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Comparison of the impacts of acid and nitrogen additions on carbon fluxes in European conifer and broadleaf forests

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Abstract

Increased reactive nitrogen (N) loadings to terrestrial ecosystems are believed to have positive effects on ecosystem carbon (C) sequestration. Global “hot spots” of N deposition are often associated with currently or formerly high deposition of sulphur (S); C fluxes in these regions might therefore not be responding solely to N loading, and could be undergoing transient change as S inputs change. In a four-year, two-forest stand (mature Norway spruce and European beech) replicated field experiment involving acidity manipulation (sulphuric acid addition), N addition (NH₄NO₃) and combined treatments, we tested the extent to which altered soil solution acidity or/and soil N availability affected

25 the concentration of soil dissolved organic carbon (DOC), soil respiration (Rs), microbial community
26 characteristics (respiration, biomass, fungi and bacteria abundances) and enzyme activity. We
27 demonstrated a large and consistent suppression of soil water DOC concentration driven by chemical
28 changes associated with increased hydrogen ion concentrations under acid treatments, independent
29 of forest type. Soil respiration was suppressed by sulphuric acid addition in the spruce forest,
30 accompanied by reduced microbial biomass, increased fungal:bacterial ratios and increased C to N
31 enzyme ratios. We did not observe equivalent effects of sulphuric acid treatments on Rs in the beech
32 forest, where microbial activity appeared to be more tightly linked to N acquisition. The only changes
33 in C cycling following N addition were increased C to N enzyme ratios, with no impact on C fluxes (either
34 Rs or DOC). We conclude that C accumulation previously attributed solely to N deposition could be
35 partly attributable to their simultaneous acidification.

36 **Capsule:** Acidification alters soil carbon fluxes in spruce and beech forests due to the altered soil pH.
37 Limited effect of N addition on soil C cycling in both forest types

38
39 **Keywords:** nitrogen, fertilization, acidity, sulphur, organic carbon, respiration, DOC, hydrolytic
40 enzymes, fungi, bacteria

41 **1. Introduction**

42 Human activities have been altering element cycling for several thousand years (Ruddiman, 2003;
43 Shotyky, 1998). Since the Industrial Revolution, emissions of greenhouse gases, nitrogen (N) and sulphur
44 (S) have been increasing (Kopáček and Posch, 2011; Lamarque et al., 2013) and their negative impacts
45 on ecosystems led to the implementation of various international agreements to abate these
46 emissions (United Nations Framework Convention on Climate Change, Convention on Long-range
47 Transboundary Air Pollution, Kyoto Protocol, Paris Agreement). Besides exerting significant negative
48 pressures on global biodiversity (Butchart et al., 2010), atmospheric pollutant deposition has impacted
49 natural biogeochemical cycles of terrestrial ecosystems, including the ecosystem C balance (Hyvönen
50 et al., 2007; Lamersdorf and Borken, 2004; Thomas et al., 2013).

51 Nitrogen limitation is an inherent feature of many temperate terrestrial ecosystems, thus fertilization
52 by atmospheric N deposition may have increased C pool in biomass and soils of terrestrial ecosystems
53 and thus stimulate the terrestrial CO₂ sink (de Vries et al., 2009). Increased N availability may affect
54 terrestrial C balances via increased net primary productivity (NPP) and/or decreased litter
55 decomposition rates (Janssens et al., 2010; Liu and Greaver, 2010). The potential of a negative effect
56 of increased N availability on decomposition rates was suggested by many experimental N addition
57 studies (e.g. Berg and Matzner 1997; Olsson et al. 2005; Hobbie 2008). Stimulation of litter
58 decomposition is expected at sites with low ambient N deposition and high-quality litter (low lignin),
59 whereas suppression of litter decomposition is suggested at sites with higher ambient N deposition or
60 at sites with low quality (high lignin) litter (Knorr et al., 2005). Despite current studies have focused on
61 the impact of N deposition on C balance, there is clear evidence that acidification (either from S or N
62 deposition) may also affects a range of key processes, including slowdown of litter decomposition
63 (Chen et al., 2015; Liang et al., 2013; Oulehle et al., 2011), as well as mobility and leaching of dissolved
64 organic carbon (Evans et al., 2012). Suppression of litter decomposition at low pH may occur through
65 reduction in the availability of C for microbial community (Persson et al., 1989) and/or by direct effect
66 on the microorganisms themselves (Błońska et al., 2017, 2016; Pennanen et al., 1998; Wei et al., 2013).
67 Input of N to the soil, either from deposition or fertilization, can have an acidifying effect, depending
68 on the fate (immobilization vs. leaching, nitrification) of mineral N in the soil. Thus, it is likely that
69 experimentally induced responses to N fertilisation might in fact be due to changes in soil acid-base
70 status (Evans et al., 2008; Chen et al., 2016). A better, integrated understanding of soil C and N
71 responses to changing inputs of both S and N is needed to predict whether soils will in future act as C
72 sinks or C sources as atmospheric pollutant loadings change. Accounting for, that much of Europe and
73 North America have experienced high historic depositions of S, and receive high N inputs, whilst other
74 regions including large parts of Asia continue to receive high depositions of both S and N (Duan et al.,
75 2016; Larssen et al., 2006; Posch et al., 2015), it is likely that recent measurements of C fluxes from
76 these regions might not be in long-term balance with climatic conditions, but represent ecosystems

77 subjected to multiple anthropogenic pressures, and in many cases undergoing transient change as
78 deposition levels decline or rise.

79 Here, we present a set of replicated acidity/N availability manipulation experiments in adjacent mature
80 broadleaved (beech forest) and coniferous (spruce forest) sites. Spruce and beech represent the
81 dominant tree species in the Central European forests. Both stands are influenced by similar abiotic
82 conditions (climate, deposition, bedrock, original soil type) but differ in vegetation cover as a result of
83 long-term differences in land-use. Furthermore, long-term ambient monitoring from spruce site
84 provided evidence of C soil loss as a consequence of acidification recovery (Oulehle et al., 2011) and
85 both sites were intensively studied to assess the rates of acidification recovery in formerly highly
86 polluted areas (Oulehle et al., 2006). We examined whether there has been a shift in the ecosystem C
87 fluxes in either (dissolved) organic or (gaseous) inorganic forms, as a function of altered acidity and N
88 availability. *In situ* decomposition experiments and measurements of microbial characteristics and
89 enzyme activity provided further insights into C metabolism under changing acidity and N availability.

90 **2. Material and Methods**

91 **2.1. Sites and experimental design**

92 In 2013 two parallel acidity and N availability manipulation experiments were established, at two
93 adjacent forests (50.59 N, 13.26 E) selected to represent major European forest types – mature Norway
94 spruce forest (*Picea abies* (L.) Karst.) and mature European beech forest (*Fagus sylvatica* L.). The spruce
95 stand consists of Norway spruce monoculture planted in the early 1930s, and the beech stand consists
96 of European beech monoculture (Oulehle et al., 2016b). The distance between stands is approximately
97 1 km and both stands are underlined by gneiss bedrock. The dominant soil type is dystric cambisol
98 (Dambrine et al., 1993). Soils at both sites were acidified by acid deposition during the 20th century,
99 which resulted in low soil pH and low base saturation. Despite similar bulk S deposition, total
100 deposition is higher in the spruce stand compared to the beech stand. Throughfall fluxes of inorganic
101 nitrogen are similar in both stands (

102 ; Oulehle et al. 2011, 2016b). Ground cover vegetation in the spruce stand (60%) is formed mainly by
 103 *Avenella flexuosa* L. Drejer with admixture of *Vaccinium myrtillus* L., *Dryopteris dilatata* (Hoffm.) A.
 104 Gray, *Calamagrostis villosa* (Chaix) J. F. Gmel. and *Galium saxatile* L. The ground cover vegetation in
 105 the beech stand (1.5%) comprises mainly of *Calamagrostis villosa* (Chaix) J. F. Gmel. and *Avenella*
 106 *flexuosa* L. Drejer.

107 At each forest stand, soil acidity and N availability were manipulated for 3 years by systematic addition
 108 of S and N treatment solutions. At each experimental site, sixteen 3 x 3 m plots were assigned to control
 109 (Ctrl), nitrogen (N), acid (S) and acid + nitrogen (S+N) treatments in a randomized blocked design, thus
 110 four replicates per treatment were available. In October 2012 plots were established (mineral soil
 111 lysimeter installation), and in April 2014 treatments began following a period of pre-treatment
 112 measurements (April 2013 – March 2014). Treatments consists of monthly additions (April –
 113 November) of sulphuric acid (H₂SO₄) and ammonium nitrate (NH₄NO₃), mixed with rainwater collected
 114 at sites and applied using 15 l watering cans evenly across the plots, followed by addition of extra 15 l
 115 rainwater to properly seep in soil all added chemicals. The treatment dose was equivalent to 50 kg S
 116 ha⁻¹ year⁻¹, while the additional water addition was equivalent to 2.5 % of average annual rainfall.
 117 Nitrogen treatments gave an input of 50 kg N ha⁻¹ year⁻¹. NH₄NO₃ addition, albeit ammonium is a weak
 118 acid, did not significantly alter pH of added solution (applied solution pH was ≈ 5.4).

119 **Table 1. Stand characteristics, soil properties (BS – base saturation), biomass properties (net primary productivity – NPP)**
 120 **and measured sulphur (S), dissolved inorganic nitrogen (DIN) deposition and dissolved organic carbon (DOC) and**
 121 **nitrogen (DON) leaching. Data from Oulehle et al. (2016b).**

	Site description				Soil (Forest floor)					
	Latitude	Longitude	Temperature °C	Precipitation mm	Forest age years	pH	BS %	C/N g g ⁻¹	C kg m ⁻²	N
Spruce	50.591	13.253	6.5	1092	85	3.80	19	24.9	3.74	0.15
Beech	50.589	13.267			155	4.34	70	20.8	6.54	0.31
	Biomass		Ambient deposition 2013-2015				Soil leaching (Forest floor)			
	NPP kg C ha ⁻¹ year ⁻¹	Litterfall kg C ha ⁻¹ year ⁻¹	Litter C/N g g ⁻¹	Biomass C/N	Bulk S kg ha ⁻¹ year ⁻¹	Throughfall S kg ha ⁻¹ year ⁻¹	Bulk DIN kg ha ⁻¹ year ⁻¹	Throughfall DIN kg ha ⁻¹ year ⁻¹	DOC g m ⁻²	DON g m ⁻²
Spruce	9373	1650	42	133	4.0	14	9.6	14	17.9	0.56
Beech	7980	2882	52	178		7.5		16	18.2	0.52

123

124 *2.2. Soil solution sampling and analysis*

125 Every 3-5 weeks during the snow/ice free period (usually April – November) samples were taken,
126 ideally after rainfall when soil moisture conditions permitted collection of soil water. In every case soil
127 water was collected before treatment application. Soil water samples were collected using Rhizon®
128 suction samplers (Rhizosphere Research Products, Wageningen, NL), comprising 10 cm long, 2.5 mm
129 diameter porous membranes attached to 50 ml syringes. Four to six Rhizon samplers were inserted
130 randomly into the forest floor of each plot, to give one composite sample per plot. In the mineral soil,
131 Prenart Super quartz standard suction samplers (Prenart Equipment, Frederiksberg, DK) were inserted
132 into 30 cm soil depth. Soil solution was collected by applying suction into a 2 l collecting bottle.

133 Samples were stored at 4 °C, and analysed for pH using Radiometer TTT-85 with a combination
134 electrode, nitrate (NO_3^-) by high-performance liquid chromatography (Knauer 1000); ammonium
135 (NH_4^+) was determined by indophenol blue colorimetry. DOC (dissolved organic carbon) was
136 determined by a Tekmar-Dohrman Apollo 9000 analyzer (Tekmar Dohrmann, OH, USA). Samples for
137 DOC determination were filtered (0.4 μm glass fiber Macherey-Nagel GF 5) before analysis.

138 *2.3. Soil CO₂ flux measurements*

139 Soil CO₂ efflux (R_s) was measured monthly between 2013 and 2016 during the snow-free period (April–
140 November/December) using a LiCor infrared gas analyser LI-8100A (LiCor Biosciences, NE, USA)
141 attached to a LiCor Survey Chamber (8100-103 20 cm survey chamber; LiCor Biosciences, NE, USA). In
142 each plot, four soil collars (287cm²) were placed into the forest floor to a 5 cm depth (64 soil collars in
143 each forest stand) in March 2013. The CO₂ efflux was calculated based on a linear increase in chamber
144 CO₂ concentration over time.

145 We used the van't Hoff equation (Davidson et al., 2006) to model the relationship between soil
146 temperature and soil respiration:

147 $R_s = \alpha e^{\beta T}$ (where $Q_{10} = e^{\beta 10}$) Eq. 1

148 where R_s is the *in situ* measured respiration rate, α and β are fitted parameters, T is *in situ* measured
149 soil temperature and Q_{10} is the calculated temperature sensitivity of respiration. R_s was calculated on
150 an hourly basis using on site measured soil temperatures and summed for the whole year to estimate
151 an annual CO_2 efflux.

152 **2.4. Microbial characteristics and DNA extraction**

153 Soils for microbial characteristics were sampled every year in May. Microbial C (C_{mic}) and N (N_{mic}) were
154 calculated as the difference between the chloroform fumigated and non-fumigated soil samples
155 (sulphate extractable C and N), and amended using extraction efficiency factors of 0.38 for microbial
156 C (Vance et al., 1987), and 0.54 for microbial N (Brookes et al., 1985). Microbial respiration was
157 assessed as CO_2 production after one week incubation of soil in closed vials placed in the dark (10 g of
158 fresh soil, 60 % of water holding capacity, 15 °C), using gas chromatography (Agilent GC HP 6850, USA).

159 Into a FastPrep™ Lysis Matrix E tube (MP Biomedicals, Solon, OH, USA) approx. 0.5 g of soil was added.
160 Then Hexadecyltrimethylammonium bromide (CTAB) extraction buffer, containing 5% CTAB (in 0.7 M
161 NaCl, 120 mM potassium phosphate, pH 8.0) and 0.5 ml phenol-chloroform-isoamylalcohol (25:24:1),
162 was added and agitated at speed 5-6 for 45s in a FastPrep Instrument (MP Biomedicals, Solon, OH,
163 USA). Following bead beating, chloroform extracted samples were precipitated in a PEG 6000/1.6 M
164 NaCl solution. 70% ethanol was used to wash pellets and re-suspended in molecular biology grade
165 water. Quantus™ (Promega, USA) was used for fluorometrical quantification of total DNA.

166 Quantative PCR (qPCR) of bacterial and fungal genes of ribosomal small subunit rRNA (SSU rRNA) was
167 performed using the FastStart SybrGREEN Roche® Supermix and Step One system (Life Technologies,
168 USA) as described previously (Bárta et al., 2017; Gittel et al., 2014). R^2 values for the standard curves
169 were > 0.99. Slope values were -> 3.37 giving an estimated amplification efficiency of > 93%.

170 Detection limits (i.e. lowest standard concentration that is significantly different from the non-
171 template controls) were less than 100 gene copies for each of the genes per assay. All samples,
172 standards and non-template controls were analysed in duplicates.

173 *2.5. Extracellular enzyme activities*

174 Extracellular enzyme activities were assessed in soils sampled in May and October each year.
175 Microplate fluorometric assays under standard conditions were used to determine extracellular
176 enzyme activities. All hydrolytic enzyme activities were determined in 1 g soil which was suspended in
177 100 ml of distilled water and for 4 min sonicated to allow disruption of the soil particles. Then 200 μ L
178 of the soil suspension was added into 50 μ L methylumbelliferyl (MUF) of substrate solution for β -
179 glucosidase (BG), exocellulase (cellobiohydrolase - CEL), phosphatase (PME) and N-acetyl-
180 glucosaminidase (NAG) determination or to 50 μ L of 7-aminomethyl-4-coumarin (AMC) substrate
181 solution for leucine-aminopeptidase (LEU) determination (Bárta et al., 2014). 50, 100 and 300 μ M
182 concentrations of each fluorogenic substrate were tested and the one with the highest enzymatic
183 activity was picked for further analysis. Plates were incubated for 120 min at 20°C. Excitation
184 wavelength of 365 nm and emission wavelength of 450 nm were used for fluorescence quantification
185 using Infinite F200 microplate reader (TECAN, Germany). To assess enzyme activities stoichiometry we
186 calculated C_{enz}/N_{enz} , C_{enz}/P_{enz} and N_{enz}/P_{enz} ratios, where $C_{enz} = BG+CEL$, $N_{enz} = NAG+LEU$ and $P_{enz} = PME$
187 (Sinsabaugh et al., 2009).

188 *2.6. Data analysis*

189 After testing the data for normality (Kolmogorov-Smirnov), skewness and homogeneity of variances
190 using Bartlett's test, pH, DOC and soil respiration were analysed for the pre-treatment period (May
191 2013 – April 2014) and for each full year of manipulation data (2014, 2015, 2016) using repeated
192 measures ANOVA (treatment as a fixed factor) for each of the stands and for each time period. Fisher's
193 LSD multiple comparison test was used at $p < 0.05$ to detect which treatments were different. Effects
194 of forest stand, season (spring and autumn sampling - for enzyme activities only) and treatments on

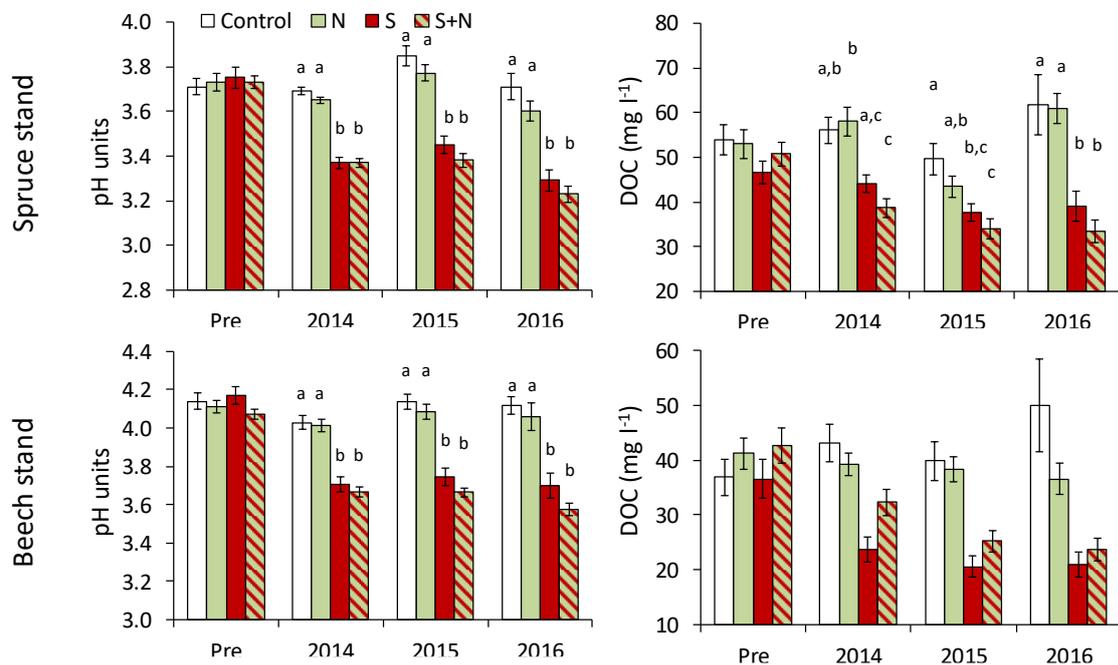
195 microbial biomass, fungi and bacteria abundances, fungi to bacteria ratios, enzymatic activity and
196 enzyme stoichiometry were determined using a general linear model (GLM) for each time period. We
197 refer only to the treatment effects in depicted figures. All data were analysed using statistical software
198 NCSS v11.

199 **3. Results**

200 **3.1. *Soil solution chemistry***

201 Pre-treatment soil water chemistry, in both forest floor and mineral soil, was similar among plots and
202 no significant differences were detected in respect of soil water pH and DOC at either the spruce or
203 beech stands (Figure 1, Suppl. Fig. 1).

204 From the onset of treatment in April 2014, immediate and significant changes in forest floor soil water
205 pH occurred at both stands. In acid treatments (S and S+N), soil water pH was reduced by 0.41-0.48 pH
206 units at the spruce stand and 0.41-0.51 pH units at the beech stand (based on 2015-2016 means).
207 Nitrogen addition alone did not induce any significant changes in soil water pH. Consistent reductions
208 of mean annual DOC concentrations in response to pH manipulation were detected at both forest
209 stands, although these were not significant at the beech stand. DOC concentrations were reduced in
210 acid treatments by 30% - 39% at the spruce stand, and by 45% - 53% at the beech stand (based on
211 2015-2016 means). Nitrogen addition treatments did not lead to significant changes in forest floor soil
212 water DOC concentrations (Figure 1). Despite the alteration of soil water chemistry in the forest floor
213 due to treatments, soil water chemistry in mineral soil (30 cm depth) revealed only limited changes of
214 pH under acid treatments in the beech stand, and no significant changes in DOC concentrations in
215 mineral soil water in either forest stand (Suppl. Fig. 1).



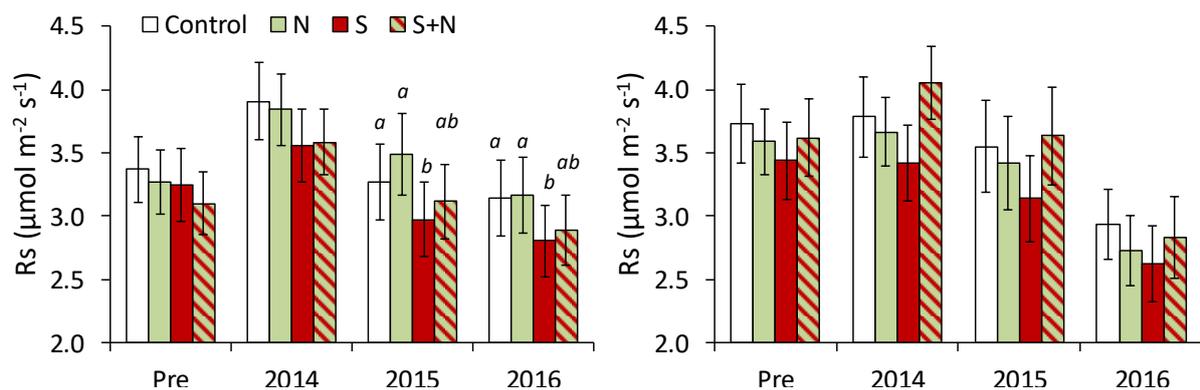
217

218 **Figure 1. Mean and standard error of forest floor soil water pH and DOC at each forest stand for control, nitrogen (N),**
 219 **sulphur (S) and sulphur + nitrogen (S+N) treatment plots (n = 4 per treatment and forest stand). Data are shown for the**
 220 **pre-treatment period (year 2013) and for the last three full years of manipulation. Significant treatment effects by site**
 221 **and time period denoted by letters (P < 0.05).**

222 3.2. Soil respiration

223 Significant exponential relationships existed between soil respiration and soil temperature thus soil
 224 temperature exerted a strong control over soil CO₂ efflux, with coefficients of determination (R²)
 225 ranging between 0.88 and 0.90 at the spruce stand and 0.79 and 0.83 at the beech stand (soil
 226 temperature ranged between 1.3 and 9.4°C during the respiration measurements). Annual mean soil
 227 respiration in the spruce stand (3.30 μmol CO₂ m⁻² s⁻¹) and beech stand (3.39 μmol CO₂ m⁻² s⁻¹) did not
 228 significantly differ. Rather high variability of the soil CO₂ efflux among individual collars, combined with
 229 strong (seasonal) temperature dependence made it difficult to detect treatment effects on total soil
 230 respiration. Nevertheless, consistent reductions of mean annual soil respiration were observed under
 231 the S only treatment in the spruce stand in 2015 and 2016 (p < 0.05), by 6 % - 10 %. A non-significant,

232 though consistent, reduction of R_s under combined treatment was measured in spruce stand. No
 233 significant effect of N addition on soil respiration was detected in the spruce stand. In the beech stand,
 234 neither acidity nor N manipulations affected soil respiration (Figure 2). Calculated mean annual C efflux
 235 was $9.9 \text{ t C ha}^{-1} \text{ year}^{-1}$ in the spruce stand and $9.7 \text{ t C ha}^{-1} \text{ year}^{-1}$ in the beech stand (based on 2013-
 236 2016 means for control plots; Eq. 1). Acid treatment in the spruce stand thus reduced C soil efflux by
 237 $0.53 - 0.95 \text{ t C ha}^{-1} \text{ year}^{-1}$.

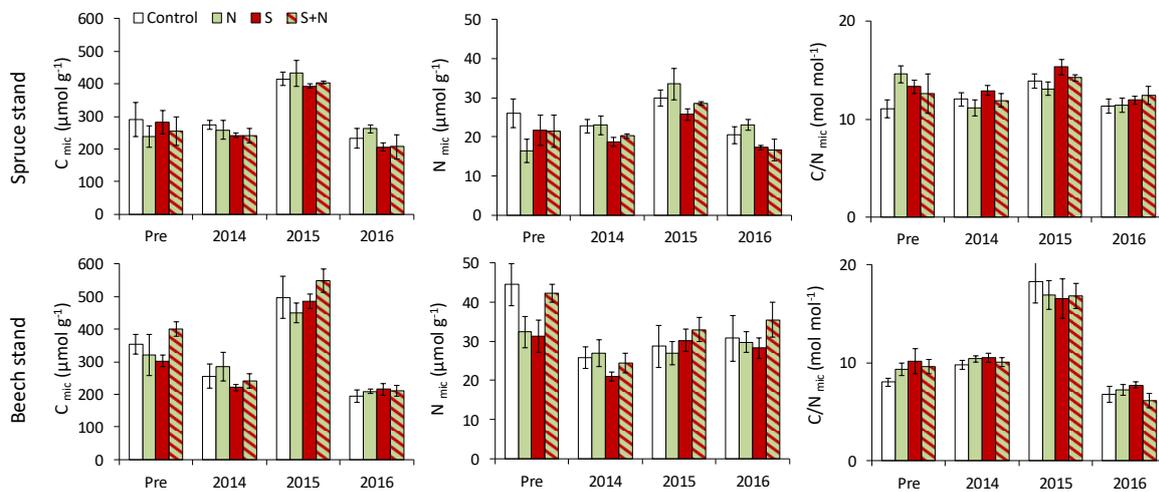


238
 239 **Figure 2.** Mean and standard error of soil respiration at each forest stand for control, nitrogen (N), sulphur (S) and
 240 sulphur + nitrogen (S+N) treatment plots ($n = 4$ per treatment and forest stand). Data are shown for the pre-treatment
 241 period (year 2013) and for the last three full years of manipulation. Significant treatment effects by site and time period
 242 denoted by letters ($P < 0.05$).

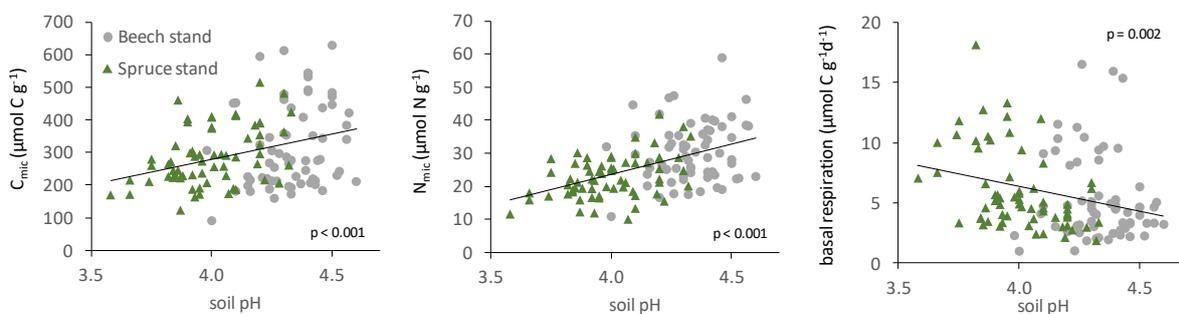
243 3.3. Microbial activity characteristics

244 Microbial biomass C, N and their ratios significantly ($P < 0.05$) differed among forest stands, with higher
 245 biomass C and N in the beech forest ($324 \mu\text{mol g}^{-1}$ and $31 \mu\text{mol g}^{-1}$, respectively) compared to the
 246 spruce forest ($290 \mu\text{mol g}^{-1}$ and $23 \mu\text{mol g}^{-1}$, respectively). Microbial C to N ratio was significantly lower
 247 in the beech stand (10.9) compared to the spruce stand (12.7) (Figure 3). Microbial respiration did not
 248 significantly differ between stands on average. Microbial biomass C and N were negatively related to
 249 soil pH across both forests (Figure 4), but in the spruce site acidic conditions led to higher respiration
 250 rates. No significant treatment effects were detected, although there was a tendency for microbial C
 251 and N to decrease under acid treatments (S, S+N) in the spruce forest, thus a shift toward higher
 252 microbial C/N was detected. Significantly ($P < 0.05$) higher fungal to bacterial ratios were measured in

253 the spruce stand (0.045) compared to the beech stand (0.031), consistently with the higher microbial
 254 C to N ratios. Despite no significant treatment effects on fungi/bacteria ratio in the spruce stand (Figure
 255 5), significantly lower bacterial gene copies were measured under acid treatments (Suppl. Table
 256 2) **Chyba! Nenalezen zdroj odkazů.** Therefore, consistent increases of fungi/bacteria ratio under acid
 257 treatments (Figure 5) appear to be attributable to decreases in bacterial abundance rather than
 258 increases in fungi. In the beech stand, a significant treatment effect on fungi/bacteria ratio was
 259 detected in 2016 (Figure 5). This resulted from decreases of bacteria and increases (albeit insignificant)
 260 in fungal gene copies (Suppl. Table 2).



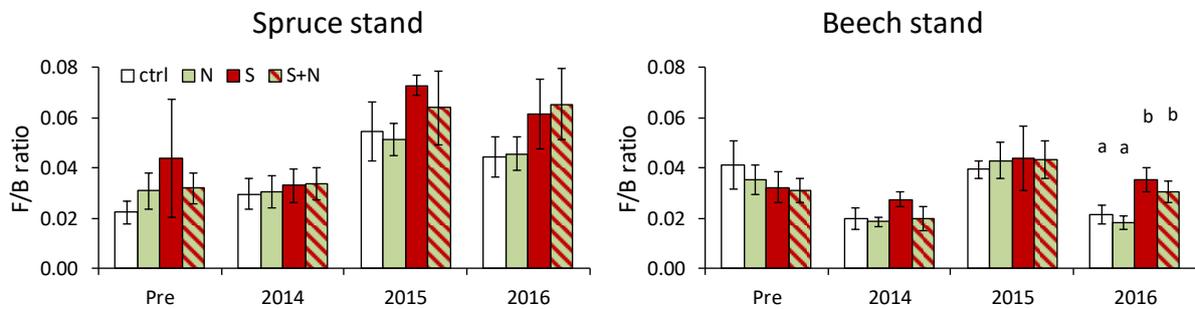
261
 262 **Figure 3. Mean and standard error of microbial C, N and C/N at each forest stand for control, nitrogen (N), sulphur (S) and**
 263 **sulphur + nitrogen (S+N) treatment plots (n = 4 per treatment and forest stand). Data are shown for the pre-treatment**
 264 **period (year 2013) and for the last three full years of manipulation. Significant treatment effects by site and time period**
 265 **denoted by letters (P < 0.05).**



266

267 **Figure 4. Soil microbial C, N and microbial respiration related to soil pH among forest stands. Linear regression and its**
 268 **significance (p value) highlighted for all data.**

269



270

271 **Figure 5. Mean and standard error of fungi to bacteria ratio at each forest stand for control, nitrogen (N), sulphur (S) and**
 272 **sulphur + nitrogen (S+N) treatment plots (n = 4 per treatment and forest stand). Data are shown for the pre-treatment**
 273 **period (year 2013) and for the last three full years of manipulation. Significant treatment effects were assessed by GLM**
 274 **analysis for each stand and time period separately.**

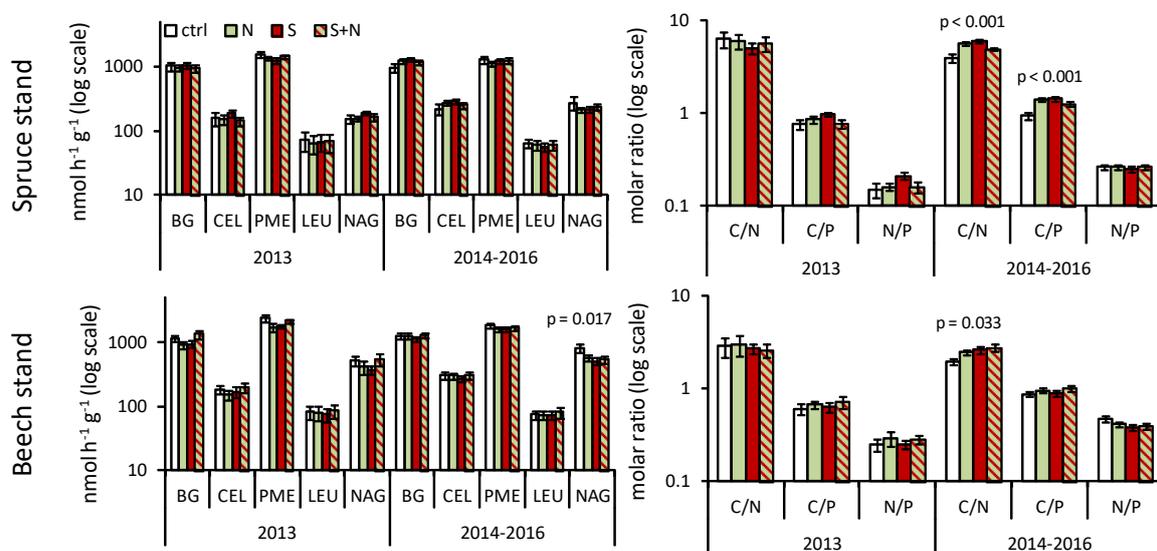
275 **3.4. Enzymatic activities**

276 Forest type significantly ($P < 0.05$) affected activity of all enzymes except these involved in C cycling (β -
 277 Glucosidase - BG and cellobiohydrolase – CEL). Consistently higher activities of enzymes primarily
 278 involved in N acquisition (leucine-aminopeptidase - LEU and N-acetyl- β -D-glucosaminidase - NAG)
 279 were measured in the beech stand compared to the spruce stand. Similarly, phosphatase (PME) activity
 280 was significantly higher in the beech stand compared to the spruce stand. Stoichiometry of enzymatic
 281 activity revealed significantly higher C_{enz}/N_{enz} and C_{enz}/P_{enz} ratios in the spruce stand compared to the
 282 beech stand ($C_{enz}/N_{enz} = 4.3$ in the spruce stand and 1.9 in the beech stand; $C_{enz}/P_{enz} = 1.0$ in the spruce
 283 stand and 0.8 in the beech stand). Conversely, a higher N_{enz}/P_{enz} ratio was measured in the beech stand
 284 compared to the spruce stand (0.43 vs 0.23) (based on control means, Figure 6).

285 Enzymatic activities among treatments did not significantly differ in the pre-treatment period. After
 286 treatment started, a tendency towards higher C enzyme activities under all treatments (N, S and S+N)
 287 was detected in the spruce stand (Suppl Table 1, Figure 6). These changes resulted in significant ($P <$

288 0.05) alteration of enzyme stoichiometry by treatments. C_{enz}/N_{enz} ratio was significantly higher under
 289 all treatments (mean of 5.3) compared to control (3.9), and C_{enz}/P_{enz} ratio significantly increased under
 290 all treatments (mean of 1.3) compared to control (1.0). No treatment effect on N_{enz}/P_{enz} ratio was
 291 detected. This suggests that treatments primarily affected C enzyme activities in the spruce stand. In
 292 the beech stand, a significant reduction of NAG activity was detected under all treatments. This
 293 resulted in a significant increase of C_{enz}/N_{enz} ratio under all treatments (mean of 2.4) compared to the
 294 control (1.8). The absence of observed change in the C_{enz}/P_{enz} ratio and a consistent (albeit insignificant,
 295 $P = 0.097$) change in N_{enz}/P_{enz} ratio suggested that treatment effects at the beech stand primarily
 296 affected N acquiring enzymes (Suppl Table 1, Figure 6). In the spruce stand, C_{enz} and N_{enz} were strongly
 297 negatively related to soil pH (Suppl. Fig. 2) while in beech no such dependence was found.

298



299

300 **Figure 6. Mean and standard error of enzyme activities and their ratios at each forest stand for control, nitrogen (N),**
 301 **sulphur (S) and sulphur + nitrogen (S+N) treatment plots (n = 4 per treatment and forest stand). Data are shown for the**
 302 **pre-treatment period (year 2013) and for the last three full years of manipulation. Significant treatment effects by site**
 303 **and time period denoted by p value. We calculated enzyme C/N, C/P and N/P ratios, where C = BG+CEL, N = NAG+LEU**
 304 **and P = PME.**

305

306 4. Discussion

307 Our results provide evidence that soil acidification alters C fluxes in both deciduous and conifer forests.
308 Dissolved organic carbon concentrations in soil solution were suppressed by acidification in both forest
309 types, and suppression of soil respiration was detected in the spruce stand. These responses have
310 significant implications for long-term soil C cycling. On the other hand, nitrogen addition alone did not
311 yet affect either DOC concentrations or soil CO₂ efflux in either forest type.

312 4.1. *Acidification alters dissolved organic carbon leaching*

313 Monthly addition of sulphuric acid immediately reduced forest floor soil solution pH by 0.4 units in
314 both stands. Due to higher initial acidity in the spruce stand, this equated to a higher post-treatment
315 hydrogen ion concentration of 270 µeq l⁻¹ in the spruce stand, compared to 130 µeq l⁻¹ in the beech
316 stand. This “acid forcing” induced a consistent decline in soil solution DOC of around 50 %. In line with
317 other lab and field experimental pH manipulations (Clark et al., 2011; Ekstrom et al., 2011; Evans et al.,
318 2012); analyses of long-term monitoring data (Finstad et al., 2016; Monteith et al., 2007); and
319 modelling assessments (Rowe et al., 2014; Sawicka et al., 2017), we therefore found clear and
320 quantitatively significant effects of soil solution acidity on soil solution DOC concentration. Albeit
321 different litter quality the effect of acid treatments on forest floor soil solution DOC concentration was
322 consistent at both sites, suggesting a prevailing chemical (rather than biotic) control over DOC
323 partitioning between soil and soil solution (Oulehle et al., 2013). Whether these effects were mediated
324 by altered dissociation of organic acids in response to changing acidity, or direct effect of changed ionic
325 strength (Hruska et al., 2009) remains unresolved. The neutralization capacity of forest mineral soils
326 at both stands was highlighted by the delayed soil solution pH decrease in deeper depth. This
327 ultimately led to a progressive acidification of mineral soil, due to the leaching of base cations
328 accompanied predominantly by sulphate anion.

329 The absence of soil solution pH and DOC responses to N addition in our experiment is consistent with
330 the interpretation that chemical processes control DOC mobility. The effect of NH₄NO₃ addition on soil

331 acidity depends on the fate of NH_4 and NO_3 ; If both NO_3 and NH_4 are taken up by the biota no
332 acidification occurs (regardless of whether nitrification has occurred or not) but if NO_3^- is leached in
333 excess of any leaching of NH_4^+ (due to preferential uptake and/or nitrification of NH_4^+) this will cause
334 acidification, and may suppress DOC leaching (Evans et al., 2008). Contrary to previously suggested
335 positive effects of N deposition on DOC leaching (e.g. Pregitzer *et al.*, 2004) we did not detect any
336 substantial DOC increase following N addition in our systems. However, long-term N inputs may
337 stimulate primary productivity (Liu and Greaver, 2010), and thus ultimately increase the supply of
338 ecosystem C available for leaching as DOC (Rowe et al., 2014), provided that other feedbacks such as
339 changes in plant species and litter quality/quantity do not offset the increase in NPP. In forests
340 subjected to long-term elevated N deposition, as in our case, response of soil solution DOC
341 concentrations may differ from N limiting systems where production of DOC may be more responsive
342 to increased N availability.

343 **4.2. Acidification effects on soil carbon dynamics**

344 Measured total soil respiration (R_s) rates were closely related to the seasonal course of soil
345 temperature, and soil temperature was the overriding control on temporal variation in R_s , with limited
346 effects of soil moisture (Oulehle et al., 2016b). A large part of R_s in forests is derived from autotrophic
347 respiration (Högberg et al., 2001), thus detection of subtle changes in soil CO_2 effluxes, driven by
348 heterotrophic decomposition of fresh litter and SOM, is difficult. Nevertheless, consistent declines in
349 R_s were detected in the spruce stand under acid treatment. The measured respiration decline of 6% -
350 10% under S treatment is similar to that found in a recent meta-analysis (7%) based on experiments
351 from China (Feng et al., 2017). The reduction of CO_2 effluxes under S treatment corresponds to the
352 27% – 49% of annual litter production (aboveground litterfall of $1937 \text{ kg C ha}^{-1} \text{ year}^{-1}$ between 2013-
353 2016, Oulehle *et al.*, 2016b). Based on a total S addition of $50 \text{ kg S ha}^{-1} \text{ year}^{-1}$ this corresponds to the
354 suppression of decomposition by 10.6 to $19 \text{ kg C kg S}^{-1}$. The reduction in decomposition per unit S
355 deposition may not be linear, however, as this would require linear relationships between S input and
356 soil water hydrogen ion concentration, and subsequently with R_s . In a previous study (Oulehle et al.,

2011) we demonstrated a continuous loss of forest floor C pool, at a rate of 1160 kg C ha⁻¹ year⁻¹, since precipitation acidity declined in the early 1990s as a consequence of “acid rain” abatement policy (total loss of 18.6 t C ha⁻¹ between 1994 and 2010). We proposed that recent loss of soil C carbon was derived from the soil C accumulated during the acidification period. Based on our calculations with the lower estimate of S deposition effect on C decomposition (10.6 kg C ha⁻¹ S ha⁻¹), a C pool of roughly 20 t C ha⁻¹ could have been accumulated between the 1960s and 1990s (the time period during which precipitation acidity fell below 4.3; Kopáček *et al.*, 2016). Our experiment thus provides a plausible explanation for observed rates of C accumulation in forest soils exposed to acidification stress and could account for major changes in forest carbon stocks in the forests of Central Europe and other acidification-impacted regions.

Acid treatments in our experiment induced decreases in soil solution pH and DOC concentration in both forests, but only appeared to suppress Rs in the spruce stand. The two forest ecosystems can be characterized by: i) higher soil pH and base saturation in the beech stand; 2) higher substrate concentration and enzyme C to N ratios in the spruce stand; 3) an apparently higher degree of interactions among the microbial community in the spruce stand (Bárta *et al.*, 2017); and 4) contrasting fungal life strategies, with saprotrophic fungi more abundant in the beech soil, and ectomycorrhizal fungi more abundant in the spruce soil (Bárta *et al.*, 2017). Taken together, these features suggest a much stronger connection of soil metabolism to C acquisition in the spruce stand compared to the beech stand. Consequently, a decrease of soil solution pH and following suppression of labile C availability would be expected to have a much stronger impact on SOM decomposition in the spruce stand compared to the beech stand (Klotzbücher *et al.*, 2011). We found a strong positive relation between soil pH and microbial biomass and high microbial respiration and C mining activity under low pH in the spruce stand. It coincided with a consistent (though non-significant, Figure 3) decrease in microbial biomass following acid treatments in the spruce stand (roughly 11 % based on 2016 data) and observed decreases in Rs. These results support our conclusion that the microbial (or at least prokaryotic) community in the spruce stand was most probably more stressed due to more acidic

383 conditions (from the beginning of the experiment, and intensified by S treatments) than the beech
384 community. The only acid treatment effects which were consistent across both forest types were
385 significant changes in enzyme stoichiometry, and partially significant changes in fungi to bacteria
386 ratios. Acid treatments significantly changed the C_{enz} to N_{enz} ratio, with increases following acid
387 treatments in both stands. However, in the spruce stand the increase in C_{enz}/N_{enz} ratio appeared to be
388 driven by increase of C_{enz} activities, which was further supported by the negative relationship between
389 soil pH and C_{enz} activity. In addition, C_{enz}/P_{enz} ratio increased in response to acid treatments in spruce
390 stand, with no change in the beech stand; this supports the assumption of predominantly enhanced C
391 enzyme activities in the spruce forest (Figure 6). More specific studies are needed to unravel the
392 feedback between altered soil pH and microbial biomass and enzyme activities. It is also important to
393 link actual qualitative changes among microbial communities caused by altered soil chemistry with
394 corresponding C fluxes.

395 It has been proposed that microbes adjust their enzymatic activity to gain the limiting element from
396 the substrate that is rich in this element and that the production of enzymes increases with substrate
397 complexity (Mooshammer et al., 2014). As shown previously, acidity affects the solubility and
398 bioavailability of organic matter (Evans et al., 2012; Scheel et al., 2007). We infer that shifts towards
399 enzymatic C mining after acidification were most probably connected with organic matter stabilization
400 and thus reduced availability to microbes, causing them to produce more C-acquiring enzymes to
401 sustain C supply. This interpretation is also supported by observed negative relationships between BG
402 and CEL and soil pH in the spruce stand and found shifts to fungal prevalence in both stands.

403 Based on our data, qualitative changes in the microbial community under acid treatments can be
404 attributed to decreases of bacteria in both forests, and to an increase of fungi in the beech forest,
405 resulting in higher fungi to bacteria ratios. According to the literature, bacteria are generally more
406 impacted by low pH than fungi (Rousk et al., 2009; Strickland and Rousk, 2010). In addition, fungi on
407 average have higher C to N biomass stoichiometry, slower growth and turnover (Rousk and Bååth,

408 2011) and probably also higher carbon use efficiency than bacteria (Allison et al., 2005; Six et al., 2006).
409 It means that they can store more C in their biomass per unit of consumed C than bacteria. Thus, the
410 observed shifts in F/B ratio (either due to bacterial decreases and/or fungal increases) might partly
411 contribute to decreases of soil microbial respiration in the field under acid treatments in both stands.

412 **4.3. N addition effects on soil organic matter dynamics**

413 In contrast to the rather limited number of experimental studies of soil acidity effects on SOM cycling
414 in natural or semi natural ecosystems (Chen et al., 2016; Pennanen et al., 1998; Persson et al., 1989;
415 Wu et al., 2016), many studies have linked enhanced soil N availability to the soil C cycle (e.g. Olsson
416 *et al.*, 2005; Hobbie, 2008; Janssens *et al.*, 2010; Liu & Greaver, 2010). In general, meta-analyses of N
417 addition experiments have shown stimulation of biomass productivity or reduction of soil respiration,
418 leading to potentially increased C sequestration. Although ¹⁵N tracing studies have suggested only
419 minor N deposition contributions to C sequestration in temperate forests (Nadelhoffer et al., 1999),
420 many field-scale N addition studies have suggested that N deposition could substantially enhance soil
421 carbon sequestration by both forests and heathlands (Field et al., 2017; Prescott, 2010). Our plot
422 experimental design did not allow testing of N addition effects on plant biomass productivity, therefore
423 allocation of C to belowground was assumed constant over the duration of our experiment. Based on
424 our data we did not see any significant effects of N addition on Rs (Figure 2), microbial biomass (Figure
425 3) and respiration (data not shown). There was no evidence of N addition effects on fungi to bacteria
426 ratios (Figure 5), but we did observed shifts in enzyme C to N stoichiometry (Figure 6), similar to those
427 in the acid treatments, suggesting ongoing qualitative changes in microbial processes yet not
428 detectable on a quantitative basis.

429 The lack of pronounced effects of N addition on the C cycle at our sites may be attributable to the high
430 pre-existing levels of available N in these systems; both forests have been subjected to elevated N
431 deposition since the 1950s, peaking in the 1980s with bulk deposition estimates of ca. 17 kg ha⁻¹ year⁻¹
432 ¹ (Oulehle *et al.* 2016b; see

433 for current N fluxes). Microbial communities may therefore already be adapted to increased N
434 availability at our sites, whereas systems receiving lower levels of ambient N deposition may be more
435 responsive to additional N inputs (Phoenix et al., 2012). Nevertheless many other temperate forested
436 ecosystems, especially those in Europe and North America, have been exposed to high levels of historic
437 N deposition, and our data suggest that heterotrophic processes in these systems may be relatively
438 unresponsive to additional N inputs. Forest ecosystems in parts of the world (apart of P-limited tropical
439 forests) where N deposition is currently increasing, such as parts of Asia, Africa and South America
440 (Dentener et al., 2006; Lamarque et al., 2013), but where fertility (i.e. soil C/N ratio) remains low may
441 be expected to show greater change, potentially including effects on ecosystem C sequestration
442 (Hyvönen et al., 2008).

443 With regard to dissolved carbon losses (and to a lesser extent also gaseous C fluxes) it is also possible
444 that the absence of responses to N addition also reflects the lack of resulting changes in soil acidity.
445 Previous work has shown that changes in DOC leaching in many N addition studies can be explained
446 by experimentally-induced shifts in acid-base chemistry (Evans et al., 2008), but in our study the
447 addition of NH_4NO_3 over 3 years was insufficient to alter soil solution pH, because little of the added N
448 was leached as NO_3 (ca. 6% of added N in the spruce forest floor and ca. 11% of added N in the beech
449 forest floor, data not shown). Thus the clear, chemically-mediated changes in DOC concentration
450 observed in the acidification treatments were not replicated in the N-only treatment.

451

452 *4.4. Implications for understanding of ecosystem responses to multiple drivers*

453

454 A high proportion of biogeochemical research to date has taken place within areas of Europe and North
455 America that have, as a result of industrialisation and agricultural intensification, been exposed to
456 multiple environmental stressors, including the simultaneous deposition of a range of atmospheric
457 pollutants. Nevertheless, experiments carried out within these ecosystems often focus on a single

458 driver, whilst many analyses of monitoring data have considered only a partial range of potential
459 driving variables. Studies of forest carbon sequestration in response to atmospheric deposition have
460 overwhelmingly focused on N deposition, yet our results suggest a potentially similar rate of soil C
461 accumulation in response to S deposition ($> 10 \text{ kg C kg S}^{-1}$) as that estimated for N deposition ($10 - 15$
462 kg C kg N^{-1} ; de Vries et al. 2009). Furthermore, S deposition to European and North American forests
463 has changed more rapidly and more dramatically over the last 50 years than N deposition, implying
464 that S may have had a greater influence on the soil C cycle during this period, and also raising the
465 possibility that some observed changes may have been incorrectly attributed to N deposition (Kolář et
466 al., 2015). In order to correctly interpret existing records, and to accurately predict future changes in
467 ecosystem biogeochemistry in response to continuing environmental change, we argue that a more
468 holistic approach to the impacts of multiple environmental drivers is needed.

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- 476 Allison, V.J., Miller, R.M., Jastrow, J.D., Matamala, R., Zak, D.R., 2005. Changes in Soil
477 Microbial Community Structure in a Tallgrass Prairie Chronosequence. *Soil Sci. Soc. Am.*
478 *J.* 69, 1412. <https://doi.org/10.2136/sssaj2004.0252>
- 479 Bárta, J., Šlajsová, P., Tahovská, K., Pícek, T., Šantrůčková, H., 2014. Different temperature
480 sensitivity and kinetics of soil enzymes indicate seasonal shifts in C, N and P nutrient
481 stoichiometry in acid forest soil. *Biogeochemistry* 117, 525–537.
482 <https://doi.org/10.1007/s10533-013-9898-1>
- 483 Bárta, J., Tahovská, K., Šantrůčková, H., Oulehle, F., 2017. Microbial communities with
484 distinct denitrification potential in spruce and beech soils differing in nitrate leaching.
485 *Sci. Rep.* 7, 9738. <https://doi.org/10.1038/s41598-017-08554-1>
- 486 Berg, B., Matzner, E., 1997. Effect of N deposition on decomposition of plant litter and soil
487 organicmatter in forest systems. *Environ. Rev.* 5, 1–25.
- 488 Błońska, E., Lasota, J., Gruba, P., 2017. Enzymatic activity and stabilization of organic matter
489 in soil with different detritus inputs 63, 242–247.
490 <https://doi.org/10.1080/00380768.2017.1326281>
- 491 Błońska, E., Lasota, J., Gruba, P., 2016. Effect of temperate forest tree species on soil
492 dehydrogenase and urease activities in relation to other properties of soil derived from
493 loess and glaciofluvial sand. *Ecol. Res.* 31, 655–664. [https://doi.org/10.1007/s11284-](https://doi.org/10.1007/s11284-016-1375-6)
494 [016-1375-6](https://doi.org/10.1007/s11284-016-1375-6)
- 495 Brookes, P.C., Landman, A., Pruden, G., Jenkinson, D.S., 1985. Chloroform fumigation and
496 the release of soil nitrogen: A rapid direct extraction method to measure microbial
497 biomass nitrogen in soil. *Soil Biol. Biochem.* 17, 837–842. [https://doi.org/10.1016/0038-](https://doi.org/10.1016/0038-0717(85)90144-0)
498 [0717\(85\)90144-0](https://doi.org/10.1016/0038-0717(85)90144-0)
- 499 Butchart, S.H.M., Walpole, M., Collen, B., van Strien, A., Scharlemann, J.P.W., Almond, R.E.A.,
500 Baillie, J.E.M., Bomhard, B., Brown, C., Bruno, J., Carpenter, K.E., Carr, G.M., Chanson, J.,
501 Chenery, A.M., Csirke, J., Davidson, N.C., Dentener, F., Foster, M., Galli, A., Galloway,
502 J.N., Genovesi, P., Gregory, R.D., Hockings, M., Kapos, V., Lamarque, J.-F., Leverington,
503 F., Loh, J., McGeoch, M.A., McRae, L., Minasyan, A., Morcillo, M.H., Oldfield, T.E.E.,
504 Pauly, D., Quader, S., Revenga, C., Sauer, J.R., Skolnik, B., Spear, D., Stanwell-Smith, D.,
505 Stuart, S.N., Symes, A., Tierney, M., Tyrrell, T.D., Vie, J.-C., Watson, R., 2010. Global
506 Biodiversity: Indicators of Recent Declines. *Science* (80-.). 328, 1164–1168.
507 <https://doi.org/10.1126/science.1187512>
- 508 Clark, J.M., van der Heijden, G.M.F., Palmer, S.M., Chapman, P.J., Bottrell, S.H., 2011.
509 Variation in the sensitivity of DOC release between different organic soils following
510 H₂SO₄ and sea-salt additions. *Eur. J. Soil Sci.* 62, 267–284.
511 <https://doi.org/10.1111/j.1365-2389.2010.01344.x>
- 512 Dambrine, E., Kinkor, V., Jehlicka, J., Gelhaye, D., 1993. Fluxes of dissolved mineral elements
513 through a forest ecosystem submitted to extremely high atmospheric pollution inputs
514 (Czech Republic). *Ann. des Sci. For.* 50, 147–157.
515 <https://doi.org/10.1051/forest:19930203>
- 516 Davidson, E.A., Janssens, I.A., Luo, Y., 2006. On the variability of respiration in terrestrial
517 ecosystems: moving beyond Q₁₀. *Glob. Chang. Biol.* 12, 154–164.
518 <https://doi.org/10.1111/j.1365-2486.2005.01065.x>

519 de Vries, W., Solberg, S., Dobbertin, M., Sterba, H., Laubhann, D., van Oijen, M., Evans, C.,
520 Gundersen, P., Kros, J., Wamelink, G.W.W., Reinds, G.J., Sutton, M.A., 2009. The impact
521 of nitrogen deposition on carbon sequestration by European forests and heathlands.
522 *For. Ecol. Manage.* 258, 1814–1823. <https://doi.org/10.1016/j.foreco.2009.02.034>
523 Dentener, F., Drevet, J., Lamarque, J.F., Bey, I., Eickhout, B., Fiore, A.M., Hauglustaine, D.,
524 Horowitz, L.W., Krol, M., Kulshrestha, U.C., Lawrence, M., Galy-Lacaux, C., Rast, S.,
525 Shindell, D., Stevenson, D., Van Noije, T., Atherton, C., Bell, N., Bergman, D., Butler, T.,
526 Cofala, J., Collins, B., Doherty, R., Ellingsen, K., Galloway, J., Gauss, M., Montanaro, V.,
527 Müller, J.F., Pitari, G., Rodriguez, J., Sanderson, M., Solmon, F., Strahan, S., Schultz, M.,
528 Sudo, K., Szopa, S., Wild, O., 2006. Nitrogen and sulfur deposition on regional and global
529 scales: A multimodel evaluation 20. <https://doi.org/10.1029/2005GB002672>
530 Duan, L., Yu, Q., Zhang, Q., Wang, Z., Pan, Y., Larssen, T., Tang, J., Mulder, J., 2016. Acid
531 deposition in Asia: Emissions, deposition, and ecosystem effects. *Atmos. Environ.* 146,
532 55–69. <https://doi.org/10.1016/j.atmosenv.2016.07.018>
533 Ekstrom, S.M., Kritzberg, E.S., Kleja, D.B., Larsson, N., Nilsson, P.A., Graneli, W., Bergkvist, B.,
534 2011. Effect of Acid Deposition on Quantity and Quality of Dissolved Organic Matter in
535 Soil-Water. *Environ. Sci. Technol.* 45, 4733–4739. <https://doi.org/10.1021/Es104126f>
536 Evans, C., Goodale, C., Caporn, S., Dise, N., Emmett, B., Fernandez, I., Field, C., Findlay, S.,
537 Lovett, G., Meesenburg, H., Moldan, F., Sheppard, L., 2008. Does elevated nitrogen
538 deposition or ecosystem recovery from acidification drive increased dissolved organic
539 carbon loss from upland soil? A review of evidence from field nitrogen addition
540 experiments. *Biogeochemistry* 91, 13–35. <https://doi.org/10.1007/s10533-008-9256-x>
541 Evans, C.D., Jones, T.G., Burden, A., Ostle, N., Zieliński, P., Cooper, M.D.A., Peacock, M.,
542 Clark, J.M., Oulehle, F., Cooper, D., Freeman, C., 2012. Acidity controls on dissolved
543 organic carbon mobility in organic soils. *Glob. Chang. Biol.* 18, 3317–3331.
544 <https://doi.org/10.1111/j.1365-2486.2012.02794.x>
545 Feng, J., Wang, J., Ding, L., Yao, P., Qiao, M., Yao, S., 2017. Meta-analyses of the effects of
546 major global change drivers on soil respiration across China. *Atmos. Environ.* 150, 181–
547 186. <https://doi.org/10.1016/j.atmosenv.2016.11.060>
548 Field, C.D., Evans, C.D., Dise, N.B., Hall, J.R., Caporn, S.J.M., 2017. Long-term nitrogen
549 deposition increases heathland carbon sequestration. *Sci. Total Environ.* 592, 426–435.
550 <https://doi.org/10.1016/j.scitotenv.2017.03.059>
551 Finstad, A.G., Andersen, T., Larsen, S., Tominaga, K., Blumentrath, S., de Wit, H.A.,
552 Tømmervik, H., Hessen, D.O., 2016. From greening to browning: Catchment vegetation
553 development and reduced S-deposition promote organic carbon load on decadal time
554 scales in Nordic lakes. *Sci. Rep.* 6, 31944. <https://doi.org/10.1038/srep31944>
555 Gittel, A., Bárta, J., Kohoutová, I., Mikutta, R., Owens, S., Gilbert, J., Schneckner, J., Wild, B.,
556 Hannisdal, B., Maerz, J., Lashchinskiy, N., Čapek, P., Šantrůčková, H., Gentsch, N.,
557 Shibistova, O., Guggenberger, G., Richter, A., Torsvik, V.L., Schleper, C., Urich, T., 2014.
558 Distinct microbial communities associated with buried soils in the Siberian tundra. *ISME*
559 *J.* 8, 841–853. <https://doi.org/10.1038/ismej.2013.219>
560 Hobbie, S.E., 2008. Nitrogen effects on decomposition: A five-year experiment in eight
561 temperate sites. *Ecology* 89, 2633–2644.
562 Högberg, P., Nordgren, A., Buchmann, N., Taylor, A.F.S., Ekblad, A., Högberg, M.N., Nyberg,
563 G., Ottosson-Löfvenius, M., Read, D.J., 2001. Large-scale forest girdling shows that
564 photosynthesis drives soil respiration. *Nature* 411, 749–752.
565

566 Hruska, J., Kram, P., McDowell, W.H., Oulehle, F., Hruška, J., Krám, P., 2009. Increased
567 Dissolved Organic Carbon (DOC) in Central European Streams is Driven by Reductions in
568 Ionic Strength Rather than Climate Change or Decreasing Acidity. *Environ. Sci. Technol.*
569 43, 4320–4326. <https://doi.org/10.1021/es803645w>

570 Hyvönen, R., Ågren, G.I., Linder, S., Persson, T., Cotrufo, M.F., Ekblad, A., Freeman, M.,
571 Grelle, A., Janssens, I.A., Jarvis, P.G., Kellomäki, S., Lindroth, A., Loustau, D., Lundmark,
572 T., Norby, R.J., Oren, R., Pilegaard, K., Ryan, M.G., Sigurdsson, B.D., Strömberg, M., van
573 Oijen, M., Wallin, G., 2007. The likely impact of elevated [CO₂], nitrogen deposition,
574 increased temperature and management on carbon sequestration in temperate and
575 boreal forest ecosystems: a literature review. *New Phytol.* 173, 463–480.
576 <https://doi.org/10.1111/j.1469-8137.2007.01967.x>

577 Hyvönen, R., Persson, T., Andersson, S., Olsson, B., Ågren, G.I., Linder, S., Hyvonen, R., 2008.
578 Impact of long-term nitrogen addition on carbon stocks in trees and soils in northern
579 Europe. *Biogeochemistry* 89, 121–137. <https://doi.org/10.1007/s10533-007-9121-3>

580 Chen, D., Li, J., Lan, Z., Hu, S., Bai, Y., 2016. Soil acidification exerts a greater control on soil
581 respiration than soil nitrogen availability in grasslands subjected to long-term nitrogen
582 enrichment. *Funct. Ecol.* 30, 658–669. <https://doi.org/10.1111/1365-2435.12525>

583 Chen, D., Wang, Y., Lan, Z., Li, J., Xing, W., Hu, S., Bai, Y., 2015. Biotic community shifts
584 explain the contrasting responses of microbial and root respiration to experimental soil
585 acidification. *Soil Biol. Biochem.* 90, 139–147.
586 <https://doi.org/10.1016/j.soilbio.2015.08.009>

587 Janssens, I.A., Dieleman, W., Luysaert, S., Subke, J.A., Reichstein, M., Ceulemans, R., Ciais,
588 P., Dolman, A.J., Grace, J., Matteucci, G., Papale, D., Piao, S.L., Schulze, E.D., Tang, J.,
589 Law, B.E., 2010. Reduction of forest soil respiration in response to nitrogen deposition.
590 *Nat. Geosci.* 3, 315–322. <https://doi.org/10.1038/ngeo844>

591 Klotzbücher, T., Kaiser, K., Guggenberger, G., Gatzek, C., Kalbitz, K., 2011. A new conceptual
592 model for the fate of lignin in decomposing plant litter. *Ecology* 92, 1052–1062.
593 <https://doi.org/10.1890/i0012-9658-92-5-1052>

594 Knorr, M., Frey, S.D., Curtis, P.S., 2005. Nitrogen additions and litter decomposition: A meta-
595 analysis. *Ecology* 86, 3252–3257. <https://doi.org/10.1890/05-0150>

596 Kolář, T., Čermák, P., Oulehle, F., Trnka, M., Štěpánek, P., Cudlín, P., Hruška, J., Büntgen, U.,
597 Rybníček, M., 2015. Pollution control enhanced spruce growth in the “Black Triangle”
598 near the Czech-Polish border. *Sci. Total Environ.* 538, 703–11.
599 <https://doi.org/10.1016/j.scitotenv.2015.08.105>

600 Kopáček, J., Hejzlar, J., Krám, P., Oulehle, F., Posch, M., 2016. Effect of industrial dust on
601 precipitation chemistry in the Czech Republic (Central Europe) from 1850 to 2013.
602 *Water Res.* 103, 30–37. <https://doi.org/10.1016/j.watres.2016.07.017>

603 Kopáček, J., Posch, M., 2011. Anthropogenic nitrogen emissions during the Holocene and
604 their possible effects on remote ecosystems. *Global Biogeochem. Cycles* 25, GB2017.
605 <https://doi.org/10.1029/2010GB003779>

606 Lamarque, J.F., Dentener, F., McConnell, J., Ro, C.U., Shaw, M., Vet, R., Bergmann, D.,
607 Cameron-Smith, P., Dalsoren, S., Doherty, R., Faluvegi, G., Ghan, S.J., Josse, B., Lee, Y.H.,
608 Mackenzie, I.A., Plummer, D., Shindell, D.T., Skeie, R.B., Stevenson, D.S., Strode, S.,
609 Zeng, G., Curran, M., Dahl-Jensen, D., Das, S., Fritzsche, D., Nolan, M., 2013. Multi-
610 model mean nitrogen and sulfur deposition from the atmospheric chemistry and
611 climate model intercomparison project (ACCMIP): Evaluation of historical and projected
612 future changes 13, 7997–8018. <https://doi.org/10.5194/acp-13-7997-2013>

613 Lamersdorf, N.P., Borken, W., 2004. Clean rain promotes fine root growth and soil
614 respiration in a Norway spruce forest. *Glob. Chang. Biol.* 10, 1351–1362.
615 <https://doi.org/10.1111/j.1365-2486.2004.00811.x>

616 Larssen, T., Lydersen, E., Tang, D., He, Y., Gao, J., Liu, H., Duan, L., Seip, H.M., Vogt, R.D.,
617 Mulder, J., Shao, M., Wang, Y., Shang, H., Zhang, X., Solberg, S., Aas, W., Økland, T.,
618 Eilertsen, O., Angell, V., Liu, Q., Zhao, D., Xiang, R., Xiao, J., Luo, J., 2006. Acid Rain in
619 China. *Environ. Sci. Technol.* 40, 418–425.

620 Liang, G., Liu, X., Chen, X., Qiu, Q., Zhang, D., Chu, G., Liu, J., Liu, S., Zhou, G., 2013. Response
621 of Soil Respiration to Acid Rain in Forests of Different Maturity in Southern China. *PLoS*
622 *One* 8, e62207. <https://doi.org/10.1371/journal.pone.0062207>

623 Liu, L.L., Greaver, T.L., 2010. A global perspective on belowground carbon dynamics under
624 nitrogen enrichment. *Ecol. Lett.* 13, 819–828. [https://doi.org/10.1111/j.1461-](https://doi.org/10.1111/j.1461-0248.2010.01482.x)
625 [0248.2010.01482.x](https://doi.org/10.1111/j.1461-0248.2010.01482.x)

626 Monteith, D.T., Stoddard, J.L., Evans, C.D., de Wit, H.A., Forsius, M., Hogasen, T., Wilander,
627 A., Skjelkvale, B.L., Jeffries, D.S., Vuorenmaa, J., Keller, B., Kopacek, J., Vesely, J.,
628 Høgåsen, T., Skjelkvåle, B.L., Kopáček, J., 2007. Dissolved organic carbon trends
629 resulting from changes in atmospheric deposition chemistry. *Nature* 450, 537–U9.
630 <https://doi.org/10.1038/Nature06316>

631 Mooshammer, M., Wanek, W., Zechmeister-Boltenstern, S., Richter, A., 2014. Stoichiometric
632 imbalances between terrestrial decomposer communities and their resources:
633 mechanisms and implications of microbial adaptations to their resources. *Front.*
634 *Microbiol.* 5, 22. <https://doi.org/10.3389/fmicb.2014.00022>

635 Nadelhoffer, K.J., Emmett, B.A., Gundersen, P., Kjonaas, O.J., Koopmans, C.J., Schleppei, P.,
636 Tietema, A., Wright, R.F., Kjønaas, O.J., 1999. Nitrogen deposition makes a minor
637 contribution to carbon sequestration in temperate forests. *Nature* 398, 145–148.

638 Olsson, P., Linder, S., Giesler, R., Hogberg, P., 2005. Fertilization of boreal forest reduces
639 both autotrophic and heterotrophic soil respiration. *Glob. Chang. Biol.* 11, 1745–1753.
640 <https://doi.org/10.1111/j.1365-2486.2005.001033.x>

641 Oulehle, F., Evans, C.D., Hofmeister, J., Krejci, R., Tahovska, K., Persson, T., Cudlin, P., Hruska,
642 J., 2011. Major changes in forest carbon and nitrogen cycling caused by declining
643 sulphur deposition. *Glob. Chang. Biol.* 17, 3115–3129. [https://doi.org/10.1111/j.1365-](https://doi.org/10.1111/j.1365-2486.2011.02468.x)
644 [2486.2011.02468.x](https://doi.org/10.1111/j.1365-2486.2011.02468.x)

645 Oulehle, F., Hofmeister, J., Cudlin, P., Hruska, J., Cudlín, P., Hruška, J., 2006. The effect of
646 reduced atmospheric deposition on soil and soil solution chemistry at a site subjected
647 to long-term acidification, Nacetin, Czech Republic. *Sci. Total Environ.* 370, 532–544.
648 <https://doi.org/10.1016/j.scitotenv.2006.07.031>

649 Oulehle, F., Kopáček, J., Chuman, T., Černohous, V., Hůnová, I., Hruška, J., Krám, P.,
650 Lachmanová, Z., Navrátil, T., Štěpánek, P., Tesař, M., Evans, C.D., 2016a. Predicting
651 sulphur and nitrogen deposition using a simple statistical method. *Atmos. Environ.* 140,
652 456–468. <https://doi.org/10.1016/j.atmosenv.2016.06.028>

653 Oulehle, F., Růžek, M., Tahovská, K., Bárta, J., Myška, O., 2016b. Carbon and Nitrogen Pools
654 and Fluxes in Adjacent Mature Norway Spruce and European Beech Forests. *Forests* 7,
655 282. <https://doi.org/10.3390/f7110282>

656 Pennanen, T., Fritze, H., Vanhala, P., Kiikkila, O., Neuvonen, S., Baath, E., 1998. Structure of a
657 microbial community in soil after prolonged addition of low levels of simulated acid
658 rain. *Appl. Environ. Microbiol.* 64, 2173–2180.

659 Persson, T., Lundkvist, H., Wiren, A., Hyvonen, R., Wessen, B., 1989. Effects of acidification

707 <https://doi.org/10.1016/j.soilbio.2010.05.007>
708 Thomas, R.B., Spal, S.E., Smith, K.R., Nippert, J.B., 2013. Evidence of recovery of *Juniperus*
709 *virginiana* trees from sulfur pollution after the Clean Air Act. *Proc. Natl. Acad. Sci.* 110,
710 15319–15324. <https://doi.org/10.1073/pnas.1308115110>
711 Vance, E.D., Brookes, P.C., Jenkinson, D.S., 1987. An extraction method for measuring soil
712 microbial biomass C. *Soil Biol. Biochem.* 19, 703–707. [https://doi.org/10.1016/0038-](https://doi.org/10.1016/0038-0717(87)90052-6)
713 [0717\(87\)90052-6](https://doi.org/10.1016/0038-0717(87)90052-6)
714 Wei, C., Yu, Q., Bai, E., Lü, X., Li, Q., Xia, J., Kardol, P., Liang, W., Wang, Z., Han, X., 2013.
715 Nitrogen deposition weakens plant-microbe interactions in grassland ecosystems. *Glob.*
716 *Chang. Biol.* 19, 3688–3697. <https://doi.org/10.1111/gcb.12348>
717 Wu, J., Liang, G., Hui, D., Deng, Q., Xiong, X., Qiu, Q., Liu, J., Chu, G., Zhou, G., Zhang, D.,
718 2016. Prolonged acid rain facilitates soil organic carbon accumulation in a mature forest
719 in Southern China 544, 94–102. <https://doi.org/10.1016/j.scitotenv.2015.11.025>
720