Modelling the combined impacts of climate change and direct anthropogenic drivers on the ecosystem of the northwest European continental shelf

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9 European Continental Shelf; POLCOMS-ERSEM

10 Highlights

- 11 1. Climate change could potentially drive substantial ecosystem changes by 2098
- 12 2. Direct anthropogenic divers impacts locally dominate over climate change effects
- 13 3. Direct drivers may mitigate climate change impacts on the European shelf ecosystem

14

15 Abstract

The potential response of the marine ecosystem of the northwest European continental shelf to 16 climate change under a medium emissions scenario (SRES A1B) is investigated using the 17 18 coupled hydrodynamics-ecosystem model POLCOMS-ERSEM. Changes in the near future 19 (2030-2040) and the far future (2082-2099) are compared to the recent past (1983-2000). The 20 sensitivity of the ecosystem to potential changes in multiple anthropogenic drivers (river nutrient 21 loads and benthic trawling) in the near future is compared to the impact of changes in climate. 22 With the exception of the biomass of benthic organisms, the influence of the anthropogenic drivers only exceeds the impact of climate change in coastal regions. Increasing river nitrogen 23 24 loads has a limited impact on the ecosystem whilst reducing river nitrogen and phosphate 25 concentrations affects net primary production (netPP) and phytoplankton and zooplankton biomass. Direct anthropogenic forcing is seen to mitigate/amplify the effects of climate change. 26 27 Increasing river nitrogen has the potential to amplify the effects of climate change at the coast by

- 28 increasing netPP. Reducing river nitrogen and phosphate mitigates the effects of climate change
- 29 for netPP and the biomass of small phytoplankton and large zooplankton species but amplifies
- 30 changes in the biomass of large phytoplankton and small zooplankton.

39 **1. Introduction**

40 Marine ecosystems are in continual adjustment responding to changes in the climate, both from

41 natural variability and long term anthropogenic climate change. Global climate change may

42 impact on ecosystems through large scale changes in temperature, stratification and circulation

43 [e.g. Bopp et al., 2001; Chust et al., 2014; Holt et al., 2012; Sarmiento et al., 2004].

44 Additionally, there is a direct impact from human activities, such as fishing, waste water

discharge, dredging, leisure, fossil fuel extraction and off-shore energy generation [e.g.

46 UKMMAS, 2010; Ducrotoy and Elliott, 2008; Halpern et al., 2008]. These direct effects tend to

47 be largest in coastal and shelf seas, where changes in the ecosystem are also likely to impact

48 most directly on humans.

49 The physical climate influences the ecosystem through temperature, which affects chemical and 50 physiological rates [Boyd et al., 2013], and by controlling the availability of nutrients [Holt et al., 51 2012] through horizontal transport and vertical mixing (which is inhibited by stratification). The 52 nutrient supply, and particularly its vertical distribution, combines with the light climate (light 53 diminishes with depth and the presence of organic and non-organic matter) to control the productivity of the lower trophic level (LTL) marine ecosystem. Carbon is taken up through 54 55 photosynthesis by phytoplankton, which are grazed by zooplankton; the plankton are transported 56 by physical processes and also sink under gravity through the water column; bacteria act to decompose dead plankton and nutrients are released through remineralisation. The LTL 57 ecosystem has a direct effect on the environment (e.g. through the development of harmful algal 58 59 blooms [Anderson et al., 2002]) and supplies food for higher trophic levels (e.g. fish) used for

60 human consumption or industry.

The northwest European continental shelf is part of the northeast Atlantic and exposed to changes in the global atmospheric and oceanic climates. It also experiences direct anthropogenic pressure from close proximity to the highly populated industrial regions of northern Europe. Fishing activity (demersal trawling) and riverine nutrient input to the ocean are two such direct anthropogenic processes that combine with changes in the climate to impact on the shelf sea ecosystem. Demersal trawling acts to disturb the seabed and induces mortality in benthic fauna, leading to disruptions of food webs and biogeochemical cycles in the vicinity of the disturbance 68 [*Kaiser et al.*, 2006]. Changes to nutrient loads impact the primary production and community

- 69 composition, particularly in hydrodynamic regimes directly connected to the riverine sources. In
- 70 particular high nutrient loads or imbalance in the ratio of nutrients may lead to deleterious
- eutrophication impacts such as high biomass [*Cadée and Hegeman*, 2002], toxic algal blooms
- 72 [Anderson et al., 2002] and near bed hypoxia [Dethlefsen and Von Westernhagen, 1983].
- 73 The policy driver behind this work is the Marine Strategy Framework Directive (MSFD:
- 74 Directive 2008/56/EC¹) of the European Commission, which requires member states to develop
- strategies to achieve a healthy marine environment and make ecosystems more resilient to
- climate change in all European marine waters by 2020 at the latest. The MSFD identifies 11 high
- 177 level descriptors, 5 of which are considered here (D1 Biodiversity, D4 Foodwebs, D5
- Eutrophication, D6 Sea floor integrity and D7 Hydrography). Each descriptor comprises a set of
- ⁷⁹ indicators which characterise marine ecosystems and requires an understanding of the possible
- 80 pressures and impacts on them. For instance, Good Environmental Status (GES) is achieved
- 81 when biodiversity is maintained (D1); the food web ensures long-term abundance of species
- 82 (D4); eutrophication is minimised (D5); benthic ecosystems are not adversely affected by
- changes in the sea floor (D6) and the ecosystem is not adversely affected by changes in
- 84 hydrographical conditions (D7). Numerical models such as applied here provide a valuable tool
- to improve the knowledge base on marine ecosystems and input to the development of innovative
- 86 tools for understanding and assessing GES in marine waters in European regional seas to inform
- 87 the implementation of the MSFD. For examples, see decision support tools developed from
- 88 numerical model simulations during the Marine Ecosystem Evolution in a Changing
- 89 Environment (MEECE) $project^2$.
- 90 Here we investigate the relative and combined effects of climate change and changes in the
- 91 direct anthropogenic drivers of benthic fishing and river nutrients, and study the impacts on the
- 92 LTL ecosystem over the northwest European shelf. A key question is the relative balance
- 93 between climate change and direct drivers, whether they act synergistically or antagonistically,

¹ http://ec.europa.eu/environment/marine/eu-coast-and-marine-policy/marine-strategy-framework-directive/index_en.htm

² http://meece.eu/

and whether management measures leading to changes in direct drivers can mitigate the effectsof climate change.

A major driver of primary production is the availability of nutrients, particularly nitrogen and 96 phosphate. Origins of these nutrients include natural (e.g. the open ocean) as well as 97 98 anthropogenic sources such as the release from industry, urban waste water treatment and 99 agriculture, through rivers and groundwater, to the ocean. For the European shelf as a whole, the 100 largest source of nutrients is the open ocean, although regionally, such as in the English Channel 101 and Irish Sea, river sources are significant [Huthnance, 2010]. For the North Sea, Thomas et al. 102 [2010] estimated ~80% of nitrogen input comes through the northern boundary, ~4% from the Baltic outflow, ~9% from rivers and ~6% from the atmosphere. In coastal regions of the North 103 104 Sea for a contemporary period (2000 onwards), Artioli et al. [2008] calculated the contribution of nitrogen from rivers to be 16%, with 26% coming from horizontal transport, 5% from 105 106 atmospheric deposition and 53% from sediments (mainly through resuspension). Elevated 107 nutrient concentrations imply that some coastal areas of the North Sea are at risk of 108 eutrophication, giving rise to increases in biomass and the potential for detrimental effects such 109 as oxygen depletion near the sea bed. During ~1950–1990, coastal waters of the North Sea 110 experienced increases in nutrient loads of $\sim 62\%$ for nitrogen and $\sim 45\%$ for phosphate [Vermaat 111 et al., 2008]. Management action to reduce riverine nutrient sources to the North Sea started in 112 the 1980s. The OSPAR Convention for the protection of the marine environment in the North-East Atlantic (http://www.ospar.org) is supported by 15 European governments and came into 113 114 force in 1998; one of its aims is to tackle all sources of pollution affecting the maritime area, 115 including the release of nutrients leading to eutrophication in the North Sea. The Paris 116 Commission (PARCOM) made recommendations in 1988, 1989 and 1992 on reducing nutrient inputs to the North Sea [OSPAR, 1988; 1989; 1992]. 117

The European shelf is also an important fishing region used by major international fishing fleets, with reported catches in the range 2.5–3.1 million tonnes/year for 2006–2012 [*ICES*, 2014]. For 1993–1996, *Rijnsdorp et al.* [1998] estimated that, in the most heavily fished regions in the southern North Sea, the Dutch beam trawler fleet covered 47–71% of the surface area up to five times per year. There is significant effort in regulating fishing effort through the Common Fisheries Policy (CFP) [*European Commission*, 2011], regulations for managing European

fishing fleets and for conserving fish stocks, first introduced in the 1970s. Recent regulations
[*European Commission*, 2014] are aimed at reducing the wasteful practice of discarding catch in
the North Sea.

The effects of climate change on the ecosystem are difficult to predict due to uncertainties in 127 emissions scenarios, climate forcing and models. Sources of uncertainty in atmospheric climate 128 forcing include the trajectory of greenhouse gas emissions, their conversion into atmospheric 129 concentrations and the responses of the global climate system to this radiative forcing [IPCC, 130 131 2001]. In studying the impact of climate change on ecosystems the additional impact of direct 132 anthropogenic drivers, such as those controlled by government policies, can also be considered. 133 Foreseen and unforeseen events (e.g. changes in governments, demographics, economic 134 recession and war) make long-term (>~50 years) projections of direct drivers highly uncertain. However, in some circumstances, the sign of change in a direct driver can be estimated: e.g., 135 136 under environmental policies nutrient emissions from agriculture, waste water and industry might be restricted. The time frame considered by governments is typically ~5-6 years (e.g. 137 138 duration of the European parliament). However, a time horizon of ~20-30 years is more useful 139 for policymakers to legislate for climate change adaptation.

140 Models are a useful tool to study potential conditions under possible future climate, and are particularly suited to sensitivity experiments where the response of the system to changes in 141 forcing can be assessed. There have been several recent studies downscaling global climate 142 change projections to the ecosystem of European regional seas. Skogen et al. [2014] investigated 143 144 eutrophication in the North Sea, Skagerrak, Kattegat and the Baltic Sea under a medium emissions scenario (A1B) and found little change in the North Sea eutrophication status due to 145 146 changes in climate; river nutrients were kept at present day values. Also using a medium emissions scenario, *Holt et al.* [2012] demonstrated that the supply of nutrients from the open 147 ocean is an important control of primary production on the northwest European shelf, 148 149 particularly for the Irish shelf and the central and northern North Sea, which are exposed to 150 exchange with the open ocean. Using a 1D water column model at three contrasting locations in the North Sea, Van der Molen et al. [2013] found that gross primary production increased and 151 152 zoobenthos biomass and sea-bed oxygen decreased under climate change conditions; there was

little interaction between the climate signal and the addition of demersal trawling indicating thatreducing demersal trawling might mitigate the effects of climate change on benthic biomass.

155 Nutrient inputs from rivers and their relationship to eutrophication has substantial policy interest in this region owing to uncertainties in whether undesirable effects arise [e.g. Gowen et al., 156 2008], costs of amelioration and the transnational nature of the problem. There have been several 157 modelling studies on the effects of reducing concentrations of nutrients released from rivers into 158 159 the North Sea under present day and recent past conditions. Using a coupled river and multi-box 160 model of the Southern Bight of the North Sea and eastern English Channel over the last 50 years, Lancelot et al. [2007] showed increases in Phaeocystis and diatom production with increasing 161 river nitrogen and phosphate, and decreasing production when river phosphate loads fell. Lacroix 162 163 et al. [2007] used a 3D model of the southern North Sea and showed that reducing river nitrogen 164 loads led to an increase in diatom biomass, whereas decreasing river phosphate reduced both 165 diatom and *Phaeocystis* biomass; in addition, changes in open-ocean nutrient concentrations transported eastwards through the English Channel also have an impact on primary production in 166 167 the Belgian Exclusive Economic Zone. Skogen and Mathisen [2009] studied the long term impact of river nutrient reductions on chlorophyll-a concentrations using a 3D physical-168 169 chemical-biological model of the North Sea and found that the response was largest near the 170 coast, although oxygen levels did not react strongly to river nutrient changes. In contrast, using 171 result from six models covering the North Sea, Lenhart et al. [2010] showed that areas experiencing near bed oxygen concentrations below the "oxygen depletion" threshold defined by 172 173 OSPAR (i.e. 6 mg O_2/l , [OSPAR, 2003]) could improve their state by reducing river nutrients. 174 For the Baltic Sea, *Meier et al.* [2012] used a coupled physical-biogeochemical model for 1961

to 2099 to study the combined future impacts of climate change and nutrient loads from
industrial and agricultural sources. Using present day nutrient loads they found that water quality
deteriorates in the future and that, for moderate reductions in nutrients consistent with current
legislation, the climate effect dominates the impact of nutrient changes.

179 In this paper we use the Proudman Oceanographic Laboratory Coastal Ocean Modelling System,

180 POLCOMS [Holt and James, 2001], coupled to the European Regional Seas Ecosystem Model,

181 ERSEM [Baretta et al., 1995; Blackford et al., 2004] to examine the effects of changes in

182 climate, benthic fishing in the North Sea and river nutrient loads on the ecosystem of the 183 European shelf. The potential ecosystem responses to climate change in the near future (2030-184 2040) and the far future (2082-2099) are compared to the recent past (1983-2000). For the near future time period, the relative sensitivity of the ecosystem to changes in multiple drivers 185 186 (climate, river nutrient loads and benthic trawling) is also studied. Since long-term projections of direct drivers are highly uncertain and the appropriate time scale for policymakers is of the order 187 188 of decades, or shorter, we concentrate on the potential impacts of direct drivers in the near future time slice. The uncertainty in the change in direct drivers is large and we focus on qualitative 189 190 effects. The model setup and experiments are described in section 2. In section 3, the effects of 191 climate and anthropogenic changes on the ecosystem are explored and regions are defined where 192 the impact of the direct anthropogenic drivers exceeds that due to changes in the large scale 193 oceanic and atmospheric climate. The cumulative effects of the direct drivers and climate change 194 are also studied to investigate whether the impact of changing the direct drivers amplifies or 195 mitigates the effects of climate change.

196

197 **2. Model experiments**

198 2.1 Model description

199 We use the coupled hydrodynamic-ecosystem model POLCOMS-ERSEM on the Atlantic 200 Margin domain (Figure 1) covering the northwest European continental shelf and the adjacent deep ocean (20°W to 13°E, 40°N to 65°N). The model has horizontal resolution of ~12km and 201 202 uses 42 s-coordinate levels [Song and Haidvogel, 1994] in the vertical. POLCOMS is a 3dimensional finite difference model able to model shelf and deep ocean processes. ERSEM is a 203 204 lower trophic level biogeochemical model which explicitly resolves carbon, nitrogen, oxygen, 205 phosphorous and silicon cycles in a coupled pelagic-benthic system. ERSEM uses four 206 phytoplankton types (flagellates, picoplankton, diatoms and dinoflagellates), three zooplankton types (heterotrophic nanoflagellates and micro and meso zooplankton) and bacteria. The model 207 setup is described by *Holt et al.*, [2012] except that here we have turned off the resuspension of 208 particulate matter from the sea bed and added an explicit representation of sea bed trawling in the 209 North Sea (see below). Validation for the present day climate scenario simulation CNTRL (see 210

- below) showed that there is no systematic increase in errors [Holt et al., 2012] compared to a
- 212 hindcast using realistic atmospheric forcing from the European Centre for Medium Range
- 213 Weather Forecasting (ECMWF) Reanalysis product (ERA40).



Figure 1: Location plot showing the model domain and numbered regions. The box drawn with a dashed line is the region for which results are presented.

217

We consider two types of trawling gear working in the North Sea, namely, beam trawls and otter 218 trawls. Data on trawling effort consist of hours fishing per year averaged over the period 1997-219 220 2004 [Greenstreet et al., 2007]. The data are converted into number of trawlers per day expected 221 in each model grid cell. The number of trawlers, T, actively fishing in each cell is generated 222 randomly each day from a Poisson distribution characterized by this expected value. The impact, 223 I, of each trawler on the biomass of deposit and filter feeders (i.e. the % mortality induced by a 224 single event) is parameterized depending on the type of sediment (mud, sand or gravel) and the type of trawler (Table 1) using data from Allen and Clarke [2007] and Kaiser et al. [2006]. The 225 biomass of aerobic bacteria removed from the benthos is added to the biomass of pelagic bacteria 226 227 in the grid cell just above the bed. Changes are scaled by a factor representing the proportion, P_r ,

of the total grid cell area that is actively trawled, calculated assuming that a trawler covers an

area of 1.3 km² each day [*Dounas et al.*, 2007]. The decrease in biomass in a grid cell, relative to

230 the original biomass is then: $(1 - 0.01 \times I) \times T \times P_r$.

231

Sediment	Deposit Feeders	Filter feeders	Meiobenthos	Aerobic bacteria		
		Beam Trawlers				
sand	-23	-73	-67	-67		
gravel	-67	-15	-42	-42		
		Otter Trawlers				
mud	-18	-31	-29	-29		
sand	-23	-4	-15	-15		

Table 1: Mortality rates (%) on benthic functional types due to trawling [from *Allen and Clarke*,
2007].

234

Freshwater river fluxes for the present day are from a climatology of daily discharge data for 250

rivers from the Global River Discharge Data Base [Vörösmarty et al., 2000] and from data

prepared by the Centre for Ecology and Hydrology as used by *Young and Holt* [2007]. A mean

annual cycle of nutrient concentrations is derived for each river from data used by *Lenhart et al.*

[2010], including data processed by van Leeuwen (CEFAS, UK) for the UK, Northern Ireland,

240 Ireland, France, Norway, Denmark and the Baltic and by Pätsch and Lenhart [2004] for

241 Germany and the Netherlands. The river nutrient climatology is calculated using data for 1984 to

242 1994, covering a period of high nutrient concentrations and the start of the impact of nutrient

reduction policies; it represents conditions in the recent past.

Non-biotic light absorption is simulated by using an annual cycle of SeaWiFS climatology of

sediment particulate matter and coloured dissolved organic matter [*Smyth et al.*, 2006], as

described by *Wakelin et al.*, [2012].

247 Open ocean (hydrodynamic and ecosystem) and atmospheric forcing are derived from the Institut

248 Pierre Simon Laplace Climate model (IPSL-CM4), [Marti et al., 2005] a global coupled ocean-

249 atmosphere general circulation model (OA-GCM) run using the Special Report on Emissions

250 Scenarios (SRES) A1B "Business as Usual" scenario [Nakicenovic and Swart, 2000]. On a 251 global scale, IPSL-CM4 exhibits results similar to other OA-GCM models used in the IPCC Fourth Assessment Report [Meehl et al., 2007]. Steinacher et al. [2010] compared results for 252 four OA-GCMs (IPSL, MPIM, CSM1.4 and CCSM3) under the SRES A2 emissions scenario for 253 254 1860-2100 and showed that the IPSL model results were in the middle of the range of response 255 for global mean sea surface temperatures and primary production. In the North Atlantic, sea 256 surface temperature increases between 1860-1869 and 2090-2099 for IPSL are approximately average compared to the other models, whereas IPSL exhibits the second strongest reduction in 257 primary production over the same period. The IPSL model underestimates primary production in 258 the northeast Atlantic compared to SeaWiFS climatology for 1997 to 2005, with better 259 260 agreement than MPIM (which also underestimates primary production) but generally lower values than the NCAR CSM1.4 simulation [Schneider et al., 2008]. Using a skill score metric to 261 assess the skill of the four OA-GCMS to reproduce the satellite based estimates of primary 262 production Steinacher et al. [2010] showed that, for the northeast Atlantic region, CSM1.4 263 scored best, with IPSL and CCSM3 having similar values to one another and MPIM performing 264 265 worst.

266 Tidal boundary data for 15 constituents are provided by a tidal model of the North Atlantic267 [*Flather*, 1981].

We use a time-slice approach whereby mean conditions in an experiment are compared with mean conditions in a reference simulation to give a measure of the climatic/anthropogenic change, on the assumption that conditions in both time-slices are approximately stationary. Here we consider a single potential realization of the future and study the effect of substantially changing direct anthropogenic drivers.

To allow the model to adjust to its lateral boundary and surface forcing and changes in anthropogenic drivers, all simulations have a 5 year spin up period (described below) before calculating the average results presented here. Five years is ample spin-up time for the physics and pelagic biology on the Atlantic Margin domain and errors in benthic fluxes due to uncertainties in the initial conditions are significantly reduced after five years [*Wakelin et al.*, 2012]. 279 We consider five experiments: CNTRL, A1B, BASE, WM and GC, outlined below and

summarized in Table 2. The CNTRL, A1B and BASE simulations are used to investigate the

impacts of large scale changes in the atmosphere and ocean under a projected near- and far-

future climate in the absence of changes in direct anthropogenic drivers. For the near future

283 period, the final two simulations (WM and GC) explore the additional impact of changes in the

anthropogenic drivers of river nutrient concentrations and benthic fishing effort.

285 2.2 Scenario descriptions

286 2.2.1 Climate change scenarios: CNTRL, BASE and A1B

The CNTRL simulation is a present day simulation for the nominal period 1983-2000. A1B is a 287 future climate scenario representative of possible conditions in 2082-2099 under the SRES A1B 288 289 "business as usual" emissions scenario. The baseline simulation for the near future, BASE, is for the period 2030 to 2040 under the same emissions scenario. Physical ocean boundary and 290 atmospheric forcing are taken directly from the OA-GCM for the relevant time slice. Open-291 292 boundary nutrient (nitrate, silicate and phosphate) and dissolved inorganic carbon (DIC) values 293 for CNTRL are from climatologies [Garcia et al., 2006; Key et al., 2004]; while for the future 294 time slices, the CNTRL values are perturbed by the fractional change in nutrients in the PISCES ecosystem model [Aumont et al., 2003], included in the IPSL-CM4 model, between the time-295 296 slice and 1983-2000. IPSL-CM4 forcing data were available for the periods 1980-2000, 2030-297 2040 and 2080-2099; five years of spin up prior to the starts of the analysis periods (1982, 2030 298 and 2082, respectively) were achieved by using the first year of available forcing in each period 299 repeatedly as necessary. For example, the CNTRL simulation comprises two initial years using 300 forcing data for 1980 followed by data for the relevant years for the remainder of the spin-up 301 time: 1980-1982. Initial temperature and salinity fields are interpolated from the IPSL-CM4 302 model for the start month of each simulation. Initial ecosystem fields for CNTRL are from a spin-up simulation where homogeneous initial values corresponding to the average bulk 303 304 properties of the shelf have been spun up for five years; for BASE and A1B the nutrient fields in the CNTRL initial data are perturbed by the fractional change in nutrients in the PISCES 305 306 ecosystem model.

- Given the 2.5°×1.27° resolution of the IPSL-CM4 model, the precipitation at the outflow grid 307 308 cell is assumed to be representative of the precipitation over each river catchment. River flows 309 are changed in proportion to the change in precipitation at the outflow grid cell on average 310 during each time slice compared to 1983-2000; the effect of change in precipitation on river 311 outflow, integrated over the regions in Figure 1, varies between -7.5% and 9.7% for BASE and -22.2% and 20.7% for A1B (Table 3). River nutrient concentrations are held constant and 312 313 therefore changes in river volumes impact on the total loads of nitrogen, phosphate and silicate 314 being released to the ocean. In the absence of reliable estimates of how non-biotic light 315 absorption might change in the future, the present day SeaWiFS climatology is used in all experiments. 316
- For the CNTRL, A1B and BASE simulations, the direct anthropogenic drivers of fishing effortand river nutrient concentration are held at present day levels.

2	20	
3	20	

Name	Scenario Description		Large scale climate	River nutrient concentrations	Fishing effort	Years	
CNTRL	control run	recent past control run	recent past	climatology	climatology	1983-2000	
A1B	climate	far future climate forcing	far future	climatology	climatology	2082-2099	
BASE	climate	near future climate forcing	near future	climatology	climatology	2030-2040	
WM	climate and anthropogenic	World Markets ^a scenario	near future	50% increase in nitrogen; climatology for phosphate	climatology	2030-2040	
GC	climate and Global anthropogenic Community ^b scenario		near future	50% decrease in nitrogen and phosphate	50% decrease in beam trawling; 75% decrease in otter trawling	2030-2040	

321 **Table 2:** Summary of experiments: climatologies are averages representative of present day

322 conditions; all % changes are relative to the relevant climatology. Large scale climate comprises

323 atmospheric forcing, physical and ecosystem ocean boundaries and river volume outflows.

324 Anthropogenic driver scenarios relate to ^aconditions of rapid economic growth and limited

325 environmental policies and ^benvironmental policies constraining socioeconomic growth.

	volume			nitrogen				phosphate				silicate			
	$\begin{array}{c} \text{CNTRL} \\ \times \ 10^3 \\ \text{m}^3 \text{s}^{\text{-1}} \end{array}$	BASE PC %	A1B PC %	CNTRL molN s ⁻¹	BASE PC %	GC PC %	WM PC %	A1B PC %	CNTRL molP s ⁻¹	BASE PC %	GC PC %	A1B PC %	CNTRL molS s ⁻¹	BASE PC %	A1B PC %
1. Southern North Sea	3.99	5.6	-6.1	19.88	6.6	-46.7	60.0	-5.2	1.48	6.7	-46.7	-4.4	14.18	6.3	-6.3
2. Central North Sea	0.58	-6.4	5.4	1.39	-5.1	-52.6	42.3	6.1	0.09	-4.3	-52.2	6.1	2.03	-7.2	5.2
3. Northern North Sea	0.52	-7.5	2.9	0.54	-10.9	-55.4	33.7	4.5	0.02	-8.8	-54.4	4.2	2.39	-9.5	3.8
4. English Channel	0.79	2.4	-8.8	5.15	2.4	-48.8	53.6	-9.1	0.11	6.8	-46.6	-3.1	3.07	3.1	-7.2
5. Skagerrak/ Kattegat	2.21	-6.6	3.0	0.95	-6.9	-53.4	39.7	2.5	0.05	-5.6	-52.8	3.2	4.62	-6.7	3.6
6. Norwegian Trench	0.71	0.3	3.5	0.13	0.3	-49.9	50.4	3.4	0.01	0.3	-49.8	3.4	1.56	1.1	3.7
7. Shetland Shelf	0.03	-4.2	1.6	0.01	-4.4	-52.2	43.4	1.7	0.00	-4.4	-52.2	1.7	0.06	-4.4	1.7
8. Irish Shelf	1.52	-2.8	2.8	0.85	-2.8	-51.4	45.8	4.1	0.09	-3.5	-51.7	3.2	0.74	-4.5	0.4
9. Irish Sea	0.85	-1.1	3.4	2.01	0.5	-49.7	50.8	3.7	0.14	0.4	-49.8	3.6	2.19	-0.9	3.4
10. Celtic Sea	0.88	5.5	7.0	4.02	5.7	-47.2	58.5	6.2	0.18	5.6	-47.2	5.8	2.32	7.4	7.1
11. Armorican Shelf	1.63	-5.2	-22.2	5.05	-5.3	-52.7	42.0	-22.2	0.28	-6.3	-53.2	-30.1	7.54	-5.2	-22.4
12. NE Atlantic	0.06	9.7	20.7	0.01	9.7	-45.1	64.6	20.7	0.00	9.7	-45.1	20.7	0.17	9.7	20.7
Baltic Outflow	20.83	7.2	13.4	1.18	7.2	-46.4	60.8	13.4	0.36	7.2	-46.4	13.4	8.33	7.2	13.4

Table 3: Inflows of water, nitrogen, phosphate and silicate from the Baltic Sea and from rivers into different regions (Figure 1) for the CNTRL simulation and the percentage changes for the scenario simulations defined as $PC = (EXP/CNTRL - 1) \times 100\%$ where EXP =

BASE, GC, WM or A1B. For phosphate, the WM PC is identical to the value for BASE and, for silicate, the WM and GC PCs are

329 both identical to BASE.

330 2.2.2 Anthropogenic driver scenarios: WM and GC

331 The final two simulations explore the impact of changes in anthropogenic drivers in addition to

the climate change signal for the period 2030 to 2040. The direct driver simulations are the same

as BASE but with changes in the river nutrient loads and trawling drivers consistent with the

334 World Market and Global Community scenarios of the European Lifestyles and Marine

Ecosystems (ELME) project³. In the ELME project, socio-economic drivers for future scenarios

are used to define environmental pressures impacting on ecosystems.

337 In the World Markets scenario there is rapid economic growth and limited environmental

policies; an increase in the use of nitrogen fertiliser in agriculture combined with no changes in

339 urban waste water treatment (UWWT) leads to an increase in riverine nitrogen whereas river

340 phosphate concentrations remain unchanged. Trawling effort is also unchanged from present day

341 levels.

342 In the Global Community scenario economic growth is constrained by environmental objectives;

343 reductions in the use of phosphate fertiliser together with increases in UWWT and a reduction in

industrial discharge lead to a decrease in both nitrogen and phosphate released into rivers. There

is also a reduction in trawling effort.

Between 1985 and 2002, the Netherlands, Germany, UK and France reduced their river nutrient outputs to the North Sea by 10-90% for ammonium and 0-70% for phosphate, with little change in levels of nitrate+nitrite from Germany, the UK and France and a reduction ~20% from the Netherlands [*Lenhart et al.*, 2010]. Given the magnitude of these changes already experienced for rivers flowing into the North Sea, we perturb all river nutrient concentrations by 50% from their climatological values as being a potentially realistic change, that also corresponds to the 1988 OSPAR recommendation to reduce nutrient loads by 50% [*OSPAR*, 1988].

353 To study the potential effects on the ecosystem of the changes under the World Market scenario,

the WM simulation uses river concentrations of total nitrogen (nitrate and ammonium) that are

355 50% larger than the present day values used in the CNTRL, A1B and BASE simulations. River

³ http://cordis.europa.eu/documents/documentlibrary/127824671EN6.pdf

phosphate loads and trawling effort are unchanged from the present day. The effect on the
amount of nitrogen being discharged into different regions varies between 33.7% and 64.6%
compared to CNTRL (Table 3), while the phosphate load changes vary between -6.3% and 9.7%

due to changes in river volume flow.

360 The possible effects of the Global Community scenario are studied using the GC simulation

361 where river concentrations of total nitrogen and phosphate are reduced by 50% compared to the

362 present day values, resulting in area changes of between -55.4% and -45.1% for nitrogen

363 discharge and -54.4 % and -45.1% for phosphate discharge compared to CNTRL. For the GC

simulation, trawling effort is reduced by 50% for beam trawlers and 75% for otter trawlers; we

assumed that the recent declining trend of fishing effort [*Greenstreet et al.*, 2007] will double its
slope under the GC condition.

367

368 3. Results

The impact of the direct anthropogenic drivers is generally confined to the continental shelf and so, for clarity, we focus on the region denoted by the dashed line in Figure 1.

In order to study the integrated response of the ecosystem to changes in forcing, mean values of

biomass (phytoplankton, zooplankton and benthic) and net primary production (netPP, the

373 difference between gross primary production and respiration by phytoplankton, zooplankton and

bacteria) are calculated for each simulation. By using mean values, changes in both the

abundance of organisms and the timing of blooms are accounted for. The total availability of

nutrients in the system is represented by using winter (December to February) means.

377 To study the relative magnitude and direction of change of the ecosystem variables compared to

378 conditions in the recent past, the fractional change (FC) is used, where $FC = \langle EXP \rangle / \langle CNTRL \rangle$

379 -1, <EXP> represents time-averaged conditions in the BASE, GC, WM or A1B simulations and

380 <CNTRL> the time-averaged conditions in the CNTRL simulation. FC typically lies in the range

- -1 to +1: FC = 0 when the future results are the same as for CNTRL; FC = -1 when the future
- variable is zero (e.g. nitrogen is absent); FC = +1 when future values are exactly twice those in
- 383 CNTRL. In circumstances where the CNTRL values are near to zero, FC can become large even

384 for only moderate future increases: in Figures 2-4, the FC is limited to a maximum value of 10. 385 In Figures 2-6, regions where the change between two simulations is small compared with inter-386 annual variability determined by the Kruskal-Wallis test [Kruskal and Wallis, 1952], are masked to grey. The Kruskal-Wallis test uses one-way analysis of variance by ranks to compare the 387 388 medians of two time series and determine if the samples come from the same population. We use 389 time series of annual means at each model point and mask in grey points where the Kruskal-390 Wallis p-value (the probability that a chi-squared distribution is at least as extreme as the 391 Kruskal-Wallis test statistic) exceeds 0.05.

392 **3.1 Ecosystem changes under climate change and direct anthropogenic driver experiments**

393 Fractional changes in near-surface winter-mean nutrients in the BASE, A1B, WM and GC 394 experiments compared to the recent past CNTRL simulation are shown in Figure 2. In the near 395 future (BASE) simulation there are local increases in nitrogen and phosphate on the shelf 396 compared to CNTRL, related to increases in river nutrient input (Table 3). In particular, river 397 volume changes due to higher precipitation in the future time slice increases nutrient outflows in 398 the southern North Sea, English Channel, the Irish Sea and the Celtic Sea. These have a 399 relatively large impact on the nitrogen and phosphate concentrations since nutrient outflows are already comparatively high in these regions, e.g. 19.88 mol N s⁻¹ and 1.48 mol P s⁻¹ for the 400 401 CNTRL simulation in the southern North Sea. In the near future, silicate, recycled from detritus, increases on the shelf. In the open ocean, the surface nutrients decrease in the future simulations, 402 with the impact increasing into the far future (A1B). This is a consequence of increased 403 404 stratification in the future and also a response to reduced nutrients in the IPSL-CM4 boundary data, in turn due to increased stratification and slowed thermohaline circulation in the global 405 model [Steinacher et al., 2010]. In the same version of POLCOMS-ERSEM as used here, Holt et 406 al. [2012] used nutrient boundary data from the present day and boundary data projected for the 407 future to show the impact of the boundary on netPP and winter nitrogen in the far future time 408 slice. They showed that using projected future nutrient boundary data affects the whole model 409 410 domain by reducing the availability of dissolved inorganic nitrogen leading to lower netPP; the 411 Irish and Celtic Seas and the southern Bight of the North Sea are affected less than the central 412 and northern North Sea and the open ocean. By the far future (Figure 2), the open-ocean decrease in nutrients has spread across the shelf, although the Celtic and Irish Seas show no significant 413

change due to larger interannual variability in these regions compared to changes from the open

415 ocean. The increase in river nitrogen in the WM simulation causes a significant increase in

416 nitrogen over most of the shelf compared to CNTRL, while reductions in river nitrogen and

417 phosphate in the GC simulation lead to strong decreases at the coast.

418 The mean annual netPP for CNTRL for the period 1983 to 2000 is shown in Figure 3a.

Comparison with observations [Holt et al., 2012] shows that the model netPP values fall within 419 the observed ranges. Fractional changes between mean netPP in the experiments compared to the 420 CNTRL simulation (Figure 3b) show a range of responses. The projected climate-forced increase 421 422 in the near future BASE simulation is significantly different from the CNTRL values only in the 423 southern North Sea, the Celtic and Irish Seas, the English Channel and west of Scotland, where 424 there in an increase in netPP; and in some isolated patches off-shelf. By the far future, the A1B simulation shows a general decline in netPP compared to the recent past CNTRL simulation; 425 426 regions of increasing netPP in the southern North Sea and Celtic Sea are no longer significantly 427 different from CNTRL as they were in the BASE simulation, although rates of netPP in coastal 428 regions of the English Channel and the Irish and Celtic Seas are still significantly larger than in 429 CNTRL. The decreases in netPP are caused by reductions in the available surface nutrients 430 (Figure 2) in the future open ocean and in the northern North Sea by 2082-2099. In these regions, 431 changes in oceanic nutrients are a first order factor in determining primary production changes 432 [Holt et al., 2012]. On-shelf increases in netPP correspond to regions with increased nutrient availability, particularly phosphate and silicate. Other processes affecting netPP under climate 433 434 change, such as increased nutrient recycling in both the pelagic and benthic systems due to 435 warmer waters, and increases in on-shelf stratification are discussed in detail by Holt et al. 436 [2012].



439 **Figure 2:** Fractional changes in surface winter-mean a) nitrogen, b) phosphate and c) silicate

440 under the BASE, WM, GC and A1B climate and anthropogenic driver experiments compared to

441 CNTRL. A colour version of this figure is available online.



Figure 3: a) mean depth integrated net primary production for CNTRL (1983-2000) and b) the
fractional change in mean net primary production for BASE, A1B, WM and GC compared to
CNTRL. A colour version of this figure is available online.

448 For the near future period, 2030-2040, the changes under the direct driver experiments (WM and 449 GC) are compared with the changes due to climate alone from the BASE experiment. The 450 fractional change in netPP for the WM simulation (Figure 3b) appears the same as for the BASE simulation: the increase in river nitrogen load has little impact on netPP, which must be limited 451 by other factors, such as light, or phosphate or silicate availability that are (almost) identical in 452 453 the BASE and WM simulations. In contrast, the GC simulation shows a reduction in the climateinduced increase in netPP in the southern North Sea, where increases that occur in the BASE 454 455 simulation become not significant or decrease in value under the GC scenario. NetPP calculated by POLCOMS-ERSEM shows lower sensitivity to reductions in river nutrients compared to 456

457 other regional models covering the southern North Sea [Lenhart et al., 2010]. In coastal seas, the 458 ratio of carbon, nitrate and phosphate in phytoplankton can vary significantly: in contrast to most 459 models, which use a fixed carbon to nutrient (C:N) ratio, ERSEM allows the C:N ratio to vary and, through this buffering effect, is less sensitive to changes in nutrient availability [Allen, 460 461 1997]. The decrease in netPP in the southern North Sea for reduced concentrations of phosphate and nitrate in GC (Figure 2) suggest that phosphate and/or nitrate limit primary production in this 462 463 area. Observations in the Southern Bight of the North Sea and the eastern English Channel 464 indicate that silicate or phosphate are the main nutrients potentially limiting phytoplankton growth [Lefebvre et al., 2011; van der Zee and Chou, 2005] under recent climatic conditions. 465

Fractional changes in biomass due to climate and anthropogenic driver changes are shown in 466 467 Figure 4. The benthic biomass is the total of deposit feeders, filter feeders and meiobenthos; 468 phytoplankton and zooplankton biomass are depth integrated and divided into small (flagellates 469 and pico) and large (diatoms and dinoflagellates) phytoplankton and small (micro and heterotrophic nanoflagellates) and large (meso) zooplankton. In the near future, the benthic 470 471 biomass increases in the North Sea, but then decreases across the shelf by the far future time 472 slice. The biomass of small zooplankton and both sizes of phytoplankton tend to decrease, with 473 the signal being stronger in the far future than the near future. However, in the Celtic Sea the 474 biomass of large phytoplankton increases corresponding to regions where the netPP increases. 475 For the large zooplankton in the Norwegian Sea biomass increases, with the change increasing 476 between the near and far future. On the shelf, increases in large zooplankton biomass in the near 477 future (due to changes in silicate) reduce in extent in the far future although still significantly 478 exceed present day values. There is no coherent transfer of the climate change signal between 479 trophic levels, for example, significant increases in netPP in the southern North Sea and Irish Sea 480 in the near future (BASE) do not translate into a significant change in phytoplankton biomass in these regions, although concentrations of detritus do increase (not shown). Also, the biomass of 481 zooplankton changes over large areas of the shelf where there is no significant change in either 482 483 size class of the phytoplankton biomass. In the future, large zooplankton increase at the expense of small zooplankton, either by out-competing for phytoplankton or by consuming the small 484 485 zooplankton.



Figure 4: Fractional changes in a) benthic biomass and the depth integrated biomass of b) small
phytoplankton, c) large phytoplankton, d) small zooplankton and e) large zooplankton under the
BASE, WM, GC and A1B climate and anthropogenic driver experiments compared to CNTRL.

490 A colour version of this figure is available online.

For changes in anthropogenic drivers, increases in river nitrogen concentrations in the WM 492 493 simulation have no effect on biomass. The reduction in river nitrogen and phosphate in the GC 494 simulation leads to decreases in small phytoplankton and small and large zooplankton biomass in 495 coastal regions. The reduction in fishing effort in the North Sea in the GC simulation leads to 496 higher benthic biomass, demonstrating the potential to mitigate the climate-forced decrease in 497 benthic biomass projected for the far future. Comparing the anthropogenic driver simulations GC 498 and WM with the near future climate simulation (BASE) demonstrates the sensitivity of the 499 system to anthropogenic driver changes in the near future (not shown). Consistent with the differences between BASE/CNTRL-1 and GC/CNTRL-1 in Figure 4, reductions in fishing 500 effort and river nutrients in GC compared to BASE lead to an increase in benthic biomass in the 501 502 North Sea and decreases in zooplankton and small phytoplankton biomass near the coast, 503 especially in the southern North Sea. There are no significant changes in the biomass for WM

504 compared to BASE.

505 Given the semi-quantitative nature of these scenarios it is now appropriate to focus on the sign of 506 the potential change.

507 **3.2 Relative impact of climate-induced and direct anthropogenic driver effects**

508 In the near future time period, Figure 4 suggests that the impact of the direct anthropogenic 509 driver changes in river nutrients and benthic trawling effort compared to the impacts of changes in ocean and atmospheric forcing are small (in the case of the WM simulation) or limited to near 510 511 coastal regions and the benthos (in the GC simulation). To consider this in more detail, the model 512 simulations are used to define regions where the climate or direct driver impacts dominate: for 513 the WM simulation, the change in the ecosystem due to climate forcing $CF = \langle BASE \rangle -$ <CNTRL> is compared to the change due to the direct drivers DD = <WM> - <BASE>. For 514 515 regions with |CF| > |DD|, the effect of the climate forcing dominates, and for |CF| < |DD|, the effect of the direct drivers dominates. Where neither <BASE> nor <WM> are significantly 516 517 different from <CNTRL>, measured using the Kruskal-Wallis test [Kruskal and Wallis, 1952],

518 there is no significant change.

519 For the WM simulation, the influence of the increase in river nitrogen load is evident in surface

520 nitrogen (Figure 5a), with the direct driver dominating changes at the coast, in the southern and

western North Sea and in the Norwegian Coastal Current; the climate change signal dominates in the open ocean reflecting changes in ocean boundary forcing and seasonal stratification. For the other parameters, the climate response dominates over the effect of the direct driver impact. Surface phosphate has a similar open ocean response to the climate forcing as surface nitrogen and experiences a climate dominated change in the southern North Sea as a response to increases in river sources of phosphate of 6.7% (Table 3) in that region.

527 A similar analysis for the GC simulation (Figure 5b) shows that reducing the river nitrogen and phosphate levels and trawling fishing effort has more of an impact relative to climate-induced 528 529 changes than increasing river nitrogen levels in the WM simulation. The region of surface 530 nitrogen significantly changed by the change in direct driver is similar to WM but does not 531 extend as far west and north in the North Sea. The region of surface phosphate impacted is more 532 restricted to very near coastal areas, where the netPP, the small phytoplankton biomass and the 533 biomass of both large and small zooplankton are also dominated by changes in the direct driver. This demonstrates that the availability of phosphate in these regions is a stronger control on 534 535 netPP than nitrogen availability, there being sufficient nitrogen in the water column but limited phosphate to support production. The large phytoplankton (including diatoms) are not affected 536 537 by phosphate reductions as they grow early in the season when phosphate is still abundant. 538 Except for off the west coast of Denmark, the reduction in benthic fishing effort in the North Sea 539 has a larger impact on the benthic biomass than changes in the climate.



Figure 5: Regions of climate or anthropogenic driver change dominance on the ecosystem for a)

544 World Market (WM) and b) Global Community (GC) scenarios for 2030-2040.

545 Although the direct anthropogenic drivers effects exceed the climate change impacts only locally (and the benthic biomass for changes in benthic fishing), the influence extends to other areas too. 546 547 In Figure 6 we examine whether the direct anthropogenic drivers act to amplify or mitigate the effects of climate change on the ecosystem in the near future (2030-2040). For the WM 548 549 simulation, for example, the total change $TC = \langle WM \rangle - \langle CNTRL \rangle$ is compared to the change due to climate forcing alone, $CF = \langle BASE \rangle - \langle CNTRL \rangle$. If $TC \rangle CF \rangle 0$, the direct driver acts 550 551 to positively amplify the climate signal and strengthens the increase in values due to climate forcing; if TC < CF < 0, the driver amplifies the climate signal but acts to strengthen the negative 552 553 signal; if CF > 0 and TC < CF or CF < 0 and TC > CF, the direct driver and the climate forced change are acting to change the ecosystem in different directions and the effect of the direct 554 555 driver is to mitigate the effects of changes in climate forcing. Regions where either there is no significant climate signal (<BASE> – <CNTRL>), measured using the Kruskal-Wallis test, or 556 the change induced by the direct driver is very small (< 5% of the change due to climate) are 557 masked in grey. 558

559





Figure 6: Regions of mitigation or amplification of the climate signal on the ecosystem by
anthropogenic drivers for a) World Market (WM) and b) Global Community (GC) scenarios for
2030-2040. Grey regions are where there is no significant affect (see text). A colour version of
this figure is available online.

561

Changes in river flow due to changes in precipitation give a small increase (6.6%, Table 3) in 567 nitrogen from the rivers with outflows in the southern North Sea, which the WM direct driver of 568 increasing river nitrogen acts to amplify. Elsewhere, climate-induced decreases (Figure 2) in 569 570 surface nitrogen from the open ocean are offset by increases from river loads. The main impact 571 of the WM driver is to amplify increases in netPP in the southern North Sea and the western Irish Sea, leading to amplification of the increase in small phytoplankton biomass in coastal areas. The 572 573 increase in large zooplankton biomass is also amplified in the same regions. However, there is no 574 impact on the biomass of large phytoplankton and little impact on small zooplankton. 575 For the GC driver, with reduced river levels of nitrogen and phosphate, the response of surface

576 nitrogen mirrors that of the WM driver: where WM amplifies (>0) the response, GC mitigates

577 and, where WM mitigates the response, GC amplifies (<0). River increases in phosphate from 578 climate change are mitigated by the GC reduction leading to mitigation of the netPP increase in 579 the southern North Sea, the Irish Sea and around UK coasts. The climate-induced increase in the large phytoplankton biomass around the UK coast is slightly amplified by the GC driver whilst 580 581 the increase in small phytoplankton biomass is mitigated. Decreases in small zooplankton 582 biomass are amplified and increases in the large zooplankton biomass are mitigated, with a 583 potential impact on high trophic levels and fisheries. Reduced fishing effort in the North Sea amplifies the increase in benthic biomass experienced under conditions of climate change in the 584 near future. 585

586

587 4. Conclusions

We studied a single realization of possible future climate conditions and investigated the effects 588 589 of additional substantial changes in direct anthropogenic drivers of benthic trawling effort in the North Sea and river nutrient loads. Climate change can potentially drive major ecosystem 590 591 changes in primary production and plankton biomass in both the near (2030-2040) and far (2082-592 2099) future. Direct driver effects dominate climate effects only locally: changes due to river 593 nutrients are confined to coastal regions and the southern North Sea, and changes due to benthic 594 trawling affect the benthic biomass. Direct drivers can mitigate or amplify the effects of climate 595 change on the ecosystem of the European shelf. In a modelling study of climate change and 596 nutrient load reductions for the Baltic Sea, Meier et al. [2012] showed that the relative impacts of 597 climate and nutrient reductions depends on the magnitude of the nutrient reduction: climate 598 change dominates changes in water quality for moderate reductions in nutrients consistent with 599 current legislation, leading to a reduction in water quality, whereas larger nutrient reductions can 600 improve water quality.

The ecosystem response to changes in the climate forcing is not linear, but responds to multiple processes mediating the climate change signal [*Holt et al.*, 2014]: for example, initial increases in netPP in the southern North Sea by 2030-2040 disappear by 2082-2099. However, in general, the changes in plankton and benthic biomass become larger and more widespread in the far future than in the near future. The importance of oceanic effects appears to increase with time as larger reductions in concentrations of oceanic nutrients by the end of the 21st century give rise to
significant changes further on the shelf, particularly into the North Sea. This might potentially
offset the more local effects due to increases in temperature on the shelf, e.g. causing a reduction
in silicate concentrations in the North Sea in the far future.

610 By comparing changes in the average conditions due to changes in climate forcing to direct anthropogenic driver experiments, our results show the potential impacts and limits of 611 government and environmental policies affecting river nutrient loads and benthic trawling effort. 612 In the northeast Atlantic, away from the European continental shelf, such policy measures have 613 little impact on the ecosystem and the climate change effects dominate. In our results, these 614 615 mainly reduce netPP. On the continental shelf, the impact of climate change on netPP and 616 plankton biomass may be mitigated to some extent by environmental policies that reduce river 617 nutrients, particularly in near coastal regions. The mitigation of increases in netPP supports the 618 MSFD GES indicator D5 on minimizing eutrophication. However, such environmental policies 619 could amplify the effects of climate change on the biomass of large phytoplankton and small 620 zooplankton, with a potential detrimental impact on the GES indicator D4 through changing food 621 webs. Policies that allow river nitrogen loads to increase in the absence of any increase in river 622 phosphate loads have little impact on net primary production and phytoplankton and zooplankton 623 biomass, although they will cause a stronger imbalance in the N:P ratio with potential increase in 624 toxicity of harmful algal bloom [*Glibert et al.*, 2014], and negative impacts on biodiversity (D1) and food webs. There is also a small amplifying effect on netPP and the biomass of small 625 626 phytoplankton and large zooplankton in coastal regions. Reducing trawling effort in the North 627 Sea leads to an increase in benthic biomass. Perturbations in the benthic fishing effort and river 628 nutrient loads that are realistic in magnitude potentially alter the response of the ecosystem by more than 5% of the response due to changes in the climate. This has a potential impact on the 629 GES indicator D6 by changing the benthic community structure and increasing biomass. To 630 effectively manage the marine ecosystem under climate change requires a quantitative 631 632 assessment of the combined impacts of climate and anthropogenic changes using improved models (see below) and better knowledge of anthropogenic pressures and their likely 633 634 magnitudes.

635 An assessment of the skill of the recent past (CNTRL) simulation [Holt et al., 2012] shows similar results to a simulation forced by atmospheric reanalysis data. However, it is difficult to 636 637 assess the likelihood or skill of the near and far future simulations. To understand the uncertainty in climate projections needs a large number of simulations forced by different climate models 638 639 spanning a range of responses. For the atmosphere, Christensen et al. [2009] compiled ensembles of high resolution regional climate models for Europe and concluded that the 640 641 minimum number of simulations needed to sample uncertainty is two Regional Climate Models that are forced by two Global Climate Models. We assume that the IPSL-CM4 'business as 642 usual' scenario climate forcing exhibits a possible 'middle of the road' climate response and 643 study average conditions over 2030-2040 and 2082-2099 compared to the recent past period 644 1983-2000. The averaging periods of 11 and 18 years respectively are not sufficient to 645 completely remove the effects of decadal variability from the long-term climate signal but 646 instead give a snapshot of possible average conditions during these periods. Additionally, the 647 \sim 12km resolution of this model does not accurately resolve all the regions where primary 648 production is expected to be high, especially as many of the dominant effects of anthropogenic 649 drivers are constrained to coastal regions. To address these issues requires a series of higher 650 resolution, transient simulations using a range of climate models, which is the subject of a future 651 652 project.

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659 6. References

Allen, J. I. (1997), A modelling study of ecosystem dynamics and nutrient cycling in the Humber plume, UK, *Journal of Sea Research*, *38*(3–4), 333-359. doi: <u>http://dx.doi.org/10.1016/S1385-</u>

662 <u>1101(97)00050-6</u>.

- Allen, J. I., and K. R. Clarke (2007), Effects of demersal trawling on ecosystem functioning in
- the North Sea: a modelling study, *Marine Ecology Progress Series*, 336, 63-75. doi:
- 665 10.3354/meps336063.
- Anderson, D., P. Glibert, and J. Burkholder (2002), Harmful algal blooms and eutrophication:
- 667 Nutrient sources, composition, and consequences, *Estuaries*, 25(4), 704-726. doi:
- 668 10.1007/bf02804901.
- 669 Artioli, Y., J. Friedrich, A. J. Gilbert, A. McQuatters-Gollop, L. D. Mee, J. E. Vermaat, F. Wulff,
- 670 C. Humborg, L. Palmeri, and F. Pollehne (2008), Nutrient budgets for European seas: A measure
- 671 of the effectiveness of nutrient reduction policies, *Marine Pollution Bulletin*, 56(9), 1609-1617.
- 672 doi: <u>http://dx.doi.org/10.1016/j.marpolbul.2008.05.027</u>.
- Aumont, O., E. Maier-Reimer, S. Blain, and P. Monfray (2003), An ecosystem model of the
- global ocean including Fe, Si, P colimitations, *Global Biogeochemical Cycles*, *17*(2), 1060. doi:
 10.1029/2001gb001745.
- Baretta, J. W., W. Ebenhöh, and P. Ruardij (1995), The European regional seas ecosystem
- model, a complex marine ecosystem model., *Netherlands Journal of Sea Research*, *33*, 233–246.
- 678 Blackford, J. C., J. I. Allen, and F. J. Gilbert (2004), Ecosystem dynamics at six contrasting sites:
- a generic modelling study, *Journal of Marine Systems*, 52(1–4), 191-215. doi:
 http://dx.doi.org/10.1016/j.jmarsys.2004.02.004.
- Bopp, L., P. Monfray, O. Aumont, J.-L. Dufresne, H. Le Treut, G. Madec, L. Terray, and J. C.
- 682 Orr (2001), Potential impact of climate change on marine export production, *Global*
- 683 Biogeochemical Cycles, 15(1), 81-99. doi: 10.1029/1999gb001256.
- Boyd, P. W., T. A. Rynearson, E. A. Armstrong, F. Fu, K. Hayashi, Z. Hu, D. A. Hutchins, R. M.
- Kudela, E. Litchman, M. R. Mulholland, U. Passow, R. F. Strzepek, K. A. Whittaker, E. Yu, and
- 686 M. K. Thomas (2013), Marine Phytoplankton Temperature versus Growth Responses from Polar
- to Tropical Waters Outcome of a Scientific Community-Wide Study., *PLoS ONE*, 8(5):
- 688 *e63091*. doi: 10.1371/journal.pone.0063091.
- 689 Cadée, G. C., and J. Hegeman (2002), Phytoplankton in the Marsdiep at the end of the 20th
- 690 century; 30 years monitoring biomass, primary production, and Phaeocystis blooms, *Journal of* 691 *Sea Research*, 48(2), 97-110. doi: <u>http://dx.doi.org/10.1016/S1385-1101(02)00161-2</u>.
- 692 Christensen, J. H., M. Rummukainen, and G. Lenderink (2009), Formulation of very-high-
- 693 resolution regional climate model ensembles for Europe, in: ENSEMBLES: Climate Change and
- 694 its Impacts: Summary of research and results from the ENSEMBLES project. Edited by: van der
- Linden, P. and Mitchell, J. F. B., Met Office Hadley Centre, Exeter, UK, 47–58.
- 696 Chust, G., J. I. Allen, L. Bopp, C. Schrum, J. Holt, K. Tsiaras, M. Zavatarelli, M. Chifflet, H.
- 697 Cannaby, I. Dadou, U. Daewel, S. L. Wakelin, E. Machu, D. Pushpadas, M. Butenschon, Y.
- 698 Artioli, G. Petihakis, C. Smith, V. Garçon, K. Goubanova, B. Le Vu, B. A. Fach, B. Salihoglu, E.

699 Clementi, and X. Irigoien (2014), Biomass changes and trophic amplification of plankton in a 700 warmer ocean, *Global Change Biology*, *20*(7), 2124-2139. doi: 10.1111/gcb.12562.

Dethlefsen, V., and H. Von Westernhagen (1983), Oxygen deficiency and effects on bottom
 fauna in the eastern German Bight 1982., *Meeresforschung*, 30(42-53).

Dounas, C., I. Davies, G. Triantafyllou, P. Koulouri, G. Petihakis, C. Arvanitidis, G. Sourlatzis,
 and A. Eleftheriou (2007), Large-scale impacts of bottom trawling on shelf primary productivity,
 Continental Shelf Research, 27(17), 2198-2210. doi: http://dx.doi.org/10.1016/j.csr.2007.05.006.

Ducrotoy, J.-P., and M. Elliott (2008), The science and management of the North Sea and the Baltic Sea: Natural history, present threats and future challenges, *Marine Pollution Bulletin*,

Baltic Sea: Natural history, present threats and future challenges, *Marine*57(1–5), 8-21. doi: http://dx.doi.org/10.1016/j.marpolbul.2008.04.030.

European Commission (2011), Communication from the Commission to the European

710 Parliament, the Council, the European Economic and Social Committee and the Committee of

the Regions, Reform of the Common Fisheries Policy, Brussels. 13.7.2011 COM (2011) 417

712 final.

European Commission (2014), Commission Delegated Regulation (EU) No 1395/2014 of 20

- 714 October 2014 establishing a discard plan for certain small pelagic fisheries and fisheries for
- 715 industrial purposes in the North Sea.

Flather, R. A. (1981), Results from a model of the northeast Atlantic relating to the Norwegian

717 Coastal Current, in The Norwegian Coastal Current, edited by R. Saetre and M. Mork, pp. 427-

718 458, Bergen University, Bergen, Norway.

Garcia, H. E., R. A. Locarnini, T. P. Boyer, and J. I. Antonov (2006), World Ocean Atlas 2005,

- Volume 4: Nutrients (phosphate, nitrate, silicate). S. Levitus, Ed. NOAA Atlas NESDIS 64, U.S.
- 721 Government Printing Office, Washington, D.C., 396 pp.

722 Glibert, P. M., J. Icarus Allen, Y. Artioli, A. Beusen, L. Bouwman, J. Harle, R. Holmes, and J.

Holt (2014), Vulnerability of coastal ecosystems to changes in harmful algal bloom distribution

in response to climate change: projections based on model analysis, *Global Change Biology*. doi:

- 725 10.1111/gcb.12662.
- Gowen, R. J., P. Tett, K. Kennington, D. K. Mills, T. M. Shammon, B. M. Stewart, N.
- 727 Greenwood, C. Flanagan, M. Devlin, and A. Wither (2008), The Irish Sea: Is it eutrophic?,
- 728 Estuarine, Coastal and Shelf Science, 76(2), 239-254. doi:
- 729 <u>http://dx.doi.org/10.1016/j.ecss.2007.07.005</u>.
- 730 Greenstreet, S., L. Robinson, G. Piet, J. Craeymeersch, R. Callaway, H. Reiss, S. Ehrich, I.
- 731 Kröncke, H. Fraser, J. Lancaster, L. Jorgensen, and A. Goffin (2007), The ecological disturbance

caused by fishing in the North Sea. Fisheries Research Services Collaborative Report No 04/07,

733 169pp.

- Halpern, B. S., S. Walbridge, K. A. Selkoe, C. V. Kappel, F. Micheli, C. D'Agrosa, J. F. Bruno,
- 735 K. S. Casey, C. Ebert, H. E. Fox, R. Fujita, D. Heinemann, H. S. Lenihan, E. M. P. Madin, M. T.
- Perry, E. R. Selig, M. Spalding, R. Steneck, and R. Watson (2008), A Global Map of Human
- 737 Impact on Marine Ecosystems, *Science*, *319*(5865), 948-952. doi: 10.1126/science.1149345.
- Holt, J., M. Butenschön, S. L. Wakelin, Y. Artioli, and J. I. Allen (2012), Oceanic controls on the
- primary production of the northwest European continental shelf: model experiments under recent
- past conditions and a potential future scenario, *Biogeosciences J1 BG*, 9(1), 97-117. doi:
- 741 10.5194/bg-9-97-2012.
- Holt, J., C. Schrum, H. Cannaby, U. Daewel, I. Allen, Y. Artioli, L. Bopp, M. Butenschon, B. A.
- Fach, J. Harle, D. Pushpadas, B. Salihoglu, and S. Wakelin (2014), Physical processes mediating
- climate change impacts on regional sea ecosystems, *Biogeosciences Discuss.*, *11*(2), 1909-1975.
- 745 doi: 10.5194/bgd-11-1909-2014.
- Holt, J. T., and I. D. James (2001), An s coordinate density evolving model of the northwest
- European continental shelf: 1. Model description and density structure, *J. Geophys. Res.*,
- 748 *106*(C7), 14015-14034. doi: 10.1029/2000jc000304.
- Huthnance, J. M. (2010), The Northeast Atlantic margins, in *Carbon and nutrient fluxes in*
- 750 continental margins: a global synthesis, edited by K.-K. Liu, L. Atkinson, R. Quiñones and L.
- 751 Talaue-McManus, pp. 215-234, Springer, Berlin Heidelberg. DOI 10.1007/978-3-540-92735-8.
- ICES (2014), ICES Catch Statistics 2006-2012. Version 17-10-2014. Accessed 17-10-2014 via
 <u>http://ices.dk/marine-data/dataset-collections/Pages/Fish-catch-and-stock-assessment.aspx</u>,
 ICES, Copenhagen.
- 755 IPCC (2001), Climate Change 2001: The Scientific Basis. Contribution of Working Group I to
- the Third Assessment Report of the Intergovernmental Panel on Climate Change [Houghton,
- J.T.,Y. Ding, D.J. Griggs, M. Noguer, P.J. van der Linden, X. Dai, K. Maskell, and C.A. Johnson
 (eds.)]. Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA,
 881pp.
- Kaiser, M. J., K. R. Clarke, H. Hinz, M. C. V. Austen, P. J. Somerfield, and I. Karakassis (2006),
 Global analysis of response and recovery of benthic biota to fishing, *Marine Ecology Progress*
- 762 Series, 311, 1-14. doi: 10.3354/meps311001.
- Key, R. M., A. Kozyr, C. L. Sabine, K. Lee, R. Wanninkhof, J. L. Bullister, R. A. Feely, F. J.
- 764 Millero, C. Mordy, and T. H. Peng (2004), A global ocean carbon climatology: Results from
- Global Data Analysis Project (GLODAP), *Global Biogeochemical Cycles*, *18*(4), GB4031. doi:
 10.1029/2004gb002247.
- 767 Kruskal, W. H., and W. A. Wallis (1952), Use of ranks in one-criterion variance analysis,
- 768 Journal of the American Statistical Association, 47(260), 583-621. doi:
- 769 doi:10.1080/01621459.1952.10483441.

- 770 Lacroix, G., K. Ruddick, N. Gypens, and C. Lancelot (2007), Modelling the relative impact of
- rivers (Scheldt/Rhine/Seine) and Western Channel waters on the nutrient and
- diatoms/Phaeocystis distributions in Belgian waters (Southern North Sea), Continental Shelf
- 773 *Research*, 27(10–11), 1422-1446. doi: <u>http://dx.doi.org/10.1016/j.csr.2007.01.013</u>.
- Lancelot, C., N. Gypens, G. Billen, J. Garnier, and V. Roubeix (2007), Testing an integrated
- river–ocean mathematical tool for linking marine eutrophication to land use: The Phaeocystis-
- dominated Belgian coastal zone (Southern North Sea) over the past 50 years, *Journal of Marine*
- 777 Systems, 64(1–4), 216-228. doi: <u>http://dx.doi.org/10.1016/j.jmarsys.2006.03.010</u>.
- ⁷⁷⁸ Lefebvre, A., N. Guiselin, F. Barbet, and F. L. Artigas (2011), Long-term hydrological and
- phytoplankton monitoring (1992–2007) of three potentially eutrophic systems in the eastern
- 780 English Channel and the Southern Bight of the North Sea, *ICES Journal of Marine Science:*
- 781 *Journal du Conseil*, *68*(10), 2029-2043. doi: 10.1093/icesjms/fsr149.
- Lenhart, H.-J., D. K. Mills, H. Baretta-Bekker, S. M. van Leeuwen, J. v. der Molen, J. W.
- 783 Baretta, M. Blaas, X. Desmit, W. Kühn, G. Lacroix, H. J. Los, A. Ménesguen, R. Neves, R.
- 784 Proctor, P. Ruardij, M. D. Skogen, A. Vanhoutte-Brunier, M. T. Villars, and S. L. Wakelin
- 785 (2010), Predicting the consequences of nutrient reduction on the eutrophication status of the
- North Sea, Journal of Marine Systems, 81(1–2), 148-170. doi:
- 787 <u>http://dx.doi.org/10.1016/j.jmarsys.2009.12.014</u>.
- 788 Marti, O., P. Braconnot, J. Bellier, R. Benshila, S. Bony, P. Brockmann, P. Cadule, A. Caubel, S.
- 789 Denvil, J.-L. Dufresne, L. Fairhead, M.-A. Filiberti, M.-A. Foujols, T. Fichefet, P. Friedlingstein,
- H. Gosse, J.-Y. Grandpeix, F. Hourdin, G. Krinner, C. Lévy, G. Madec, I. Musat, N. de Noblet,
- J. Polcher, and C. Talandier (2005), *The new IPSL climate system model: IPSL-CM4.*, 84 pp.,
- 792 Note du Pole de Modélisation, Institut Pierre-Simon Laplace (IPSL), France.
- Meehl, G. A., T. F. Stocker, W. D. Collins, P. Friedlingstein, A. T. Gaye, J. M. Gregory, A.
- Kitoh, R. Knutti, J. M. Murphy, A. Noda, S. C. B. Raper, I. G. Watterson, A. J. Weaver, and Z.-
- C. Zhao (2007), Global Climate Projections. In: Climate Change 2007: The Physical Science
- 796 Basis. Contribution of Working Group I to the Fourth Assessment Report of the
- 797 Intergovernmental Panel on Climate Change [Solomon, S., D. Qin, M. Manning, Z. Chen, M.
- 798 Marquis, K.B. Averyt, M. Tignor and H.L. Miller (eds.)]. Cambridge University Press,
- 799 Cambridge, United Kingdom and New York, NY, USA.
- Meier, H. E. M., R. Hordoir, H. C. Andersson, C. Dieterich, K. Eilola, B. G. Gustafsson, A.
- Höglund, and S. Schimanke (2012), Modeling the combined impact of changing climate and
- changing nutrient loads on the Baltic Sea environment in an ensemble of transient simulations
- 603 for 1961–2099, *Clim Dyn*, *39*(9-10), 2421-2441. doi: 10.1007/s00382-012-1339-7.
- 804 Nakicenovic, N., and N. Swart (Eds.) (2000), Special Report on Emissions Scenarios. A Special
- 805 *Report of Working Group III of the Intergovernmental Panel on Climate Change.*, 599 pp.,
- 806 Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA.
- 807 OSPAR (1988), PARCOM Recommendation 88/2 on the Reduction in Inputs of Nutrients to the
 808 Paris Convention Area.<u>www.ospar.org</u>.

- 809 OSPAR (1989), PARCOM Recommendation 89/4 on a Coordinated Programme for the
- 810 Reduction of Nutrients.<u>www.ospar.org</u>.
- 811 OSPAR (1992), PARCOM Recommendation 92/7 on the Reduction of Nutrient Inputs from
- 812 Agriculture into Areas Where these Inputs are Likely, Directly or Indirectly, to Cause Pollution.
- 813 <u>www.ospar.org</u>.
- 814 OSPAR (2003), Integrated Report 2003 on the Eutrophication Status of the OSPAR Maritime
- 815 Area Based Upon the First Application of the Comprehensive Procedure, OSPAR Commission,
- 816 Eutrophication Series, Publication nr. 189, ISBN 1-904426-25- 5, download available at
- 817 <u>www.ospar.org</u>.
- Pätsch, J., and H. J. Lenhart (2004), Daily Loads of Nutrients, Total Alkalinity, Dissolved
- 819 Inorganic Carbon and Dissolved Organic Carbon of the European continental rivers for the years
- 820 1977–2002., Berichte aus dem Zentrum fur Meeres- und Klimaforschung; Reihe B:
- 821 *Oxeanographie.*, *48*, 159pp.
- Rijnsdorp, A., A. Buys, F. Storbeck, and E. Visser (1998), Micro-scale distribution of beam
- trawl effort in the southern North Sea between 1993 and 1996 in relation to the trawling
- frequency of the sea bed and the impact on benthic organisms, *ICES Journal of Marine Science:*
- 825 *Journal du Conseil*, *55*(3), 403-419.
- 826 Sarmiento, J. L., R. Slater, R. Barber, L. Bopp, S. C. Doney, A. C. Hirst, J. Kleypas, R. Matear,
- U. Mikolajewicz, P. Monfray, V. Soldatov, S. A. Spall, and R. Stouffer (2004), Response of
- ocean ecosystems to climate warming, *Global Biogeochemical Cycles*, *18*(3), GB3003. doi:
- 829 10.1029/2003gb002134.
- 830 Schneider, B., L. Bopp, M. Gehlen, J. Segschneider, T. L. Frölicher, P. Cadule, P. Friedlingstein,
- 831 S. C. Doney, M. J. Behrenfeld, and F. Joos (2008), Climate-induced interannual variability of
- marine primary and export production in three global coupled climate carbon cycle models,
- 833 *Biogeosciences*, 5(2), 597-614. doi: 10.5194/bg-5-597-2008.
- 834 Skogen, M. D., and L. R. Mathisen (2009), Long-term effects of reduced nutrient inputs to the
- North Sea, *Estuarine, Coastal and Shelf Science*, 82(3), 433-442. doi:
- 836 <u>http://dx.doi.org/10.1016/j.ecss.2009.02.006</u>.
- 837 Skogen, M. D., K. Eilola, J. L. S. Hansen, H. E. M. Meier, M. S. Molchanov, and V. A.
- 838 Ryabchenko (2014), Eutrophication status of the North Sea, Skagerrak, Kattegat and the Baltic
- 839 Sea in present and future climates: A model study, *Journal of Marine Systems*, *132*(0), 174-184.
- 840 doi: <u>http://dx.doi.org/10.1016/j.jmarsys.2014.02.004</u>.
- 841 Smyth, T. J., G. F. Moore, T. Hirata, and J. Aiken (2006), Semianalytical model for the
- 842 derivation of ocean color inherent optical properties: description, implementation, and
- 843 performance assessment, *Applied Optics*, *45*(31), 8116-8131. doi: 10.1364/AO.45.008116.

- Song, Y., and D. Haidvogel (1994), A Semi-implicit Ocean Circulation Model Using a
- 845 Generalized Topography-Following Coordinate System, Journal of Computational Physics,
- 846 *115*(1), 228-244. doi: <u>http://dx.doi.org/10.1006/jcph.1994.1189</u>.
- 847 Steinacher, M., F. Joos, T. L. Frölicher, L. Bopp, P. Cadule, V. Cocco, S. C. Doney, M. Gehlen,
- K. Lindsay, J. K. Moore, B. Schneider, and J. Segschneider (2010), Projected 21st century
- decrease in marine productivity: a multi-model analysis, *Biogeosciences*, 7(3), 979-1005. doi:
- 850 10.5194/bg-7-979-2010.
- Thomas, H., Y. Bozec, H. de Baar, K. Elkalay, M. Frankignoulle, W. Kühn, H. Lenhart, A. Moll,
- J. Pätsch, G. Radach, L.-S. Chiettecatte, and A. Borges (2010), Carbon and nutrient budgets of
- the North Sea, in *Carbon and nutrient fluxes in continental margins*, edited by K.-K. Liu, L.
- Atkinson, R. Quiñones and L. Talaue-McManus, pp. 346-355, Springer, Berlin Heidelberg. DOI
- 855 10.1007/978-3-540-92735-8.
- van der Molen, J., J. Aldridge, C. Coughlan, E. Parker, D. Stephens, and P. Ruardij (2013),
- 857 Modelling marine ecosystem response to climate change and trawling in the North Sea,
- 858 Biogeochemistry, 113(1-3), 213-236. doi: 10.1007/s10533-012-9763-7.
- van der Zee, C., and L. Chou (2005), Seasonal cycling of phosphorus in the Southern Bight of
 the North Sea, *Biogeosciences*, 2(1), 27-42. doi: 10.5194/bg-2-27-2005.
- Vermaat, J. E., A. McQuatters-Gollop, M. A. Eleveld, and A. J. Gilbert (2008), Past, present and
- future nutrient loads of the North Sea: Causes and consequences, *Estuarine, Coastal and Shelf*
- 863 Science, 80(1), 53-59. doi: <u>http://dx.doi.org/10.1016/j.ecss.2008.07.005</u>.
- Vörösmarty, C. J., B. M. Fekete, M. Meybeck, and R. B. Lammers (2000), Global system of
- rivers: Its role in organizing continental land mass and defining land-to-ocean linkages, *Global*
- 866 Biogeochemical Cycles, 14(2), 599-621. doi: 10.1029/1999gb900092.
- 867 Wakelin, S. L., J. T. Holt, J. C. Blackford, J. I. Allen, M. Butenschön, and Y. Artioli (2012),
- 868 Modeling the carbon fluxes of the northwest European continental shelf: Validation and budgets,
- 869 J. Geophys. Res., 117(C5), C05020. doi: 10.1029/2011JC007402.
- 870 Young, E. F., and J. T. Holt (2007), Prediction and analysis of long-term variability of
- temperature and salinity in the Irish Sea, J. Geophys. Res., 112(C1), C01008. doi:
- 872 10.1029/2005JC003386.
- 873
- 874