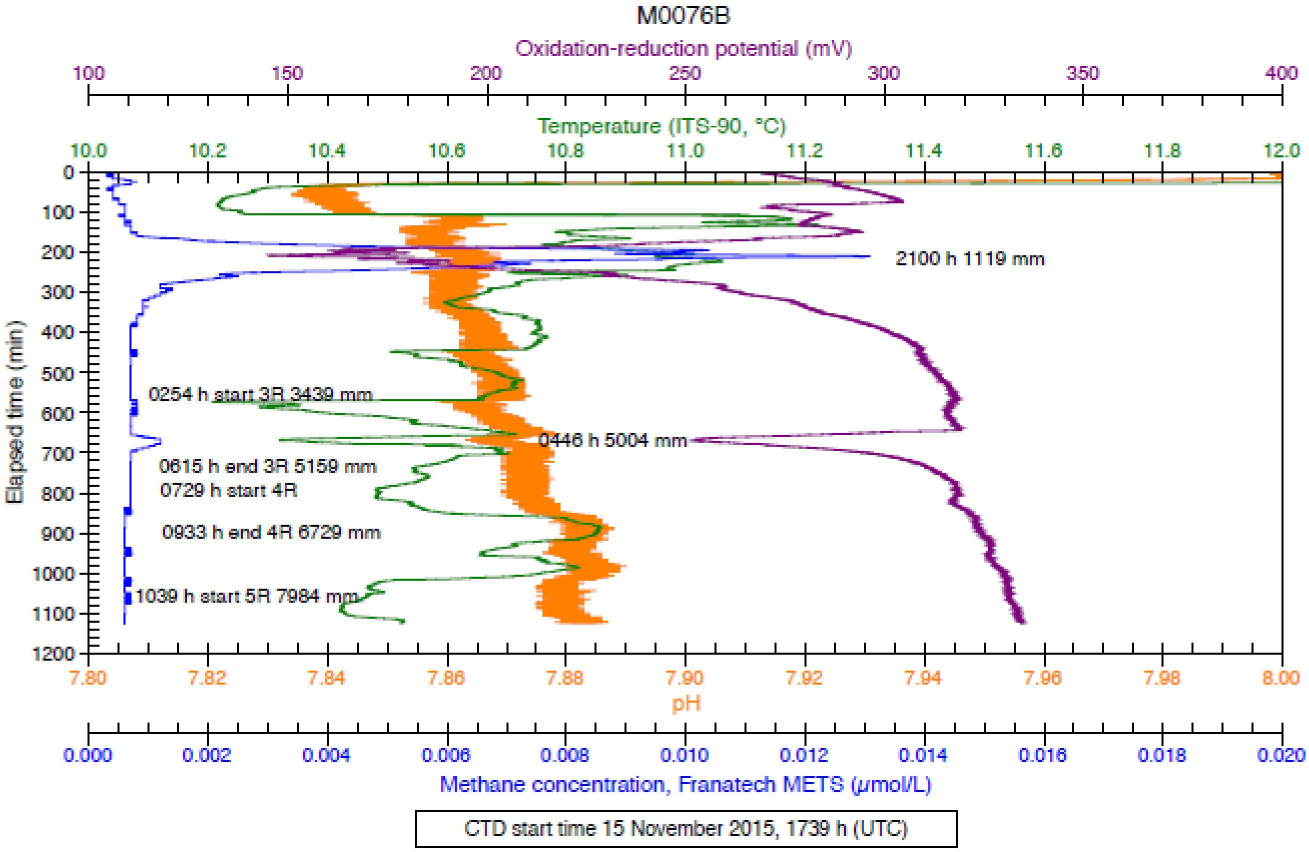

Figure 12. Example of variations in fluid chemistry during drilling operations and correlations of geochemical signatures recorded by the sensor packages on the rock drills from sensor data for Hole M0076B, Cores 1R–5R. Elapsed time = time since the start of the sensor package data file. Penetration depth (in mm) was reconstructed from drill logs.

Figure 13. Frame-grab photograph from drilling video of bubbles (black arrows) that were observed issuing from Hole M0070C and around the drill base during operations at this site, even when coring had stopped.

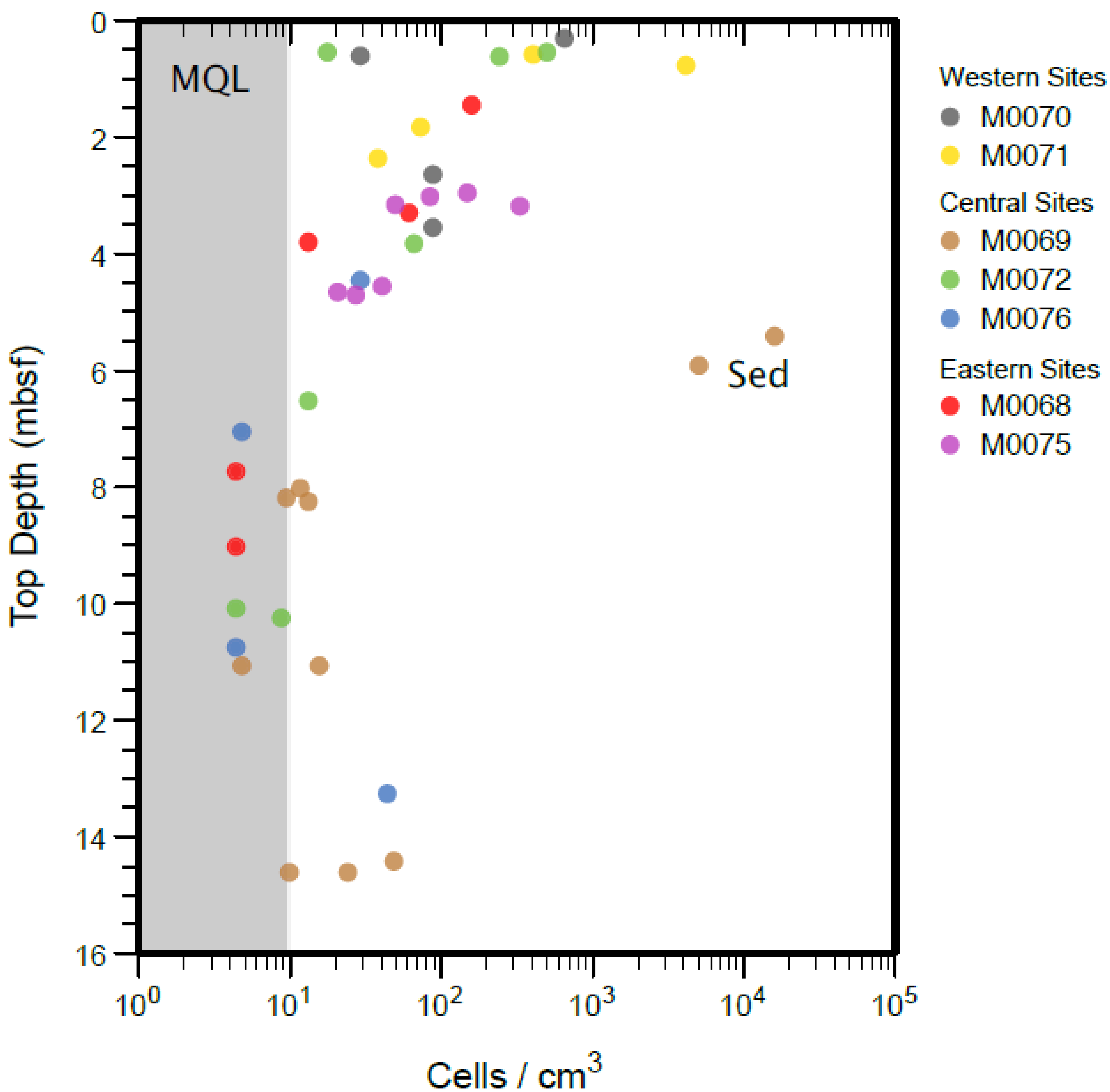

Figure 14. Downhole variations in cell counts from interior portions of whole round cores of the basement rock samples (and two sediment (Sed) samples from Hole M0069A) taken onboard during IODP Exp. 357. Data from Hole M0074 not included due to extensive damage to this short sediment core. The shaded region shows the range of counts below the minimum quantification limit (MQL) of 9.8 cells cm^{-3}.