


# Big in the benthos: Future change of seafloor community biomass in a global, body size-resolved model

Andrew Yool  | Adrian P. Martin | Thomas R. Anderson | Brian J. Bett | Daniel O. B. Jones | Henry A. Ruhl

National Oceanography Centre, University of Southampton Waterfront Campus, Southampton, UK

## Correspondence

Andrew Yool, National Oceanography Centre, Southampton, UK.  
Email: axy@noc.ac.uk

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## Abstract

Deep-water benthic communities in the ocean are almost wholly dependent on near-surface pelagic ecosystems for their supply of energy and material resources. Primary production in sunlit surface waters is channelled through complex food webs that extensively recycle organic material, but lose a fraction as particulate organic carbon (POC) that sinks into the ocean interior. This exported production is further rarefied by microbial breakdown in the abyssal ocean, but a residual ultimately drives diverse assemblages of seafloor heterotrophs. Advances have led to an understanding of the importance of size (body mass) in structuring these communities. Here we force a size-resolved benthic biomass model, BORIS, using seafloor POC flux from a coupled ocean-biogeochemistry model, NEMO-MEDUSA, to investigate global patterns in benthic biomass. BORIS resolves 16 size classes of metazoans, successively doubling in mass from approximately 1  $\mu\text{g}$  to 28 mg. Simulations find a wide range of seasonal responses to differing patterns of POC forcing, with both a decline in seasonal variability, and an increase in peak lag times with increasing body size. However, the dominant factor for modelled benthic communities is the integrated magnitude of POC reaching the seafloor rather than its seasonal pattern. Scenarios of POC forcing under climate change and ocean acidification are then applied to investigate how benthic communities may change under different future conditions. Against a backdrop of falling surface primary production ( $-6.1\%$ ), and driven by changes in pelagic remineralization with depth, results show that while benthic communities in shallow seas generally show higher biomass in a warmed world ( $+3.2\%$ ), deep-sea communities experience a substantial decline ( $-32\%$ ) under a high greenhouse gas emissions scenario. Our results underscore the importance for benthic ecology of reducing uncertainty in the magnitude and seasonality of seafloor POC fluxes, as well as the importance of studying a broader range of seafloor environments for future model development.

## KEYWORDS

allometric, benthic, ecology, future, global, model, particulate organic carbon flux, seafloor

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## 1 | INTRODUCTION

Of the particulate organic carbon (POC) fixed annually in the surface ocean by primary producers ( $\approx 45 \text{ Pg C year}^{-1}$ ; Behrenfeld & Falkowski, 1997), almost all is remineralized back to dissolved inorganic carbon (DIC) within a short time period (months to years). However, mediated by a complex community of benthic organisms, a small fraction of this POC ( $0.05 \text{ Pg C year}^{-1}$ ; Hain, Sigman, & Haug, 2014) is sequestered by seafloor burial, and represents a significant flux on geological timescales (Mawbey & Lear, 2013). This same community also remineralizes POC, turning it back to DIC and nutrients that eventually resupply productive surface communities (Dunne, Sarmiento, & Gnanadesikan, 2007). Consequently, understanding the role that benthic communities play in the biogeochemical cycles of the ocean, and estimating how this role may change into the future, is of considerable interest (e.g. Kriest & Oschlies, 2013; Moodley et al., 2011).

Benthic communities are an important source of commercially exploited seafood, both for direct human consumption, and for other harvested species. They are also often highly biodiverse, and in the deep-sea, may be relatively pristine compared to other marine habitats more directly impacted by human activities (Jackson et al., 2001; Lotze et al., 2006; Ramirez-Llodra et al., 2011). As such, they may be indicators of anthropogenic change in deep-sea habitats (Rees, Boyd, Schratzberger, & Murray, 2006), whether that be by global-scale climate change or ocean acidification (Birchenough et al., 2015; Mora et al., 2013; Sweetman et al., 2017; Yasuhara, Cronin, deMenocal, Okahashi, & Linsley, 2008), or by local seafloor mineral extraction (Halfar & Fujita, 2007).

Advances in deep-sea photography (e.g. Bett, 2003), sampling (e.g. Gooday, Bett, Shires, & Lamshead, 1998) and even in situ experimentation (e.g. Jeffreys et al., 2013; Main et al., 2015; Nomaki, Heinz, Nakatsuka, Shimanaga, & Kitazato, 2005; Witte, Aberle, Sand, & Wenzhöfer, 2003) have significantly increased understanding of how these communities are organized (Danovaro, Snelgrove, & Tyler, 2014). Evidence suggests that food web complexity and functional groupings of organisms may be of secondary importance to body size in controlling energy flow through these communities (Blanchard, Heneghan, Everett, Trebilco, & Richardson, 2017; Brown, Gillooly, Allen, Savage, & West, 2004; Dickie, Kerr, & Boudreau, 1987; Peters, 1983). As body mass is readily quantified and well correlated with metabolic processes (e.g. Brey, 2010) as well as community biomass and abundance (e.g. Hildrew, Raffaelli, & Edmonds-Brown, 2007), this is attractive on both practical and theoretical grounds. However, there is still considerable debate among biologists and ecologists as to the causes, and universality, of the mass-scaling of metabolism (e.g. Glazier, 2005, 2010; Hildrew et al., 2007; Hunt & Roy, 2006). Nevertheless, the continuing development of the Metabolic Theory of Ecology (McClain, Allen, Tittensor, & Rex, 2012; Schramski, Dell, Grady, Sibly, & Brown, 2015) demonstrates the practical value of size-based assessments of ecosystem function. Allometry has proven a useful concept in both terrestrial

and aquatic ecology (e.g. Hildrew et al., 2007; Schmidt-Nielsen, 1984), and its application to benthic communities, and marine systems in general, has clear potential (e.g. Blanchard et al., 2009, 2017).

As a preliminary step towards this, Kelly-Gerreyn et al. (2014) introduced an allometry-based model of the seafloor metazoan community that reproduced the distribution of biomass and abundance at three contrasting locations: (i) Faroe-Shetland Channel, North Atlantic deep-sea; (ii) Fladen Ground, North Sea continental shelf; and (iii) Oman Margin, Arabian Sea continental slope (see section 2.1.1 for more details). They employed high-quality size-resolved observational data on the meio- and macrobenthos from these three sites. Their model was tuned for these locations individually and subsequently returned to all three sites simultaneously to provide a more broadly applicable, or “unified” parameterization (Ichino et al., 2015; see section 2.1.1 and Appendix S1). The three study locations are highly contrasting in terms of water depth, habitat temperature, and the magnitude and seasonality of productivity. Consequently, Kelly-Gerreyn et al. (2014) consequently suggested that the model was likely to be robust and of broad application to marine benthic communities.

Future forecast simulations indicate that primary production—the key driver of the Kelly-Gerreyn et al. (2014) model—may be seriously impacted by climate change during the 21st century. The Coupled Model Intercomparison Project (CMIP5) used in Assessment Report 5 (AR5) of the Intergovernmental Panel on Climate Change (IPCC) found that, under a high greenhouse gas emissions scenario (RCP 8.5; see section 2.2.1), primary production changed by an average of  $-8.1\%$ , with falls in export production being even greater, ranging from  $-7\%$  to  $-16\%$  (Bopp et al., 2013). Similar results were found by Yool, Popova, Coward, Bernie, and Anderson (2013), who inferred a link with ocean acidification (OA), which led to drops in primary production ( $-6.3\%$ ) translating to much larger declines in export flux to the deep ocean ( $-40.7\%$  at 1000 m water depth). Using simulated changes in export flux and an empirical seafloor biomass model, Jones et al. (2014) estimated that an average change in seafloor POC flux of  $-11.4\%$  translated to a decline in total benthic biomass of  $-5.2\%$  at the global scale. These changes suggest that benthic communities will experience substantial impact in the near future.

In addition to changes in the food supply available to them, although not examined here, it is anticipated that benthic communities will experience other stresses. Temperature has been already identified as a key stressor from palaeological records (Yasuhara, Okahashi, Cronin, Rasmussen, & Hunt, 2014), especially where it is naturally low or high and organisms are most vulnerable to change (Yasuhara & Danovaro, 2016). More broadly, much as with pelagic communities (Gruber, 2011; Popova et al., 2016), benthic communities are also vulnerable to additional anthropogenically driven stressors, including acidification, deoxygenation and contamination (Levin & Le Bris, 2015). Exposure to stressors such as temperature and acidification would likely be strongly depth-dependent because of their surface sources, but the other stressors may

be communicated more directly to the deep ocean through the biological pump (deoxygenation) or gravitational sinking (contamination).

Here we take Kelly-Gerreyn et al. (2014)'s allometry-based model of the benthic community and force it with POC fluxes derived from the pelagic ecosystem model of Yool, Popova, Coward, et al. (2013). The model's unperturbed behaviour is examined, both in terms of global geographical variability and the role of temporal variability (especially seasonality) in POC flux. How these features may change into the future is then investigated using two end-member IPCC AR5 scenarios for the 21st century (RCPs 2.6 and 8.5). We focus on the outcomes for integrated seafloor biomass and its distribution across the modelled size classes, and how both of these properties vary in space and time.

## 2 | MATERIALS AND METHODS

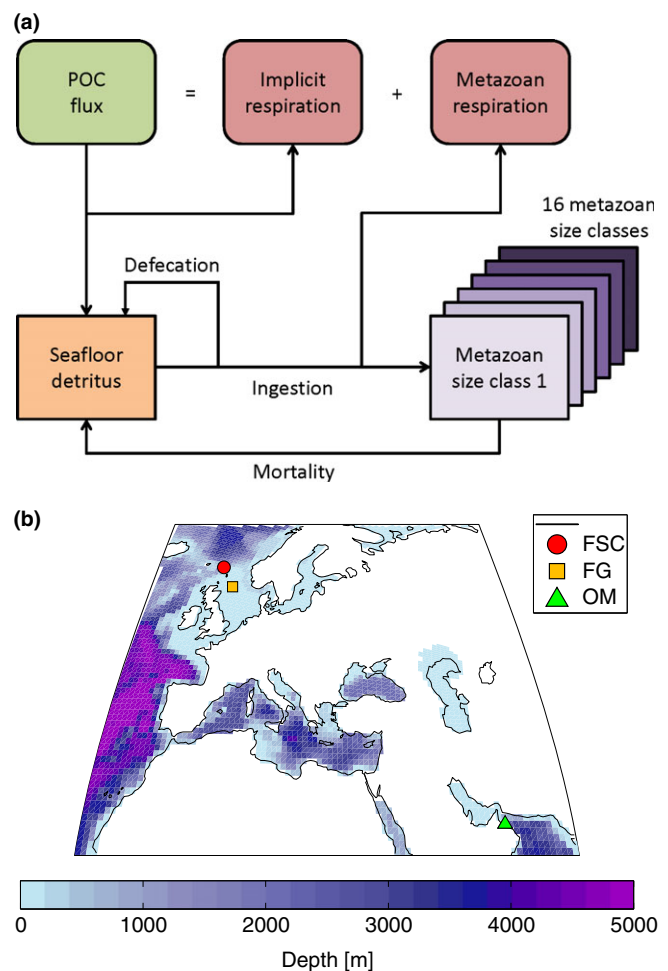
### 2.1 | BORIS model

The benthic community is modelled using the Benthic Organisms Resolved In Size v1.0 (BORIS-1; henceforth BORIS) model (Kelly-Gerreyn et al., 2014),<sup>1</sup> which represents the biomasses of seafloor metazoans in the meio- to macrofaunal size range (0.9  $\mu\text{g}$  wet wt to 30 mg wet wt; see Table S1). This size range was selected based on the availability of high-quality size-resolved field data. POC flux consumption and subsequent respiration by smaller and larger organisms are represented implicitly via the  $f_{\text{other}}$  parameter (see Appendix S1). As illustrated in Figure 1a, the model is driven by the POC flux reaching the seafloor which, after a fixed fraction is consumed by the implicit respiration of other members of the benthos, enters a detrital reservoir that is accessed by the modelled metazoans, both as a source of food (ingestion) and a sink for losses (mortality, defecation). Ultimately, all of the POC that arrives at the seafloor is consumed and respired; that is, burial sequestration is assumed to be minimal, and the model provides a biomass size distribution of the modelled metazoan community. Long-term observations of POC flux to the seabed and corresponding sediment community respiration in the deep ocean suggest that this is a reasonable simplification (Smith & Kaufmann, 1999; Smith et al., 2009). Similarly, global-scale assessments of the deep ocean (>2,000 m water depth) POC burial flux suggest a value of only 3% of seafloor POC flux (Dunne et al., 2007).

Appendix S1 presents a detailed description of the structure, parameterization and evaluation of BORIS. The computational implementation of BORIS is described in Appendix S1.

### 2.2 | Default performance

In Kelly-Gerreyn et al. (2014), the eight model parameters in BORIS were tuned using a microgenetic algorithm approach (Ward, Friedrichs, Anderson, & Oschlies, 2010). This was carried out separately



**FIGURE 1** (a) Schematic diagram of the components and interactions in the BORIS model. “Implicit respiration” refers to respiration by microbes and other unmodelled benthic organisms. (b) Map of the locations of the three sites used to originally parameterize the model. The grid cells and water depths shown are those of the NEMO model used in this study

for each of the three field locations. The procedure is fully described in Kelly-Gerreyn et al. (2014) and summarized in Appendix S1.

The field data used to calibrate BORIS came from: Faroe-Shetland Channel (FSC; 61.92°N, 2.80°W; 1623 m water depth), Fladen Ground (FG; 58.27°N, 0.88°E; 153 m) and Oman Margin (OM; 23.38°N, 59.00°E; 507 m). Figure 1b shows these locations. Table S2 lists the values of the parameters derived by Kelly-Gerreyn et al. (2014) for the three sites, together with the ranges from which they were drawn. Also listed is a “unified” parameter set (column All) that was produced using the same approach as in Kelly-Gerreyn et al. (2014), but fitting the model to all three sites simultaneously. This parameter set was previously employed by Ichino et al. (2015), and is used throughout this study.

Fig. S1a shows the resulting biomass distributions across the modelled size classes at equilibrium for the three locations. The model reproduces the observed general rise in standing stock with body size. The performance of the model tuned for individual locations ( $r^2 = .96, .67, .07$ ) is better than that of the unified model ( $r^2 = .76, .41, .04$ ).

<sup>1</sup>In Kelly-Gerreyn et al. (2014), the model was not explicitly given a name, but we adopt the moniker BORIS here to facilitate reference and discussion.

## 2.3 | NEMO-MEDUSA

Simulations of a coupled ocean-biogeochemistry model, NEMO-MEDUSA, provide the geographical and seasonal flux of POC to the seafloor. The physical ocean component is the Nucleus for European Modelling of the Ocean (NEMO) (Madec, 2008), run at 1° horizontal resolution for the global domain. The biogeochemical component is the Model of Ecosystem Dynamics, nutrient Utilisation, Sequestration and Acidification (MEDUSA; Yool, Popova, & Anderson, 2013; Yool, Popova, Coward, et al., 2013). POC flux to the seafloor comprises two separate size classes of detrital particles, the balance of which means that shallow regions are typically dominated by “small,” slow-sinking particles, while deep regions are dominated by “large,” fast-sinking particles. NEMO-MEDUSA agrees well with observed deep-sea POC fluxes (Honjo, Manganini, Krishfield, & Francois, 2008; Yool, Popova, & Anderson, 2013; Yool, Popova, Coward, et al., 2013). Note that POC in NEMO-MEDUSA is in molar units, while POC in BORIS is instead in wet weight units, and that these are converted using linear factors for carbon mass, dry weight and water content (Brey, Müller-Wiegmann, Zittier, & Hagen, 2010).

Appendix S2 presents more details concerning the structure, parameterization and validation of NEMO-MEDUSA.

## 2.4 | Forcing scenarios

We use extant, multicentennial simulations of NEMO-MEDUSA as a source of seafloor POC flux for both (i) an unperturbed, control state and (ii) an anthropogenically forced, historical state (1860–2005). The latter was extended into the future (2006–2099) under two contrasting scenarios of 21st century climate change, RCP 2.6 and RCP 8.5 (Rogelj, Meinshausen, & Knutti, 2012). These simulations were forced at the ocean surface using output from the UK Meteorological Office's (UKMO) HadGEM2-ES Earth system model (Jones et al., 2011). The control simulation used a repeated 30-year cycle of forcing from the unperturbed portion of the historical simulation, with constant, pre-industrial  $p\text{CO}_2$  concentrations.

In terms of POC fluxes to the seafloor, Yool, Popova, Coward, et al. (2013) found that these declined into the future in parallel with the decline in primary production driven by increasing ocean stratification and decreasing nutrient availability. However, while the decline in global primary production is modest (–6%), the corresponding decline in POC flux to the deep ocean is greater (–40%; 1,000 m water depth), as a result of the role of ocean acidification. Driven by oceanic uptake of anthropogenic  $\text{CO}_2$ , this both shoals the modelled calcite compensation depth and decreases NEMO-MEDUSA's production of biogenic calcium carbonate, with the latter the dominant factor for POC fluxes.

Appendix S2 presents more details concerning these scenarios and the forcing output.

## 2.5 | Simulations

Monthly average POC flux ( $\text{g C m}^{-2} \text{ day}^{-1}$ ) to the seafloor was extracted from NEMO-MEDUSA for both the control and the

historical portion of the RCP simulations described above (1980–1999). A monthly climatology was produced using the output from the control simulation, that is consisting of average January, average February. This climatology was used to force simulations of BORIS in order to explore its seasonal behaviour, initially at the three calibration sites.

These simulations were started from analytically derived initial conditions (see Appendix S1), and then spun-up for a period of 100 years to ensure that BORIS's constituent tracers exhibited a stable seasonal cycle.

This control climatological forcing was then applied at the global scale to explore large-scale geographical patterns in the behaviour of BORIS for unperturbed conditions. Simulations of 100-year duration were performed separately at each location in the NEMO-MEDUSA grid (65238 cells in total; 62% of the  $360 \times 292$  domain), and analysis made use of the average seasonal cycle in the final decade of the simulation. Each grid location was assumed to be independent of adjacent locations and dependent solely on POC falling from above.

This analysis was then extended using historical (1980–2005) and future (2006–2099) forcing to investigate the response of the modelled benthic ecosystem to changes in the POC flux driven by climate change and ocean acidification. In this case, BORIS was initialized at each geographical location from the corresponding end-state of the climatological control simulation, and then run for the period 1980–2099, with parallel RCP 2.6 and RCP 8.5 scenarios for the 2006–2099 period. Unlike the situation with climatological forcing, because of the time-varying nature of the forcing in these latter cases, analysis of the results used output from the full period rather than only the end of the simulation.

In all simulations of BORIS, the seafloor POC fluxes provided by NEMO-MEDUSA were first amended by subtracting a fixed and geographically uniform proportion of POC—the parameter  $f_{\text{other}}$ . As noted above, this is designed to account for the fraction consumed by unmodelled seafloor organisms (e.g. microbes and megabenthos). The remaining POC entered the seafloor detritus pool at which point it became available for BORIS's modelled metazoan size classes. This parameterization is discussed further later, and details concerning it can be found in Appendix S1. As noted earlier, organisms outside the size range considered by BORIS (larger and smaller) consume a portion of the POC flux, and the feeding relationship of all of these organisms on seafloor POC is much more complex than the simple transfer to respiration represented here.

## 3 | RESULTS

In the following, BORIS only implicitly represents both smaller and larger size classes that contribute to complete seafloor biomass. Consequently, the term “total biomass” is used in the presentation of results from BORIS to refer specifically to the total biomass of the modelled size classes resolved by BORIS and not to the complete total seafloor biomass of living organisms.

### 3.1 | Site simulations

Table 1 presents the observed and modelled seafloor POC fluxes for the three sites alongside both those observed, and those from the optimized BORIS model (per Table S2). The input POC flux that is available to the modelled size classes of BORIS,  $Q$ , is a fraction of the total flux reaching the seafloor, and the observed fluxes in Table 1 are accompanied by inferred values of parameter  $f_{\text{other}}$  that determine the fraction of the seafloor POC flux that is consumed by unmodelled size classes. In the case of all three sites,  $f_{\text{other}}$  is  $>0.9$  and, in the case of FG, exceeds 0.98, that is  $<2\%$  of the seafloor POC flux is available to the modelled benthic metazoans. This partitioning of food consumption is consistent with studies such as those of Pfannkuche and Soltwedel (1998) which found continental slope benthic systems to be dominated by bacteria and other small size class biomass ( $>90\%$ ). Similarly, in situ deep-sea measurements (e.g. Witte et al., 2003) have suggested that as microbes account for 95% of benthic biomass, they likely dominate total seafloor community respiration.

In a similar fashion, Table 1 also reports the seafloor POC fluxes of NEMO-MEDUSA accompanied by inferred values of  $f_{\text{other}}$  for both control and historical periods. While these model-inferred values agree well with the observation-inferred values for the FG site, that is are consistently close to 0.98, those for the other two sites are lower—in the case of the FSC site, inferred  $f_{\text{other}}$  values are only around 0.3. As Appendix S3 describes, this discrepancy remains even where a larger  $3 \times 3$  area around the sites is considered to account for grid mismatches, although the low value of  $f_{\text{other}}$  inferred at FSC (0.34) is consistent with a low productivity bias in the subpolar North Atlantic in NEMO-MEDUSA, noted by Yool, Popova, & Anderson, 2013; Yool, Popova, Coward, et al., 2013.

The NEMO-MEDUSA values of seafloor POC flux reported in Table 1 are multidecadal averages, but the model, as well as reality, exhibits significant temporal variability. Variability occurs at all three stations, ranging from seasonal, through interannual to decadal-scale patterns (section 3.3), as illustrated in Fig. S2, and considered further in Appendix S3.

Taking the seasonal climatology of seafloor POC fluxes at the three geographical sites from the control simulation (Fig. S2), and applying them to BORIS at the three sites using a uniform value of  $f_{\text{other}}$  of 0.9 (Appendix S3), Fig. S1b shows the resulting seasonal patterns of biomass for the modelled size classes and the seafloor detrital pool,  $R$ . Seabed flux at the three sites exhibits strong seasonal variability, and BORIS responds to this variability differently between the three sites. This is seen in Table S3 by the variable lags to POC flux shown by different benthic components. Total biomass is found to lag POC flux least at FG (52 days), is intermediate at FSC (90 days), but lags by 200 days at OM. While this suggests an interesting complexity in the relationship between POC flux and benthic communities, the three calibration sites alone cannot address this detail.

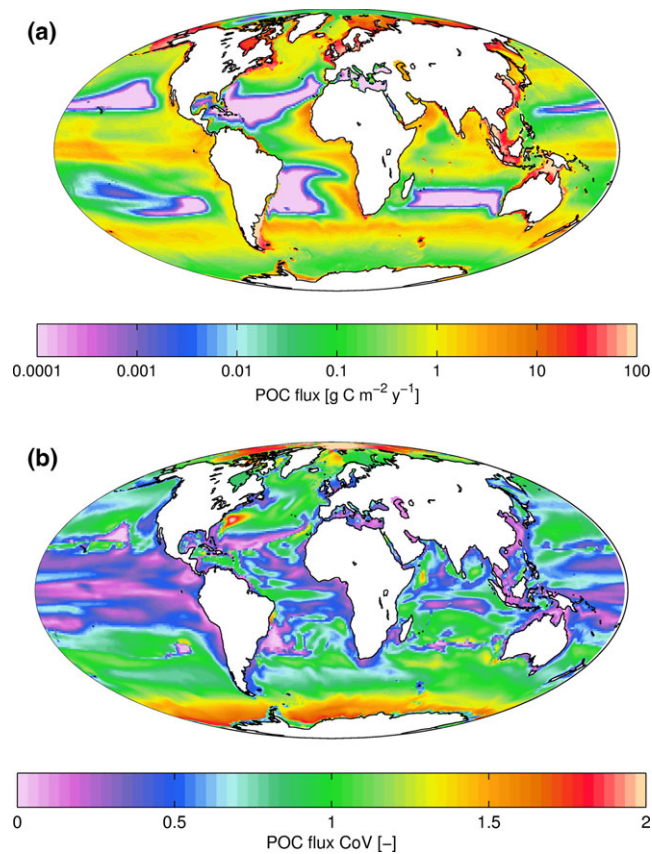
**TABLE 1** Estimated and modelled seafloor POC fluxes ( $\text{mg C m}^{-2} \text{ day}^{-1}$ ) and corresponding  $f_{\text{other}}$  values. Row 1 lists the optimized POC fluxes ( $Q$ ; Equation A4) consumed by modelled metazoans (Table S2). Row 2 lists the field POC fluxes to the seafloor (Table 1; Kelly-Gerreyn et al., 2014), together with the implied  $f_{\text{other}}$  values ( $-$ ) required to produce the  $Q$  values above. The remaining rows list NEMO-MEDUSA output for the grid cells in which the geographical sites are located, and for the surrounding  $3 \times 3$  cell neighbourhood. Averages for the period 1980–1999 from both the control and historical simulations are shown. In all NEMO-MEDUSA cases, the  $f_{\text{other}}$  values are those that would be required to produce the  $Q$  values at the top of the table

Property	POC flux			Estimated $f_{\text{other}}$		
	FSC	FG	OM	FSC	FG	OM
Optimized $Q$	3.2	2.3	1.1	–	–	–
Observed POC flux	40	120	24	0.92	0.98	0.96
Control POC flux	4.8	100	3.7	0.34	0.98	0.70
$3 \times 3$ cell region	4.8	85	4.2	0.34	0.97	0.74
History POC flux	4.2	91	4.5	0.25	0.98	0.75
$3 \times 3$ cell region	4.5	77	4.8	0.29	0.97	0.77

### 3.2 | Global patterns

Although the three calibration sites examined above vary in both total POC flux and in seasonal pattern, they represent only a fraction of the global range in these properties. Figure 2 shows the annual mean and coefficient of variation (CoV) of the climatological POC flux of NEMO-MEDUSA. This flux is primarily a function of export production (itself related to local primary production) and seafloor depth, with the highest values in the productive shallow waters of coastal regions, and the lowest values in the oligotrophic deep-water areas of the open ocean. Water depth plays a critical role in setting the seafloor flux, as it effectively sets the timescale over which POC sinking through a water column experiences remineralization. In shallow-water regions, this is typically a short period of time, so much of the POC exported reaches the seafloor, while in deep-water regions, sinking and remineralization periods are extended and seafloor POC fluxes are a much smaller fraction of export (see also Appendix S2). Across the global ocean, NEMO-MEDUSA simulates a range in seafloor POC flux that spans 12 orders of magnitude (Figure 3b). Patterns in seasonality are somewhat more complicated, with CoV values that are independent of both productivity and water column depth. With the exception of the high latitude Antarctic and Arctic, where POC fluxes are understandably strongly seasonal, there are no discernible zonal patterns comparable to those in the seafloor POC flux itself. Areas of similar productivity can exhibit quite different seasonality, based on prevailing hydrography, such as major currents, or on features in surface forcing driven by local weather patterns.

Figures 3a and S3 show the resulting patterns of annual mean seafloor detritus and total biomass (the sum of all explicitly modelled biomass classes) when forced by these POC fluxes. As expected, the



**FIGURE 2** NEMO-MEDUSA simulated seafloor POC flux, as (a) mean and (b) coefficient of variation (CoV; based on monthly means to illustrate seasonality)

relationships between POC flux and seafloor detritus or biomass are very close (Figure 2a), with the same patterns of lows in the oligotrophic gyres and highs in productive shelf regions. To illustrate the strength and consistency of these relationships, Figure 3b plots annual mean seafloor detritus and biomass for each size class against annual mean POC flux. Clearly, the patterns are exceptionally strong, with nearly perfectly linear (in a log-log sense) relationships between seafloor POC flux and all of the components of BORIS. There is a low level of spread around the general correspondence. This variability is potentially caused by differences in seasonal patterns of POC flux, but this is very modest by comparison with the mean annual response. For comparison, Fig. S4d illustrates the influence of seasonality in these patterns by repeating Figure 3b with monthly data. This illustrates that while BORIS is well constrained by POC flux on a mean annual basis, its seasonal behaviour can be considerably more complex.

Seasonality and temporal lags in the global model are examined in Fig. S5. The majority of model grid cells lie towards higher POC fluxes though are well spread on the CoV (seasonality) axis.

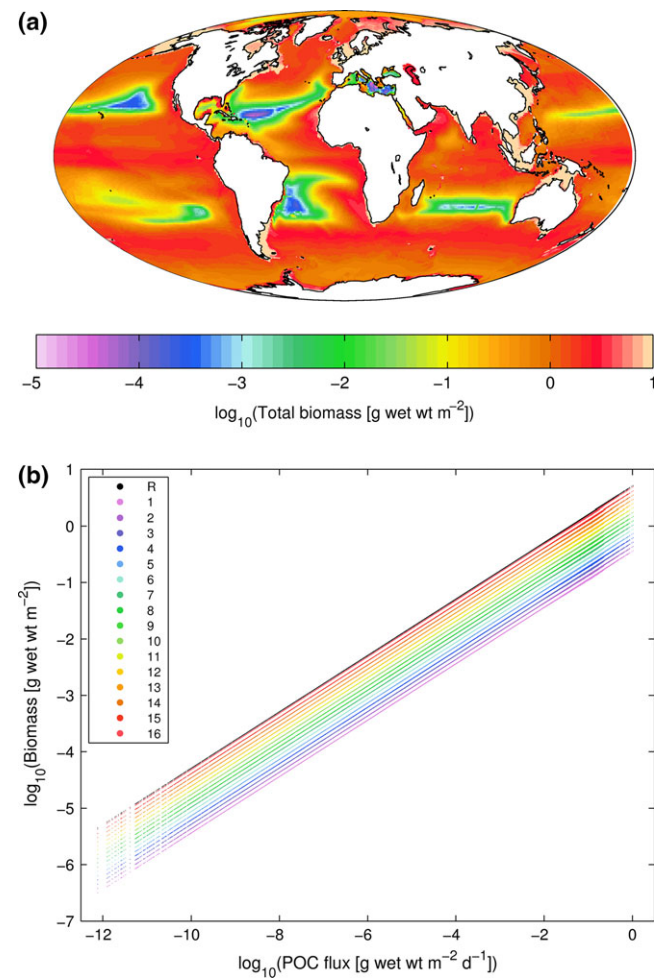
Projected across this POC flux domain, Fig. S5b, c show the average time delay between annual maximum POC flux and the corresponding annual maxima in seafloor detritus and total biomass. The general pattern in average temporal lags between peak annual

seafloor POC flux and seafloor detritus (Fig. S5b) or modelled total biomass (Fig. S5c) is for the lag to increase with decreasing POC flux, with a tendency for shorter lags where CoV is higher (more seasonal). Fig. S6 presents the corresponding lags for size classes 1 and 16, and illustrates a broadly consistent pattern of greater lags in the larger size classes (see also Fig. S7).

### 3.3 | Future change

To investigate the impact of changes in seafloor POC flux for benthic communities, Figures 4, S8 and Table 2 detail the change in globally integrated seafloor biomass across the 21st century for the RCP 2.6 and RCP 8.5 scenarios described earlier (Rogelj et al., 2012).

Table 2 reports globally integrated primary production, its translation to export production and the flux of POC to the seafloor, and the resulting modelled total biomass at a series of water depth intervals from shallow seas (to 0.1 km) to the full depth of the ocean (to



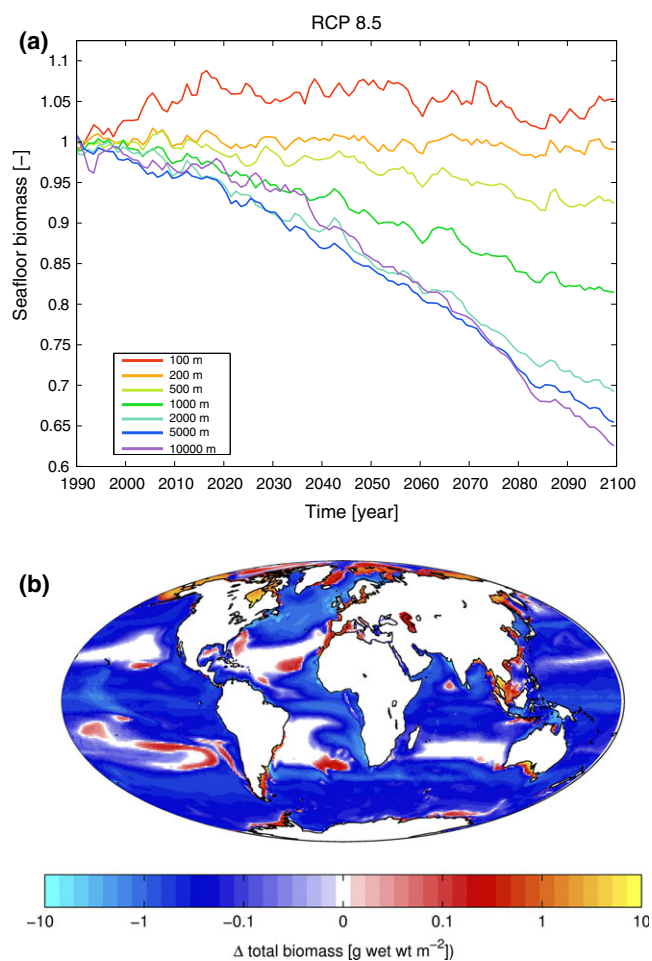
**FIGURE 3** (a) Mean annual field of total modelled seafloor biomass (Fig. S3 shows the corresponding field of seafloor detritus). (b) The relationship between mean annual POC and mean annual seafloor detritus (R; black) and biomass (classes 1–16; colours). Note that POC flux is expressed as  $\text{g wet wt m}^{-2} \text{ day}^{-1}$ , and that logarithmic scales are used throughout

10 km). In terms of primary production, while the extreme RCP 8.5 scenario experiences a 6.1% decline by the 2090s, control of emissions in RCP 2.6 results in a small increase in global productivity by the end of the century (in part related to slightly elevated temperatures enhancing phytoplankton growth rates). Export production shows a small decline of  $-1.3\%$  under RCP 2.6 (in part related to slightly elevated temperatures enhancing remineralization rates), but a much larger decline of  $-11.4\%$  under RCP 8.5's more extreme change. Integrated to the global scale, POC fluxes to the seafloor are increased for RCP 2.6 ( $+8.5\%$ ), while under RCP 8.5 there is a decline ( $-3.9\%$ ). This seeming disparity with export production stems from the seafloor POC flux being biased towards shallow regions (where a greater fraction of export survives to the seafloor), and the exclusion of regions with depths  $<100$  m from the reported export production.

To separate the impact of changes in seafloor POC flux and water depth, Table 2 divides the seafloor into seven water depth bands. For the shallow seas ( $<100$  m water depth), both scenarios result in an increase in seafloor biomass, largely driven by warmer conditions that increase growth rates in shallow regions (Yool, Popova, Coward, et al. 2013). For both scenarios, this increase declines and then reverses (by 200 m water depth under RCP 8.5) with increasing water depth. By 5,000 m water depth, seafloor biomass is lower for RCP 2.6 ( $-7.0\%$ ) and much lower for RCP 8.5 ( $-32.0\%$ ). Integrating globally, the total modelled seafloor biomass in BORIS declines slightly under RCP 2.6 ( $-1.1\%$ ), contrary to the corresponding increases in both primary production and seafloor POC flux. Under RCP 8.5, seafloor biomass decreases substantially ( $-17.6\%$ ), far more than either primary production or seafloor POC flux.

To distinguish the influences of changing primary productivity and water depth, Table 3 separates seafloor POC flux and the resulting modelled seafloor total biomass into depth bands. The pattern of increased seafloor POC flux in shallow seas and decreasing flux with increasing water depth is clear. Changes in seafloor POC flux at all water depths are greater than the corresponding changes in modelled total biomass. While the shallowest 5% of the ocean's seafloor experiences increased POC flux ( $+10.2\%$ ) and seafloor biomass ( $+3.2\%$ ), more than 83% of the ocean seafloor is at water depths  $>2$  km depth, where there are large decreases in POC flux ( $-53.4\%$  or greater) and modelled total biomass ( $-32.0\%$  or greater). The relationship between change in seafloor POC flux and change in modelled total biomass is strongly linear with a slope of approximately 0.6, that is biomass changes less than the POC flux. This contrasts with the pelagic situation of NEMO-MEDUSA where a decline of primary production of  $-6.1\%$  was much more closely paralleled by a corresponding decline of  $-5.7\%$  in the biomass of surface phytoplankton (Yool, Popova, Coward, et al. 2013).

Table 2 and Fig. S9 indicate how changes in primary production, export production, seafloor POC flux and seafloor biomass are correlated. In the case of export and primary production (Fig. S9a), this is almost linear, with a slight bias towards decreased export; that is, a  $-10\%$  decline in primary production leads to a large decline in export, while a 10% increase in primary production leads to a slightly



**FIGURE 4** Temporal and spatial distributions of seafloor total biomass under scenario RCP 8.5. (a) Time-series of mean annual seafloor total biomass by water depth bands (note: all lines are normalized to their 1990s averages). (b) Changes in seafloor total biomass between the 1990s and 2090s (note: the colour scale is logarithmic). Fig. S8 details results under scenario RCP 2.6

smaller increase in export. Similar, but less well-correlated and much less linear, relationships are apparent between seafloor POC flux and both export (Fig. S9b) and primary (Fig. S9d) production. However, the correlation between seafloor biomass and POC flux is strongly linear, with the variability introduced by seasonality practically insignificant when viewed across the global range.

Figures 4a and S8a show corresponding plots of temporal change in globally integrated modelled seafloor total biomass for the two scenarios. Both show relatively small changes in total biomass in shallow seas ( $<200$  m water depth), with larger changes generally occurring at greater water depths. In the case of RCP 2.6, these changes have generally saturated—or have even reversed direction—by the middle of the 21st century, in keeping with the strong mitigation characteristic of this scenario. In marked contrast, RCP 8.5 shows strong declines in seafloor total biomass that continue throughout the century, particular at greater water depths, also in keeping with the business-as-usual nature of this scenario.

To place these changes in a geographical context, Figures 4b and S8b show patterns of absolute change in biomass at the global scale.

**TABLE 2** Average model forcing (primary and export production and seafloor POC flux), and average simulated seafloor biomass for the 1990s and 2090s under scenarios RCP 2.6 and RCP 8.5. Seafloor biomass is listed for seven water depth bands. For both scenarios, the percentage change between the 1990s and 2090s is indicated in brackets

Property	Units	1990s	2090s, RCP 2.6	2090s, RCP 8.5
Primary production	g C m <sup>-2</sup> day <sup>-1</sup>	0.314	0.318 (+1.4%)	0.295 (-6.1%)
Export production, 100 m	g C m <sup>-2</sup> day <sup>-1</sup>	0.065	0.064 (-1.3%)	0.058 (-11.4%)
Seafloor POC flux	g C m <sup>-2</sup> day <sup>-1</sup>	0.029	0.032 (+8.4%)	0.028 (-3.9%)
Biomass, → 0.1 km	g wet wt m <sup>-2</sup>	8.831	9.594 (+8.6%)	9.109 (+3.2%)
Biomass, → 0.2 km	g wet wt m <sup>-2</sup>	7.603	7.735 (+1.7%)	7.545 (-0.8%)
Biomass, → 0.5 km	g wet wt m <sup>-2</sup>	4.142	4.106 (-0.9%)	3.839 (-7.3%)
Biomass, → 1 km	g wet wt m <sup>-2</sup>	1.887	1.806 (-4.3%)	1.551 (-17.8%)
Biomass, → 2 km	g wet wt m <sup>-2</sup>	1.248	1.175 (-5.8%)	0.887 (-28.9%)
Biomass, → 5 km	g wet wt m <sup>-2</sup>	1.067	0.992 (-7.0%)	0.725 (-32.0%)
Biomass, → 10 km	g wet wt m <sup>-2</sup>	0.780	0.765 (-2.0%)	0.514 (-34.2%)
Globally integrated biomass	Pg wet wt	0.607	0.600 (-1.1%)	0.500 (-17.6%)

Under RCP 8.5, seafloor biomass declines across the globe, with exceptions in some shallow seas, and where the boundaries of oligotrophic gyre areas have shifted. RCP 2.6 repeats much of this pattern, but with large areas of increased seafloor total biomass, principally in the Pacific and Indian basins, where they occur in both shallow and deep-water regions. Seafloor total biomass in the Atlantic Ocean—noted for its productivity decline in Yool, Popova, Coward, et al. (2013)—broadly declines regardless of the scenario. The large shallow-water continental shelves of the Arctic Ocean show a marked increase in seafloor biomass in line with that of primary production (cf. Yool, Popova, & Coward, 2015; Yool, Popova, Coward, et al. 2013).

## 4 | DISCUSSION

We have explored the spatial (global scale) and temporal (seasonal to interannual) response of a body mass size-resolved model of seafloor metazoan biomass, BORIS (Figure 1a), to variability in the seafloor POC fluxes that drive it.

Building on the constant seafloor POC flux forcing in Kelly-Gerrey et al. (2014) using the NEMO-MEDUSA model's seasonally resolved forcing, we find diversity in the seasonal response to

sedimenting POC, with some regions responding rapidly to inputs of POC, while others respond more slowly (Fig. S1b). Our three calibration sites represent a limited diversity in POC flux, and we extended our analysis globally giving a much broader range of POC forcing, in terms of both magnitude (Figure 2a) and seasonality (Figure 2b). This revealed further diversity in the phenological responses of BORIS (Figures 3a, S3 and S5) as well as strong, conservative patterns of biomass tied to POC supply to the seafloor (Figure 3b).

Earlier studies have noted that anthropogenic change may influence ocean productivity and thereby standing stocks of the benthos (e.g. Jones et al., 2014; Mora et al., 2013). Here we have examined the impact on seafloor biomass under end-member low (RCP 2.6) and high (RCP 8.5) IPCC emissions scenarios. While the impact on seafloor biomass was greater in the higher emissions case, both scenarios showed a similar pattern of change in which shallow seas increased in biomass while there was an increasingly negative impact with increasing water depth (Figure 4). As a result of ocean acidification-mediated change in surface-ocean communities, the high scenario used here is particularly severe for seafloor communities and translates relatively modest change in primary productivity into substantial decreases in seafloor POC fluxes (Figure 4a, b; Tables 2 and 3).

The strongly seasonal response of BORIS to the model-derived POC forcing used here has potentially important implications for both observational and modelling studies.

For observational studies, the model results underscore the importance of understanding POC flux variability (seasonality) as this may drive biomass cycles. However, the difficulty of obtaining high quality body mass distribution time-series from benthic ecosystems currently severely limits the number of datasets available to further development of BORIS. As such, when direct measurements of seafloor POC flux at any location are infrequent, it is important that seasonality is investigated by examination of the surface ocean system that drives it. For instance, through the use of surface samples of nutrients and plankton or production measurements, or estimated via remote-sensing (e.g. Henson, Sanders, & Madsen, 2012).

For modelling, our results suggest that the use of observational data in models needs to be based on an appropriate consideration of their temporal context. In particular, does the observation represent a stable body mass distribution or simply a single time-slice of a dynamic cycle. Model validation—and, especially, tuning—needs to ensure that model output and observational data are compared in a like-for-like manner.

While BORIS does exhibit substantial variability in response to temporally varying POC forcing, it is also conservative with respect to the average magnitude of that forcing. Despite the wide variety in seasonal patterns of POC forcing from NEMO-MEDUSA, average seafloor detritus and total biomass in BORIS is predictable from local annual POC flux. This suggests that, when characterizing seafloor dynamics, an accurate annual estimate of seafloor POC flux is critical.

It is notable that the magnitude of seasonal variability varies strongly between the model components. Seasonality is greatest for



Domain	Area	POC flux			Total biomass		
		1990s	RCP 2.6	RCP 8.5	1990s	RCP 2.6	RCP 8.5
Global	365.4	0.029	+8.4%	-3.9%	1.664	-1.1%	-17.6%
Seafloor → 0.1 km	18.7	0.340	+17.3%	+10.2%	8.831	+8.6%	+3.2%
Seafloor → 0.2 km	7.5	0.232	+1.9%	-1.0%	7.603	+1.7%	-0.8%
Seafloor → 0.5 km	9.5	0.074	-2.3%	-13.0%	4.142	-0.9%	-7.3%
Seafloor → 1 km	8.2	0.017	-8.7%	-29.3%	1.887	-4.3%	-17.8%
Seafloor → 2 km	15.2	0.008	-12.2%	-47.3%	1.248	-5.8%	-28.9%
Seafloor → 5 km	256.1	0.006	-12.5%	-53.4%	1.067	-7.0%	-32.0%
Seafloor → 10 km	50.2	0.003	-7.2%	-56.3%	0.780	-2.0%	-34.2%

**TABLE 3** Seafloor area ( $10^6$  km<sup>2</sup>), POC flux to the seafloor ( $\text{mg C m}^{-2} \text{ day}^{-1}$ ) and modelled total biomass ( $\text{g wet wt m}^{-2}$ ) listed at the global scale and for seven water depth bands for the 1990s and 2090s under scenarios RCP 2.6 and RCP 8.5. For both scenarios, the percentage change between the 1990s and 2090s is indicated in brackets

the detrital pool that the model benthos feeds on, and declines with individual body mass in the modelled size classes, such that larger organisms have noticeably lower CoV than smaller organisms (Fig. S10), presumably a function of their longer life-spans and slower physiological rates (McClain et al., 2012). This has interesting implications for field sampling programmes. For instance, in areas with highly seasonal productivity and resultant seafloor POC fluxes, while the temporal nature of these fluxes needs to be well characterized in order that the total annual flux is constrained, this may be less pressing for the benthic communities, and particularly the largest members of those communities. The largest benthic organisms (megabenthos) have until recently been difficult to quantify; however, new techniques have dramatically enhanced current capability (Morris et al., 2016). If the model's behaviour well describes that of real communities, lower frequency sampling of the larger body mass components of the community may be sufficient to capture a representative picture in an efficient manner. This is significant given the technical difficulty and expense of sampling the seafloor.

Temporal variability can also be considered in terms of the approach to equilibrium of body mass distribution. Through idealized perturbations of the seafloor detritus pool (half and double), Fig. S11 shows that the time to equilibrium for BORIS increases with decreasing seafloor POC flux, that is communities with lower total biomass take longer to return to equilibrium, and that it may take a very long time ( $\approx 1$  My for the lowest POC fluxes). While this extreme value is no doubt highly questionable, times ranging from years to decades of recovery have been suggested in field observations of the deep-sea (e.g. Huvenne, Bett, Masson, Le Bas, & Wheeler, 2016). This has obvious implications for appropriate environmental assessment and monitoring procedures, because recovery periods may be inversely related to total seafloor biomass. Although open-ocean benthic systems have broadly—so far—escaped major perturbations on the scales that pelagic and terrestrial ecosystems have experienced to date, the changes that may accompany global warming, ocean acidification and increasing human exploitation of the deep-sea floor are pervasive (Halfar & Fujita, 2007; Huvenne et al., 2016).

As seafloor POC fluxes vary in time, seafloor communities do not have a fixed equilibrium to converge to and will instead track a

“moving target.” Fig. S12 shows dynamically simulated and equilibrium total biomass at the three calibration sites, averaged on both monthly and annual timescales. On short timescales, the dynamic simulations struggle to follow the detailed steady states calculated analytically, though to varying degrees that correlate with POC flux magnitude; that is, site FG has highest flux and most closely tracks the analytical steady state. However, averaged over annual timescales, the dynamic simulations relatively closely track the equilibrium states.

Fig. S13 presents a global-scale examination of the relationship between dynamic simulations and equilibrium states. In general, the dynamic state is out of equilibrium to a slightly positive degree, both for the 1990s and the 2090s. The oligotrophic gyres are regions of larger disequilibrium although the biomass in these regions is much lower (Fig. S13a). At the end of the century, there is a greater positive disequilibrium in most areas, and noticeably larger areas of positive disequilibrium in the gyres (Fig. S13b). This general shift to more positive disequilibrium stems from a combination of declining ocean productivity and the system lags shown in Fig. S11, especially for the low biomass/slow equilibration gyre regions. Fig. S13c, d show the most “out of equilibrium” state variable of BORIS for the 1990s and the 2090s. Broadly, the global benthos is split between locations with excess seafloor detritus, that is more than at equilibrium, and those where the largest size class is most out of equilibrium. By the 2090s, the balance of this situation has shifted somewhat, such that the largest size class is most out of equilibrium over a larger portion of the world.

In general, intra- and interannual variations in seafloor POC flux introduce deviations, but as the three calibration sites illustrate, the dynamical simulation time-averages the equilibrium state quite closely (Fig. S12b). Consequently, the future behaviour of BORIS can be accurately estimated using time-averaged seafloor POC fluxes and steady-state calculations, at least for areas with non-negligible POC fluxes and shorter equilibration times.

A key result from this study is a forecast decline in benthic biomass substantially greater than the corresponding decline in surface ocean productivity that drives it (see also Yool, Popova, Coward, et al. 2013). Jones et al. (2014) also examined the fate of benthic biomass in a multimodel study, finding a decline in productivity and near-seafloor POC flux (2000s → 2090s) of more than -11% that resulted in a decline in

seafloor biomass of more than –5%. Our results appear more marked, so to make a direct comparison, we have forced the empirical model of Jones et al. (2014) with the same NEMO-MEDUSA output used in the present study (Appendix S4). Table 4 indicates that when subject to common forcing regimes, the two models produce very similar outcomes. Comparing the broadly similar body size ranges of total biomass in BORIS with macrofauna biomass in Jones et al. (2014), the respective changes are –6.2% and –7.3% under RCP 2.6 and –31% and –37% under RCP 8.5. These closely matched results derived from substantially different approaches point to the details of the POC flux forcing as being more important than the type of the benthic model being forced by that flux.

This study couples two models, each with their own limitations. Use of additional models of POC forcing could increase confidence in results; however, the performance of NEMO-MEDUSA has been extensively validated (Yool, Popova, & Anderson, 2013; Yool, Popova, Coward, et al., 2013), and the present results suggest that the behaviour of BORIS is a relatively simple function of seafloor POC flux. Only monthly average POC flux forcing has been applied in the present case, and it is possible that shorter-term variability may operate in the field (Smith et al., 2014; Witte et al., 2003).

The BORIS model makes a number of assumptions and simplifications that influence its representation of the benthos and their ecology. As currently configured, BORIS omits organisms outside the reliable range of the calibration data. BORIS effectively represents seafloor microbes as “external” to the modelled system; that is, they are ignored after they take a portion of the POC flux, and this overlooks the potential role of meio- and macrobenthos in preconditioning detrital substrates (Rowe & Deming, 2011). A practical consequence is the requirement for the  $f_{\text{other}}$  parameter, which makes specific assumptions and must be separately derived. It is possible to extrapolate the optimized model to both larger and smaller size classes (Ichino et al., 2015), but this was not carried out here in order to maintain model traceability with Kelly-Gerreyn et al. (2014).

BORIS also omits any consideration of the ecological roles played by temperature (Hunt & Roy, 2006), oxygen (Mosch et al., 2012) and ocean acidification (Andersson, Mackenzie, & Gattuso, 2011). For example, the Metabolic Theory of Ecology (Brown et al., 2004) would predict an inverse relationship between habitat temperature and standing stock biomass, with a potential effect factor of 2 between extremes of present-day deep-sea temperatures. Similarly, recent work has highlighted the role of increasing temperature in

reductions of body size (Reuman, Holt, & Yvon-Durocher, 2014). These factors co-occur in climate change scenarios, with the potential for antagonistic effects, for instance, decreasing seafloor POC flux with enhanced habitat temperature (Mora et al., 2013). As Fig. S14 illustrates, although NEMO-MEDUSA forecasts much of the deep seafloor to be only slightly affected by temperature change during the 21st century (mean change of 0.1°C), large temperature changes are predicted to occur in the most productive, shallow seas.

BORIS simplifies trophic relationships such that the modelled metazoans feed commensally from a common detrital pool. More generally, in representing organisms only by their body mass class, BORIS ignores mass-independent biological factors such as lifestyle (sessile, burrowing, swimming), feeding mechanism (deposit, filter, predation) and developmental mode (direct, lecithotrophic, planktotrophic), all of which may play a role in benthic system dynamics.

A significant limitation has been our use of globally uniform value of parameter  $f_{\text{other}}$ . Our calibration data suggest a value of 0.9 or more, but it remains unclear what the most appropriate value is. Consumption of incoming seafloor POC flux by microbes, and indeed the megabenthos, probably reflects both simple (e.g. POC quality) and complex factors (e.g. local community structure). We expect that  $f_{\text{other}}$  will vary systematically with the number of doubling body size classes that are explicitly modelled, in a manner comparable to the “energy equivalence rule” / “Damuth’s rule” (White, Ernest, Kerkhoff, & Enquist, 2007). Ideally, a “universal benthic model,” grounded within the framework of allometry, should replace this parameter with a more explicit representation of the processes that it simplifies. A consistent way forward may be to extend the size range of organisms modelled to include microbes and megabenthos, and dispense with the  $f_{\text{other}}$  parameter.

BORIS (and other models; Wei et al., 2010) is formulated around the bulk flux of POC to the seafloor and does not consider factors such as the type or quality of the material reaching benthic communities. Benthic metazoans of a broad size range certainly directly utilize energy from high-quality carbohydrates in rapidly sedimenting phytodetritus (Dunlop et al., 2016; FitzGeorge-Balfour, Billett, Wolff, Thompson, & Tyler, 2010). However, the flux of material from the surface ocean will also include organic matter that has been extensively reworked in the surface and/or mesopelagic realms, which will be of much lower quality (Valls et al., 2014). In NEMO-MEDUSA, detritus is produced by a number of different processes acting on different model groups (phytoplankton, zooplankton), but BORIS only considers

**TABLE 4** Direct comparison of the BORIS model and the Jones et al. (2014) empirical analysis when driven by common NEMO-MEDUSA POC flux, at the seabed in the former, and at 500 m above bottom (mab) in the latter (Martin et al., 1987). Predicted POC flux and corresponding benthic biomass for the 1990s and the 2090s under the RCP 2.6 and RCP 8.5 scenarios detailed in the text. BORIS total biomass (summed mass of modelled size classes) is taken to be broadly equivalent to the macrofauna category of Jones et al. (2014)

Model	Property	Units	1990s	2090s, RCP 2.6	2090s, RCP 8.5
BORIS	POC seafloor flux	g wet wt m <sup>-2</sup> day <sup>-1</sup>	0.021	0.019 (–11.5%)	0.010 (–50.7%)
	“Total biomass”	g wet wt m <sup>-2</sup>	1.059	0.993 (–6.2%)	0.728 (–31.2%)
Jones et al. (2014)	POC flux at 500 mab	g wet wt m <sup>-2</sup> day <sup>-1</sup>	0.032	0.029 (–10.1%)	0.018 (–42.8%)
	Macrofauna	g wet wt m <sup>-2</sup>	1.373	1.270 (–7.3%)	0.855 (–37.3%)

the first-order seafloor POC flux. As such, it overlooks nutritional factors that may play a role in real systems; however, knowledge of the significance of such aspects for benthic ecology is incomplete.

Despite these limitations, BORIS represents a dynamic, time-dependent alternative to empirical models such as Wei et al. (2010), and several avenues for future refinement are clear: extending the size range of organisms; including explicitly resolving microbes and their role in remineralization; ecological relationships between metazoans and microbes (microbial gardening; Mayor, Sanders, Giering, & Anderson, 2014) and between metazoans (predation).

A key driver of future work will be the assimilation of additional observational data sets, ideally from diverse sites that differ from those used in the initial development of BORIS. Our results on equilibration times would suggest, in particular, examination of sites with very low seafloor POC fluxes to provide end members that can constrain model behaviour under conditions of extreme oligotrophy. Additionally, similar time-lagged responses to climatic and upper ocean processes have been detected or inferred in abyssal benthic systems (Laguionie-Marchais, Paterson, Bett, Smith, & Ruhl, 2016; Ruhl & Smith, 2004). Incorporation of data from such sites will increase confidence when extrapolating to future situations in which local conditions have significantly departed from current. Such sites may also highlight environmental factors that are currently omitted from BORIS and which may drive its further development. Similarly, the assimilation of new seafloor POC flux data sets also represents a critical future avenue, especially where high temporal frequency sampling is achieved (Smith et al., 2014).

From a model-focused perspective, a potential extension to the work reported here would be to utilize BORIS in simulations forced with the output from a broader range of pelagic ecosystem models. For instance, from either the existing CMIP5 database (Jones et al., 2014), or those models participating in the upcoming CMIP6 exercise. In addition to evaluating BORIS under a broader range of potential present-day and future conditions, exposure to different models may provide an impetus to consider factors beyond simple bulk POC flux, that some of these models, including NEMO-MEDUSA, can provide.

In summary, here we apply a body mass allometry-based model of benthic biomass, BORIS, to global scale and force it with time-varying seafloor POC fluxes from a model of pelagic biogeochemistry, NEMO-MEDUSA. We examined present-day patterns of seafloor biomass, and investigated how these may change into the future, using end-member climate change scenarios. For the present-day, forced by temporally varying POC fluxes, BORIS exhibited strong seasonal and interannual behaviour, with distinct patterns of timing for different components under different forcings. Although there was considerable diversity in seasonal behaviour, mean annual benthic biomass was found to be strongly dependent on the seafloor POC flux, highlighting the importance of good quantification of this key flux. Compared with seasonality in seafloor POC flux, modelled seasonality in the seafloor detritus pool and benthic biomass was found to be much smaller, particularly with increasing individual body mass. These findings have potentially important implications

both for which benthic components should be monitored, and for the frequency at which that monitoring should occur. Under future climate change scenarios, modelled benthic biomass was found to decrease to a greater degree than that of surface ocean productivity (−18% as compared to −6%), with deep-water communities experiencing greater declines than those in shallow seas. This is in accordance with patterns of seafloor POC fluxes in NEMO-MEDUSA, which are increasingly attenuated with water depth by future change, and with the dominant role of seafloor POC flux magnitude in BORIS. We note that absent forcing factors (temperature, oxygen, pH), and our simplified ecological assumptions (implicit microbes, trophic relations), represent important aspects for future development of BORIS, with the acquisition of observations from a broader range of seafloor environments of key importance.

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## SUPPORTING INFORMATION

Additional Supporting Information may be found online in the supporting information tab for this article.

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